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Cadre structurel, déformations et exhumation des Schistes du Santa Marta : accumulation et histoire de déformation d'un terrain caraïbe au nord de la Sierra Nevada de Santa Marta

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THÈSE

Pour obtenir le grade de

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préparée au sein des **Laboratoires ISTERre et Departament des
Geosciences**

dans les **Écoles Doctorales de l'Université Grenoble Alpes et
Universidad Nacional de Colombia**

CADRE STRUCTUREL, DÉFORMATIONS ET EXHUMATION DES SCHISTES DU SANTA MARTA: ACCUMULATION, ET HISTOIRE DE DÉFORMATION D'UN TERRAIN CARAÏBE AU NORD DE LA SIERRA NEVADA DE SANTA MARTA

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Schistes du Santa Marta: accumulation,
et histoire de déformation d'un terrain Caraïbe au nord
de la Sierra Nevada de Santa Marta*

*Marco estructural deformaciones y exhumación de los
Esquistos de Santa Marta : la acreción e historia de
deformación de un terreno Caribeño al norte de la Sierra
Nevada de Santa Marta*

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Año 2016

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Nevada de Santa Marta***

***Structural Framework, deformation and exhumation of
the Santa Marta Schists: accretion and deformational
history of a Caribbean Terrane at the north of the Sierra
Nevada de Santa Marta***

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Tesis presentada como requisito parcial para optar al título Ph. D.
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Thèse

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pour l'obtention du grade de

Docteur ès Sciences
de l'Université Grenoble Alpes

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Géologie structurale, Tectonique et Geodynamique
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Grenoble, France

2017

*To see a world in a grain of sand and a
heaven in a wild flower, hold infinity in
the palm of your hand and eternity in an
hour.*

William Blake

*Je laisse Sisyphe au bas de la montagne !
On retrouve toujours son fardeau. Mais
Sisyphe enseigne la fidélité supérieure
qui nie les dieux et soulève les rochers.
Lui aussi juge que tout est bien. Cet
univers désormais sans maître ne lui
paraît ni stérile ni futile. Chacun des
grains de cette pierre, chaque éclat
minéral de cette montagne pleine de
nuit, à lui seul, forme un monde. La lutte
elle-même vers les sommets suffit à
remplir un coeur d'homme. Il faut
imaginer Sisyphe heureux.*

Le Mythe du Sisyphe
Albert Camus

Thesis Abstract

The Sierra Nevada de Santa Marta (SNSM) is perhaps the most complex crustal massif in the Northern Andes. Its unique situation as an isolated triangular massif segmented from the continuity of the 7000 km long Andes as the last standing mountain before the domains of the younger Caribbean plate, places the SNSM as an island separated from all surrounding mountain ranges of the continental margin. A prominent relief characterizes this mountain reaching the highest altitude in the entire Caribbean realm at 5750 m, and defines, the SNSM as the highest coastal mountain range in the world. For this reason, the SNSM is a unique geological feature that embraces an outstanding biodiversity from its coral reefs in the Caribbean Sea passing through heavily vegetated tropical rainforests, high cloud forests, and moorlands, until its magnificent summit capped by glaciers. Such an extraordinary place is inhabited by humans since at least 6000 years ago. These people are: the Arhuaco, Wiwa, Kogi and Kankuamo had called the Sierra Nevada de Santa Marta, as “The Heart” of the World: *Gonawindúa*, “This ancestral name refers in the ancient inhabitants cosmogony to the most important peak in the Sierra Nevada de Santa Marta, which in the moment that the spiritual world transcended into the material world was the first element that emerged for remind us of our tasks, is the subsistence of everything, *Gonawindúa* is where our sense of responsibility, our physical and material sense, intelligence, wisdom and strength are codified and kept.”

Beyond its historical and mystic background, the SNSM hides an equally fascinating geological history as well. One of the most interesting topics to address when studying the evolution of this massif resides in the correlations between the Northern Andes, Central American and Mexican blocks because the SNSM not only possesses a complex geometry, but a prolific assembly of igneous and metamorphic rocks with ages that range from the Mesoproterozoic to the Eocene, and collateral sediments and volcanoclastic deposits from Devonian to Miocene times. By its position on the northwestern margin of South America the study of the SNSM provides the opportunity to resolve important questions on the evolution of super-continental cycles since Grenvillian times through the Neoproterozoic Pan-African orogeny, the late Paleozoic Ouachitan-Appalachian orogeny that led to Pangæa assembly, and Triassic Pangæa break-up followed by the Jurassic Central Atlantic Rift and more recently by the start of the Caribbean plate accretion/subduction since the Late Cretaceous against northwestern South America, which was one of the most important factors operating during the Andean orogeny. In the SNSM, rocks from all the previously described Wilson cycles are preserved. During these vast periods, the

SNSM massif occupied the outboard position in an accretionary margin of NW Gondwana and in this position accumulated remnants of the collisional and orogenic episodes, which are preserved as low to high metamorphic grade crustal slivers, separated by sutures.

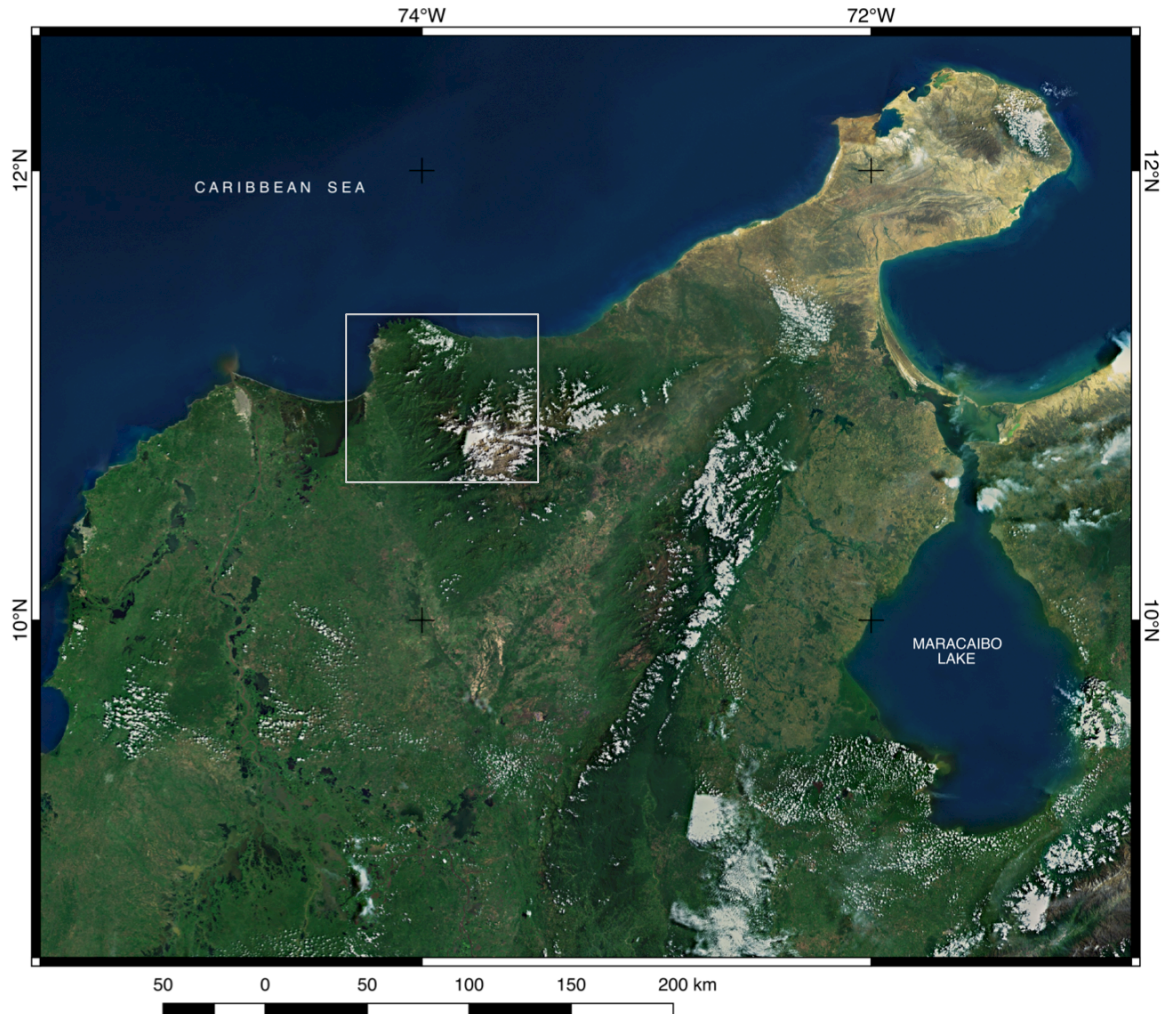


Figure 1. Satellite imagery from the Northern Andes of Colombia, the center of the image focuses on the triangular massif of the Sierra Nevada de Santa Marta. The gray box indicates the studied area. Taken from http://server.arcgisonline.com/arcgis/rest/services/ESRI_Imagery_World_2D/MapServer.

Pioneer expeditions shed light on the geology of the Sierra Nevada de Santa Marta. The classic work of Gansser, 1955, advanced in a first reconnaissance of the rocks that integrate the massif and produced a first geological map accompanied by petrological and structural observations. About 20 years later MacDonald et al., 1971; Macdonald and Hurley, 1969; Tschanz et al., 1974, 1969 updated a first full inventory of the geology of the Sierra Nevada de Santa Marta, even performing radiometric K/Ar dating

that permitted for the first time to investigate thermal processes that affected the rocks. Recently several studies have attempted to reconstruct the geological history of this massif, relying on the foundations provided by Tschanz et al., 1969, these works had outlined the palaeogeographic implications of the SNSM through different epochs and gradually augmented the resolution on mapping, geochemistry and isotopic geology (Bayona et al., 2010; Cardona et al., 2010b, 2010c, 2011, 2011a; Cordani et al., 2005; Doolan, 1970; Duque, 2009; Ordóñez et al., 2002; Restrepo-Pace et al., 1997; Villagómez et al., 2011b; Zuluaga and Stowell, 2012), however, most of the SNSM remains unknown for geologists.

In this investigation I attempt to unravel the geological history of the Sierra Nevada de Santa Marta Massif using the most advanced geochronological, thermochronological geochemical and isotopic techniques that allowed gathering a significant amount of new data to add to the existent database on the SNSM. Although given the intrinsically complex geologic history these analyses were only conducted after several months of fieldwork, in which a detailed geological and metamorphic map was constructed and sample collection was performed following the procedures and techniques in modern structural geology (Fossen, 2010; Ramsay and Huber, 1987).

Our results include a reevaluated geological map 1:25000, in which I define 4 new stratigraphic units, accompanied by two crustal-scale cross sections of 320 km length that dissect the massif, and 8 parallel cross sections at the NW corner of the SNSM metamorphic belt. The geochemical and isotopic dataset includes: i) 17 igneous and metamorphic rocks and 6 detrital samples dated by laser-ablation induced-coupled-plasma mass-spectrometry (LA-ICP-MS), U-Pb zircon geochronology that resulted in 2790 new dates and in-situ trace element analyses, ii) 16 igneous and metamorphic rocks that yielded 31 new thermochronometric ages as follows: 12 zircon fission track ages, 11 Apatite fission track ages and 7 (U-Th)/He in apatite ages, iii) Whole rock geochemistry from 10 samples and iv) Microprobe mineral chemistry in spot analyses and x-ray maps from 4 samples that yielded zoned and peritectic garnet. These data were acquired from the units of the northwestern metamorphic suite of the SNSM massif. With these data we investigated i) The units that conform the SNSM metamorphic belts, their chronological and stratigraphic relationships from the Precambrian to the Eocene; ii) The time span and P-T conditions of a late Paleozoic-early Mesozoic metamorphic event (Chapter 1), iii) The timing of igneous activity accretion and exhumation of oceanic and continental terranes during the Late Cretaceous to late Miocene. iv) A mechanism for explaining how this exhumation occurred under a collisional regime by a climate influenced process at elevated erosion and thermal gradients (Chapter 2); v) The late processes of denudation and sedimentation controlled by tectonics in two marginal basins since the early Miocene under decreased erosion rates and thermal gradients (Chapter 3).

Chapter 1. Convergence between Laurentia and Gondwana eventually led to collision and formation of Pangæa during the late Paleozoic. This long lasting plate interaction generated HT-HP rocks, which characterize a large part of the northwest portion of the Sierra Nevada de Santa Marta Massif (SNSM) in the Northern Andes of Colombia, where a tilted metamorphic lower crustal section is preserved. Field and petrological observations coupled with zircon U-Pb LA-ICP-MS geochronologic data from 17 new samples of igneous and metamorphic rocks together with whole rock and trace element geochemistry are used for reconstructing the tectonic evolution of this strongly remobilized crustal section juxtaposed onto low-grade metamorphic sequences during the late Paleozoic. The metamorphic evolution was investigated by means of several geothermobarometers (Grt+Bt, Hbl+Grt, Hbl+Pl+Qtz, GASP and Ti-in-Zr) for obtaining reliable estimates on *P-T* conditions during at least two metamorphic phases identified by their mineral paragenesis. Our results provide new constraints for palaeogeographic reconstructions between Laurentia and Gondwana during middle Permian to Early Jurassic times. This reconstruction is defined by the interaction of the NW margin of Gondwana with Central American and Mexican terranes during Pangæa

amalgamation. An evolutionary scenario for this supercontinental cycle is proposed in several phases defined by arc inception, timing of metamorphism, and late exhumation of lower crustal segments. The closure of the Rheic Ocean (ca. 290 Ma) marks the start of interaction of crustal blocks during Pangæa assembly that caused 1) post-collisional anatectic melts at ca. 278 ± 0.5 Ma and maximum crustal thickening, 2) Barrovian type metamorphism peak *P-T* conditions (13 kb, 840°C) at ≈ 254 Ma followed by 3) slab retreat and delamination, 4) orogenic collapse, 5) thermal weakening, crustal extension and underplating of mafic crust that caused *LP-HT* metamorphism (5.6 kb, 550°C) at ≈ 225 Ma marking a rifting phase. These processes were accompanied by the infilling of a back-arc basin that accumulated at least 2000 m of carbonate and volcanoclastic sediments provided by the surrounding basement highs. By ~ 216 Ma most of the conjugated margin was removed by drifting, but remnants of this Permian arc are still preserved in Oaxaquia and Acatlán complexes and the Maya block. The crustal attenuation responsible for rifting and drift contributed to an eventual isostatic rebound of the subducted Pacific slab that reactivated arc activity during the Jurassic. By this time, the arc migrated into a continent-ward position, and the former back-arc basin

turned into a fore-arc depocenter accumulating at least an additional burden of ≈ 5000 m of volcanoclastic deposits. Our new reconstruction defines an age span of 60 Myr from the Alleghenides collision to Pangaea break-up at this portion of the Northern Andes.

Chapter 2. New low-temperature thermochronological data from the Precambrian, Permian-to Jurassic, Late Cretaceous and Paleocene-Eocene rocks that make up the northern metamorphic belt of the Sierra Nevada de Santa Marta (SNSM) of the Northern Andes of Colombia reveal exhumation trends in this part of the continental margin during Paleocene to Miocene times. Our analysis involves apatite and zircon fission tracks, as well as apatite (U-Th)/He dating, which allowed to constrain the thermal histories (~ 240 °C – ~ 60 °C) of an inverted crustal sequence that exposes greenschist to granulite facies metamorphic rocks. Temperature-time paths evidence that cooling occurred diachronously in the massif through the activation of NW verging thrusts with the onset of rapid erosional exhumation after collision and TTG series pluton emplacement under an initially high geothermal gradient of 40°C/km. In this regime, the massif exhumed at elevated rates ~ 0.9 km/Myr during 50-45 Ma. These rates incremented to 2 km/Myr for the interval 45-40 Ma with a decreasing geothermal gradient of 30°C/km, reaching

the highest rates of 2.7 km/Myr during 37-35 Myr. We interpret the exhumation of this inverted lower crustal sequence because of coupling of ductile extrusion of a thickened continental crust and surface denudation. Post-Eocene exhumation events show subsequent cooling under a decreased exhumation rates of ~ 0.2 km/Myr at ca. 25 Ma, related to the start of a composite transpressive-transtensive regime in the southern Caribbean plate boundary as a result of changes in convergence between the Caribbean and South American plates. This reorganization led to the activation of the NW trending Santa Marta - Bucaramanga (SMBF) and E-W trending Oca (OF) faults. These faults contributed to the partitioning of a subsequent oblique convergence and transferred strain from the Caribbean Plate into the Santander Massif, the Eastern Cordillera of Colombia, the Perijá Range, and the Merida Andes until the late Miocene. We interpret that this new tectonic regime inhibited further NW thrust propagation and caused a NE tilting of the SNSM massif due to extensional activity of the SMBF coupled with the tectonic drag caused by the transit of the Caribbean plate. These processes resulted in attenuated exhumation rates of ~ 0.09 km/Myr that persisted until the late Miocene during the SNSM thermal re-equilibration.

Chapter 3. Geo-thermochronologic data from sedimentary rocks, coupled with

stratigraphic analyses, unravel an early Neogene exhumation of the Sierra Nevada de Santa Marta massif (SNSM). This fault-bounded triangular block in the northern Andes of Colombia, exposes Precambrian to Paleozoic metamorphic rocks and suites of a Triassic-Jurassic magmatic arc. We compare the Neogene basin fills of two marginal basins, a western one located along the Santa Marta-Bucaramanga fault (SMBF) and the northern one occupying the southern margin of the Oca fault. The western sequence consists of gravitational conglomeratic deposits \approx 1200 m that define a progradational Gilbert-type delta emanating from a scarp coinciding with the SMBF. By its stratigraphic and structural position this sequence compares to the northern transgressive sequence. Provenance analyses

of representative sections of these two sequences, evidence a sourcing from the underlying basement. This finding is in concordance with zircon U-Pb age spectra that lack pre-Grenvillian and Ordovician signals of coeval sediments of the Magdalena basin, demonstrating, that they were deposited in a local basins disconnected to the trunk system of the proto-Magdalena river. In order to reconcile the different organization of the two marginal basins we propose a regional-scale tilting of the SNSM towards the northeast during the early Neogene. This model takes account of the coeval progradational sedimentary cycle at the fault-bounded western flank and the transgressive onlapping of the northern conglomerates on a tilted surface during an overall exhumation of the massif.

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Resumen

La Sierra Nevada de Santa Marta (SNSM) es quizás el macizo cortical más complejo encontrado en los Andes del Norte. Su situación es única pues se trata de un macizo triangular aislado, segmentado de la continuidad de 7000 km de longitud de los Andes como la última montaña antes de los dominios de la placa Caribe más joven. Estas circunstancias particulares ubican a la SNSM como una isla separada de todas las cadenas montañosas circundantes del margen continental. Un relieve prominente caracteriza a esta montaña alcanzando la mayor altitud del Caribe a 5750 msnm y define la SNSM como la mayor elevación costera del mundo. Por esta razón es un elemento geológico único que involucra una excepcional biodiversidad desde sus arrecifes de coral en el Mar Caribe pasando a través de selvas tropicales lluviosas, bosques altos de niebla, y páramos, hasta su magnífica cumbre cubierta de nieve. Este lugar extraordinario ha sido habitado por seres humanos desde hace al menos 6000 años, estos pobladores los: Arhuacos, Wiwa, Kogi y Kankuamo han llamado a la Sierra Nevada de Santa Marta, “el corazón” del mundo: *Gonawindúa*, “Este nombre ancestral se refiere en su cosmogonía al pico más importante de la Sierra Nevada de Santa Marta, el cual en el momento que el mundo espiritual trascendió hacia el mundo material fue el primer elemento que emergió para recordarnos de nuestras tareas, es el sostén de todo, Gonawindúa es donde nuestro sentido de responsabilidad, nuestro sentido físico y material, nuestra inteligencia, sabiduría y fuerza son codificadas y mantenidas.

Más allá de su trasfondo histórico y místico, la Sierra Nevada de Santa Marta también oculta una historia geológica igualmente fascinante. Uno de los temas más interesantes a estudiar durante la evolución de este macizo, reside en las correlaciones existentes con los Andes del Norte, y los bloques de Centro América y México, debido a que la SNSM no solo posee una geometría compleja, pero también un prolífico ensamblaje de rocas ígneas y metamórficas con edades que varían desde el Mesoproterozoico hasta el Eoceno, así como series sedimentarias y depósitos volcánico-clásticos desde el Devónico hasta el Mioceno Superior. Por su posición en el margen noroccidental de Suramérica el estudio de la SNSM nos da la oportunidad de resolver importantes preguntas acerca de la evolución de los ciclos súper-continenciales desde la época del orógeno Grenvillian, a través de la orogenia Neoproterozoica Brasileña/Pan Africana, la orogenia Ouachita-Appalachia del Paleozoico tardío que culminó con la aglutinación de Pangaea, y la posterior separación de Pangaea seguida por el *rift* Jurásico que marco la apertura del Océano Atlántico central, y más recientemente durante el Cretácico Superior el inicio de la

interacción de la placa del Caribe y el Noroccidente de Suramérica, el cual fue uno de los factores mas importantes operando durante la orogenia Andina.

En la Sierra Nevada de Santa Marta rocas pertenecientes a todos los ciclos de Wilson descritos previamente se encuentran preservadas. Durante estos vastos periodos de tiempo el macizo de la SNSM ocupó una posición frontal en un margen acrecional del NW de Gondwana y en esta posición acumuló remanentes de episodios orogénicos y colisionales, los cuales se encuentran preservados como, escamas corticales de rocas metamórficas de bajo a alto grado, separadas por suturas.

Expediciones pioneras dieron las primeras observaciones de la geología de la Sierra Nevada de Santa Marta, el trabajo clásico de Gansser, 1955, avanzó en el primer reconocimiento sistemático de las rocas que integran el macizo y produjo el primer mapa geológico acompañado de observaciones petrológicas y estructurales. Cerca de 20 años mas tarde MacDonald et al., 1971, MacDonald and Hurley, 1969; Tschanz et al., 1974, 1969 realizaron una actualización completa de la geología de la Sierra Nevada de Santa Marta, e incluso llevaron a cabo dataciones K/Ar lo cual permitió por primera vez investigar los procesos termales que afectaron a estas rocas. Recientemente varios estudios han intentado realizar una reconstrucción la historia geológica del macizo de la SNSM, apoyándose en las bases cimentadas por Tschanz et al., 1969, estos trabajos han delineado las implicaciones paleogeograficas de la SNSM a través de diferentes épocas y día a día han aumentado la resolución en cartografía, geoquímica y geología isotópica (Bayona et al., 2010; Cardona et al., 2011, 2010b, 2010c, 2011a; Cordani et al., 2005; Doolan, 1970; Duque, 2009; Ordóñez et al., 2002; Restrepo-Pace et al., 1997; Villagómez et al., 2011b; Zuluaga and Stowell, 2012). Sin embargo la mayor parte de la SNSM continua siendo un territorio desconocido para los geólogos y otros visitantes extranjeros, y durante mi experiencia trabajando en el área, no fue posible alcanzar varias locaciones enigmáticas que deberán esperar por futuras exploraciones.

En esta investigación he intentado revelar la historia geológica de la Sierra Nevada de Santa Marta utilizando las técnicas isotópicas, geocronológicas, termocronológicas y geoquímicas mas avanzadas, las cuales permitieron adquirir una cantidad significativa de nuevos datos que hemos añadido a la base de datos existente en la SNSM. Aunque dada la historia geológica intrínsecamente compleja estos análisis fueron conducidos únicamente después de varios meses de trabajo de campo, durante los cuales un nuevo mapa geológico y metamórfico fue construido y la colección de muestras fue realizada siguiendo los procedimientos y técnicas modernas en geología estructural (Fossen, 2010; Ramsay and Huber, 1987).

Nuestros resultados incluyen un mapa geológico reevaluado a escala 1:25000, el cual incluye la definición de cuatro nuevas unidades estratigráficas acompañado por dos secciones estructurales a escala cortical de 320 km de longitud que atraviesan el macizo y 8 secciones seriadas en la esquina NW del cinturón metamórfico de la SNSM. la base de datos geoquímica e isotópica incluye: i) 17 muestras de rocas ígneas y metamórficas y 6 muestras detríticas datadas por geocronología de U-Pb en zircones mediante espectrometría de masa por plasma inducido por ablación con láser (LA-ICP-MS), cuyos resultados suman 2790 edades nuevas y análisis de elementos traza in-situ, ii) 16 rocas ígneas y metamórficas que resultaron en 31 nuevas edades termocronológicas de la siguiente manera: 12 edades de trazas de fisión en zircones, 11 edades de trazas de fisión en apatitos y 7 edades (U-Th)/He en apatitos, iii) Geoquímica de roca total se realizó en 10 muestras y finalmente iv) Análisis de la química mineral en granates zonados y peritéticos en 4 secciones delgadas se realizaron utilizando la microsonda electrónica. Estos datos fueron adquiridos de las unidades de la serie metamórfica del cinturón noroccidental de la SNSM, y de rocas sedimentarias en cuencas adyacentes al macizo. Con estos datos hemos investigado: i) Las unidades que conforman los cinturones metamórficos de la SNSM, y su relaciones cronológicas y estratigráficas desde el Precámbrico hasta el Eoceno, ii) el intervalo de tiempo y condiciones P-T de un evento metamórfico Paleozoico tardío a Mesozoico temprano (Capítulo 1), iii) la temporalidad de actividad ígnea, acreción y exhumación de terrenos oceánicos y continentales durante el Cretácico Tardío hasta el Mioceno Tardío, iv) un mecanismo para explicar como esta exhumación ocurrió durante un régimen colisional debido a procesos influenciados por el clima a elevados gradientes termales y tasas de erosión (Capítulo 2), v) los procesos tardíos de denudación y sedimentación controlados por tectónica en dos cuencas marginales desde el Mioceno Inferior bajo gradientes geotérmicos y tasas de erosión decrecientes (Capítulo 3).

Capítulo 1. La convergencia entre Laurentia y Gondwana eventualmente culminó con la colisión y posterior aglutinación de Pangæa durante el Paleozoico Tardío. Durante la temporalmente extensa interacción de estas placas tectónicas, rocas de alta presión y alta temperatura fueron originadas (HP-HT), dichas rocas caracterizan una gran parte de la porción noroccidental del macizo de la Sierra Nevada de Santa Marta (SNSM) en los

Andes del norte de Colombia, donde una sección imbricada de la corteza inferior se encuentra preservada. Observaciones de campo y petrología acopladas junto con geocronología de zircones (LA-ICP-MS) de 17 nuevas muestras de rocas ígneas y metamórficas junto con geoquímica de roca total y mediciones de elementos traza en zircones fueron usadas para reconstruir la evolución tectónica de esta sección cortical fuertemente removilizada, yuxtapuesta

sobre rocas metamórficas de bajo grado durante el Paleozoico Tardío. La evolución metamórfica fue investigada utilizando diferentes termobarómetros (Grt-Bt, Hbl+Grt, Hbl+Pl+Qtz, GAS y Ti-in-Zr) con el fin de obtener estimaciones confiables de las condiciones de P-T durante al menos dos fases metamórficas identificadas por su paragenesis mineral. Nuestros resultados proporcionan nuevos constreñimientos para las reconstrucciones paleogeográficas entre Laurentia y Gondwana durante el Pérmico medio y hasta el Jurásico Temprano. Esta reconstrucción se define por la interacción del margen NW de Gondwana con los terrenos Centro-Americanos y Mexicanos durante la formación de Pangæa. Un escenario evolutivo para este ciclo supercontinental se propone en diferentes fases definidas por el inicio de un arco, el tiempo de metamorfismo, y exhumación tardía de segmentos de la corteza inferior. El cierre del Océano Reico (ca. 290 Ma) marca el inicio de la interacción de los diferentes bloques corticales durante la formación de Pangæa que causó: 1) fundidos anatócticos post-colisionales ca. 278 ± 0.5 Ma acompañados de un máximo engrosamiento cortical, 2) metamorfismo de tipo Barrovian con condiciones *P-T* máximas de (13 kb, 840°C) en ≈ 254 Ma seguido por 3) retroceso del segmento de placa subducido y delaminación cortical, 4) colapso orogénico, 5) debilitamiento térmico, extensión cortical y *underplating* de corteza máfica en la base

de la corteza continental que causó metamorfismo *LP-HT* en condiciones (5.6 kb, 550°C) a ≈ 225 Ma marcando el comienzo de una fase de *rift*. Estos procesos fueron acompañados por el relleno de una cuenca sedimentaria de tras-arco que acumuló al menos 2000m de sedimentos calcáreos y volcánico-clásticos suministrados por los altos de basamento circundantes. En ≈ 216 Ma la mayor parte del margen conjugado fue removida, pero los remanentes de este arco Pérmico aun se encuentran preservados en los complejos Oaxaca y Acatlán, y en el Bloque Maya. La atenuación cortical responsable del proceso de *rifting* y posterior deriva contribuyo a un eventual rebote isostático del segmento de placa del Pacífico subducido que reinició la actividad del arco durante el Jurásico. Para esta época el arco migró hacia el interior del continente, y la previa cuenca de tras-arco se redefinió como una depocentro de ante-arco el cual acumuló una carga adicional de ≈ 5000 m de depósitos volcánico-clásticos. Nuestra nueva reconstrucción define un rango de tiempo de 60 Ma desde la colisión de los Alleghenides, hasta la separación de Pangæa para este sector de los Andes del norte.

Capítulo 2. Nuevos datos termocronológicos de baja temperatura de rocas del Precámbrico, Pérmico a Jurásico, Cretácico Tardío y Paleoceno-Eoceno las cuales componen el cinturón metamórfico noroccidental de la sierra Nevada de Santa

Marta (SNSM) en los Andes del norte de Colombia revelan las tendencias de exhumación en esta parte del margen continental durante el intervalo Paleoceno-Mioceno. Nuestros análisis involucran trazas de fisión en zircones y apatitos y dataciones (U-Th)/He en apatitos, las cuales permitieron establecer historias termales (~240 °C – ~60°C) de una secuencia cortical invertida que expone rocas metamórficas desde facies esquistos verdes hasta granulita. Las trayectorias de temperatura-tiempo evidencian que el enfriamiento ocurrió de manera diacrónica en el macizo a través de la activación de fallas de cabalgamiento con vergencia NW con el inicio de una rápida exhumación erosional posterior a la colisión y el emplazamiento de plutones de la serie TTG bajo un gradiente geotérmico inicialmente alto de 40°C/km. Bajo este régimen el macizo se exhumó a tasas ~0.9 km/Myr durante 50-45 Ma. Estas tasas se incrementaron a 2 km/Myr para el intervalo 45-40 Ma con un gradiente geotérmico en disminución de 30°C/km, alcanzando los mayores valores de 2.7 km/Myr entre 37-35 Myr. Interpretamos la exhumación de esta secuencia de corteza inferior invertida como resultado del acoplamiento de la extrusión dúctil de una corteza continental engrosada junto con denudación superficial. Los eventos de exhumación post-Eoceno muestran un enfriamiento posterior y un decrecimiento de las tasas de exhumación hasta valores de ~0.2 km/Myr para ca. 25

Ma, esta variación se relaciona al inicio de un régimen compuesto transpresivo-transtensivo en el límite sur de la placa del Caribe como resultado de cambios en la convergencia entre las placas del Caribe y Sur América. Esta reorganización condujo a la activación de la falla de Santa Marta-Bucaramanga (SMBF) de rumbo NW y de la falla de Oca de rumbo E-W. Estas fallas asumieron la convergencia particionada y transfirieron deformación desde la placa del Caribe hacia el Macizo de Santander, la Cordillera Oriental de Colombia, la Serranía del Perijá, y los Andes de Mérida hasta el Mioceno Tardío. Interpretamos que este nuevo régimen tectónico inhibió la propagación de cabalgamientos convergencia NW y fue la causa de un basculamiento en dirección NE del macizo de la SNSM debido a la actividad extensional de la SMBF en conjunción con el arrastre tectónico producido por la placa del Caribe. Estos procesos resultaron en tasas de exhumación disminuidas de ~0.09 km/Myr que persistieron hasta el Mioceno Tardío durante la re-equilibración termal de la SNSM

Capítulo 3. Datos geo-termocronológicos provenientes de rocas sedimentarias junto con análisis estratigráficos, revelan una exhumación del macizo de la Sierra Nevada de Santa Marta (SNSM) durante el Neógeno temprano. este bloque triangular delimitado por fallas en los Andes del norte de

Colombia, expone rocas metamórficas Precámbricas a Paleozoicas, y restos de un arco magmático Triásico-Jurásico. En este trabajo comparamos el relleno de cuencas Neógenas de dos cuencas marginales, una occidental localizada junto a la falla de Santa Marta-Bucaramanga (SMBF) y una cuenca norte ocupando el margen sur de la falla de Oca.

La secuencia del occidente consiste de depósitos gravitacionales de conglomerados de hasta ≈ 1200 m de espesor que definen un delta tipo Gilbert progradacional emanando de un escarpe que coincide con la SMBF. Debido a su posición estructural y estratigráfica esta secuencia se compara a una secuencia transgresiva en la cuenca norte. Análisis de proveniencia de dos secciones representativas de estas dos secuencias, evidencian una fuente derivada

del basamento subyacente. Este hallazgo es concordante con el espectro de edades U-Pb en circones que carece de señales pre-Grenvillianas y Ordovícicas de sedimentos coetáneos de la cuenca del Magdalena, demostrando, que estos sedimentos fueron depositados en cuencas locales desconectadas del eje del sistema fluvial del proto-Magdalena. Con el fin de reconciliar la diferente organización de las dos cuencas marginales proponemos un basculamiento de escala regional de la SNSM hacia el noreste durante el Neógeno temprano. Este modelo considera la contemporaneidad entre un ciclo progradacional en el flanco occidental delimitado por una falla y un traspaso transgresivo de los conglomerados de la cuenca norte en una superficie inclinada durante una exhumación generalizada del macizo.

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Résumé de la thèse

La Sierra Nevada de Santa Marta (SNSM) est peut-être le massif de la croûte terrestre le plus complexe trouvé dans les Andes du Nord. Sa situation unique comme un massif triangulaire isolé séparé de la continuité de 7000 km de long de la chaîne des Andes, et comme le dernier massif avant les domaines de la plaque Caraïbes plus jeune, place la SNSM comme une île séparée de toutes les chaînes de montagnes environnantes de la marge continentale. Un relief important caractérise cette montagne qui atteint l'altitude la plus élevée dans tout le domaine des Caraïbes à 5750 m, et fait de la SNSM la plus grande chaîne de montagnes côtières du monde. Pour cette raison, la SNSM est un objet géologique unique qui comprend une biodiversité exceptionnelle depuis ses récifs coralliens dans la mer des Caraïbes, passant des forêts tropicales humides, aux forêts de nuage d'altitude puis aux bruyères, jusqu'à son magnifique sommet couronné de glaciers. Un tel endroit extraordinaire est habité par les humains depuis 6000 ans. Ces peuples les Arhuaco, Wiwa, Kogi et Kankuamos avaient appelé la Sierra Nevada de Santa Marta, comme "Le Coeur" du Monde: *Gonawindúa*, «*Ce nom ancestral fait référence dans leur cosmogonie au pic le plus important de la Sierra Nevada de Santa Marta, qui quand le monde spirituel a transcendé le monde matériel, a été le premier élément qui a émergé pour nous rappeler nos tâches, qu'il est le pilier de tout, Gonawindúa est le lieu où notre sens de la responsabilité, notre sentiment physique et matériel, notre intelligence, notre sagesse et notre force sont codifiés et conservés.* »

Au-delà du contexte historique et mystique la SNSM cache une histoire géologique également fascinante. L'un des sujets les plus intéressants à aborder au moment de l'étude de l'évolution de ce massif réside dans ses corrélations avec les Andes du Nord, d'Amérique centrale et des blocs Mexicains, puisque la SNSM non seulement possède une géométrie complexe, mais un ensemble prolifique de roches ignées et métamorphiques, dont les âges varient du Mésoprotérozoïque à l'Eocène, comprenant des sédiments et les dépôts volcanoclastiques depuis le Dévonien jusqu'au Miocène. En raison de sa position sur la marge nord-ouest de l'Amérique du Sud, l'étude de la SNSM nous donne l'occasion de résoudre des questions importantes sur l'évolution des cycles continentaux depuis l'époque de l'orogène Grenvillien, en passant par les orogènes Neoprotérozoïque-Pan-Africain, Ouachitan-Appalaches du Paléozoïque tardif, qui ont conduit à l'assemblage de la Pangaea, puis à la fragmentation de la Pangaea, suivie par le Rift au Jurassique de l'Atlantique central, et plus récemment, depuis le Crétacé supérieur, par le début de l'accrétion / subduction de la plaque des Caraïbes contre le nord-ouest de Amérique du Sud, qui a été l'un des moteurs les plus importants de l'orogénèse Andine-Alpine.

Dans la SNSM, des roches de tous les cycles de Wilson décrits précédemment sont conservées. Pendant ces vastes périodes de temps, la SNSM occupait une position externe sur la marge d'accrétion du NW du Gondwana et a donc accumulé des témoins des épisodes collisionnels et orogéniques, qui sont conservés comme écailles crustales de roches métamorphiques de bas à haut grade, séparées par des sutures. Pionier dans les expéditions explorant la géologie de la SNSM, le travail classique de Ganser (1955), effectua la première identification systématique des roches constituant ce massif et produisit la première carte géologique accompagnées d'observations pétrographiques et structurales. Près de 20 ans plus tard, MacDonald et al. (1971), Macdonald et Hurley (1969), Tschanz et al. (1974, 1969) réalisèrent une mise à jour complète de la géologie de la Sierra Nevada de Santa Marta, effectuèrent même des datations radiométriques K/Ar qui permirent pour la première fois d'étudier les processus thermiques qui affectèrent ces roches. Récemment, plusieurs études ont tenté de reconstituer l'histoire géologique de la SNSM, en s'appuyant sur les travaux fondamentaux de Tschanz et al. (1969), et ont exposé les implications paléogéographiques de l'évolution de la SNSM à travers les âges, et jour après jour, ont augmenté la résolution de la cartographie, de la géochimie et de la géologie isotopique (Bayona et al., 2010; Cardona et al., 2011, 2010b, 2010c, 2011a; Cordani et al., 2005; Doolan, 1970; Duke 2009; Ordóñez et al., 2002; Restrepo Pace et al., 1997; Villagomez et al., 2011b; Zuluaga et Stowell, 2012). Cependant, la plus grande partie de la SNSM reste inconnue des géologues et d'autres visiteurs étrangers, et au cours de mon expérience de travail dans la région, je ne pus pas atteindre certaines zones énigmatiques qui attendent de nouvelles recherches.

Dans cette enquête, je tenté de démêler l'histoire géologique de la Sierra Nevada de Santa Marta Massif utilisant les techniques les plus avancées en géochronologie, thermochronologie, géochimie et géochimie isotopique, techniques qui ont permis de recueillir une quantité importante de nouvelles données qui s'ajoutent à la base de données existante sur la SNSM. Mais compte tenu de l'histoire géologique intrinsèquement complexe, ces analyses n'ont été réalisées qu'après plusieurs mois de travail sur le terrain, au cours desquels une carte géologique et métamorphique détaillée a été construite et la collecte des échantillons a été réalisée suivant les procédures et techniques modernes de la géologie structurale (Fossen, 2010; Ramsay et Huber, 1987). Nos résultats comprennent une carte géologique révisée au 1:25 000, comprenant la définition de quatre nouvelles unités stratigraphiques, accompagnée de deux sections crustales de 320 km de long traversant le massif, et 8 coupes sériées de l'angle nord-ouest de la ceinture métamorphique de la SNSM. L'ensemble des données géochimiques et isotopiques comprend: i) 17 échantillons de roches ignées et métamorphiques et six échantillons de roches détritiques datées par U-Pb sur zircon par ablation laser induite par couplage plasma spectrométrie de masse (LA-ICP-MS), qui ont abouti à 2790 nouveaux âges et analyses in-situ d'élément-traces, ii) 16 roches ignées et métamorphiques qui ont donné 31 nouveaux âges thermochronométriques:

12 âges par traces de fission sur zircon, 11 âges par traces de fission sur apatites et 7 âges (U-Th)/He sur apatites, iii) Géochimie de la roche totale de 10 échantillons et finalement iv) analyses à la microsonde sur quatre lames minces de grenats zonés et péritectiques pour analyser leur chimie minérale.

Ces données ont été données acquises sur les unités du complexe métamorphique de la bordure nord-ouest de la SNSM et de roches sédimentaires voisines du massif. Avec ces données, nous avons étudié i) Les unités qui constituent la ceinture métamorphique nord-occidentale de la SNSM, et leurs relations chronologiques et stratigraphiques du Précambrien à l'Eocène; ii) L'intervalle de temps et les conditions P-T de l'événement métamorphique Paléozoïque tardif à Mésozoïque précoce (chapitre 1), iii) La chronologie de l'activité magmatique, des accrétions et de l'exhumation des terrains océaniques et continentaux au Crétacé supérieur et jusqu'à la fin du Miocène, iv) Un mécanisme pour expliquer comment l'exhumation a eu lieu sous un régime collisionnel influencé par un climat à forts gradients thermiques et par des taux d'érosion élevés (chapitre 2); v) Les processus tardifs de dénudation et de sédimentation contrôlés par la tectonique dans deux bassins sédimentaires depuis le Miocène inférieur soumis à des taux d'érosion et des gradients thermiques décroissants (chapitre 3).

Chapitre 1. La convergence entre Laurentia et Gondwana a finalement conduit à la collision et à la formation de la Pangée au Paléozoïque supérieur. De ces longues interactions de plaques, a résulté la création de roches de Haute Pression (HP) et de Haute Température (HT), qui caractérisent une grande partie du nord-ouest de la Sierra Nevada de Santa Marta Massif (SNSM) dans les Andes du Nord de la Colombie, où une section imbriquée de croûte inférieure est préservée.

Les observations de terrain et l'étude pétrographique, couplées aux âges U-Pb (LA-ICP-MS) sur zircon de 17 nouveaux échantillons de roches ignées et métamorphiques et l'analyse des élément-traces sur zircons et de la géochimie sur roche totale sont utilisés pour la

reconstruction de l'évolution tectonique de cette section crustale fortement remobilisée, charriée sur des séquences métamorphiques de bas grade à la fin du Paléozoïque. L'évolution métamorphique a été étudiée au moyen de plusieurs thermobaromètres (Grt + Bt, Hbl + Grt, Hbl + Pl + Qtz, GASP et Ti-in-Zr) pour obtenir des estimations fiables sur les conditions *P-T* pendant au moins deux phases métamorphiques identifiés par leurs paragenèses minérales.

Nos résultats fournissent de nouvelles contraintes pour les reconstructions paléogéographiques entre Laurentia et Gondwana au Permien moyen et jusqu'au début du Jurassique. Cette reconstruction est définie par l'interaction de la marge nord-ouest de Gondwana avec les terrains

d'Amérique centrale et du Mexique actuels lors de la formation de la Pangaea.

Un scénario évolutif pour ce cycle supercontinental est proposé en plusieurs phases comprenant la création d'un arc, l'épisode métamorphique, et l'exhumation tardive des segments de la croûte inférieure. La fermeture de l'océan Rhéique (ca. 290 Ma) marque le début de l'interaction des blocs crustaux lors de l'assemblage de la Pangaea, qui provoqua 1) une fusion anatectique post-collisionnelle et un épaississement maximal de la croûte vers 278 ± 0.5 Ma, 2) un métamorphisme de type Barrovian culminant à 13 kb et 840°C , vers 254 Ma, suivie par 3) le retrait de la plaque subduite et sa délamination, 4) un effondrement orogénique 5) affaiblissement thermique, extension crustale et sous-placage de croûte mafique qui a causé un métamorphisme *LP-HT* (5,6 kb, 550) vers 225 Ma, marquant le début d'une phase de rifting.

Ces processus sont associés au remplissage d'un bassin d'arrière-arc dans lequel se sont accumulés au moins 2000 m de sédiments carbonatés et volcanoclastiques fournis par les massifs de socles exhumés.

Vers 216 Ma la plus grande partie de la marge conjuguée a été retiré, mais les restes de cet arc Permien sont encore conservés dans les complexes Oaxaca et Acatlán et dans le bloc Maya. L'amincissement cortical responsable du rifting et la dérive qui a suivi a contribué au rebond isostatique de la plaque Pacifique subduite qui a réactivé l'arc

magmatique au Jurassique. A cette époque, l'arc a migré vers le continent, et l'ancien bassin d'arrière-arc passa en position d'avant-arc dans lequel se sont accumulés \approx 5000 m de dépôts volcano-clastiques. Cette nouvelle reconstitution établit une durée de 60 Myr d'âge entre la collision de les Alleghenides et l'éclatement du Pangaea dans cette partie des Andes du Nord.

Chapitre 2. De nouvelles données thermochronologiques à basse température sur des roches du Précambrien, Permien à Jurassique, Crétacé supérieur et Paléocène-Eocène qui forment la ceinture métamorphique nord-occidentale de la SNSM des Andes du Nord de la Colombie révèlent les tendances à l'exhumation de cette partie de la marge continentale entre le Paléocène et le Miocène. Nos analyses comprennent des traces de fission sur apatites et zircons et des âges (U-Th)/He sur apatite, qui ont permis de contraindre l'histoire thermique ($\sim 240^\circ\text{C}$ - $\sim 60^\circ\text{C}$) d'une séquence crustale inversée qui expose des roches métamorphiques de faciès schistes verts à granulite.

Les trajets température-temps prouvent que ce refroidissement est diachrone dans le massif en raison de l'activation de chevauchements vers le NW, avec le début d'une exhumation rapide liée à l'érosion après collision, et l'intrusion de plutons de la série TTG sous gradient géothermique initialement élevé de $40^\circ\text{C} / \text{km}$.

Dans ce régime, le massif a été exhumée à des taux élevés ~ 0.9 km/Myr de 50 à 45 Ma ; ce taux augmenta à 2 km/Myr pour l'intervalle 45-40 Ma avec un gradient géothermique diminuant à 30°C/km, et atteignit les taux les plus élevés pendant 2.7km/Myr de 37 à 35 Myr.

Nous interprétons l'exhumation de cette séquence de croûte inférieure inversée comme liée à l'association de l'extrusion ductile d'une croûte continentale épaissie et de la dénudation du surface. Les événements d'exhumation post-éocènes montrent une diminution des taux de refroidissement et d'exhumation à ~ 0.2 km/Myr vers 25 Ma, en rapport avec le début d'un régime composite de transpression-transension à la limite sud de la plaque Caraïbes en raison d'un changement de la convergence entre les plaques Caraïbes et Amérique du Sud. Cette réorganisation a conduit à l'activation de la Faille Santa Marta-Bucaramanga (SMBF) de direction NW et de la faille de Oca de direction E-W.

Ces failles ont accommodé le partitionnement de la convergence et ont transféré la déformation depuis la plaque Caraïbe vers le Massif de Santander, la Cordillère orientale de Colombie, la Serranía de Perija, et les Andes de Merida jusqu'à la fin du Miocène.

Nous proposons que ce nouveau régime tectonique a inhibé la propagation vers le NW de chevauchements et a provoqué le basculement vers le NE de la SNSM en raison

de l'activité distensive de la SMBF, associée à l'arrachement tectonique causé par le transit de la plaque Caraïbes. Ces processus ont entraîné des taux d'exhumation réduits de ~ 0.09 km/Myr qui ont persisté jusqu'à la fin du Miocène, lors de la rééquilibration thermique de la SNSM.

Chapitre 3. Geo-thermochronologie de roches sédimentaires avec analyses stratigraphiques, révèlent une exhumation du massif de la Sierra Nevada de Santa Marta (SNSM) au début du Néogène. Ce bloc triangulaire délimité par des failles dans les Andes du nord de la Colombie, exposée roches métamorphiques du Précambrien et Paléozoïque et reste d'un Trias-Jurassique arc magmatique. Dans cet article, nous comparons les bassins Néogène remplissent deux bassins marginaux, localisés ouest le long de la faille de Santa Marta-Bucaramanga (SMBF) et occupant le nord du bassin de la marge sud de la faille de Oca.

La séquence de l'Ouest est constituée de dépôts de conglomérats de gravité jusqu'à une épaisseur ≈ 1200 m définissant un delta type Gilbert progradante émanant d'un escarpement qui correspond à la SMBF. En raison de sa position structurale et stratigraphique cette séquence est comparée à une séquence de transgression dans le bassin nord. Analyse de provenance des deux sections de ces deux séquences montrent une source dérivée du socle sous-jacent.

Ce résultat est cohérent avec le spectre des âges zircons U-Pb dépourvus de signaux pré-Grenvillienne et Ordovicien des sédiments contemporains dans la bassin de Magdalena, montrant que ces sédiments ont été déposés dans les bassins versants locaux déconnectés de l'axe du système fluvial proto - Magdalena. Afin de concilier l'organisation différente des deux bassins marginaux nous proposons une inclinaison a échelle

régionale de la SNSM vers le nord-est au début du Néogène. Ce modèle tient compte de la contemporanéité entre un cycle progradante sur le flanc ouest délimitée par un faille et un conglomérats transgressive transfert du bassin nord dans une surface inclinée pendant exhumation généralisée du massif.

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Contents

Contents.....	XXIX
INTRODUCTION	1
Northern Andean Block.....	6
Grenvillian inliers in the NW margin of South America.....	7
Permo-Triassic Tahamí Terrane	8
Late Cretaceous-Paleogene Caribe Terrane	11
Objectives	19
Thesis Outline	21
CHAPTER 1	22
Permo-Triassic evolution in the Sierra Nevada de Santa Marta, from the Alleghenides collision to Pangæa break-up	22
Abstract	22
1. INTRODUCTION	23
2. GEOLOGICAL SETTING	26
2.1 The Massif of the Sierra Nevada de Santa Marta.....	26
3. METHODS	34
3.1 LA-ICP-MS U–Pb Geochronology	34
3.2 Microprobe analyses	36
3.3 Whole rock geochemistry	36
4. RESULTS	37
4.1 Field relationships, petrography and mineralogy	37
4.1.1 Sevilla Metamorphic Belt	37
4.1.2 Inner Santa Marta Metamorphic Belt.....	38
4.1.3 Outer Santa Marta Metamorphic Belt	47
4.2 Whole rock geochemistry	51
4.2.1 High-grade metamorphic rocks from the Sevilla Metamorphic Belt.....	51
4.2.2 Amphibolite-grade rocks from the IB.....	53
4.2.3 Greenschists from the OB.....	55
4.2.4 Granitoids associated with the Santa Marta Batholith	55
4.3 Zircon Morphology	55
4.4 Zircon U-Pb Geochronology.....	58
4.4.1 Sevilla Metamorphic Belt (SB).....	60
4.2 Inner Santa Marta Metamorphic Belt (IB)	63
4.4.3 The Outer Santa Marta Metamorphic Belt (OB).....	64
4.5 Trace Elements in Zircon	67
4.5.1 Sevilla Metamorphic Belt	68

4.5.2 Inner Santa Marta Metamorphic Belt.....	69
4.5.3 Outer Santa Marta Metamorphic Belt.....	71
4.6 Mineral Chemistry.....	74
4.7 Thermobarometry.....	78
4.7.1 La Secreta Mylonites - Garnet Amphibolites.....	80
4.7.2 La Secreta Mylonites - Biotite-Garnet schists.....	81
4.7.3 Ti-in zircon thermometry.....	83
5. DISCUSSION.....	85
5.1 Geochronological implications for the SNSM.....	85
5.2 Concluding remarks on zircon trace elements and mineral chemistry.....	91
5.3 Pangaea Assembly during Alleghanian Orogeny.....	95
5.4 Post-Orogenic Collapse and metamorphism during Pangaea Break-Up.....	97
6. CONCLUSIONS.....	102
7. REFERENCES.....	105
CHAPTER 2.....	112
Cenozoic exhumation of the Sierra Nevada de Santa Marta: From a collisional margin to a transpressive boundary.....	112
Abstract.....	112
1. INTRODUCTION.....	113
2. TECTONIC SETTING.....	115
3. METHODS.....	122
3.1 Bedrock fission-track thermochronology.....	122
3.1.1 Zircon fission track.....	122
3.1.2 Apatite fission track.....	122
3.1.3 (U-Th)/He in Apatites.....	122
4. RESULTS.....	123
4.2 Sevilla Metamorphic Belt (SB).....	124
4.2.1 ZFT.....	124
4.2.2 AFT.....	124
4.2.3 Apatite (U-Th)/He.....	125
4.3 Inner Santa Marta Metamorphic Belt (IB).....	125
4.3.1 ZFT.....	125
4.3.2 AFT.....	126
4.2.3 Apatite (U-Th)/He.....	126
4.4 Outer Santa Marta Metamorphic Belt (OB).....	126
4.4.1 ZFT.....	127
4.3.2 AFT.....	127
5. DISCUSSION.....	130
5.1 Accretion of Caribbean Terranes and onset of magmatism.....	130
5.2 Erosion rates.....	131
5.3 Exhumation driven by crustal channel flow.....	137

5.4 Exhumation along major faults	139
5.5 Implications on the evolution of the Nor-Andean Block.....	144
7. CONCLUSIONS	147
8. REFERENCES	150
CHAPTER 3	156
Early Neogene unroofing of the Sierra Nevada de Santa Marta determined from detrital geo-thermochronology and the petrology of clastic basin sediments.....	156
Abstract	156
1. INTRODUCTION	157
2. TECTONIC SETTING	160
2.1 Sierra Nevada de Santa Marta massif.....	162
3. STRATIGRAPHY OF THE ARACATACA AND PALOMINO MARGINAL BASINS	164
3.1 Aracataca Marginal basin (Western SNSM).....	164
3.2 Palomino Marginal Basin (Northern SNSM).....	165
4. METHODS	168
4.1 Clast Counting.....	168
4.2 LA-ICP-MS U-Pb Geochronology	168
4.3 Detrital fission-track thermochronology	170
5. RESULTS	171
5.1 Stratigraphy	171
5.1.1 Aracataca Conglomerates.....	174
5.1.2 Palomino Conglomerates	182
5.2 Clast Counting.....	190
5.3 Heavy Mineral Analysis.....	192
5.4 Detrital Zircon U-Pb Dating Results of the Conglomerate Matrix.....	195
5.5 Detrital Fission Track Thermochronology.....	198
5.5.1 Zircon fission-track results	198
5.5.2 Apatite fission-track results	200
6. DISCUSSION	201
6.1 Sediment Provenance	201
6.2 Regional Stratigraphic Correlations	205
6.3 Tectonic Implications for the Caribbean Realm	211
7. CONCLUSIONS	215
8. ACKNOWLEDGMENTS	217
9. REFERENCES	218
OVERALL CONCLUSIONS	225
ACKNOWLEDGMENTS	235

Figures List

INTRODUCTION

Figure 1. South America and main plate boundaries..	5
Figure 2. Regional tectonic map of the Caribbean realm showing the most relevant geological features intervening in the plate boundary configuration....	7
Figure 3. Simplified terrane map of Colombia after Colombian Geological survey (Gomez et al., 2015	9

CHAPTER 1

Figure 1. Regional tectonic map of the Caribbean realm showing the most relevant geological features intervening in the plate boundary configuration...	24
Figure 2. Geological Map of the Sierra Nevada de Santa Marta Massif..	29
Figure 3. Regional cross sections of the SNSM, locations in fig 2 labeled y letters A-A' and B-B'	30
Figure 4. Geological Map of the Sierra Nevada de Santa Marta Massif (SNSM). C-C' and D-D' locates cross sections of figures 5 & 6.....	32
Figure 5. Sections C-C' and D-D' illustrating the structural style of the SNSM ...	33
Figure 6. Aguja Creek cross section D-D'.	42
Figure 7. Main structural features of the low to medium metamorphic grade units at the SNSM	50
Figure 8. Protholith discrimination diagrams.....	52
Figure 9. Whole rock geochemistry: Green represents metavolcanic rocks and red samples are related to granitoids(a) AFM diagram (Irvine and Baragar, 1971); (b) A/CNK- A/NK plot (Shand, 1943); (c) TAS diagram (Bas et al., 1986); (d) Geotectonic classification of volcanic rocks (Schandl and Gorton, 2002); (e) REE diagrams normalized to chondrite (Nakamura, 1974) and Primitive Mantle (McDonough and Sun, 1995).	54
Figure 10. Panchromatic cathodoluminescence imagery of zircons analyzed in this study..	59
Figure 11. Zircon U-Pb LA-ICP-MS probability density plots obtained from the Precambrian units of the Sevilla Metamorphic Belt.....	61
Figure 12. Concordia plots (left) and trace element composition of zircons from the Precambrian units of the Sevilla Metamorphic Belt south of the Orihueca fault.	62
Figure 13. Zircon U-Pb LA-ICP-MS probability density plots obtained from the Permian to Jurassic metasedimentary and metaigneous units of the IB. Data sets labeled by star correspond are taken from (Cardona et al., 2010b, 2010c).	65
Figure 14. Concordia plots and trace element composition of zircons from the Permo- Triassic metaigneous and metasedimentary units of the IB.	66

Figure 15. Zircon U-Pb LA-ICP-MS probability density plots obtained from the Cretaceous to Eocene metabasites and granitoids of the OB and SB.	72
Figure 16. Concordia plots (left) and trace element composition (right) of zircons from the Cretaceous to Eocene metabasites and granitoids of the OB and SB....	73
Figure 17. Garnet composition discrimination diagrams (left) image, extract from the Grt maskfile generated in XMapTools® (Lanari et al., 2014).	79
Figure 20. P-T paths for garnet amphibolites	81
Figure 21. P-T paths calculated for Biotite+Garnet schists	84
Figure 22. Stratigraphic summary of the Sierra Nevada Province, Sevilla Belt, Inner Santa Marta Metamorphic Belt, and Outer Santa Marta Metamorphic Belt according to the field relationships and stratigraphy section	86
Figure 23. Palaeogeographic reconstruction at the Neoproterozoic- Cambrian transition \approx 650 - 520 Ma.	88
Figure 24. a) Schematic multi-stage evolution for the conjugate NW Gondwana margin and related arcs during Laurentia-Gondwana interaction. b) Palaeogeographic reconstruction of Pangaea late Permian \approx 250 Ma	98
Figure 25. a) Schematic not-to-scale multi-stage tectono-metamorphic evolution for the SNSM massif since Late Permian to Late Cretaceous times, involving metamorphism, anatexia, and sedimentation cycles.	101

CHAPTER 2

Figure 1. Regional tectonic map of the Caribbean realm showing the most relevant geological features intervening in the plate boundary configuration.	114
Figure 2. Digital elevation map (DEM) of the Sierra Nevada de Santa Marta Massif, including sample localities from previous studies and the new acquired samples.	116
Figure 3. Geological Map of the NW metamorphic belt of the Sierra Nevada de Santa Marta Massif (SNSM) with location of samples analyzed in this study ...	120
Figure 4. Bedrock ZFT data from the SNSM, probability density plots elaborated with Binomfit (Ehlers et al. 2005; Stewart & Brandon 2004).....	128
Figure 5. Bedrock AFT data from data from the SNSM, probability density plots elaborated with Binomfit (Ehlers et al. 2005; Stewart & Brandon 2004).....	129
Figure 6. Temperature- time Paths of the Sevilla Metamorphic Belt (SB), constructed with ZFT, AFT and (U-Th)/He cooling ages.....	133
Figure 7. Temperature-time Paths of the Inner Sant Marta Metamorphic Belt (IB), constructed with ZFT, AFT and (U-Th)/He cooling ages.....	135
Figure 8. Schematic diagram of kinematic relationship between channel flow and extrusion of a lower crustal channel, applied to the crustal section of the SNSM. Modified from (Godin et al., 2006)	138
Figure 9. 3D model of the Sierra Nevada de Santa Marta massif, cross section trace corresponds to the topographic profile presented in Fig. 2.	140
Figure 10. AFT and ZFT elevation profiles from samples of the IB and the Sevilla Metamorphic Belt.....	142

Figure 11. AFT, ZFT, AHe, ZHe thermochronometric ages obtained in this study and compiled with previous data (Cardona et al., 2011a; Villagómez et al., 2011),	144
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CHAPTER 3

Figure 1. Regional tectonic map of the Caribbean realm showing the most relevant geological features intervening in the plate boundary configuration.	158
Figure 2. Geological Map of the Sierra Nevada de Santa Marta Massif (SNSM) and sampling sites.....	163
Figure 3. Stratigraphic column of the sedimentary successions at the Rio Aracataca with compositional variations of clast counts. Right: location map of the Rio Aracataca section. Facies associations are described in detail in Table. 2.	172
Figure 4. Stratigraphic column of the sedimentary succession at the Rio Negro and at the Palomino basin with compositional variations from clast counts. Right: location map of the Rio Negro section. Facies associations are described in detail in Table. 3.....	173
Figure 5. Figures from the Aracataca section (facies refer to Table 2.):.....	189
Figure 6. Clast counts from The SNSM Cenozoic Aracataca and Palomino basins at the Negro and Aracataca rivers, respectively.	193
Figure 7. Panchromatic cathodoluminescence imagery of zircons analysed in this study.....	196
Figure 8. Detrital zircon U-Pb LA-ICP-MS probability density plots obtained from the Post-Eocene sediments of the western and northern foothills of the Sierra Nevada de Santa (Aracataca and Palomino basins),	197
Figure 9. Detrital ZFT data from Aracataca and Palomino basins, probability density plots elaborated with Binomfit (Ehlers et al. 2005; Stewart & Brandon 2004).	199
Figure 10. Detrital AFT data from data from Aracataca and Palomino basins, probability density plots elaborated with Binomfit (Ehlers et al. 2005; Stewart & Brandon 2004).	201
Figure 11. Schematic cross section showing the sources for the ZFT age populations found at the Aracataca Basin. Modified from (Bernet et al. 2006).	204
Figure 12. Stratigraphic correlation chart from the Plato-San Jorge basin, the adjacent Lower Magdalena Valley and San Jacinto belt; as well as the Guajira basin in the N foothills of the SNSM. Modified from (Cardona et al., 2012; Duque-Caro, 1979; Flinch, 2003; Tschanz et al., 1969).	206
Figure. 13 Conceptual model not to scale, depicting tectonic slopes associated with a simple tilt block/half-graben, involving a fault-bounded syn-extensional clastic wedge.	210
Figure 14. Schematic multi-stage evolution for the Caribbean-South American Margin, in the sense of cross section A-A' Fig 3.....	212

Figure 15. Comparative plots of detrital zircon U-Pb age spectra for Oligocene - lower Miocene sedimentary rocks from the Guajira Peninsula (Zapata et al., 2010), the Middle Magdalena Valley (Caballero et al., 2013; Horton et al., 2015), the axial Eastern Cordillera and Llanos basin (Horton et al., 2010b, 2010a), and the Aracataca and Palomino basins from this study. 213

OVERALL CONCLUSIONS

Figure 1. Sketch of a Thermal-Mechanical evolutive scenario of the crust at plate boundaries and the crustal tectonic cycle 226
Figure 2. Comparison of all U-Pb data previously published and from this work of bedrock samples from the SNSM massif represented by the blue curve and bins, the red arrows indicate the orogenic peaks recognized by Runcorn, (1962) brown shade in the background corresponds to the spectra of U-Pb ages compiled by (Hawkesworth et al., 2010). 227
Figure 3. Spectra of U-Pb data from the SNSM massif illustrates periods of crustal growth associated to tectonic events 228

INTRODUCTION

Evolution of orogens is defined by events that may involve igneous activity, metamorphism, surface uplift, and exhumation coupled with erosion and coeval sedimentation processes. The different stages of evolution can be deciphered by the integration of a growing number of geophysical and geological techniques including: field observation to mantle geophysics through metamorphic petrology, sedimentology, geochemistry, thermochronology, numerical and analogue modeling and others on various spatial and temporal scales that may reveal the transient stages of an orogen (Rolland et al., 2012).

The orogenic stages may involve a “juvenile” stage of crustal thickening and synchronous or late metamorphism that may vary in metamorphic degree depending of a given geological configuration, time span, and kinematic constraints. During this stage, maximum crustal thickening is achieved, and most of the high temperature process act during mineral phase equilibration until a metamorphic peak is reached.

The onset of exhumation marks another fundamental process in the evolution of orogens and is defined by the onset of erosion that acts over a newly exposed mountain range causing the activation of sedimentary cycles which will transport minerals to new basins or depocenters in which the removed material will accumulate to form syn-orogenic sedimentary rocks.

These orogenic stages can coexist at any given moment and overlap in their histories, depending mostly on the timing of igneous activity, metamorphism and exhumation, which are controlled by plate tectonic-scale processes, as changes in plate convergence

magnitude and direction, crustal thermal regimes and external influences to the system like plume-induced anomalies that increase thermal gradients and crustal production.

During orogenesis, an accretionary margin will evolve from an orogenic wedge to a continental plateau, if the mass flux related to tectonic accretion dominates over the redistribution caused by erosion. Crustal thickening enhances radiogenic heat production and conduction overcomes advection. Such processes may involve the increase of geothermal gradients and crustal melting. Lower crustal weakening will induce lateral flow and therefore, vertical growth of the wedge range is replaced by a rectangular continental plateau geometry indicating that the orogenic belt cannot sustain topographic gradients anymore. Orogenic plateaus can sustain steady state equilibrium for several tens of millions of years if accretion is balanced by gravity driven flow (Vanderhaeghe, 2012).

In a long-lived orogenic plateau maintaining steady state conditions of growth and maximum thickness, a transition to crustal thinning is influenced by lithospheric-scale boundary condition modifications such as plate convergence reorganization, and subduction failure.

Subduction can fail by either thickening of the oceanic plateau by accretion, or thickening of the continental plateau, or a combination on different time scales of these processes. Another external factor to consider is a redistribution of the geoid highs by mantle plume activity, which may influence an extensional regime under a thickened continental crust (Murphy and Nance, 2013). Under these circumstances, an orogen will tend to thin by its gravitational potential, and thus reach a stage of

gravitational collapse in which late crustal attenuation will increase geothermal gradients until finally rifting and opening of new oceans may occur.

During orogenic evolution, the generation of crust due to magmatism generates zircons that constitute an accessory mineral in the majority of upper crustal rocks, and by their endurance under extreme thermobarometric conditions, are ideal for studying the igneous and metamorphic evolution of an orogen. At the same time the crust is consumed along subduction zones, in many cases disappearing by tectonic erosion, the remnants of an eroded crust are incorporated in the accretionary wedge. A positive feedback of oceanic crustal thickness may enhance exhumation of the continental plateau, particularly observed in several parts of the Andean margin (Spikings and Simpson, 2014). Another consequence of the increment of the thickness of the subducted plate is a late subduction failure (Stern, 2011). During these interactions between the oceanic-continental plates, the activation of sedimentary cycles reflects the changes between exhumation and plate geodynamics. Therefore it is possible to quantify the time span and magnitude, temperature and pressure conditions of these events by studying zircons, apatites and garnet among other mineral species, found both in-situ and recycled from their sources and accumulated in sedimentary basins adjacent to the orogenic belt.

The Andes form a 7000 km long orogenic belt on the western margin of South America (Fig. 1). The current scenario is defined by the subduction of the oceanic Pacific plate, under the cold lithosphere of the South American plate. The interaction of the Pacific, South American and Caribbean Plates controls the structure of the Andes, which in the Neogene led to the individualization of the North Andean block, as a response of plate

reconfiguration kinematics during the rupture of the Pacific Farallon plate originating the Cocos and Nazca plates.

In this work, my aim is to study the tectonic history of the northern Andes focused on the Sierra Nevada de Santa Marta (SNSM) massif, the highest topographic relief of the Colombian Andes. This massif occupies a northwestern position within the Maracaibo Block (Figs. 2, 3 & 4) and is composed by the amalgamation of igneous and metamorphic rocks from different ages; my approach is directed to establish an orogenic evolution of the SNSM massif by a combination of isotopic and geochemistry analyses in detrital and bedrock minerals. In the following sections, I will describe briefly the main geological features from the Northern Andes of Colombia.

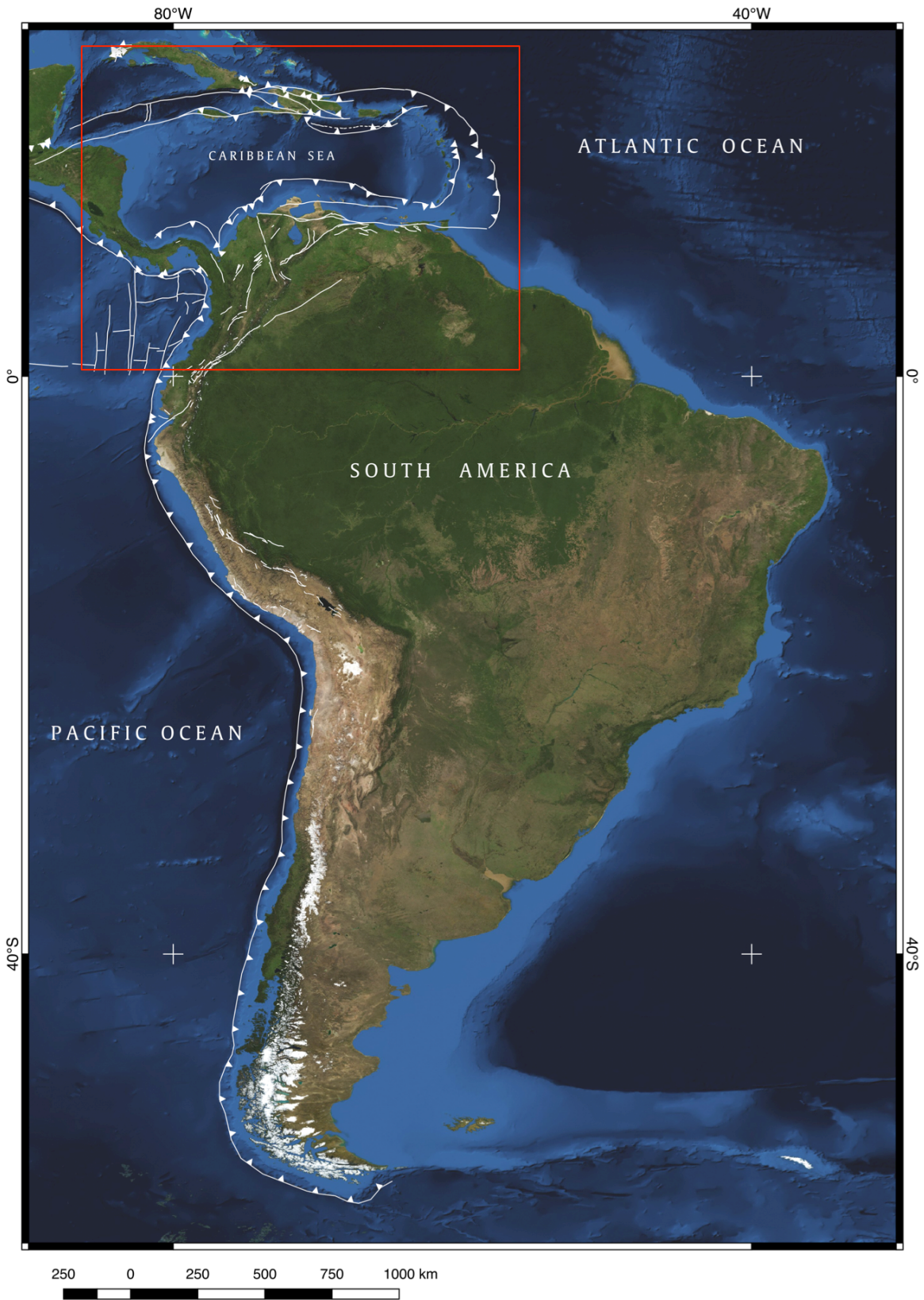


Figure 1. South America and main plate boundaries. Taken from http://server.arcgisonline.com/arcgis/rest/services/ESRI_Imagery_World_2D/.

Northern Andean Block

The Colombian Caribbean margin consists of a group of individualized tectonic blocks that represent an atypical section of the Andean margin because of its transpressive plate-tectonic configuration. In this context, the Northern Andean block can be divided into several tectonic terranes separated from the rest of the Andes through a main frontal thrust system denominated the Borde Llanero fault system that bounds the deformation front to the east against the Guyana shield (Fig. 2 & 3). From southwestern Colombia, the mountain ranges of the Colombian Cordilleras trend NE (Fig. 2). The Central and the Eastern Cordillera are separated by the Upper and Lower Magdalena basin, which reaches its major depth near the Caribbean margin, filled with Mesozoic and Cenozoic sediments. To the west of the Lower Magdalena basin is the accretional Sinú-San Jacinto fold belt (Fig. 2).

The Lower Magdalena basin resulted from a Paleocene rotational basin, separating the Sierra Nevada de Santa Marta massif (SNSM) from the foothills of the Central Cordillera and the San Lucas range (Fig. 2). The Lower Magdalena basin is attributed to transtensional tectonics, in which normal faults perpendicular to the margin were responsible for block rotation and translation of the SNSM block to its current position (Bernal-Olaya et al., 2015; Flinch, 2003; MacDonald and Opdyke, 1972; Montes et al., 2010).

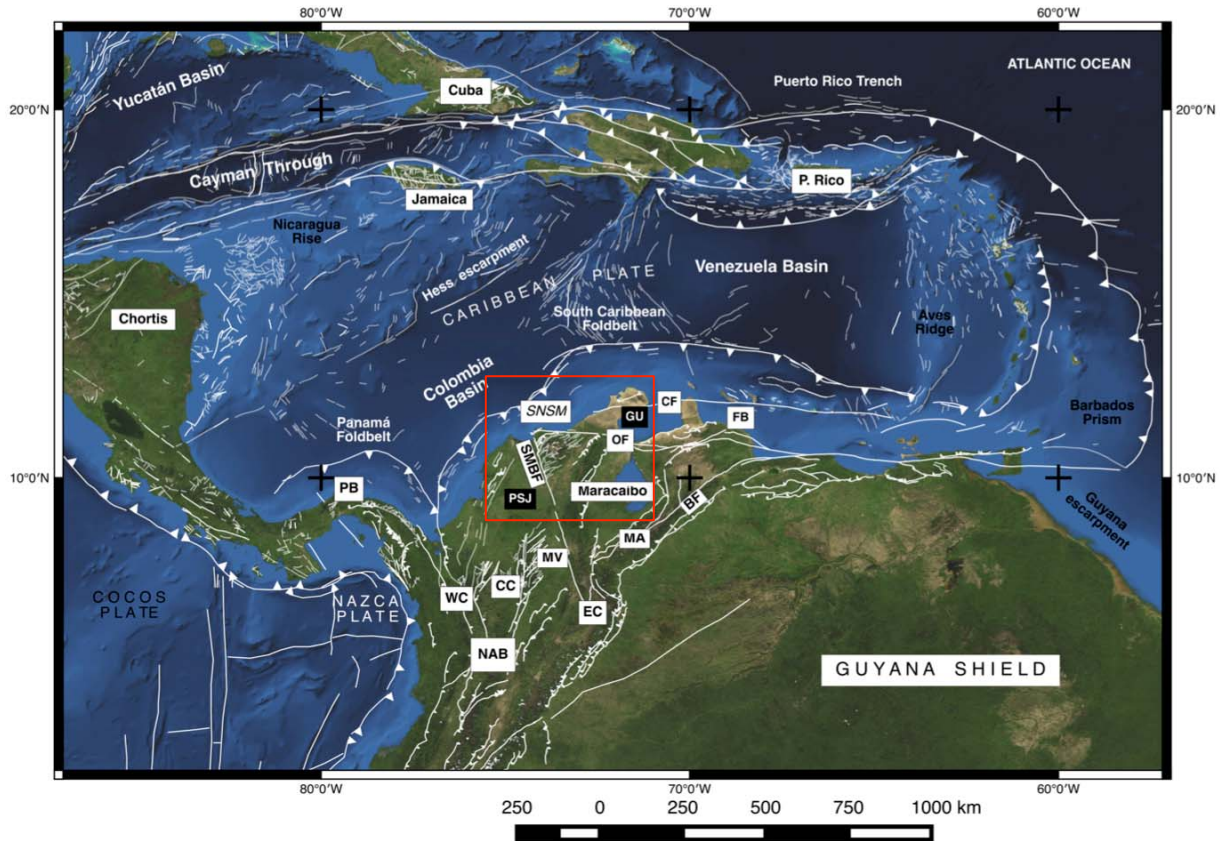


Figure 2. Regional tectonic map of the Caribbean realm showing the most relevant geological features intervening in the plate boundary configuration. (Veloza et al., 2012). Elements of the North Andean Block (NAB): Maracaibo block, SNSM: Sierra Nevada de Santa Marta, OF: Oca Fault, CF: Cuisa Fault, SMBF: Santa Marta-Bucaramanga Fault, GU: Guajira Basin, PSJ: Plato-San Jorge Basin, MA: Mérida Andes, BF: Boconó Fault, FB: Falcon Basin, WC: Western Cordillera, CC: Central Cordillera, EC: Eastern Cordillera, MV: Magdalena Valley, PB: Panamá Block. Satellite imagery taken from: http://server.arcgisonline.com/arcgis/rest/services/ESRI_Imagery_World_2D/.

Grenvillian inliers in the NW margin of South America

The easternmost morphotectonic province in the Colombian Andes corresponds to a Precambrian belt of granulites, gneisses and anorthosites (Fig. 3), derived from igneous and volcanoclastic protoliths, which chronologically can be traced up to the Grenvillian orogeny, locally defined as the Putumayo orogeny (Cardona et al., 2006; Kroonenberg, 1982; Ordóñez et al., 2002). The Putumayo orogeny was the consequence of the

collision between Amazonia and the Sveconorwegian province in Baltica during the early Neoproterozoic (Ibanez-Mejia et al., 2011).

These Precambrian inliers are denominated as the Chibcha terrane and are restricted to the east of the Otú-Pericos-Palestina fault system (Feininger, 1970; Ordóñez, 1999; Vinasco et al., 2006). They crop out in the San Lucas Range, the Santander Massif, the Guajira Peninsula, and the SNSM (Fig. 2 & 3), and constitute Grenvillian remnants that had undergone at least three metamorphic episodes (Cardona et al., 2010a).

The Precambrian rocks of the Santander Massif are intruded by Ordovician to Silurian plutons (Van Der Lelij, 2013), corresponding to a magmatic arc that fringed Gondwana. Remnants of this arc can be found in the Central Cordillera were they had been defined as the Anacona suspect terrane (Fig. 3; Martens et al., 2014). Ordovician-Silurian plutons are elusive or absent in the Guajira peninsula and SNSM, although zircons of this age interval are commonly reworked through orogenic recycling and partial melting into younger plutons as well as into the medium-grade metamorphic units of the SNSM (Cardona et al., 2010b, 2010c).

Permo-Triassic Tahamí Terrane

The central morphotectonic province of the Colombian Andes corresponds to a suite of Paleozoic meta-sedimentary and meta-igneous rocks denominated the Tahamí terrane, which includes Permian orthogneisses affected by partial melting during the Triassic (Cochrane et al., 2014a; Vinasco et al., 2006). The Tahamí terrane includes the Central

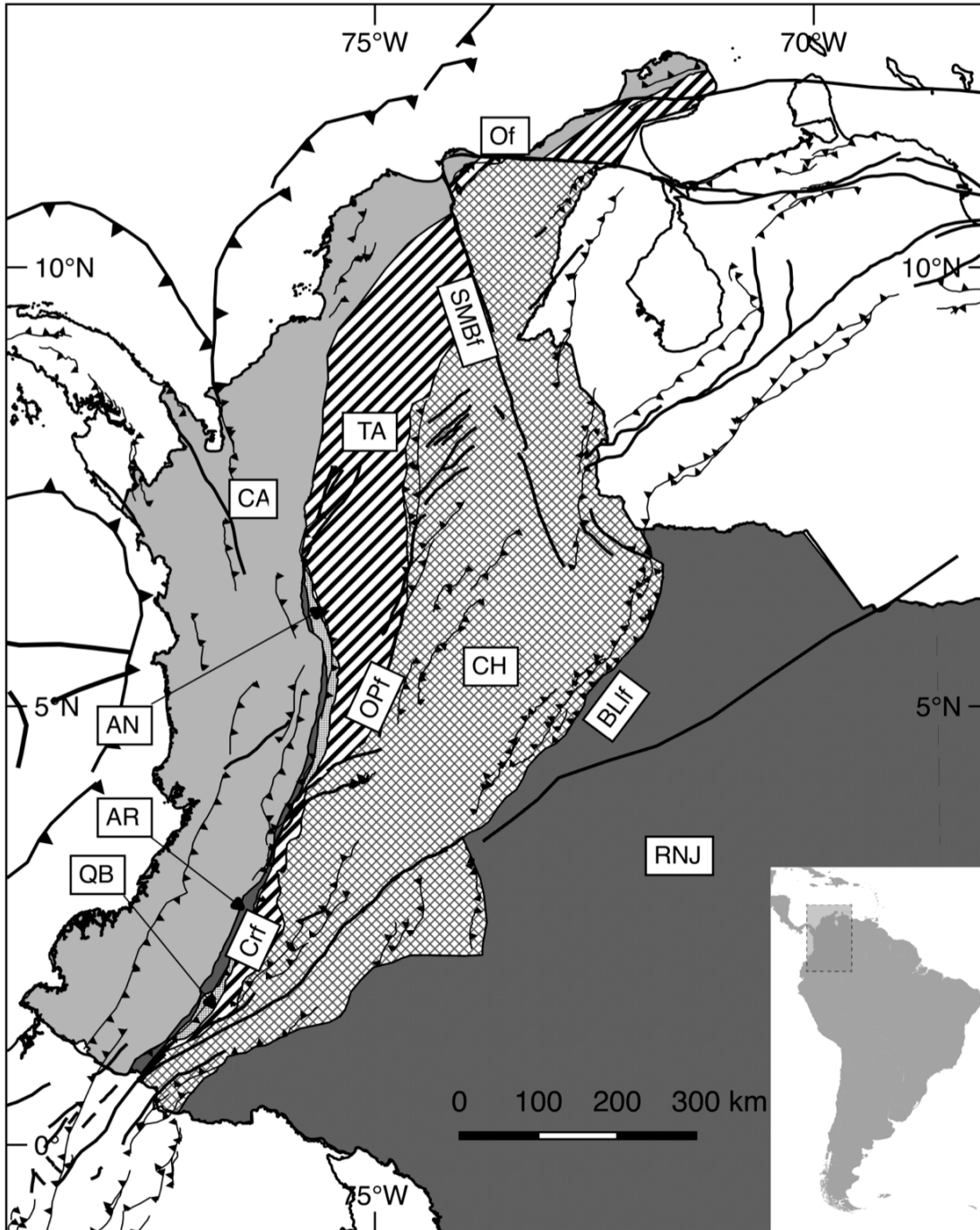


Figure 3. Simplified terrane map of Colombia after Colombian Geological survey (Gomez et al., 2015). Terrane names: RNJ: Rio Negro Juruena Province; CH: Chibcha terrane; TA: Tahamí terrane; AN: Anacona terrane; QB: Quebradagrande terrane; AR: Arquía Terrane; CA: Caribe Terrane; OPf: Otú-Pericos fault; SMBf: Santa Marta-Bucaramanga fault; CRf: Cauca Romeral fault; Of: Oca fault; BLf: Borde Llanero fault.

Cordillera of Colombia (Ordóñez Carmona and Pimentel, 2002; Restrepo et al., 2011; Spikings et al., 2015; Villagómez et al., 2011a; Villagómez and Spikings, 2013; Vinasco et al., 2006), the northernmost part of the Sierra Nevada de Santa Marta Massif (Cardona et al., 2010c), and the Guajira Peninsula (Fig. 3). Furthermore, it corresponds to the remnant of a Peri-Gondwana arc composed of meta-igneous and metasedimentary rocks that host anatectic melts. The Tahamí terrane can be correlated with other exposures of Permian to Triassic meta-sedimentary and meta-igneous rocks in the Andes, such as the Loja (Aspden and Litherland, 1992), and the Amotape or El Oro Complex (Chew et al., 2007; Riel et al., 2013), the late Paleozoic sediments and metaluminous to peraluminous intrusive rocks of the early Paleozoic Marañon Complex in Peru (Chew et al., 2007; Mišković et al., 2009) and the Chiapas massif in southern Mexico (Weber et al., 2007).

The remnants of the Permo-Triassic metamorphic belt in the SNSM are on their NW margin capped by a sequence of meta-igneous and meta-pelitic units with a Caribbean affinity (Doolan, 1970; MacDonald et al., 1971; Tschanz et al., 1969, 1974). Recent geochemical and geochronological studies (Cardona et al., 2010b, 2010c; Cordani et al., 2005) allowed establishing the accretion of the Caribbean plateau and related arcs at least since the Turonian (Cardona et al., 2011).

Paleomagnetic and geochronological data from the Cesar Ranchería basin (Bayona et al., 2006; Nova et al., 2013), and seismic stratigraphy and zircon U-Pb geochronology in well core samples from the metamorphic basement of the Lower Magdalena valley at the Plató-San Jorge depocenter (Fig. 1; Montes et al., 2010) shed light on the Cenozoic clockwise rotation of the SNSM to its current position, and permitted as well to establish a co-genetic relationship between the Central Cordillera and the SNSM basement,

where Permo-Triassic granitoids are slightly deformed at the Plat6-San Jorge basin, whereas orthogneisses of the same age in the SNSM are mylonitic. From all this it was concluded that the SNSM was segmented from a former continuous margin that was composed of a Permo-Triassic arc (Cardona et al., 2010c; Montes et al., 2005, 2010).

Late Cretaceous-Paleogene Caribe Terrane

The youngest morphotectonic province of the Colombian Andes is defined as the Caribe Terrane (Fig. 3). The Caribe Terrane is located in the western margin of the Colombian Andes and comprises the oceanic and volcanic sequences accreted through the Romeral Suture to the pre-existent Permo-Triassic terranes of the continental margin, these terranes were separated from the rocks of the continental margin through a major suture known as the Cauca-Romeral fault (Fig. 3).

The origin of Caribbean plate is widely discussed, between a Pacific realm or a segmentation and incorporation between Cocos and Nazca plates, against cratonic South America. The main boundary of this scene can be reconstructed up to Late Cretaceous ages where an active subduction process began related to Thethys closure, evidenced from the marine sequences with an important basaltic input which acted until lower Paleocene times. The accretion of an oceanic plateau during the Campanian defined as the CLIP (Hoernle et al., 2004; Kerr et al., 1997; Villag6mez and Spikings, 2013; Whattam and Stern, 2015), yields major implications on the origin and late evolution of the Caribbean Plate.

Models for Caribbean evolution involve the development of this proto Caribbean oceanic floor and its accretion against continental South America, which is defined by a

preexistent Jurassic volcanic arc and basin emplaced on older Paleozoic sedimentary and low to medium grade metamorphic rocks of the Tahamí terrane, which overly unconformably the Precambrian basement of the Chibcha terrane.

The Caribbean Colombian Margin is configured by two belts that define the continental slope intensely deformed and mountain areas in its continental portion. The first is commonly addressed as the *South Caribbean Deformed Belt (SCDB)*, which is constituted by sediments from Oligocene times to the present. The marginal high is conformed by the South American Plate Basement Cretaceous to Cenozoic platform sediments. At the Sierra Nevada de Santa Marta massif (SNSM) (Fig. 4) this structural high exhibits its major development and is composed by three provinces (Tschanz et al., 1974): 1) the south eastern flank (Sierra Nevada Province) comprises a Jurassic magmatic-volcanic arc that grew over a Grenvillian basement (Cardona et al., 2010a; Restrepo-Pace et al., 1997). 2) The intermediate Sevilla province consists of Neoproterozoic gneisses intruded by Permian to Triassic syntectonic granitoids (Cardona et al., 2010c). 3), and the Santa Marta province to the northwest consists of Triassic-Cretaceous greenschists associated with amphibolites and orthogneisses (Doolan, 1970).

In order to understand the tectonic relationships it is crucial to investigate the origin and structural style of the oceanic unit, that conforms the Santa Marta belts (Tschanz et al., 1974): the sub-units of the schists and amphibolites of Santa Marta had been object of an exhaustive geochemical and isotopic research that includes results on the provenance obtained by zircon U-Pb dating (Cardona et al., 2010b). As a set these units are characterized by an increase in the metamorphic degree from NW to SE including a low-grade sequence of metavulcanites and metapelites with MORB signature (Fm.

Concha) to the schists and amphibolites of a medium to high-pressure grade in an internal structural position (San Lorenzo Schists). For the Metavulcanites a zircon age of 82 Ma was obtained (Cardona et al., 2010b). The zircons from a continental source are scarce. For the most internal units of San Lorenzo and Rodadero (amphibolites, amphibole schists) the youngest reworked zircons showed Middle Triassic ca. \approx 235 Ma and Middle Jurassic ca. \approx 157 ages. At these units, zircon populations include an age spectrum since the Mesoproterozoic and therefore show an affinity with the continental basement. Compared with the situation at the Guajira Peninsula, these sequences do not include continental slope derived deposits (Weber et al., 2010). In contrast, the continental basement to the east of the Santa Marta belts exposes in an intrusive contact with Jurassic extrusive rocks, low deformed platform sediments from the Albian (Fig. 4 & 5). The lack of continental slope sediments implies a tectonic hiatus that would be explained by tectonic erosion of these marine sequences typical of the Guajira Peninsula.

For the Subduction event switched subduction polarity situations can be assumed, which are combined, although supposing an allocthonous origin for the Caribbean plate. A coherent synthesis for the study area assumes a subduction of the continental plate and an oceanic-transitional plate or proto-Caribbean under the Caribbean plate. A late stage of subduction failure in consequence is followed by an accretionary event for an epoch prior or coeval with the metamorphic peak. This evolution concludes with a subduction polarity reversal since the Late Cretaceous linked with the current scenario (Pindell et al., 2005; van der Lelij et al., 2010).

The main interest in the studied area resides in the tectonic relationship between units that belong originally to the domain of the Caribbean plate (Santa Marta schists) and

the continental basement of the Sevilla Metamorphic Belt (Fig. 5). For the contact between these two provinces, a suture has been speculated, but its nature and tectonic meaning is still matter of debate. The problematic of the study area can be compiled in two main questions: 1) what mechanisms juxtaposed these tectonic units? And 2) what is the cause of Cenozoic uplift of the SNSM?

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Objectives

The main objective of this dissertation is to unravel the tectonometamorphic history of the northwestern corner of the Sierra Nevada de Santa Marta massif (SNSM). During the late Paleozoic to Neogene, investigating igneous and metamorphic bedrock from the massif and the Neogene sediment series preserved on the flanks of this mountain range. In order to achieve this main objective, I:

1. studied the deformation of the schists in the northwestern corner of the SNSM massif from microtectonics to map scale, for determining the mechanisms of accretion and late exhumation of these units by means of integrating geochronology and thermochronology.
2. examined the conditions that led to a Late Cretaceous to early Paleogene subduction failure, between the South American plate and the Great Caribbean arc in the study area, and identified deformational phases for collisional tectonics, proposing a new evolutionary model.
3. identified the evolution phases of the basement of the continental margin by U-Pb dating of zircons from igneous and metamorphic rocks.
4. constrained the peak metamorphic conditions (P-T) undergone by the units of the Inner Santa Marta Metamorphic belt, in order to define their geodynamic implications in a chronological context, by analyzing mineral chemistry of paragenesis.
5. evaluated the timing and spatial distribution of magmatic and metamorphic events during the Alleghanian, and Andean orogenies. The occurrence and importance of rift,

subduction, and collision-related volcanism is evaluated by investigation of the age and chemistry of zircons and bulk rock compositions.

6. examined the exhumation history of different tectonic provinces using apatite and zircon fission-track coupled with apatite (U-Th)/He thermochronology.

7. evaluated the widely accepted Caribbean tectonomagmatic origin for the low to medium metamorphic grade units of the SNSM, their stratigraphic context, boundaries and its implications in the geodynamic evolution of northwestern SNSM.

8. investigated the detrital zircon provenance in Neogene sediment series for establishing a chronology on the exhumation of the massif.

Thesis Outline

Structural Framework, deformation and exhumation of the Santa Marta Schists: accretion and deformational history of a Caribbean Terrane at the north of the Sierra Nevada de Santa Marta

The methods and results of this work, along with interpretations and conclusions of this dissertation are presented in three chapters, which cover the proposed objectives of research. The geological context of the evolution of the SNSM is explained in each chapter as needed. I choose to present each chapter as an independent research article with individual discussion, conclusions and references. Currently this thesis includes two articles to be submitted (Chapter 1, Chapter 2), and one article under review for publication in the Geological Society of America Bulletin (Chapter 3). A regional description of the crustal remnants of the continental margin and their provenance, metamorphism and chronological relations during the orogenic cycle that led to Pangaea formation and late break-up is presented in Chapter 1. A thermochronological study constraining the early Cenozoic exhumation of the northwestern SNSM in the context of Caribbean plate subduction and collision against South America and later onset of strike-slip tectonics is discussed in Chapter 2. Finally, a combined structural and sedimentary geology investigation aimed at reconstructing the tectonic and exhumation history of the SNSM during the early Neogene is discussed in chapter 3. A conclusion chapter after Chapter 3 summarizes the main contributions of this dissertation, which was done under joint supervision between the Universidad Nacional de Colombia at Bogotá, Colombia and the Université Grenoble Alpes at Grenoble, France.

CHAPTER 1

Permo-Triassic evolution in the Sierra Nevada de Santa Marta, from the Alleghenides collision to Pangæa break-up

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Abstract

Convergence between Laurentia and Gondwana eventually led to the consolidation of Pangaea during the late Paleozoic. This long lasting plate interaction gave rise to belts of HT-HP rocks, which characterize a large part of the northwest portion of the Sierra Nevada de Santa Marta Massif (SNSM) in the Northern Andes of Colombia, where a tilted metamorphic lower crustal section is preserved. Field and petrological observations coupled with zircon U-Pb LA-ICP-MS geochronologic data from 17 samples of igneous and metamorphic rocks, together with whole rock and trace element geochemistry, are used for reconstructing the evolution of this strongly remobilized crustal section that was juxtaposed against a low-grade metamorphic terrane during the late Paleozoic. The metamorphic evolution is asserted by means of several geothermobarometers and yields (Grt+Bt, Hbl+Grt, Hbl+Pl+Qtz, GASP and Ti-in-Zr) and can be tracked by P-T conditions, allowing the differentiation of at least two metamorphic phases identified by their mineral paragenesis. Our results provide new constraints for the palaeogeographic reconstruction of the relative position of Laurentia and Gondwana during the middle Permian to Early Jurassic. This reconstruction is defined by the interaction of the NW margin of Gondwana with Central American and Mexican terranes during Pangaea amalgamation. An evolutionary scenario for this supercontinental cycle is proposed in several phases defined by arc inception, timing of metamorphism, and late exhumation of lower crustal segments. The closure of the Rheic Ocean (ca. 290 Ma) marks the start of interaction of crustal blocks during Pangaea assembly that caused 1) post-collisional anatexis melts at ca. 278±0.5

Ma and maximum crustal thickening, 2) Barrovian type metamorphism peak P - T conditions (13 kb, 840°C) at \approx 254 Ma followed by 3) slab retreat and delamination, 4) orogenic collapse, 5) thermal weakening, crustal extension and underplating of mafic crust that caused HT - LP metamorphism (5.6 kb, 550°C) at \approx 225 Ma marking a rifting phase. These processes were accompanied by the infilling of a back-arc basin that accumulated at least 2000 m of volcanoclastic sediments provided by the adjacent arc, and towards the basement highs fringed by carbonate platforms. By \sim 216 Ma most of the conjugated margin was removed, but remnants of this Permian arc are still preserved in the Oaxaquia, Acatlán, and the Maya blocks. The crustal attenuation responsible for rifting and drift contributed to an eventual isostatic rebound of the subducted Pacific slab that reactivated arc activity during the Jurassic. By this time, the arc migrated into a continent-ward position, and the former back-arc basin turned into a fore-arc depocenter accumulating at least an additional burden of \approx 5000 m of volcanoclastic deposits. This evolutionary scheme comprises an age span of 60 Myr from the Alleghenides collision to Pangaea break-up at this portion of the Northern Andes.

Key words: Lower crustal sequence, Peri-Gondwana correlations, Barrovian metamorphism, Pangaea, U-Pb Geochronology, P-T paths.

1. INTRODUCTION

Paleogeographic reconstructions and geodynamic models support the closure of the Rheic Ocean from Carboniferous to Permian times. The Rheic Ocean closure in eastern Pangaea was almost complete during the Mississippian with the onset of the Variscan orogeny (Franke, 2006; Nance et al., 2012). In western Pangaea, the closure of the Rheic Ocean and onset of crustal thickening occurred during the middle Permian. Because of this long lasting deformation during Pangaea formation, the Variscan-Appalachian-Ouachitan-Marathon and Huastecan mountain ranges constituted a continuous but diachronous late Paleozoic mountain belt, referred to as the Alleghenides (Pindell and Dewey, 1982; Pindell, 1985).

In the northwestern South American Andes (Fig. 1), remnants of the late stages of amalgamation of this colossal mountain belt are preserved in isolated blocks, in which

Precambrian Gondwanan inliers are juxtaposed against Permo-Triassic meta-igneous and meta-sedimentary rocks probably related to an accreted arc. However, little is known about the geodynamic evolution of these blocks and how they came to interact during two major phases of this supercontinental cycle: 1) timing of metamorphism and igneous activity involved during arc collision, crustal thickening, metamorphic peak, thermal re-equilibration and 2) late post-orogenic collapse crustal attenuation, rifting and drifting, which culminated in Pangaea break-up.

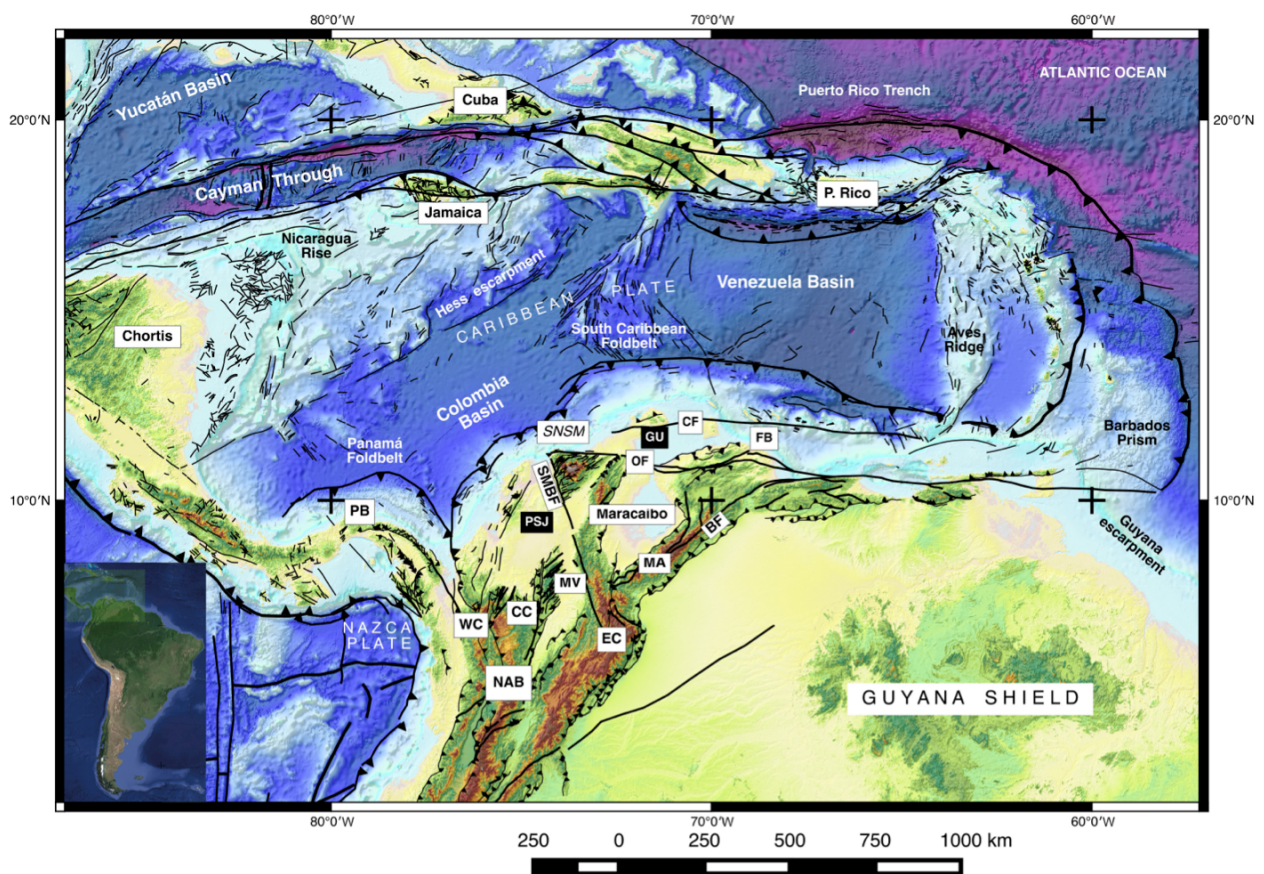


Figure 1. Regional tectonic map of the Caribbean realm showing the most relevant geological features intervening in the plate boundary configuration. After (Veloza et al., 2012). Elements of the North Andean Block (NAB): Maracaibo block, SNSM: Sierra Nevada de Santa Marta, OF: Oca Fault, CF: Cuisa Fault, SMBF: Santa Marta-Bucaramanga Fault, GU: Guajira Basin, PSJ: Plato-San Jorge Basin MA: Mérida Andes, BF: Boconó Fault, FB: Falcon Basin, WC: Western Cordillera, CC: Central Cordillera, EC: Eastern Cordillera, MV: Magdalena Valley, PB: Panamá Block.

This study presents new field, petrological, geochronological, and geochemical data, as well as thermobarometric estimates from an inverted lower crustal section of the Sierra Nevada de Santa Marta Massif in the Northern Andes of Colombia. The lithology of the Sierra Nevada de Santa Marta holds evidence of an orogenic cycle where the accretion of allochthonous terranes to NW Gondwana played a major role during Pangaea assembly. These new data provide constraints on the timing of igneous activity, the metamorphic peak during Pangaea assembly, and the later fragmentation of Western Pangaea within the western Tethys supercontinental cycle, which led to decompression and crustal thinning.

Here, we focus on the HT-HP Permo-Triassic units associated with the Inner Santa Marta Metamorphic belt, and Permian to lower Jurassic meta-sedimentary rocks of a fore-arc basin associated with the two-stage evolution of a colliding Permian arc and later arc migration into the continental domain (Fig. 2 & 3). Our aims are: 1) to redefine the stratigraphy of the provinces that make up the Sierra Nevada de Santa Marta: the Sevilla Metamorphic Belt, the Inner Santa Marta Metamorphic belt, and the Outer Santa Marta Metamorphic belt, 2) To constrain the age of the metamorphic events that occurred along the continental margin during crustal thickening that led to thermal re-equilibration during the Pangaea assembly. 3) To estimate the P-T conditions operating during the metamorphic phases, and 4) to propose a kinematic model for the tectonic evolution of the NW margin of Gondwana from the early Permian to the lower Jurassic, including terrane accretion, HT-HP metamorphism, and late decompression during extension coeval with deposition of volcanoclastic rocks in a fore-arc basin.

2. GEOLOGICAL SETTING

2.1 The Massif of the Sierra Nevada de Santa Marta

The SNSM integrates, along with the Mérida Andes, the Perijá range, and the Santander Massif, the Maracaibo block. The triangular outline of this block relates to two intracontinental border faults composed of the sinistral Santa Marta-Bucaramanga on its western side and the dextral Oca-Ancón fault on its northern side. Since the Pliocene, the Maracaibo block is separated, additionally, from stable South America by the dextral Boconó fault. The Santa Marta-Bucaramanga and Oca faults meet at the western tip of the SNSM without extending into the continental slope of the Caribbean basin, suggesting they are limited to the continental crust and accommodated a northwest-directed escape, supposing a basal detachment, as envisioned by an “orogenic float” model (Audemard & Audemard, 2002).

Along with the Mérida Andes, the Perijá range, and the Santander Massif it integrates the Maracaibo block that is separated, in turn, from stable South America by the Boconó fault (Fig. 1). In the Maracaibo block, the SNSM occupies the most NW position and is separated from the Perijá range (PR) by the Cesar-Ranchería basin, which is filled with Cretaceous to Neogene marine to continental sediments that rest unconformably above a basement that varies between Precambrian granulites to Jurassic volcanoclastic and sedimentary rocks. The massifs that belong to the Maracaibo block like the SNSM and PR are fragmented as a consequence of a complex tectonic history resulting from a Jurassic Pacific slab rollback, crustal extension, and late Cenozoic

transpressive tectonics. Within these massifs the record of Jurassic back-arc tectonics is preserved. Extensional tectonics gave origin to the Central Atlantic Magmatic Province (CAMP), which followed the orogenic collapse of the Ouachita-Appalachian-Variscan mountain range during Early Triassic Pangaea break-up. The supercontinental cycle transition from the Alleghanian orogenic mountain building and collapse to Pangaea rifting to drifting can be studied in the NW metamorphic belt of the SNSM massif, where a thick sequence of Permian to Cretaceous meta-pelites and meta-volcanic rocks are juxtaposed against Precambrian basement through an upper Paleozoic suture zone (Fig. 2 & 3).

In its internal part, the crustal organization of the SNSM approximates a continent-ward dipping monocline that bends, on meeting the Caribbean crust at its western tip, into a steeply NW-dipping attitude. Above this bend, Caribbean extrusive-intrusive sequences are obducted onto crustal units of continental affinity.

By its well-defined structural setting the SNSM exposes a complex crustal profile, comprising a suite of highly mobilized high-grade sequences at its northwestern corner or base, followed by an undisturbed Meso-proterozoic basement, overlain, in turn, by unmetamorphosed Jurassic extrusive-intrusive sequences, that are capped by Cretaceous platform sediments.

These successions have been grouped into morphotectonic provinces, according to metamorphic grade and age, deformation style and lithological associations (Tschanz et al., 1974). The southwestern Sierra Nevada Province comprises a basement composed of the meso-proterozoic Mangos Granulite, which is succeeded, above a major hiatus, by Lower Jurassic volcanic and volcanoclastic sequences and plutons of a Jurassic arc. This province is separated from the Sevilla Belt by the Sevilla Lineament,

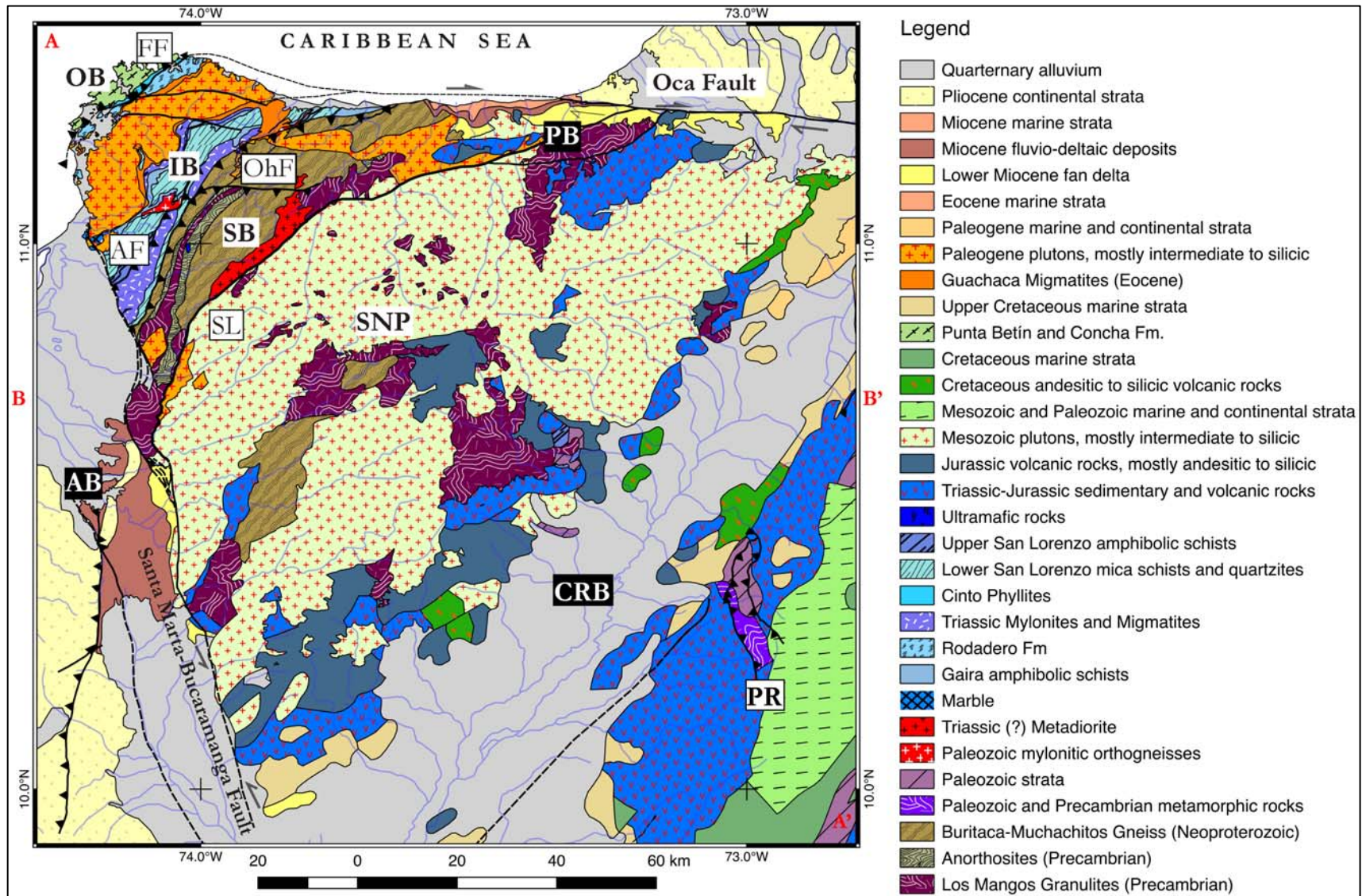


Figure 2 Geological Map of the Sierra Nevada de Santa Marta Massif. SNP: Sierra Nevada Province, SB: Sevilla Belt, IB: Inner Santa Marta Metamorphic Belt, OB: Outer Santa Marta Metamorphic Belt. AB: Aracataca Basin, PB: Palomino Basin, CRB: Cesar Rancheria Basin. Faults: SL: Sevilla Lineament, OhF: Orihueca Fault, AF: Aguja Fault, FF: Florin Fault. Updated from Ingeominas (2007a).

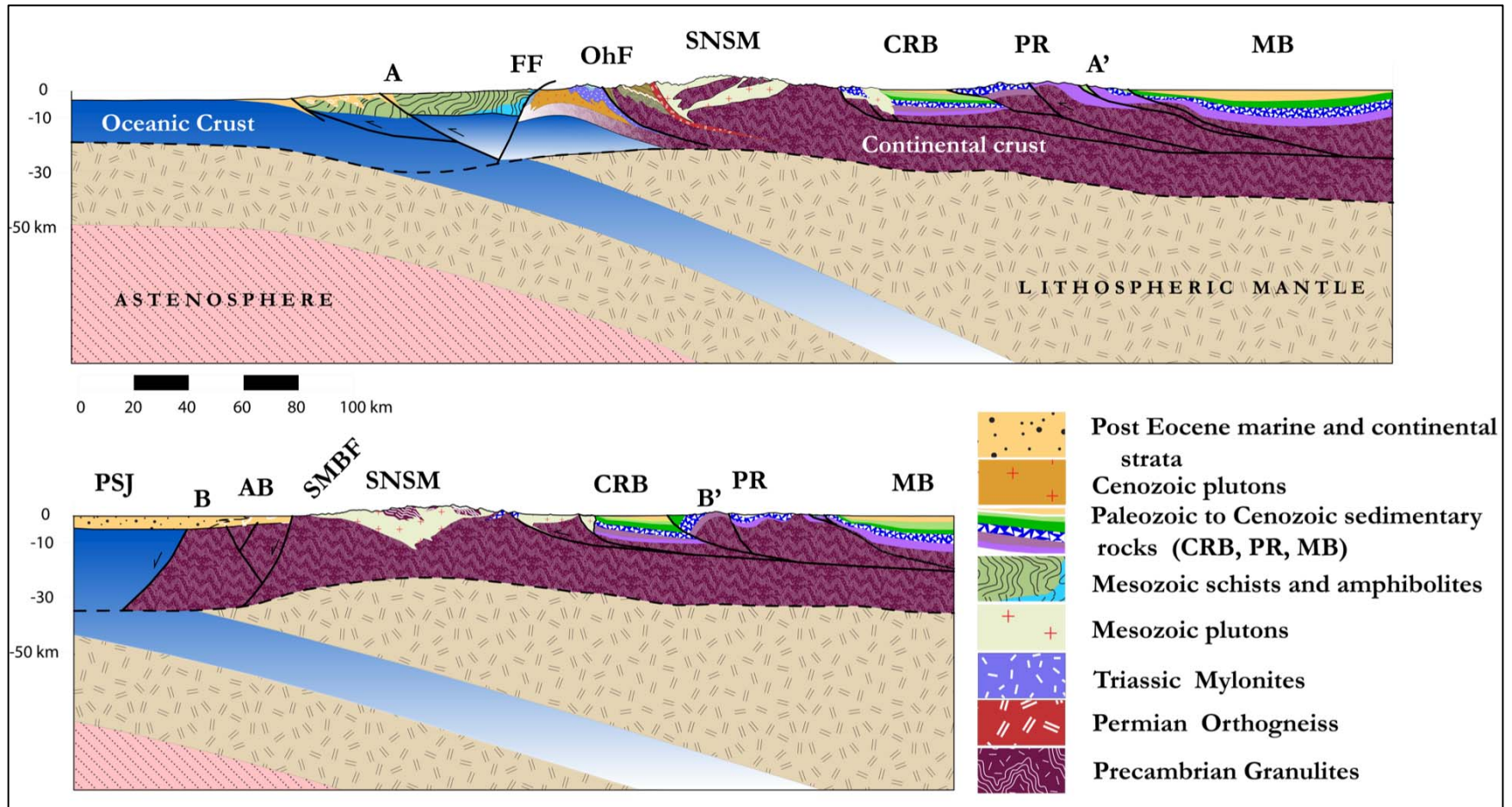


Figure 3. Regional cross sections of the SNSM, locations in fig 2 labeled by letters A-A' and B-B'. AB: Aracataca Basin, PB: Palomino Basin, MB: Maracaibo Basin, PSJ: Plat6-San Jorge Basin, CrB: Cesar Rancheria Basin, SMBF: Santa Marta Bucaramanga Fault, SP: Serranía del Perijá, AF: Aguja Fault, OhF: Orihueca Fault, FF: Florin Fault. Moho depth is the dashed line (ca. 16-25 km) constrained from inversion of gravity, magnetic data and integration with seismological data (Camargo, 2014; Sanchez and Mann, 2015).

that is intruded by elongated stocks of inferred late Paleozoic, Early Jurassic and Eocene ages (Tschanz et al., 1974; Cardona et al., 2010), attesting thus to repeated periods of magmatic activity.

The Sierra Nevada Province (Tschanz et al., 1974; Cardona et al., 2010) bounded to the northwest by the Sevilla Lineament is relatable to the Sunsas-Putumayo province that was juxtaposed against the South American Craton, during the amalgamation of Rodinia (Ibanez-Mejia et al., 2011). Several metamorphic episodes affected the NW part of Gondwana, with a major Neoproterozoic HT-HP metamorphic event (Ordóñez et al., 2002; Restrepo-Pace et al., 1997). These Precambrian granulitic and gneissic terranes compose the NW part of Gondwana and were isolated from other continental blocks by the Iapetus Ocean between Laurussia and Gondwana, and subsequently by the Rheic Ocean (Nance et al., 2012).

Most of the Sierra Nevada Province is extensively affected by Jurassic arc plutons and covered by Triassic and Jurassic volcanic deposits (Figs 2 & 3), which are more evident on the SE flank of the massif in the vicinity of the Cesar-Rancheria basin, where bimodal volcanic flows and associated felsic batholiths alternate with siliciclastic and carbonate sequences.

The Sevilla Metamorphic Belt (SB), albeit of its lithologic diversity, has not been differentiated into tectonic units. The SB contains medium to high-grade metamorphic mafic gneisses of a supposed Paleozoic age (Cardona et al., 2006; Tschanz et al., 1974; Zuluaga and Stowell, 2012).

The metamorphic assemblage at the northwestern corner of the SNSM composes the Santa Marta province (Tschanz et al., 1974) and can be subdivided into two metamorphic belts, which are separated by the Eocene Santa Marta Batholith (SMB).

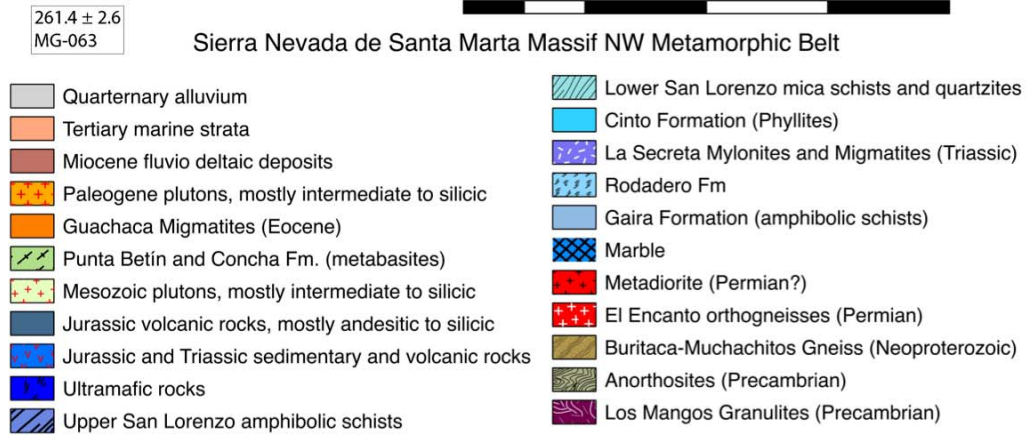
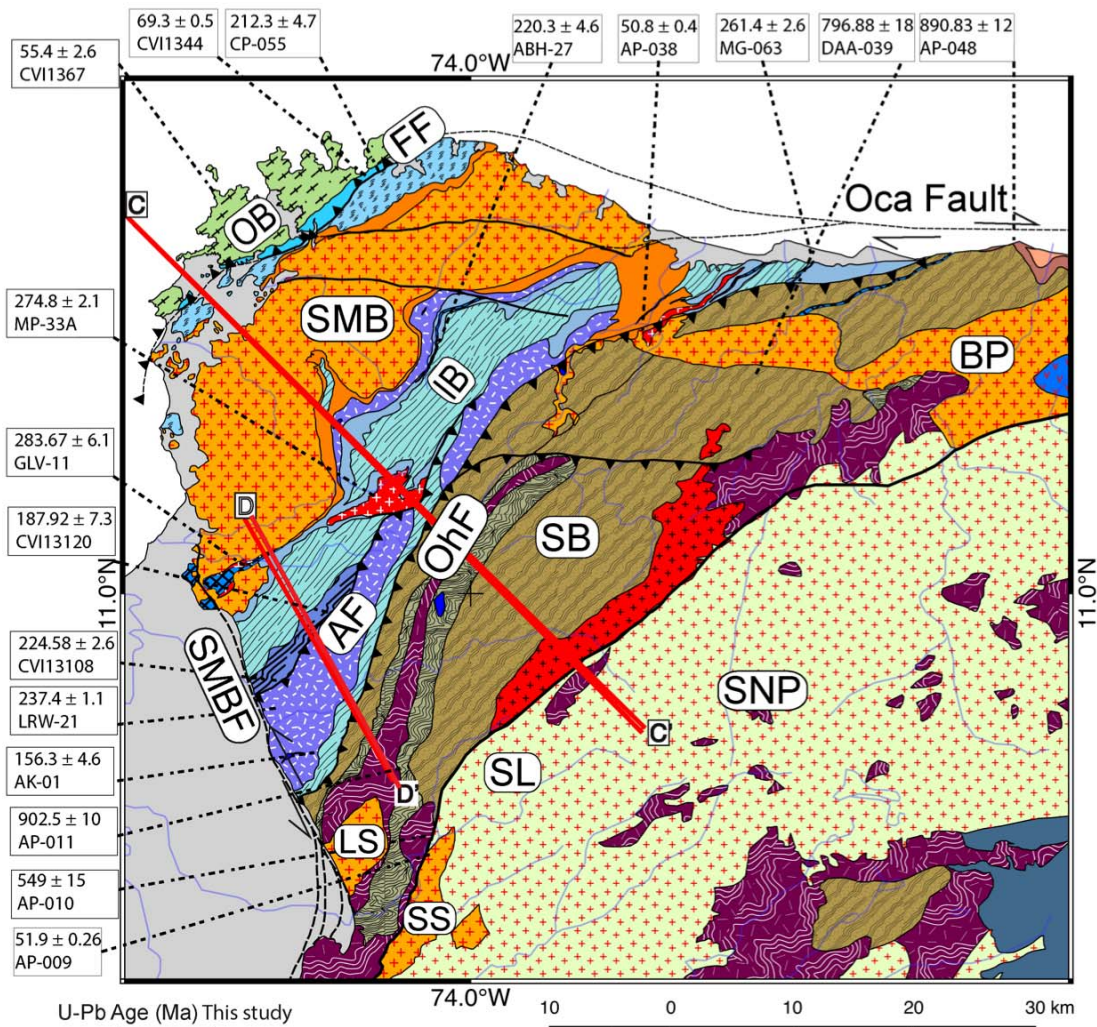


Figure 4. Geological Map of the Sierra Nevada de Santa Marta Massif (SNSM). C-C' and D-D' locates cross sections of figures 5 & 6. OB: Outer Santa Marta Metamorphic Belt, SMB: Santa Marta Batholith, IB: Inner Santa Marta Metamorphic Belt, SB: Sevilla Metamorphic Belt, SNP: Sierra Nevada Province, SL: Sevilla Lineament BP: Buritaca Pluton, SS: Sevilla Stock, LS: Latal Stock, SMBF: Santa Marta Bucaramanga Fault, AF: Aguja Fault, OhF: Orihueca Fault, FF: Florin Fault.

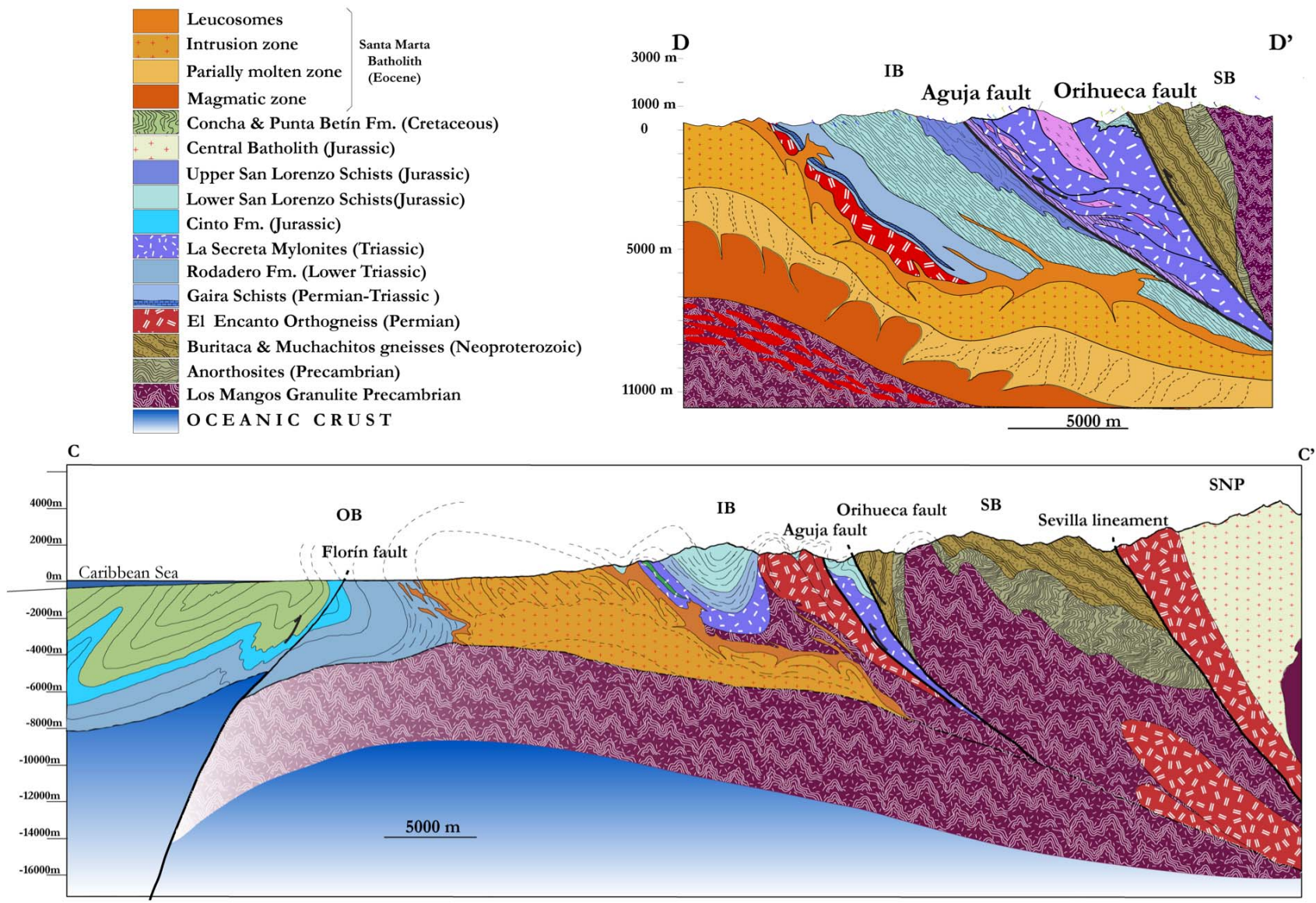


Figure 5. Sections C-C' and D-D' illustrating the structural style of the SNSM. OB: Outer Santa Marta Metamorphic Belt, IB: Inner Santa Marta Metamorphic Belt, SB: Sevilla Metamorphic Belt, SNP: Sierra Nevada Province.

The inner Santa Marta Metamorphic Belt (IB; Figs. 2 & 4) contains epidote-amphibolite to high amphibolite facies rocks from the Gaira and San Lorenzo schists and mylonites that overly Permian granitoids and are in turn intruded by Cretaceous-Paleogene granitoids. The Outer Santa Marta Metamorphic Belt (OB; Fig. 2) is defined by greenschist to epidote amphibolite facies meta-pelites and meta-volcanic rocks of Jurassic-Cretaceous age from the Rodadero, Cinto, Punta Betín and Concha formations that constitute the NW verging overturned flank of an antiform (Doolan, 1970), nucleated by the SMB (Figs. 3 & 5).

3. METHODS

3.1 LA-ICP-MS U-Pb Geochronology

Tracing the timing of metamorphism and igneous activity of the morphotectonic provinces of the SNSM was attempted by LA-ICP-MS U-Pb dating. U-Pb dating was done in conjunction with the morphological analysis of zircon crystals with cathodoluminescence imagery. Coherence of results was checked by integrating age data with other geologic information, such as field relationships, mineral paragenesis, and stratigraphy.

Zircon U-Pb ages were obtained for 17 samples. Zircons were extracted using conventional mineral separation techniques, including rock crushing, sieving (250-60 μm fraction), and mineral concentration was carried out by means of heavy liquid and magnetic separation. Detrital zircons (50 to 200 grains per sample) were manually picked, selecting clean, crack and inclusion-free grains. Subsequently polished grain epoxy mounts were coated with graphite and imaged at the Institute of Geochemistry

and Petrology of the ETH at Zürich, Switzerland using a Vega ©Tescan MV2300VP SEM, to create a detailed set of panchromatic cathode-luminescence images. Laser ablation spots were selected, in both grain cores and rims.

LA-ICP-MS were carried in zircons ablated with a NewWave UP-193nm Ar-F excimer ablation system coupled to a PerkinElmer PE SCIEX Elan 6100 ICP-MS to measure Pb/U and Pb isotopic ratios. The following parameters were applied during this process: 30 µm diameter beam size, 10 Hz repetition rate, 30–45 second signal and a beam intensity of 2.2–2.5 J/cm². Reproducibility of U-Pb data was monitored by measurements of GEMOC GJ-1 CA-ID-TIMS ²⁰⁶Pb–²³⁸U age of 608.5 ± 0.4 Ma (Jackson et al., 2004), used as a primary standard. The external reference standards Plešovice 337.13 ± 0.37 Ma (Sláma et al., 2008), and TEMORA 1 416.75±0.24 Ma (Black et al., 2003), were used to calibrate and monitor fractionation and consistency in the measured U–Pb dates. Ages were calculated using LAMTRACE (Jackson, 2008). Additional data reduction details are given by Ulianov et al. (2012). Statistical analyses of zircon data were performed using Isoplot 3.71 (Ludwig, 2003). All discordant (> 1–3%) analyses of magmatic zircons were discarded. Only zircons with concordance greater than 90% were accepted and plotted. Analytical results are presented in Table A1, Trace element compositions are presented in Table A2. Zircon trace element data obtained by LA-ICP-MS U-Pb were normalized to primitive mantle (McDonough and Sun, 1995).

3.2 Microprobe analyses

Mineral chemistry has been studied by microprobe analysis on thin-sections, using a JEOL JXA-8230 microprobe at the ISTERre, Université Grenoble Alpes, with 15 kV acceleration potential and 10 nA beam current. Representative analyses of minerals are presented in Table A3. Acquisition of the garnet x-ray maps was done on the same microprobe with 15 kV acceleration potential, 100 nA beam current and 300 ms counting time.

3.3 Whole rock geochemistry

Whole-rock analyses were performed at the ALS LTDA labs facilities in Medellin, Colombia and are presented in Table A4. Concentrations of major oxides and trace elements were determined by atomic emission spectroscopy (ICP-AES) and inductively coupled plasma – mass spectrometry (ICP-MS) respectively in accordance with the following procedure: A prepared sample (0.200 g) is added to (LiBO_2 / $\text{Li}_2\text{B}_4\text{O}_7$) lithium metaborate/lithium tetraborate flux (0.90 g), mixed and fused in a furnace at 1000°C. The resulting melt is then cooled and dissolved in 100 ml of 4% nitric acid/2% hydrochloric acid. This solution is then analyzed by ICP-AES and the results are corrected for spectral inter-element interferences. Oxide concentration is calculated from the determined elemental concentration and the result is reported in that format. If required, the total oxide content is determined from the ICP analyte concentrations and loss on Ignition (L.O.I.) values. A prepared sample (1.0 g) is placed in an oven at

1000°C for one hour, cooled and then weighed. The percent loss on ignition is calculated from the difference in weight.

We analyzed our results with Ancomp (Röhr, 1997), a DOS program that compares bulk rock analyses finding the most similar values in a database of 3591 analyses of igneous, sedimentary and pyroclastic rocks. We complemented this dataset by compiling published geochemical data from the Northern Andes (Bosch et al., 2002; Bustamante et al., 2012; Cardona et al., 2010b, 2010c; Cochrane et al., 2014a, 2014b, 2014b; Colmenares, 2007; John et al., 2010; López et al., 2014; Mamani et al., 2010; Riel et al., 2013; Spikings et al., 2015; Villagómez et al., 2011a; Vinasco et al., 2006) in order to increase statistical confidence. The output table is included as Table A5 with major oxide composition of analyzed samples and a list of the ten samples whose composition are the most similar. The results are calculated in both: absolute difference of wt% and log wt% oxides for comparison. Protolith discrimination diagrams are used for evaluating data in a petrographic and field context to obtain a coherent and reliable discrimination of the protoliths of the metamorphic rocks studied in the SNSM.

4. RESULTS

4.1 Field relationships, petrography and mineralogy

4.1.1 Sevilla Metamorphic Belt

The Sevilla Metamorphic Belt SB is a strongly deformed NE-SW striking, and NW verging gneissic complex, some 20 km thick and 75 km long (Figs. 2, 3, 4 &5). The SB is genetically related with the high-grade metamorphic rocks of the Sierra Nevada

Province. It consists of amphibole and quartz-feldspathic gneisses rich in hornblende, biotite, pyroxene, and amethyst, these gneisses commonly exhibit cataclastic textures (Muchachitos and Buritaca gneiss suites), interdigitated with undifferentiated amphibolites, mica-schists, quartzites, diopside marbles and ultramafic pockets (Tschanz et al., 1974). These metamorphic rocks had predominantly vulcano-sedimentary protoliths that experienced high amphibolite to granulite facies metamorphism.

The NE boundary of the SB with the IB has not been clearly defined, and we therefore propose that the high-grade Muchachitos and Buritaca gneiss, should be grouped into the SB, these high-grade units nucleate an antiform and expose the lower crust. The SB is separated from the SNP by the Sevilla lineament, granulites are common at the SNP and mostly intruded by Jurassic felsic plutons. The Sevilla Belt and Inner Santa Marta Belt are separated by The Orihueca fault that thrusts the high-grade complexes of the SB onto the medium-grade complexes (Figs. 5 & 6) of the IB (El Encanto orthogneisses, la Secreta Mylonites, and San Lorenzo and Gaira Schists). This limit is proposed regarding all the available field and geochronological data. Chronologic relationships will be discussed in detail in section (5.1).

4.1.2 Inner Santa Marta Metamorphic Belt

The Inner Santa Marta Metamorphic belt is the morphotectonic province that shows the most prolific lithologies in the entire SNSM, with metamorphic grade ranging from greenschist to high amphibolite facies. The older unit of orthogneisses underlies this suite of parashists. Internally, the IB is further subdivided by the Aguja fault that

separates medium grade garnet-sillimanite bearing mylonites to the SE from mica-amphibole-epidote schists to the NW (Figs. 5 & 6).

El Encanto Orthogneiss

The El Encanto Orthogneiss outcrops are isolated and show lobular and elliptical shapes. The typical texture of these granitoids is coarse-grained with bands composed of amphiboles and plagioclase. Mineral lineation is strongly developed in some locations with L-type tectonic fabrics, with a NE attitude almost parallel to the strike of the regional foliation, which shows ductile deformation in quartz veins arranged in ptigmatic folds (Fig. 7d).

The contact with paraschists is of a non-intrusive nature and clearly suggests that the orthogneisses are the oldest unit underlying the paraschists of the Gaira and San Lorenzo formations. Discrete outcrops of El Encanto Orthogneiss occur in the Valencia creek in the northern foothills of the SNSM (Cardona et al., 2010c), were it had been described but not formally named. New outcrops of this unit have been identified in the Encanto creek, about 10 km to the SW of the Valencia creek, in an intrusive contact with nebulitic diatexites from Eocene migmatization. Furthermore, outcrops exist close to Minca in the very core of the IB at Jueves Santo and Viernes Santo creeks (Figs. 4 & 5). There is a discordant contact with amphibolite, mica schists, and mylonites rich in biotite and garnet. In the vicinity of the Cienaga Marble quarry near the Córdoba River, the El Encanto orthogneiss is in contact with dark green amphibolites of the Gaira schist unit (Figs. 5 & 6).

La Secreta Mylonites

A 45 km long and 3-6 km wide belt of mylonites is present in the hanging wall of the Aguja fault. The mylonites dip in a SE direction that follows the Aguja fault plane attitude. This strongly tectonized basement overrides metasedimentary sequences of the Gaira hornblende schists and San Lorenzo mica schists, and in turn is thrust by Precambrian gneisses through the Orihueca fault (Figs. 5 & 6). The best exposure of these mylonites is found at La Secreta Creek on the western flank of the SNSM (Fig. 7b). Previously, the mylonite belt was included in the SB together with the Buritaca gneiss, and drawn a very diffuse NW boundary on the SNSM map of Tschanz et al. (1974).

Towards the SE boundary of the IB, Hbl, gneisses acquire a migmatitic appearance, as they display layers or lenses of leucocratic segregations set into a relatively homogeneous micaceous matrix. Where leuco- and mesosome produce a relatively continuous banding these migmatite-like rocks obtain a stromatic structure.

Lens-like leucosome segregations and mineral aggregates confer to these gneisses a phlebitic or ophthalmitic structure. Flow structures around felsic mineral aggregates and lenses of mafic fragments are associated with strain shadows related to these competent parts and attest to deformations that accompanied these segregation processes.

Foliation trends parallel to lithologic boundaries and major faults in the IB and dips SE. Subhorizontal to moderately inclined mineral aggregate lineations are ubiquitous in gneisses containing plagioclase porphyroblasts, that constitute the mesosome of this rock association. A top to NE dextral shear sense was locally deduced from, mineral fish, C-bands, strain shadows and mantled feldspar porphyroclasts that define a stair-stepped array and deformation bands.

Mylonitized gneiss can be subdivided into garnet-biotite schists derived from pelites and amphibolites of a probable vulcano-sedimentary origin. These medium grade mylonites show recrystallized quartz with undulate extinction, red biotite fish deviates of the mylonitic foliation around garnet porphyroblasts formed in strain shadows (Figs. 6 & 7c).

The mylonitic zone typically consists of biotite-amphibole gneiss with plagioclase and garnet forming porphyroblasts. Schists include phenocrysts of up to 2 cm of garnet embedded in a biotite matrix, the restite consists of amphibolite (Fig. 7a).

Biotite schist paragenesis consists of Bt+Grt+Qtz+Pl. The amphibole gneiss paragenesis is Hbl+Grt+Pl+Bt+Qtz+Czo+Sph. Amphibole crystals are brown in their core and grade to green with undulated pleochroism at the rim. Close to the Aguja fault, high-grade mylonite paragenesis consists of Sil+Grt+Bt+Qz+Pl+Or+Ms. Kinematic markers such as folds formed in strain shadows, porphyroblasts with stepped wings, and shearbands flanking boudins indicate a dextral shear sense (Figs. 7c). The remobilization under amphibolite facies conditions produced simple shear constriction, which is evidenced by a subhorizontal lineation. Leucocratic bands of Qtz+Pl+Grt are parallel to foliation and evidence a sub-solidus segregation remobilization, resulting in a phlebitic to stromatic metatexite texture. Sheath folds are common in shear zones (Fig. 7b).

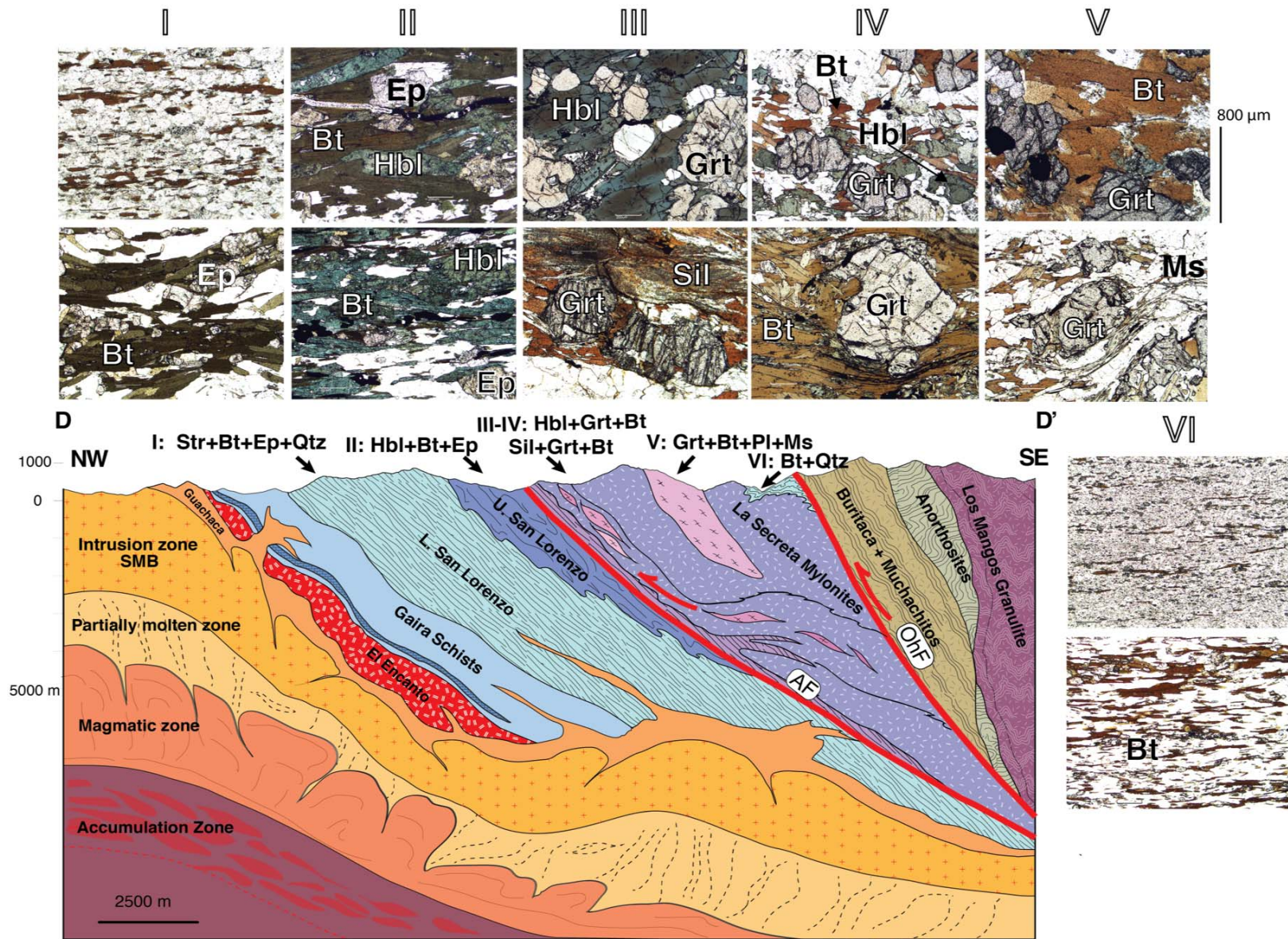


Figure 6. Aguja Creek cross section, unit labels according to the Figure 2. AF: Aguja fault, OhF: Orihueca fault. Top petrographic slides show paragenesis labeled in roman numbers on top of images and with its corresponding location at section, mineral abbreviations are indicated. I: Lower San Lorenzo Schists, II: Upper San Lorenzo Schists, III+IV+V: La Secreta Mylonites, VI: San Lorenzo Schists (quartzites). The magmatic zones at the SMB are adapted after Vanderhaeghe (2009).

Gaira Schists

The Gaira schists defined by Tschanz et al., (1974), consists mainly of dark green amphibolites and schists, including marble levels, biotite and graphite schists and ultramafic rocks. Although in the OB the Gaira Schist has been divided into several different units (Doolan, 1970), for the IB the original definition is preserved. Amphibolite foliation resembles original sedimentary layering of the protoliths and is common to find minor interlayering of muscovite-biotite schists and a thick marble level at the base, and intercalations with schists within the unit.

Gaira schists are unconformably overlain by the metapelitic San Lorenzo schists. At the lower contact with la Secreta Mylonites Bt+Grt+Pl mylonitic schists locally show stromatic metatexite textures.

For the Gaira schists an approximate thickness of ≈ 2000 m of amphibolites interbedded with marble layers and biotite schist is estimated (Figs. 5 & 6). However, the Gaira schists can easily double this thickness when including the mylonitic schists with low partial melting (metatexites) in the lower part of the schists. The mylonites are intruded by the SMB that originated diatexites that include mostly amphibolite restites as paleosomes and xenoliths. The lower contact of the Gaira Schists with El Encanto Orthogneiss is non-intrusive and indicates deposition of the volcanoclastic protoliths unconformably above a pre-existing crystalline unit.

Lower San Lorenzo Schists

The Lower San Lorenzo Schists includes all the pelitic schists where quartzites and Bt+Ms schists are predominant, we distinguish the underlying amphibolite unit with marble content as the Gaira Schists, and the upper hornblende schist unit as the Upper San Lorenzo Schists. Our criterion for such a division is the marked difference in composition represented by amphibole occurrence. The pelitic composition of the protolith is evidenced in the presence of aluminosilicates (Sil+Ky) in the Bt+Ms schists (Zuluaga and Stowell, 2012), and St+Grt in the graphite schists at the base. The Lower San Lorenzo Schists crops out along the road towards La Tagua, capping the San Lorenzo High (Figs. 4 & 5). This unit is also exposed in the Aguja creek and on the road north of the Aguja creek. This area was included originally in the Gaira Schists by Tschanz et al., (1974).

New outcrop localities of the San Lorenzo Schists were found at the hanging wall of the Aguja fault, and in the vicinity of the Orihueca fault, where quartzites are thrust by the Buritaca gneiss (Figs. 3, 4 & 5), quartzites overlay the La Secreta Mylonites dipping towards the SE, extending along strike 27 km and 1.5 km (Fig. 4). Stratigraphic relationships are very complex and interpretation of the units is mainly based on geochronological data of Cardona et al. (2010b), as well as new zircon U-Pb data described in detail in section (4.3). The top of the quartzites abuts against the Orihueca fault (Fig. 3, 4 & 5) in contact directly with amethyst bearing amphibolite gneisses of the Buritaca Gneiss. Interestingly, both units show overlapping folds with parallel orientations that crosscut contacts, indicating that deformation occurred after the coupling of the two units (Fig. 7h & i). Detailed mapping is necessary to establish if an

amphibolite unit exists at the base of the unit, or if the metapelitic sequence rests directly over mylonites and migmatites of the La Secreta Mylonites.

On a regional scope, at the base of the Lower San Lorenzo Formation muscovite-biotite, quartzite schists, and quartzites, are the most common lithotypes with subordinated amphibolite schists. The average paragenesis consists of Bt+Grt+Ep+Ms+Qzt+Pl+Sil+Ky+Sph. However, aluminosilicates are subordinate. Quartzites are gray to brown, containing Qzt+Pl+Bt+Ms, with a paragneisses similar to fine grained schists with high quartz content. Biotite crystals are mostly red and locally green due to metasomatism. Quartzites are interbedded with mica schists and graphite schists that contain Grt+St+Qtz and carbonate segregation veins parallel to S1. The Bt+Ep+Pl+Qtz schists in the middle part of the Lower San Lorenzo Schists (Figs. 5 & 6) exhibit a very strong subvertical L-type tectonic lineation fabric, which is easily recognizable in the Aguja Creek, indicating a lengthening shear orthogonal to the foliation strike (Fig. 7e & g). Schists and quartzites locally present meta-andesite sills with phenocrysts in a matrix in a porphyritic relict texture. These sills contain relict Hbl, which is absent in the other lithotypes. Along with ductile deformation of boudinage in some locations, kink bands are developed within the platy minerals. Homogeneous quartzite beds are common in the middle part of the unit.

Upper San Lorenzo Schists

The best exposure so far of the Upper San Lorenzo Schists is at the tributaries that feed the Aguja Creek from SE to NW (Fig. 4). In this section, the unit is composed of hornblende schists, amphibole schists, biotite-muscovite schists interbedded with rare quartzite beds, in contrast to the more metapelitic character of the Lower San Lorenzo

Schists. The lower part of the schists exhibits the development of boudinage coeval with a very persistent mineral elongation lineation. The contact with the Lower San Lorenzo Schists is marked by the occurrence of staurolite and aluminosilicates. The paragenesis of the Upper San Lorenzo Schists consists of Hbl+Bt+Ep+Pl+Qtz+Sph. Biotite is predominantly green, and is found with Hbl+Ep (Fig. 6). Feldspars exhibit a myrmekitic texture and epidote prophyroblasts are rather common. La Secreta Mylonites is thrust onto the Upper San Lorenzo epidote-hornblende schists through the Aguja fault (Figs. 5 & 6). Both units show very different deformation styles and metamorphic grades (Fig. 7e & f).

Guachaca Migmatites

Migmatites were caused by the intrusion of the Santa Marta Batholith granitoid suite into the units of the IB and the OB. The broad NE-SW oriented migmatite zone intrudes a sequence of mica and amphibole schists from the Gaira Schists. The best exposures on the N flank of the SNSM are located in the Guachaca River, El Encanto and El Rumbon creeks and in the central part of the IB, in the vicinity of Minca at Pozo Azul. The migmatites are always exposed in contact with the lower part of the Gaira Schists. At the Piedras River location, migmatization involved marble beds and caused skarn type contact metamorphism. To the SW the exposition of the migmatite zone is discrete and the SMB intrudes a thick sequence of marble from the Gaira Schists, with frequent mica-schist, amphibolite and orthogneiss xenoliths.

Migmatites are classified as nebulitic diatexites and are present in the border zone of the SMB. These rocks show schlieren and nebulitic textures, and eventually grade into a zone showing meso-metric leucosomes and syn-migmatitic foliation. In the more

distal parts of the pluton, clusters of granitic tabular dikes and sills crosscut the Gaira Schists in the Cinto creek at the OB. This phenomenon is observed around most of the Paleogene plutons, where they intrude the more fertile low- and medium-grade metamorphic rocks, but it is absent at the contact between the Paleogene plutons and the high-grade metamorphic gneisses of the SB, in which case a hornblende zone marks the contact instead.

4.1.3 Outer Santa Marta Metamorphic Belt

The Outer Santa Marta Metamorphic belt (OB) consists of a sequence of mainly greenschist to epidote-amphibolite facies rocks, cropping out with a regional SE dip direction. This metamorphic suite is limited to the Southeast by the SMB and internally divided into at least two tectonic blocks by the Florin fault (Figs. 4 & 5). The metamorphic suite is grouped as the Santa Marta Schists (Tschanz et al., 1974), but can be subdivided as the Concha, Punta Betín, Cinto and Rodadero formations, based on lithological affinities and metamorphic grade (Doolan, 1970; MacDonald et al., 1971).

Rodadero Formation

The Rodadero Formation (Doolan, 1970) corresponds to S0 layered amphibolites and amphibolitic schists with Hbl+Pl+Di+Bt+Ep+Czo, alternated with quartz-feldspathic Grt+Bt+Ms schists. Inherited isoclinal folds with vertical axes show two distinctive mineral lineations, NE verging dextral faults sub-perpendicular to the foliation, and NW verging normal faults that cut in a low angle the foliation affect dikes and the host

rock. A main characteristic of this unit is that is commonly affected by felsic dikes that cut through foliation in an oblique angle and quartz segregations parallel to S1, these melts are genetically related with the Santa Marta Batholith. The Rodadero and Cinto formations constitute the NW verging overturned flank of an antiform, which is nucleated by the SMB. Thermobarometry studies of the Rodadero Formation indicate pressures of 7.6 -9.5 Kb and temperatures of 565-665°C in amphibole composition is concordant with Dalradian type metamorphic belt (Bustamante et al., 2009), in contrast to 5.5 Kb and 529°C for the Cinto Formation phyllites (Cardona 2008b). The stratigraphic relationship between the two formations has therefore been interpreted as a tectonic contact through the Florin fault, because of the structural relationships and the inverse metamorphic grade.

Cinto Formation

The Cinto Formation is about 700 m thick and dips to the SE, internally the Cinto formation presents ESE verging folds. Differently from the Concha and Punta Betín formations it includes quartzites, graphite phyllites, and pschites containing Qtz+Grt+Ms+Bt. Several pschite levels with pebble-sized porphyroclasts. The layered phyllites contain regularly dissected hinges of isoclinal folds and quartz lenses parallel to S1. Graphite schists are frequently interbedded with phyllites that show relicts of cross-bedding in the more siliceous beds that indicate an overturned position.

Concha and Punta Betín formations

The phyllite units at the NW tip of the SNSM consist mainly of an array of folded slivers of oceanic crust derived material that varies in metamorphic grade from zeolite facies through greenschist facies in the coastal range (Concha Formation meta-volcanic rocks), and grade into the amphibolite facies schists of the Punta Betín Formation. Protholites of these units have a mixed oceanic and island arc signature (Cardona et al., 2010b). The base of the sequence is composed of volcanoclastic breccias with copper and calcopyrite and sulphide nodules interdigitated with Chl+Act carbonate schists, graphite schist, quartz-phyllites, derived from pelites and tuffs of basaltic to andesitic composition (Bustamante et al., 2009) and subordinated meta-gabbros (Hbl+Pl), and andesites in dykes and sills. The presence of syn-depositional folds, slumps and relict olistostroms with lenticular geometries and an oblique cleavage evidence slope deposits associated with tectonic activity. Towards the top of the Punta Betín Formation meta-gabbros, tuffs and basalts are interbedded with exhalative siliceous deposits. Although intensely folded the Punta Betín Formation preserves a regional SE dip. In the Actinolite schists, the main metamorphic foliation S1 cuts S0 in a low angle, but with the same strike. S1 exhibits a penetrative cleavage that defines a crenulation lineation L2. L2 crosscuts in an angle of 70° an inherited L1 lineation parallel to the regional shortening. L1 is associated to a first deformation phase of lengthened fabric preserved in hinges. L2 was produced in a second phase related to strike slip tectonics (Fig. 7j, k & l). Kinematic indicators in meta-gabbros in the upper part of the Punta Betín Formation close to the contact with the Cinto Formation evidence oblique shortening.

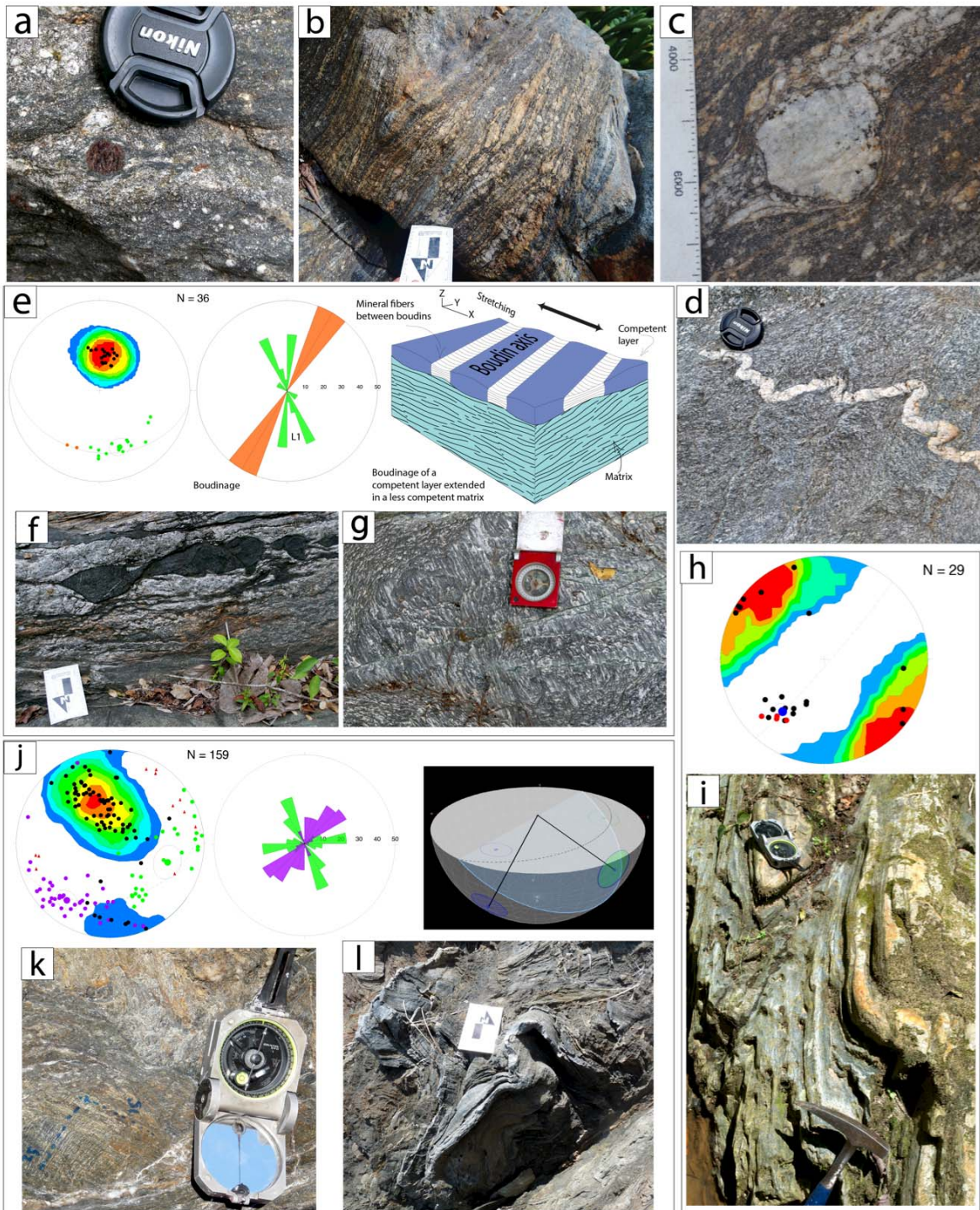


Figure 7. Main structural features of the low to medium metamorphic grade units at the SNSM . La Secreta Mylonites a) Garnet porphyroblast embedded in a Bt+Pl+Qtz matrix, b) stromatic migmatite-like structure with discontinuous leucosome segregations in an amphibolite facies Bt+Grt mylonite. c) Feldspar prophyroblast with pressure shadows showing a weak stair-stepping that indicate a dextral shear sense. d) Hbl + Pl Encanto orthogneiss; vein is affected

by ptygmatic fold. e) Stereographic projection (lower hemisphere) of foliation poles (black, contoured) and a N to northwest trending L_1 mineral aggregate lineations (green) in mica schists of the Lower San Lorenzo Schists with corresponding rose diagram with mineral aggregate lineation (green petals) and stretching direction of boudins (orange petals); inset shows boudinaged competent layer with x-direction oriented parallel to mineral lineation. f) Mafic boudin set off at shear bands and embedded within a Qzt+Pl matrix indicates a sinistral shear; Upper San Lorenzo Schists; g) L-type tectonite from the Lower San Lorenzo Schists, stretching lineation corresponds to L_1 , mineral association consists of Bt+Pl+Qz; h) Stereographic projection (lower hemisphere) of foliation poles contoured; and fold axis plunging SW of quartzites and amphibolites located NE of the Orihueca fault, i) subvertical folds at the Santa Clara creek, this deformation is found in the hanging wall and footwall of the Orihueca fault affecting the Precambrian and Jurassic units supposing that the deformation that originated this folds occurred after coupling of the tectonic blocks of the Sevilla Metamorphic Belt and the IB; j) Compilation of poles to the foliation at the OB, L_1 in green is the oldest lineation from a first folding phase, L_2 transversal to the foliation is related to an overprinted deformation. k) photograph showing the interception of L_1 and L_2 (Concha Formation); l) Asymmetric fold with an superposed event of sinistral transpression related to L_2 (Concha Formation).

4.2 Whole rock geochemistry

We included whole rock geochemistry analysis from nine samples that were also analysed by zircon U-Pb dating (Table A4). These samples were selected from the morphotectonic provinces described above. We focused on rare Earth elements (REE) as well as the relatively immobile High Field Strength (HFS) and transition elements (Ti, Zr, Hf, Nb, Th, Ta, Y, Cr, P, Ni, Sc). Classified at protolith discrimination diagrams (Fig. 8) based on major oxide ratios, and major oxides vs Cr.

4.2.1 High-grade metamorphic rocks from the Sevilla Metamorphic Belt

Major oxide and trace element data were acquired from a mafic migmatite from the Buritaca Gneiss (AP-010) that belongs to the lower calc-alkaline to tholeiitic series (Irvine and Baragar, 1971). The REE chondrite normalized plot shows a flat slope (Fig.

9e). Primitive mantle normalized multi element analysis of this sample shows a negative Th anomaly.

Protolith discrimination diagrams

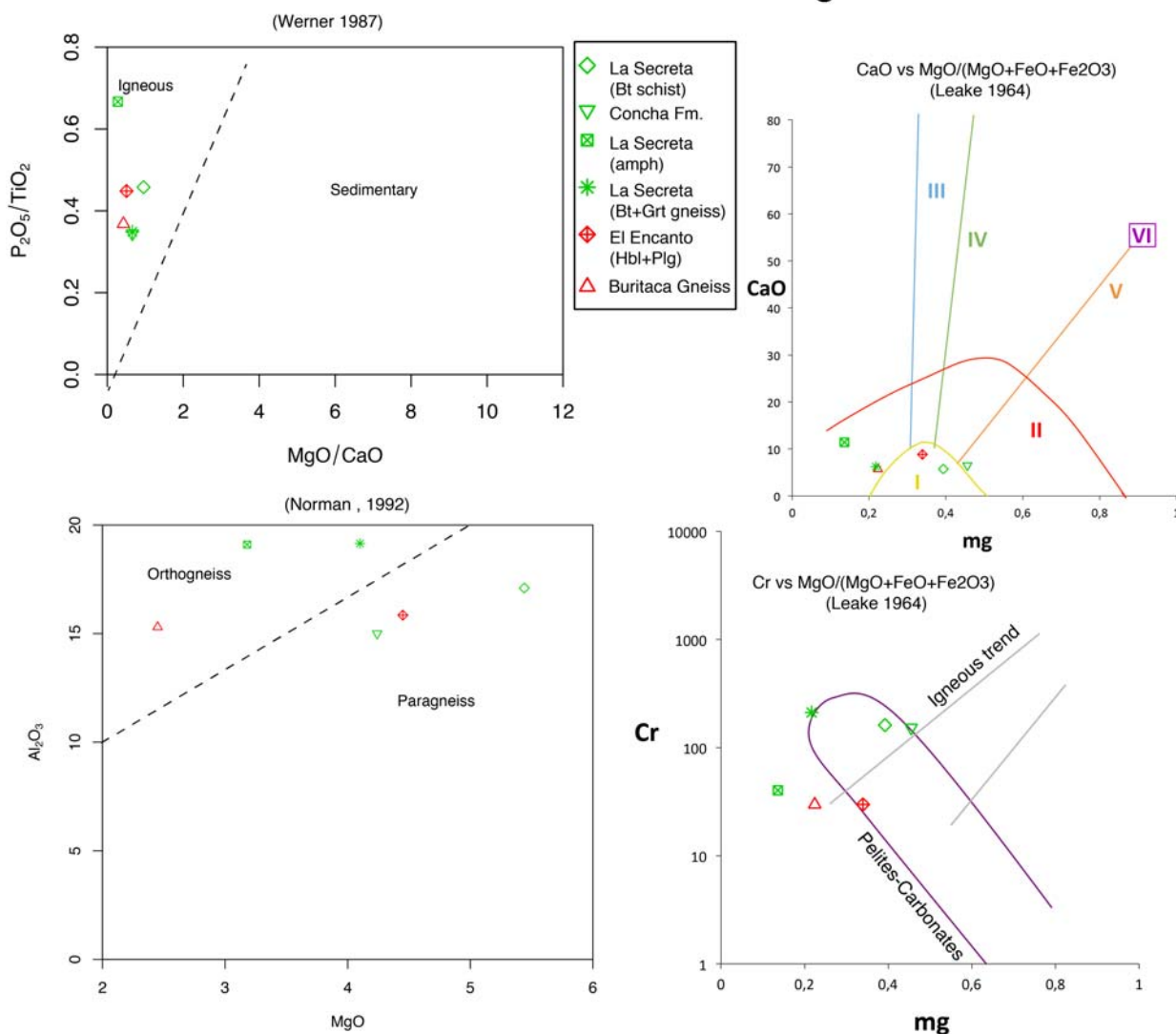


Figure 8. Protolith discrimination diagrams. (top left) discrimination between igneous and sedimentary provenance (Werner, 1987); (top right) Diagram that shows the origin of mafic rocks based on Niggli's numbers (Leake, 1964), $c = CaO$; $mg = MgO / (MgO + FeO + Fe_2O_3)$ for amphibolites and amphibolic gneisses, I. pelites, II. Igneous rocks, III. Calcareous rocks, IV. Mixing zone between limestone and pelites, V. Mixing zone between pelites and dolomites, VI. dolomites.; (Bottom left) Al_2O_3 Vs MgO (wt%) (Norman et al., 1992); (bottom right) Diagram that shows the igneous trend for amphibolites and mafic schists (Evans and Leake, 1960), $Cr = Chrome$; $mg = MgO / (MgO + FeO + Fe_2O_3)$.

4.2.2 Amphibolite-grade rocks from the IB

Four samples from the IB, which include two garnet bearing mylonites, one garnet biotite schist, and one orthogneiss were analyzed obtaining major oxide and trace element data. Samples garnet amphibolites from the La Secreta Mylonites belong in the tholeiitic field, Bt schists from la Secreta Mylonites and sample Hbl+Pl granitoid from el Encanto Orthogneiss correspond to the calc-alkaline series (Irvine and Baragar, 1971), and is classified as an orthogneiss indicated by trace element composition and high Y content as a calc-alkaline gabbro (Fig. 8).

Garnet amphibolites, plot in the field of ultrabasic rocks (Bas et al., 1986; Cox, 1979). Biotite-garnet schists according to the protolith discrimination diagrams corresponds to a volcanoclastic protolith, probably a basaltic tuff or siltstone with an important terrigenous component. Mylonites Pl+Gr+Bt is compositionally a gabbroid (Bas et al, 1986), this corresponds with a protolith discrimination that shows a mixed igneous to pelitic protolith for this sample.

The primitive mantle normalized multi element plot (McDonough and Sun, 1995) of Bt+Grt+Pl schists shows a flat to negative slope with negative Ta, Ti, Eu, Nb, and Sr anomalies. Garnet amphibolites and garnet plagioclase metaigneous sample LRW-21 show a flat profile (Fig. 9). In the mafic orthogneisses (El Encanto) Th/Ta vs Yb and Th/Hf vs Ta/Hf ratios are compatible with an active continental margin environment (ACM) (Schandl and Gorton, 2002). Garnet amphibolites correspond to the boundary between a within plate volcanic zone (WPVZ) field and ACM. Bt+Grt+Pl schists and Pl+Grt+Bt gneisses plot in the WPVZ field (Fig. 9).

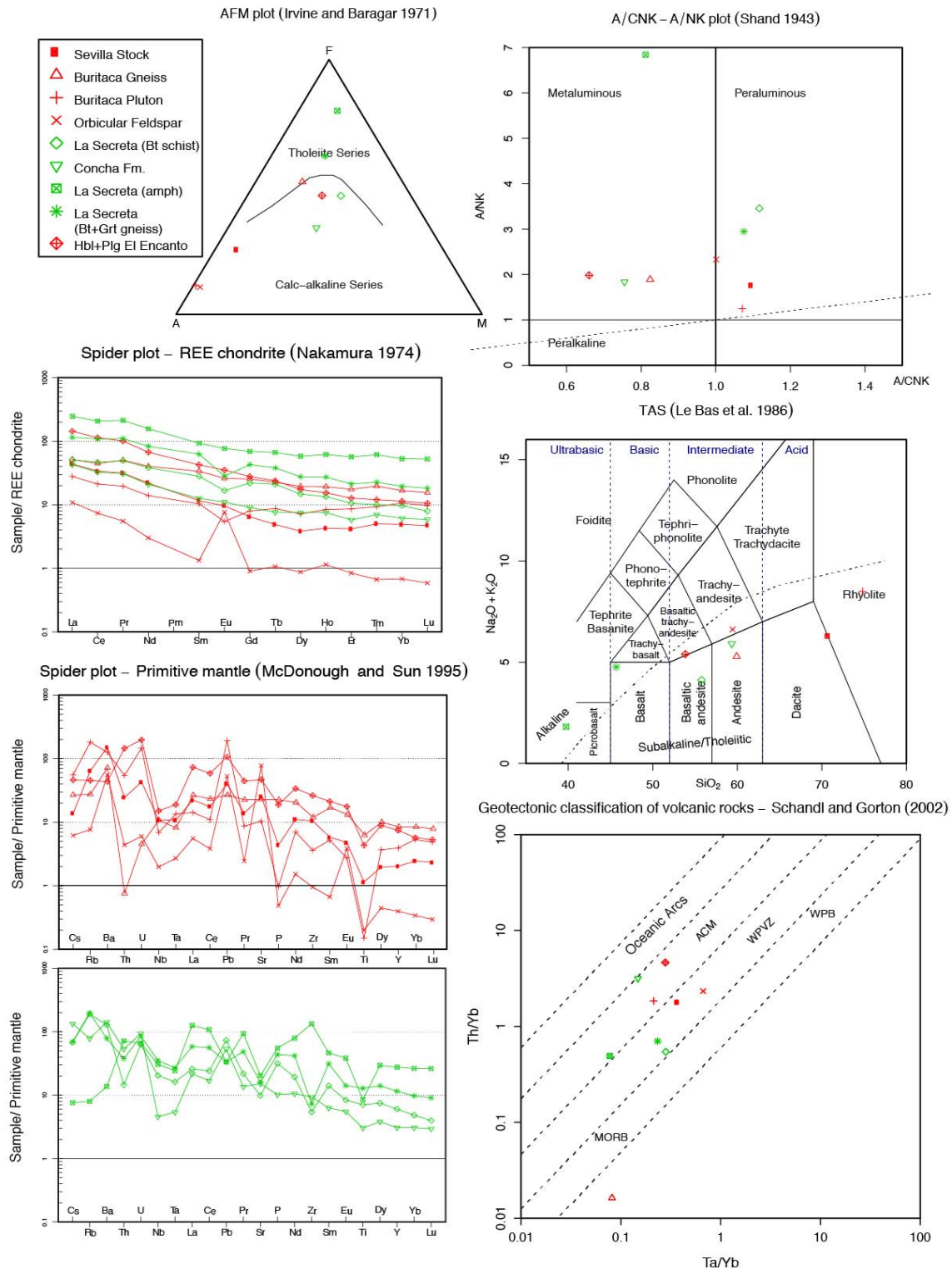


Figure 9. Whole rock geochemistry: Green represents metavolcanic rocks and red samples are related to granitoids(a) AFM diagram (Irvine and Baragar, 1971); (b) A/CNK- A/NK plot (Shand, 1943); (c) TAS diagram (Bas et al., 1986); (d) Geotectonic classification of volcanic rocks (Schandl and Gorton, 2002); (e) REE diagrams normalized to chondrite (Nakamura, 1974) and Primitive Mantle (McDonough and Sun, 1995).

4.2.3 Greenschists from the OB

Major element oxide and trace element data of the lower greenschist facies porphyritic meta-andesite point to a calc-alkaline series protolith. Th/Ta vs Yb ratios indicate an ACM to oceanic arc environment (Schandl and Gorton, 2002).

Primitive mantle and chondrite normalized multi-element plots of this sample show a negative slope with negative Ta-Ti and Nb anomalies (Fig. 9e).

4.2.4 Granitoids associated with the Santa Marta Batholith

Three samples from Eocene granitoids were analyzed. Samples AP-009 and AP-038 correspond to the Sevilla Stock rocks and the Guachaca Migmatites respectively. Sample CJJ-91 is a decimetric orbicular feldspar collected in the central part of the SMB. The Guachaca Migmatite is a very acid granitic intrusion whereas the Sevilla Stock rock resides on the boundary between granite and granodiorite. The orbicular feldspar shows an intermediate dioritic composition (Fig. 9).

Trace element compositions of the Guachaca migmatites and the Sevilla Stock revealed negative Ta, Ti, Nb, and Sr anomalies. The negative Ti anomaly was significantly more pronounced in migmatites. The orbicular feldspar shows a strong positive Eu anomaly, which was not observed in the other two granitoids.

4.3 Zircon Morphology

When it comes to discussing zircons from metamorphic rocks, it is very important to make a distinction between zircons that grew as result of a metamorphic reaction (=

“metamorphic zircons”), and zircons that crystallized from magma and recrystallized during thermal overprints (= “metamorphosed zircons”) in a regional metamorphic context, or because of magma recycling. In the SNSM, we found both metamorphosed and metamorphic crystals along with magmatic zircons. By studying their morphologies and combined core-rim age data, we gained information on the tectonic history of this complex area. Magmatic zircons are mostly euhedral translucent and prismatic with oscillatory zoning. In contrast metamorphic zircons tend to be rounded subhedral, with more complex zoning patterns that include, thick rims, patchy and convolute zoning, and collapsed cores. Although crystals are mostly colourless, as a main feature of the SNSM the Precambrian zircons present a purple hue whereas some Permian zircons present a yellowish hue. The specific traits of the crystals found at the different samples are briefly described in the following paragraphs.

Cathodoluminescence imagery analysis revealed core-rim relations that allow defining at least two episodes of rim growth in metamorphosed zircons, around inherited cores of Precambrian and Paleozoic ages (Fig. 10).

Zircons from the SB samples are subhedral, sub-rounded to sub-angular, and their color varies between translucent to purple and faintly yellow. These Precambrian crystals from high-grade metamorphic granulites, amphibolic and quartz-feldspathic gneisses, have a larger size (>250 μm) in contrast to their younger Permian-Jurassic counterparts. Purple zircons are mostly older than the translucent ones and can be associated with a 1200-1300 Ma age peak. Samples with dominant populations younger than 1200 Ma mostly yield translucent to slightly yellowish crystals, with purplish faint xenomorphic cores.

Inner textures show characteristic recrystallization under high-grade metamorphic conditions, with rims that commonly are brighter and thicker, evidencing a homogeneous composition. Xenomorphic cores occur frequently and present dark purple colors. Inherited cores sometimes show fracturing at the outer rims. Precambrian crystals found in detrital samples (Chapter 3) show the same traits, but the purple colored population is more abundant and again it corresponds to crystals older than 1200 Ma.

Zircons from orthogneisses are euhedral, angular, and prismatic, with oscillatory zoning and sharp edges, with packed, closely spaced rims attesting to a continuous stable growth. Crystals vary between translucent to yellowish colored, and are mostly clean and inclusion free (Fig. 10d). The crystal population is unimodal in terms of morphology and age, average size of the crystals is around 200 μm .

Mylonite zircons are neoformed irregular subhedral to anhedral crystals sub-rounded to sub-angular, and colorless. The typical sub-rounded to multifaceted metamorphic morphology evidence highly reabsorbed shapes, irregular concentric zoning with chaotic texture with local aspects of flow and overprint of new recrystallization (Fig. 10 j & k).

Zircons from the Permian-Jurassic meta-pelites and meta-basites in the IB and the OB show a bimodal distribution with three distinctive populations. The first population is made of typical subrounded metamorphic zircons of Precambrian age and magmatic Ordovician-Jurassic zircons recycled from plutons. The second population found at the metamorphic units, has translucent to yellow Permian euhedral to subhedral crystals. A third population is made of Jurassic tabular prismatic euhedral translucent crystals. Zoning is oscillatory in magmatic crystals and with the core-rim structures common

for high-grade metamorphic granulites and gneisses (Fig. 10 f, l, m & p). Within the metapelitic units a distinctive population of rounded zircons with Precambrian cores and subrounded Permian rims, thicker and brighter than the cores, can be observed with a maximum lower Permian depositional age (Fig. 10e & i), and is seldom found in the younger metapelite units (Fig. 10a).

Crystals from the meta-basites of the OB are mainly euhedral to subhedral, tabular prismatic and acicular, prismatic, colorless, inclusion free and with grain sizes of up to 100 μm (Fig. 10 g & q). These magmatic zircons are always smaller than all the other observed populations.

The Eocene Sevilla Stock contains translucent prismatic, inclusion free zircon crystals with oscillatory zoning in the outer rims, and inherited Precambrian cores with the typical metamorphic texture of Precambrian zircons (Fig. 10n). Zircons from the Guachaca Migmatites are subhedral, showing extensive damage and fracturing with a disturbed oscillatory zoning, abundant inclusions, collapsed cores, and patchy zoning (Fig. 10o).

4.4 Zircon U-Pb Geochronology

Results from 17 samples from the metamorphic units of the NW corner of the SNSM are presented here. One sample from a granitoid of the southern SB, four samples from high-grade metamorphic gneisses of the SB, and eight samples from meta-pelites, mica and amphibole schists, orthogneisses, and granitoids from the IB. Additionally three samples from pschites and meta-basites of the OB, and finally one leucosome

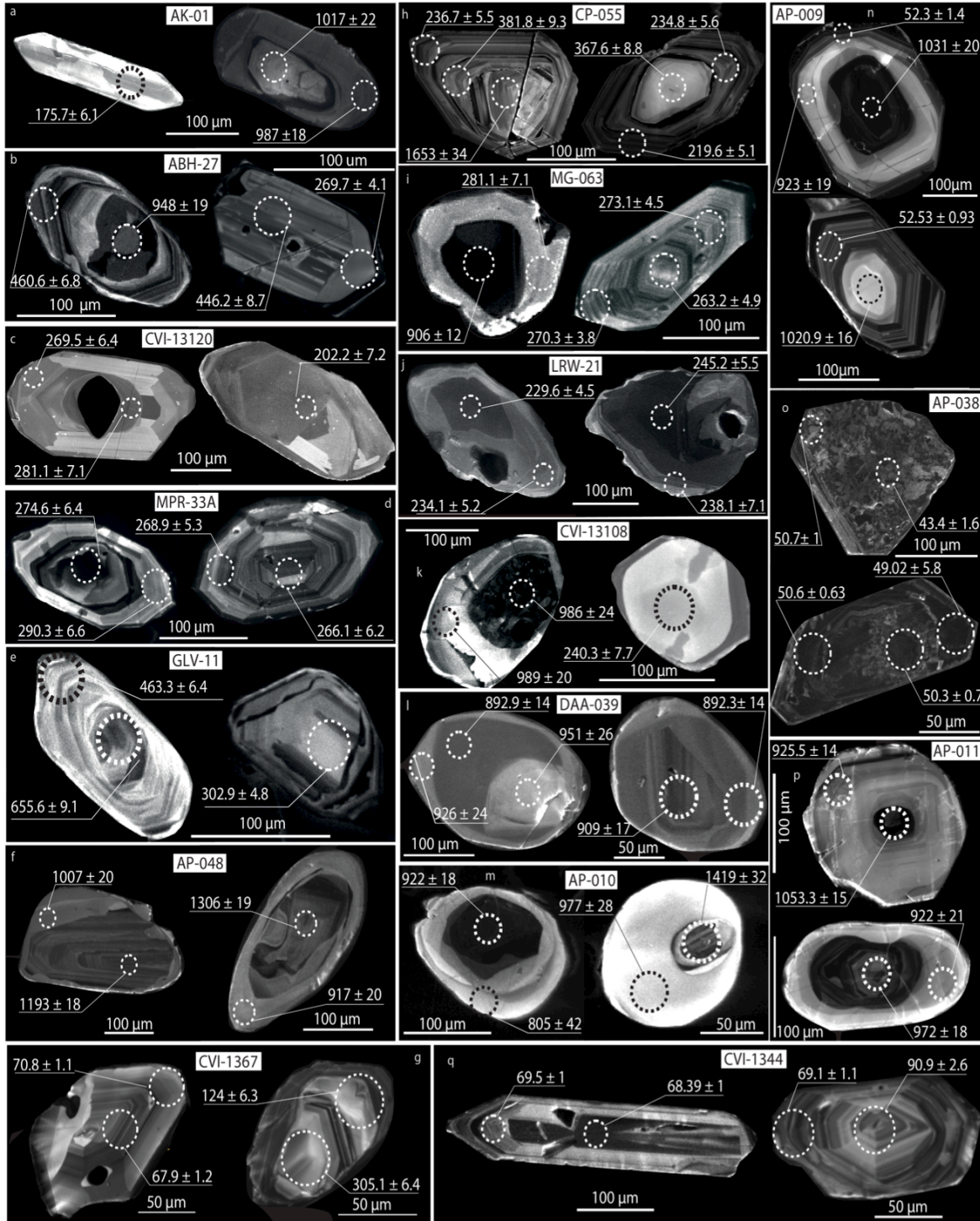


Figure 10. Panchromatic cathodoluminescence imagery of zircons analyzed in this study. a- Lower San Lorenzo Schists, b- Lower San Lorenzo Schists, c- Upper San Lorenzo Schists, d- El Encanto Orthogneiss, e- Gaira Schists, f- Muchachitos Gneiss, g- Concha Formation, h- Cinto Formation, i- Gaira Schist, j & k - La Secreta Mylonites, l & m- Buritaca Gneiss, n- Sevilla Stock, o- Guachaca Migmatite, p- Muchachitos Gneiss, q- Cinto Formation (dike). Dashed circles ablation size diameter is 30 μm.

from the southern SMB. Sample locations and general description is summarized in figures 4 & 5 and table A1, geochronological data are shown in figures 10 to 16.

4.4.1 Sevilla Metamorphic Belt (SB)

Four samples from the amphibolite gneisses within the SB yield concordant 970 Ma ages (Fig. 11 & 12). The Muchachitos Gneiss (Samples AP-048, DAA-039) yields an age spectrum between 900-1300 Ma, and 800-1080 Ma respectively with a rim recrystallization episode close to 944 ± 36 Ma. and youngest grain an ages of ca. 890 ± 12 Ma and ca. 796 ± 18 Ma (Fig. 11 & 12).

The Buritaca Gneiss (Samples AP-011, AP-010) showed similar age trends with an age spectrum between 900-1300 Ma, and a rim recrystallization event at 955 ± 24 Ma. (Fig. 11 & 12) and the youngest crystal at ca. 902.5 ± 10 Ma for the rocks that did not suffered migmatization. However, mafic migmatites from this unit showed a younger overprint with a population of ca. 797-898 Ma distinguished in cores and rims and concordant zircons between 549-772 Ma ages found at cores and rims, the age ca. 549 ± 15 Ma corresponds to a rim and belongs to the youngest of the Precambrian populations identified in this study. Younger veinlets in migmatites provide Phanerozoic crystals with Permian-Jurassic ages ca. 199.8-258.1 Ma (Fig. 11 & 12).

Zircons from the Sevilla Stock show recrystallization at ca. 51.9 ± 0.3 Ma with inherited Precambrian cores (800-1400 Ma) and some 900-1100 Ma rims. A discrete 180-200 Ma population and four Permian crystals completes the sample signature (Fig. 15 & 16).

Precambrian High-grade units

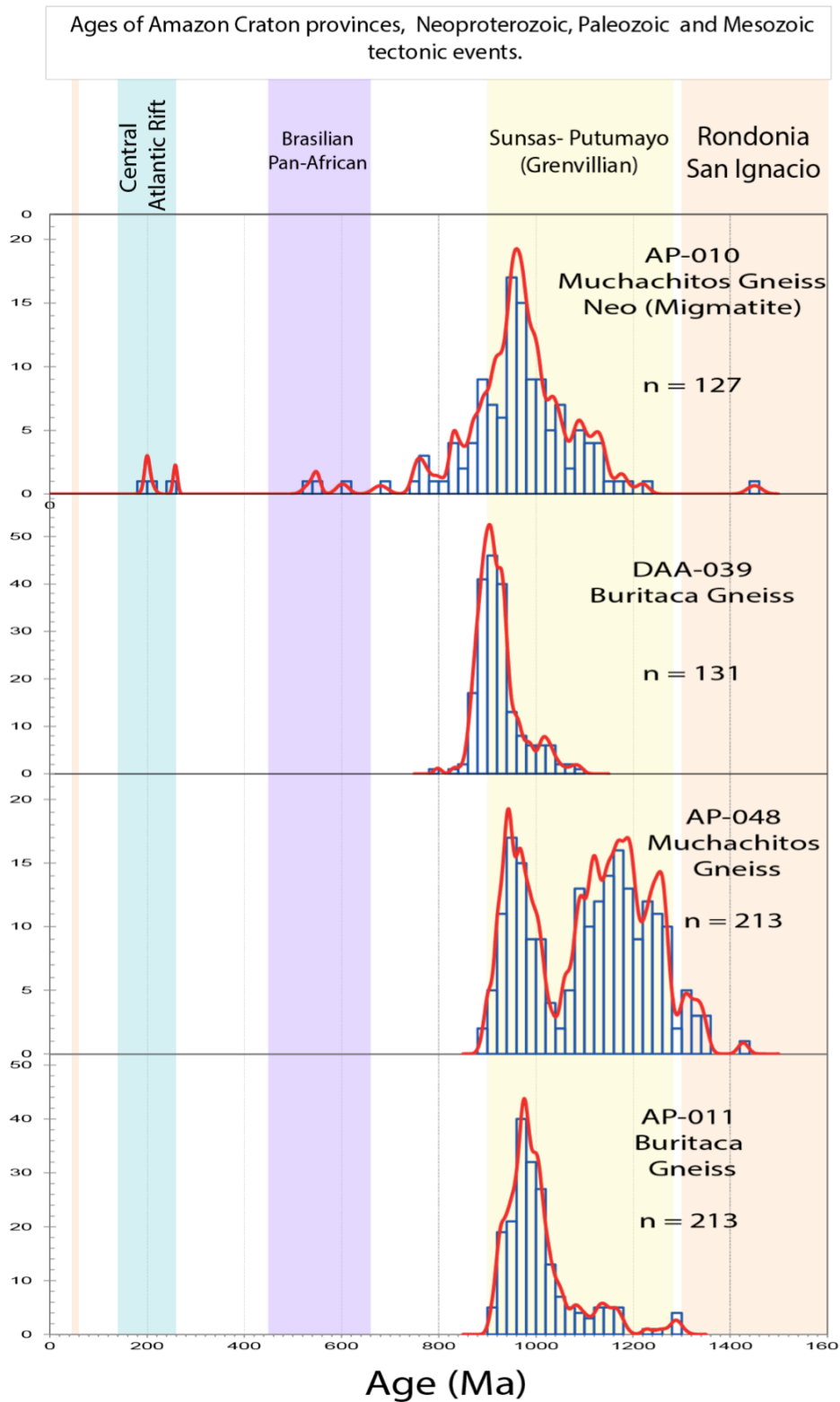


Figure 11. Zircon U-Pb LA-ICP-MS probability density plots obtained from the Precambrian units of the Sevilla Metamorphic Belt.

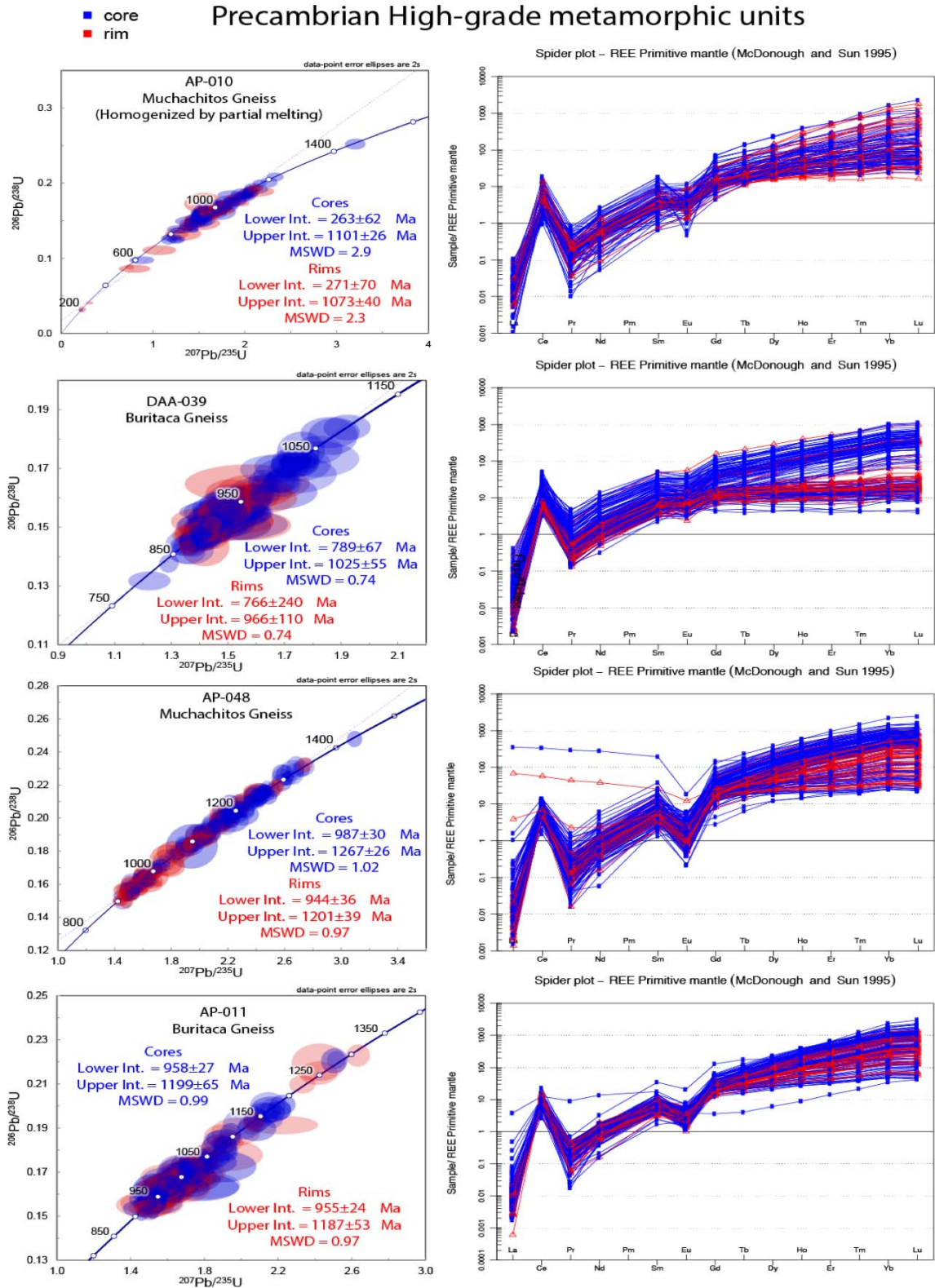


Figure 12. Concordia plots (left) and trace element composition of zircons from the Precambrian units of the Sevilla Metamorphic Belt south of the Orihueca fault.

4.2 Inner Santa Marta Metamorphic Belt (IB)

Eight samples from the SE dipping IB were analyzed. Sample MPR-33A from el Encanto Orthogneiss has an age of 274.8 ± 2.1 Ma, and mica-schist sample GLV-11 yields an age of 283.67 ± 6.1 Ma. Sample MG-063 from the Gaira Schists has an age of 261.46 ± 2.6 Ma. El Encanto Orthogneiss sample MPR-33A shows a discrete Pan-African/Brasiliano age population of 543 ± 14 Ma, an late Pennsylvanian age population between 302-310 Ma, and a dominant Cisuralian to Guadalupian age population between 292-270 Ma (Fig. 13 & 14). Garnet-mica schist sample GLV-11 shows an early Paleozoic population of 284-463 Ma of four crystals. The Permian age is present in a recrystallized rim over a Carboniferous core. Other age populations in this sample include a Pan-African/Brasiliano 522-655 Ma population, a Neoproterozoic population at 900-1200 Ma and a Paleoproterozoic population at 1500-1770 Ma. One zircon yielded an age of 2235 ± 58 Ma (Fig. 13 & 14). Mica schist from sample MG-063 contains age populations at around 950 Ma, 655-850 Ma, and 270 Ma. One zircon yielded an Ordovician age of ca. 468.9 ± 7.1 Ma. In this sample, metamorphic rounded zircons with Precambrian cores and bright homogeneous rims of Permian age are frequent (Fig. 10, 13 & 14). Sample LRW-21 Pl+Grt+Bt from La Secreta Mylonites shows an age range between 230-250 Ma with growth of cores and rims, and a weighted average age of 237.4 ± 1.1 Ma. Bt+Grt+Pl schist sample CVI13108 from La Secreta Mylonites shows an age spectrum of 224 to 1300 Ma, with the youngest crystal age of 224.58 ± 3 Ma, and most cores and rims between 225-250 Ma. Nonetheless, an early Permian age population of 270-285 Ma and a Precambrian inherited population of 1000-1300 Ma were also detected (Fig. 13 & 14). Quartzitic schists sample ABH-27 from the Lower

San Lorenzo Schists contains zircons with Neoproterozoic 563-1189 Ma inherited cores, and Middle Ordovician to middle Carboniferous rim overgrowths and external Triassic rims c.a. 220-243 Ma. Sample CVI-13120 from Bt+Ep schists from the upper San Lorenzo Schists showed a Jurassic age of 187.92 ± 7.6 Ma. In contrast, sample AK-01 from a paraschist of the Lower San Lorenzo unit showed two main age trends with a 1000-1400 Ma population, and a concordant Jurassic 184.96 ± 0.78 Ma population. Leucosome AP-038 from the Guachaca Migmatites yields an age of ca. 50.8 ± 0.4 Ma. This sample contains few inherited Permian and Precambrian zircons (Fig. 15 & 16). Eocene zircons are extensively damaged, and they have inclusions and collapsed cores (Fig. 10o).

4.4.3 The Outer Santa Marta Metamorphic Belt (OB)

Three samples were analyzed from the OB. Sample CVI1367 a meta-andesite of the Concha Formation at the OB yields a lower Paleocene age of 55.4 ± 2.6 Ma. In the hanging wall of the SE verging Florin fault, CP-055 pschist from the Cinto Formation yields a Triassic age of 212 ± 4.7 Ma product of magmatic overgrowth, this sample contains Pan African/Brasiliano zircons at ca. 611- 505 Ma, Devonian to Carboniferous zircons between 402-310 Ma and few Permian grains at 254.1 ± 5.9 Ma and 280.4 ± 6.6 Ma. At the margins of the SMB, the basal part of the Rodadero Formation is affected by migmatization that grades into granitoids, which are also intruding the Cinto Formation phyllites as leucocratic dikes (sample CVI1344) and yield a Maastrichtian age of 69 ± 0.5 Ma ages. One sample of the Eocene Sevilla Stock hosted between the Precambrian Buritaca gneiss and the Triassic to Jurassic western batholith border zone was analyzed (Fig. 15 & 16).

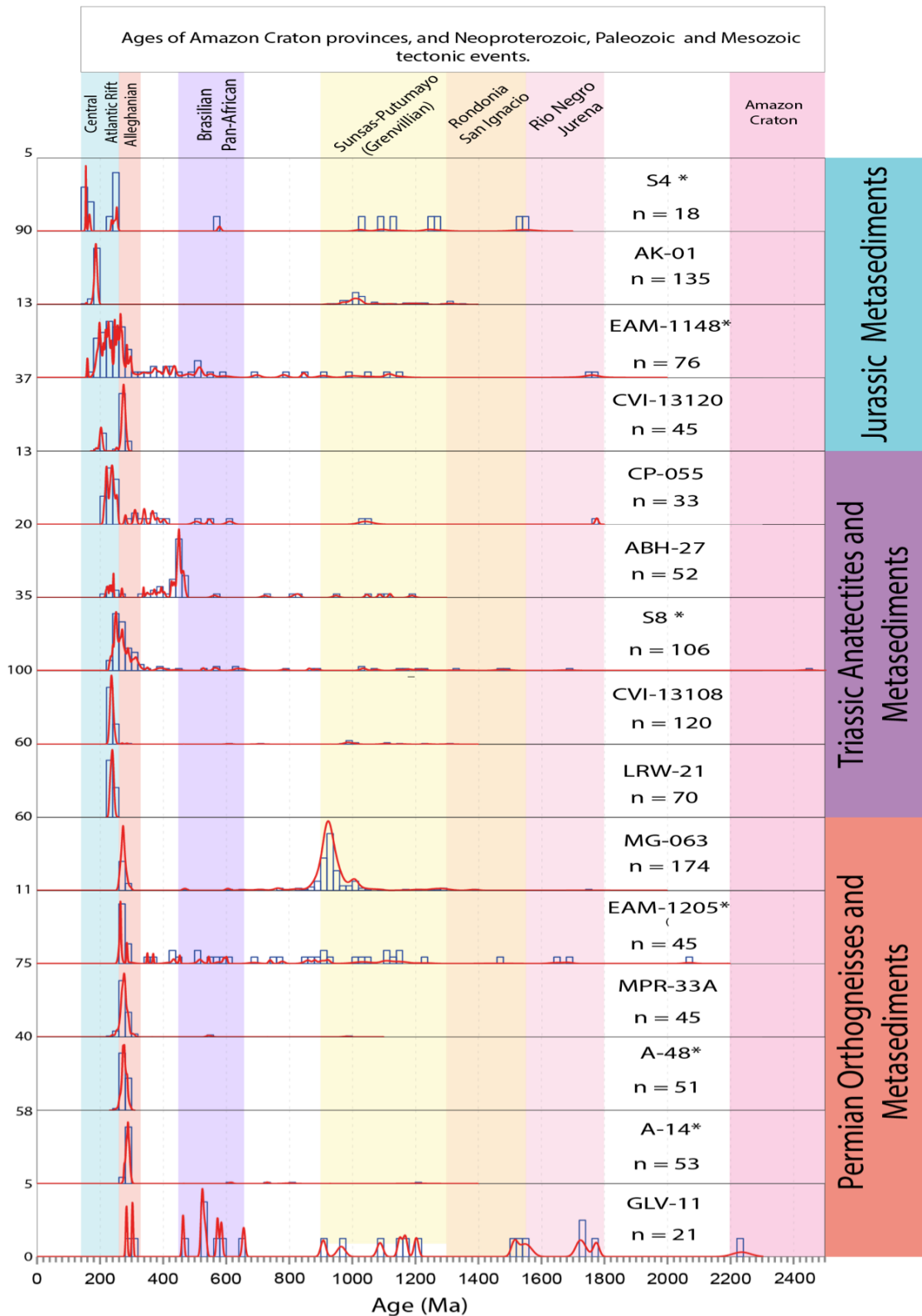


Figure 13. Zircon U-Pb LA-ICP-MS probability density plots obtained from the Permian to Jurassic metasedimentary and metaigneous units of the IB. Data sets labeled by asterisk are taken from (Cardona et al., 2010b, 2010c).

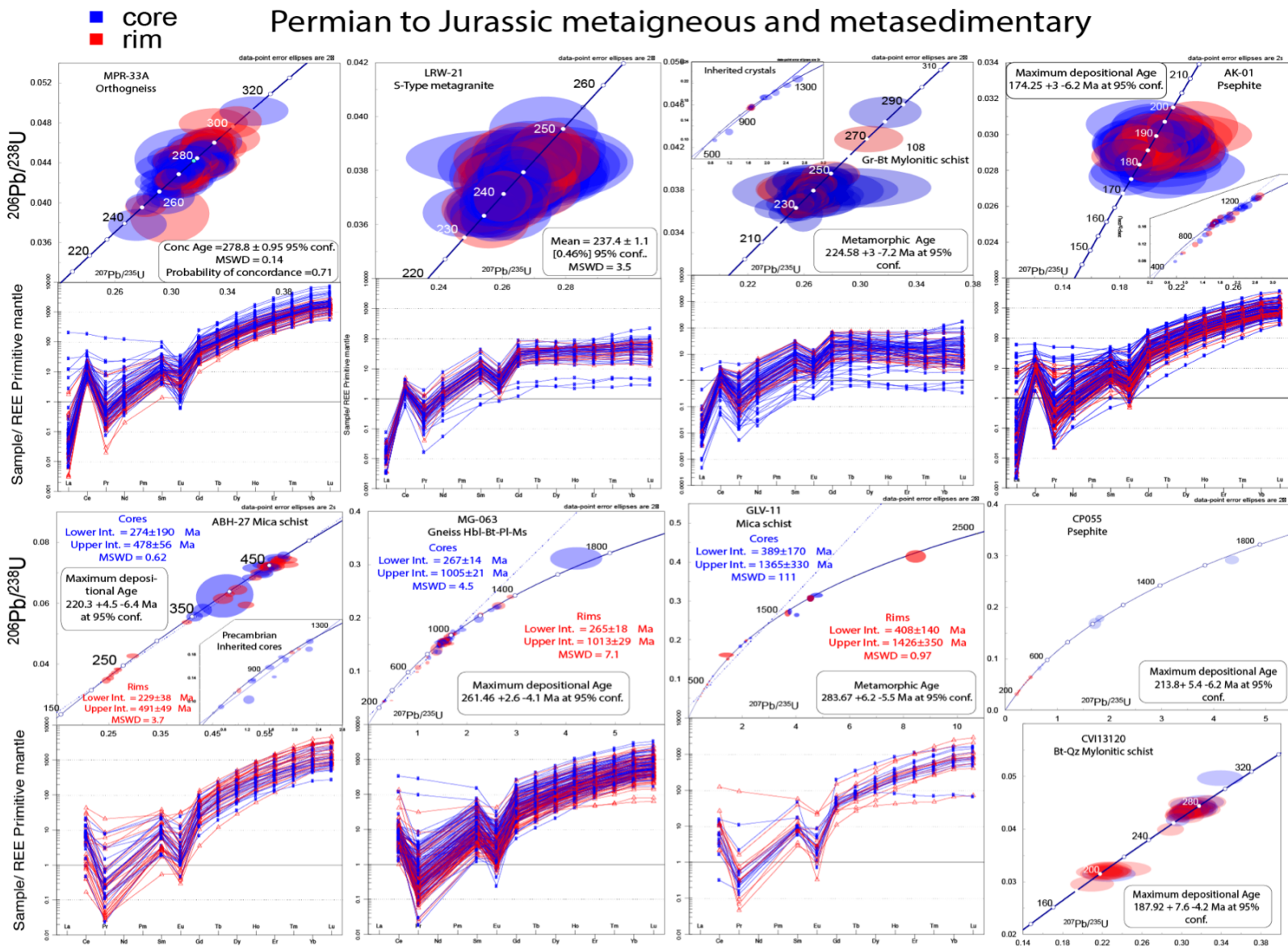


Figure 14. Concordia plots and trace element composition of zircons from the Permian to Triassic metaigneous and metasedimentary units of the IB.

4.5 Trace Elements in Zircon

REE composition in zircon reveals the original conditions in which zircons grew associated either with a magmatic or metamorphic origin. This discrimination must be done based on zircon morphology and host rock context, which is fundamental for defining a coherent history for a given set of crystals. As zircon is a common accessory mineral in the continental crust, and due to its relatively high crystallization temperature, it is obvious that its formation history is linked to other phases during magmatism or metamorphism, if temperatures are high enough for allowing phase re-equilibration and zircon recrystallization. It has been shown that under epidote-amphibolite facies zircon recrystallization does not occur (Lancelot et al., 1983), but at high amphibolite facies conditions (600-650°C) zircon can recrystallize (Peucat et al., 1985). For the metamorphic units of the SNSM we had obtained a similar T-t trend, with episodes of zircon crystallization during metamorphism, and coeval magmatism. At higher temperatures (650-1000°C), zircon can be totally reabsorbed and reset with no appreciable Pb-loss (Mezger and Krogstad, 1997). For this reason analyzing REE content in magmatic and metamorphic zircon is fundamental for defining the conditions under which recrystallization occurred. HREE content in zircon can evidence if recrystallization took place under a sub-solidus environment with limited supply of trace elements, considering the concurrent growth of feldspars and garnet in a closed system (Murali et al., 1983, Rubatto, 2002; Schaltegger et al., 1999). Garnet crystallization produces depletion on HREE in the reactive bulk and in consequence HREE patterns in zircon that grew under these conditions tend to be flat.

4.5.1 Sevilla Metamorphic Belt

The REE content in zircons from the Precambrian units of the SB is defined by a very clear compositional trend, in which cores are enriched in HREE, whereas rims are depleted in HREE (Fig. 12).

Sample AP-011 from the Buritaca Gneiss shows enriched cores, which exhibit higher HREE values ($\text{Lu}_n/\text{Sm}_n = 5 - 1221$, av. 126.8) than the depleted rims ($\text{Lu}_n/\text{Sm}_n = 4 - 288$, av. 91.7). A negative Eu anomaly is present ($\text{Eu}/\text{Eu}^* 0.05-0.61$, av. 0.25). Th/U ratios are high with values of 0.56 av. at cores and 0.5 av. at rims.

Quartz amphibole cataclastic gneisses from Sample AP-048 (Muchachitos Gneiss) exhibit a negative Eu anomaly ($\text{Eu}/\text{Eu}^* 0.05-0.95$, av. 0.15), with a HREE depletion trend towards the rims ($\text{Lu}_n/\text{Sm}_n = 4.5 - 184$, av. 56.3), and slightly higher values in the cores ($\text{Lu}_n/\text{Sm}_n = 4 - 396$, av. 92.12). As a general trend, cores younger than 980 Ma have Lu_n/Sm_n values lower than 20. Th/U ratios are high with values of 0.22 av. for the cores and 0.18 av. for the rims.

Sample DAA-39 quartz-feldspathic gneiss from the Buritaca Gneiss shows a compositional trend in cores and rims with a gap that separates a flat profile for rims and younger cores, from a steeper profile for the eldest cores. HREE element content declines from ($\text{Lu}_n/\text{Sm}_n = 1.4 - 104$, av. 18) in the cores to ($\text{Lu}_n/\text{Sm}_n = 1.5 - 45$, av. 8.7) in the rims. Th/U ratios are high and present values of 0.62 av. in the cores and 0.36 av. at the rims. Differently from the other samples, the Eu anomaly is less pronounced ($\text{Eu}/\text{Eu}^* 0.2-1.7$, av. 0.75),

For sample AP-010, a mafic migmatitic gneiss, HREE is very similar in cores ($\text{Lu}_n/\text{Sm}_n = 4 - 416$, av. 50.6) and rims ($\text{Lu}_n/\text{Sm}_n = 5 - 320$, av. 59). This minor rim enrichment

shows some melt rejuvenation due to recrystallization in a melt. Th/U ratios show high values of about 0.26 av. for the cores and 0.27 av. for the rims. A negative Eu anomaly is present (Eu/Eu^* 0.017-1.6, av. 0.5),

The Eocene Sevilla Stock shows high Y content (830 - 5200 ppm, av 1800 ppm), typical of magmatic zircons (Rubatto, 2002). This pluton shows an inverse REE profile in comparison to the Precambrian units (Fig. 16). At the Eocene Sevilla Stock core and rim populations are sharply separated. Most of the rims are Eocene and the inherited cores are Precambrian. The general trend shows a high enrichment in HREE in all the Eocene population with ($\text{Lu}_n/\text{Sm}_n = 137 - 5700$, av. 805), whereas Precambrian cores and rims show ($\text{Lu}_n/\text{Sm}_n = 1.9 - 550$, av. 108). A negative Eu anomaly is present, and significantly lower in Eocene zircons (Eu/Eu^* 0.43-1,1 av. 0,62) than in Precambrian zircons (Eu/Eu^* 0.038-0,915 av. 0,3). Th/U ratios are on average 0.36 for older population and 0.12 av. for the young crystals.

4.5.2 Inner Santa Marta Metamorphic Belt

Permian El Encanto Orthogneiss show an extremely high Y content (550 -7750 ppm, av 1900 ppm), values that are common in magmatic zircons. This is coherent with magmatic crystal morphologies. HREE elements describe a homogeneous compositional trend between cores and rims (Fig. 14), with relatively high values ($\text{Lu}_n/\text{Sm}_n = 21- 729$, av. 168) in the cores and ($\text{Lu}_n/\text{Sm}_n = 34- 264$, av. 159) in the rims. A negative Eu anomaly evidences classic magmatic zircon growth during feldspar crystallization (Eu/Eu^* 0.02-0.6, av. 0.3). Th/U ratios are high 0.8 av, and in this case are interpreted as related to a magmatic origin.

Sample GLV-11 from the Gaira Schists shows very high age variability with Precambrian and early Paleozoic crystals in very few grains ($n=20$). For this reason, it would be inaccurate to try to establish trace element compositional trends. The observation of the latest recrystallization event in this sample is recorded in a single rim age of 284 ± 5.2 Ma. The rim has low HREE, $Lu_n/Sm_n = 6.66$, and high Th/U ratio 0.8. Mica-schist sample MG-063 contains a Precambrian population that shares the same trace element signature as observed in the Precambrian Muchachitos Gneiss. The secondary Permian population shows ($Lu_n/Sm_n = 43-419$, av. 145), and high Th/U ratios av. 0.75. Samples from the La Secreta Mylonites show the most distinctive HREE profile with a general depletion in cores and rims, compared to the other lithologies analyzed in the SNSM (Fig. 11). For example, the mylonitic Pl+Grt+Bt schists sample LRW-21 shows ($Lu_n/Sm_n = 1-12$, av. 5) in the cores and ($Lu_n/Sm_n = 2.3-8.2$, av. 5.25) in the rims. The negative Eu anomaly shows values of ($Eu/Eu^* 0.15-1.2$, av. 0.26) in the rims and ($Eu/Eu^* 0.13-1.04$, av. 0.26). Th/U ratios show a general 0.22 av. in the cores 0.20 av. in the rims. Bt+Grt schist sample CVI-13108 contains some relevant Precambrian inheritance, with a trace element content that is almost equal to the HREE indexes obtained for the Precambrian units of the SB. Metamorphic zircons (Fig. 14) with depleted HREE contents in the cores of ($Lu_n/Sm_n = 0.3-48$, av. 6) and ($Lu_n/Sm_n = 0.3-7$, av. 2.08) in the rims. This ratio shows flat to negative HREE profiles. The negative Eu anomaly ($Eu/Eu^* 0.15-1.7$, av. 0.43) is concurrent and does not vary significantly from core to rim. Th/U ratios are in the range of 0.15-0.17 from core to rim. In sample ABH-27 from the Lower San Lorenzo Schists the HREE distribution of the early Paleozoic zircons shows values in the cores of ($Lu_n/Sm_n = 31-732$, av. 185) and ($Lu_n/Sm_n = 42-623$, av. 337) in the rims. In addition, it shows a negative Eu

anomaly (Eu/Eu^* 0.032-0.996, av. 0.43), Th/U ratios of 0.012-1.8, av. 0.45, and high Y content (281-4400 ppm). Permo-Triassic rims present high HREE values ($\text{Lu}_n/\text{Sm}_n = 44-1185$, av. 436.5), but low Th/U ratios of 0.0086 -0.0349. The youngest analyzed metasedimentary sample AK-01, from the San Lorenzo Schists yields two distinctive populations (Fig. 14). The Precambrian zircons show trace element compositions, and Th/U ratios akin to the Buritaca Gneiss. The Jurassic population shows high Y contents (135-4240 ppm), and elevated HREE contents ($\text{Lu}_n/\text{Sm}_n = 25-581$, av. 113) with a minor rim enrichment ($\text{Lu}_n/\text{Sm}_n = 37-263$, av. 128). The negative Eu anomaly is typical of these magmatic zircons (Eu/Eu^* av 0.28), along with high Th/U ratios 0.83 av. in the cores and 0.95 av. in the rims. At the Eocene granitic leucosome associated to the SMB intruded in the Gaira Schists amphibolites zircons show high HREE concentrations ($\text{Lu}_n/\text{Sm}_n = 16-380$, av. 127). This sample presents the highest Y values from all the analyzed samples (936 -28000 ppm) (Fig. 16). The negative Eu anomaly of these migmatitic crystals is (Eu/Eu^* av. 0.2). Th/U ratios are relatively low (0.0126 - 1.5, av 0.14).

4.5.3 Outer Santa Marta Metamorphic Belt

Zircons for both metavolcanic samples CVI-1367 and CVI1344 present high Y contents (400-6000 ppm) (Fig. 13). For sample CVI-1367, the HREE are enriched progressively from cores ($\text{Lu}_n/\text{Sm}_n = 65-559$, av. 245) to rims ($\text{Lu}_n/\text{Sm}_n = 159-433$, av. 270), Th/U ratios are on average 0.5, and a negative Eu anomaly (Eu/Eu^* av 0.5) is observed. Sample CVI-1344 shows HREE values ranging from cores ($\text{Lu}_n/\text{Sm}_n = 10-355$, av. 214) to rims ($\text{Lu}_n/\text{Sm}_n = 88-363$, av. 228). Th/U ratios are in the range of 0.3 av., and a negative Eu anomaly (Eu/Eu^* av 0.3).

Cretaceous metabasites & Paleogene Plutons

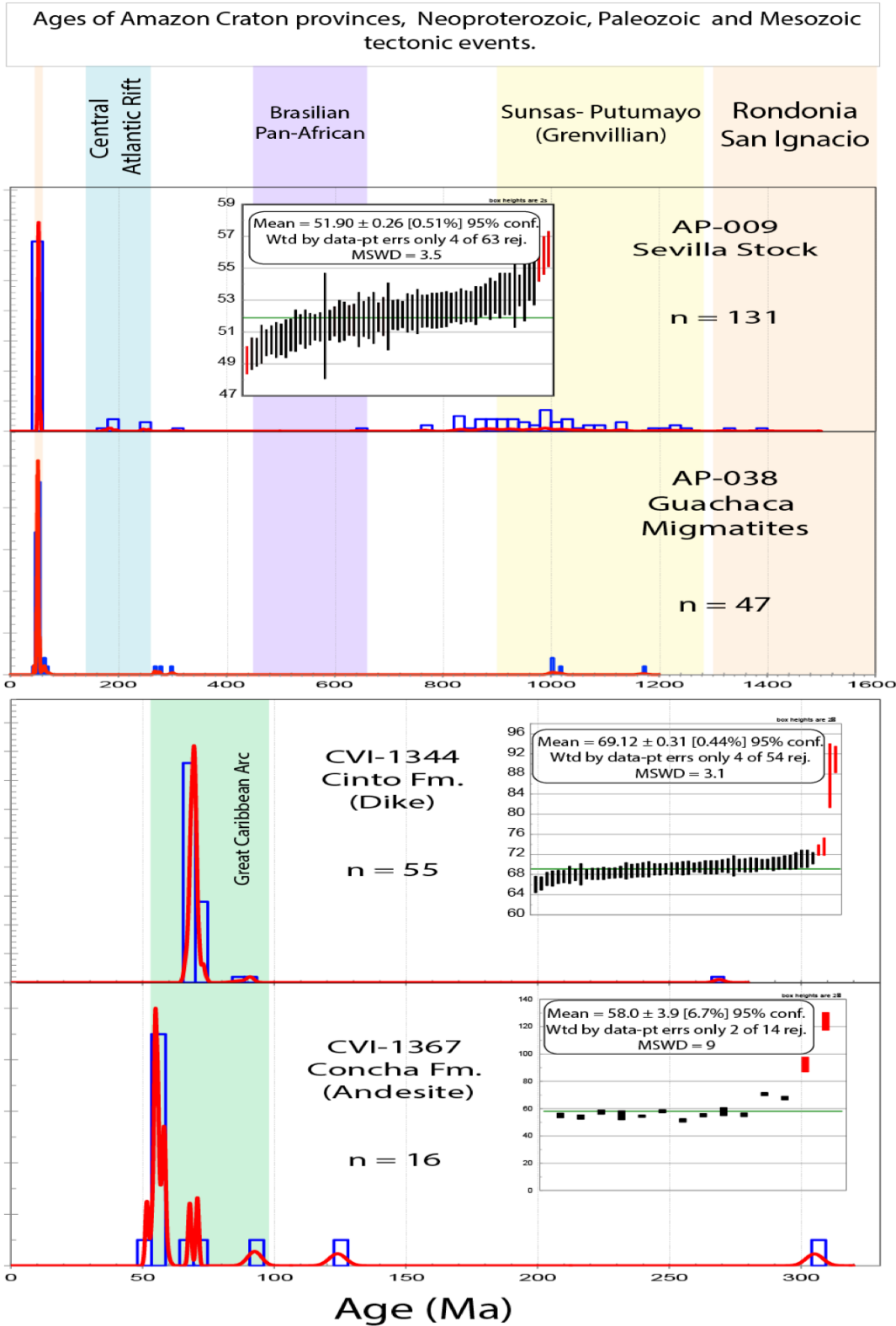


Figure 15. Zircon U-Pb LA-ICP-MS probability density plots obtained from the Cretaceous to Eocene metabasites and granitoids of the OB and SB.

Paleogene Plutons

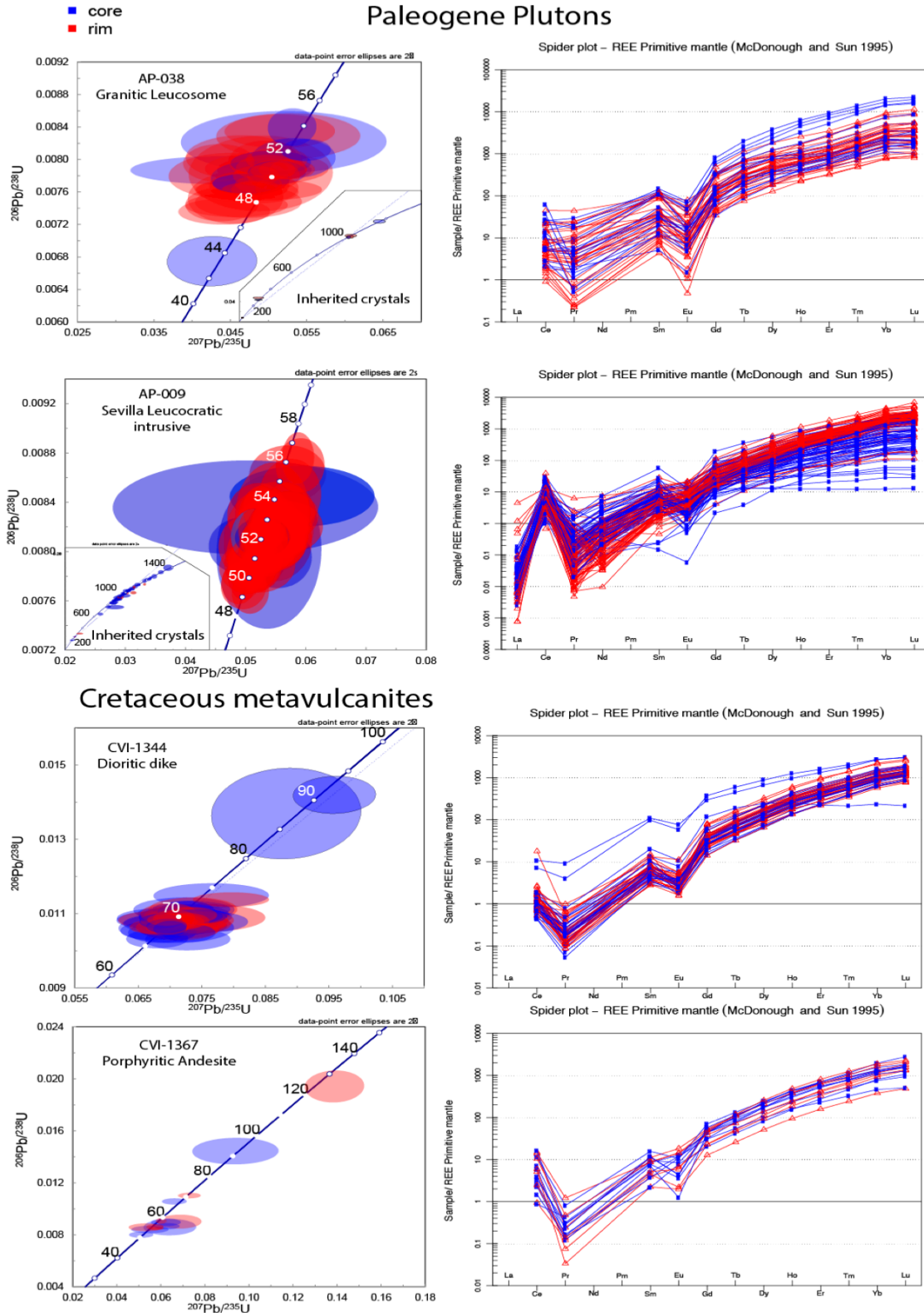


Figure 16. Concordia plots (left) and trace element composition (right) of zircons from the Cretaceous to Eocene metabasites and granitoids of the OB and SB.

4.6 Mineral Chemistry

Representative microprobe analyses of garnet, biotite, plagioclase, hornblende from three samples of the La Secreta Mylonites, are presented in table A2. Garnet exhibits high chemical variations from core to rim observed in weak to a strong zoning (Fig. 17 & 18). From major oxides, mole fractions were calculated into the structural formula for separating the end-member compositions (X_{Alm} , X_{Prp} , X_{Sps} , X_{Grt}). Each end-member was plotted in a compositional profile for visualizing Mg, Fe and Mn contents within cores and rims of the crystals (Fig. 18). The garnet amphibolite sample presents weak zoning CVI-1388 and the X_{Alm} content ranges from 0.55 to 0.65, enriched in the rim, the X_{Prp} content from 0.09 to 0.12 depleted in the rim, the X_{Grs} content from 0.01 to 0.03, and the X_{Sps} content from 0.01 to 0.04 enriched in the rims. Minor zoning is evidenced in x-ray maps by a rim Mn enrichment and Mg depletion. Fe shows as well a minor increase towards the rim.

At the transitional zone to the Bt+Grt schist strong zoning can be observed in the compositional profiles with X_{Alm} content ranging from 0.45 to 0.68, showing enrichment in the rim. X_{Prp} values vary from 0.12 to 0.31 with core enrichment, the compositional variation in the X_{Alm} , X_{Prp} end-members is abrupt and coincides with zoning observed in back-scattered images (Fig. 17) and x-ray compositional maps (Fig. 18), the X_{Grs} content varies from 0.09 to 0.2 higher at core, and X_{Sps} from 0.01 to 0.09 enriched in the rim.

Bt+Ms+Grt schist showed an X_{Alm} content ranging 0.65 to 0.72, with a minor enrichment towards the rim, X_{Prp} from 0.1 to 0.12 with rim enrichment, X_{Grs} content from 0.03-0.13, and X_{Sps} from 0.04 to 0.15 depleted in the rim. The compositional

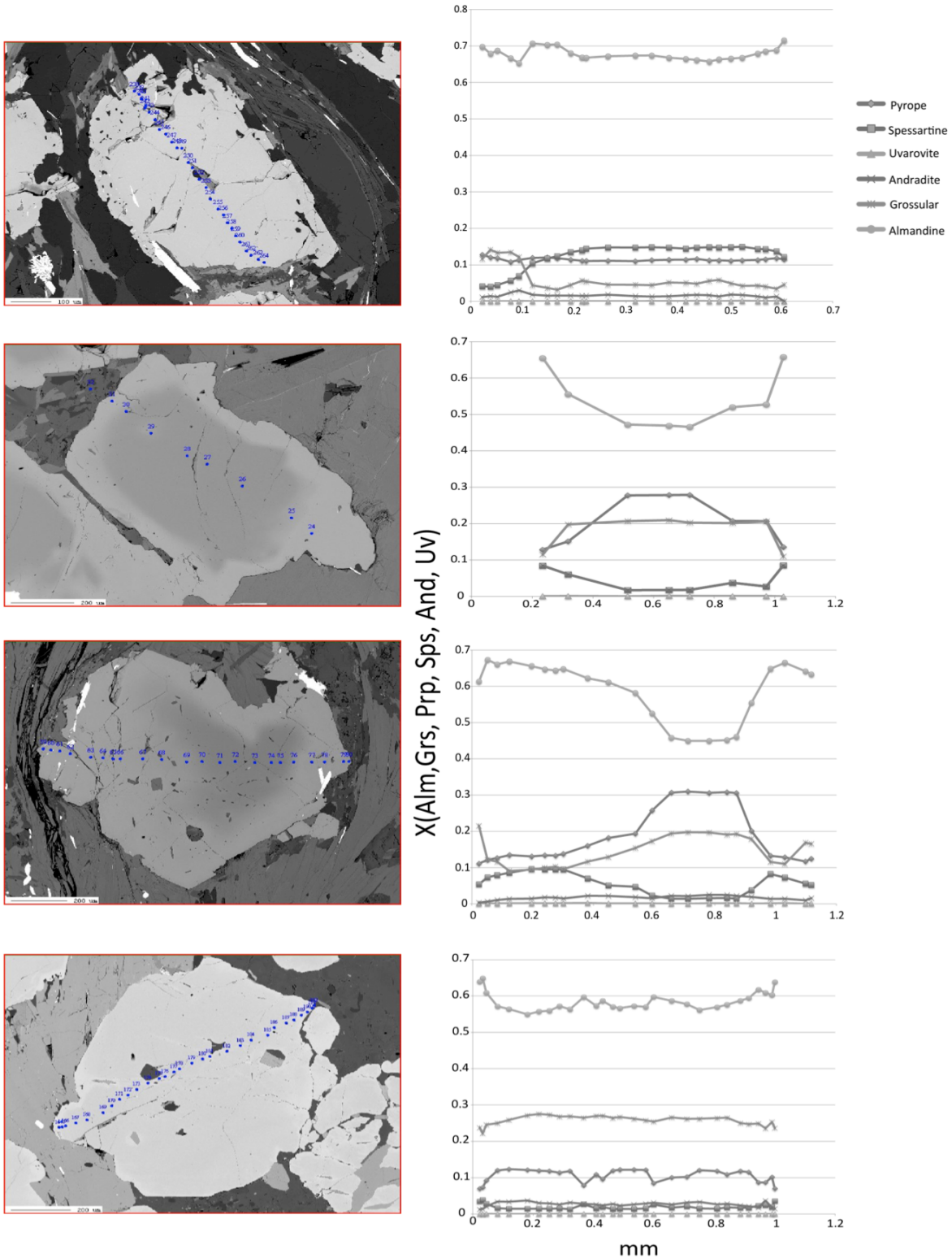


Figure 17. (a) Garnet compositional profile, sample CVI-1388. (b) Garnet compositional profile, sample CVI-13108. (c) Garnet compositional profile, sample CVI-13108. (d) Garnet compositional profile, sample CVI-1385, note strong zonation and garnet Mn enrichment towards rims.

profile is asymmetric, and the rim shows inclusions. A Mn x-ray map shows core enrichment for this element differently from Fe and Mg, which remained constant (Fig. 17 & 18). Major differences between X_{Alm} and X_{Prp} reflect differences in $MgO/(MgO+FeO)$ bulk compositions.

Representative analyses of biotite in Bt+Grt+Pl schists show X_{mg} values of 0.6 to 0.645 typically of intermediate to magnesium biotites. Biotite crystals are not zoned and mostly homogeneous.

Hornblende crystals from Hbl+Grt+Pl+Qtz sample CVI-1388 are zoned. This can be observed in thin section, like a marked color variation from brown in the cores to bluish-green in the rims (Fig. 6). This mineralogical variation corresponds to compositional trends of the end-members and is related to an increasing Ti content. In consequence of the increment in metamorphic grade, in this case describing a decrease of Ti to the crystal borders and therefore a prograde reaction for hornblende. End-member X_{Mg} varies between 0.34-0.41 with higher values in the crystal cores. X_{Fe} varies between 0.59-0.66 with higher values in crystal rims, and with Al enrichment and Mg depletion in hornblende crystal borders. In the same sample, garnet shows a discrete zoning in rims with a minor Fe and Mn increase and Mg depletion in Mg.

Alternative evidence of this retrograde path for Garnets from Hbl+Grt+Bt+Pl+Qtz amphibolites and in Bt+Grt+Qtz+Pl schists is addressed with garnet geochemical discrimination diagrams (Mange and Morton, 2007; Wright, 1938). Based on approximately 544000 analytical spots measured by x-ray map standardization, it can be demonstrated that the garnet rims crystallized in biotite schists under lower amphibolite facies metamorphic conditions, whereas the cores crystallized under high amphibolite facies conditions (Fig. 19a-c).

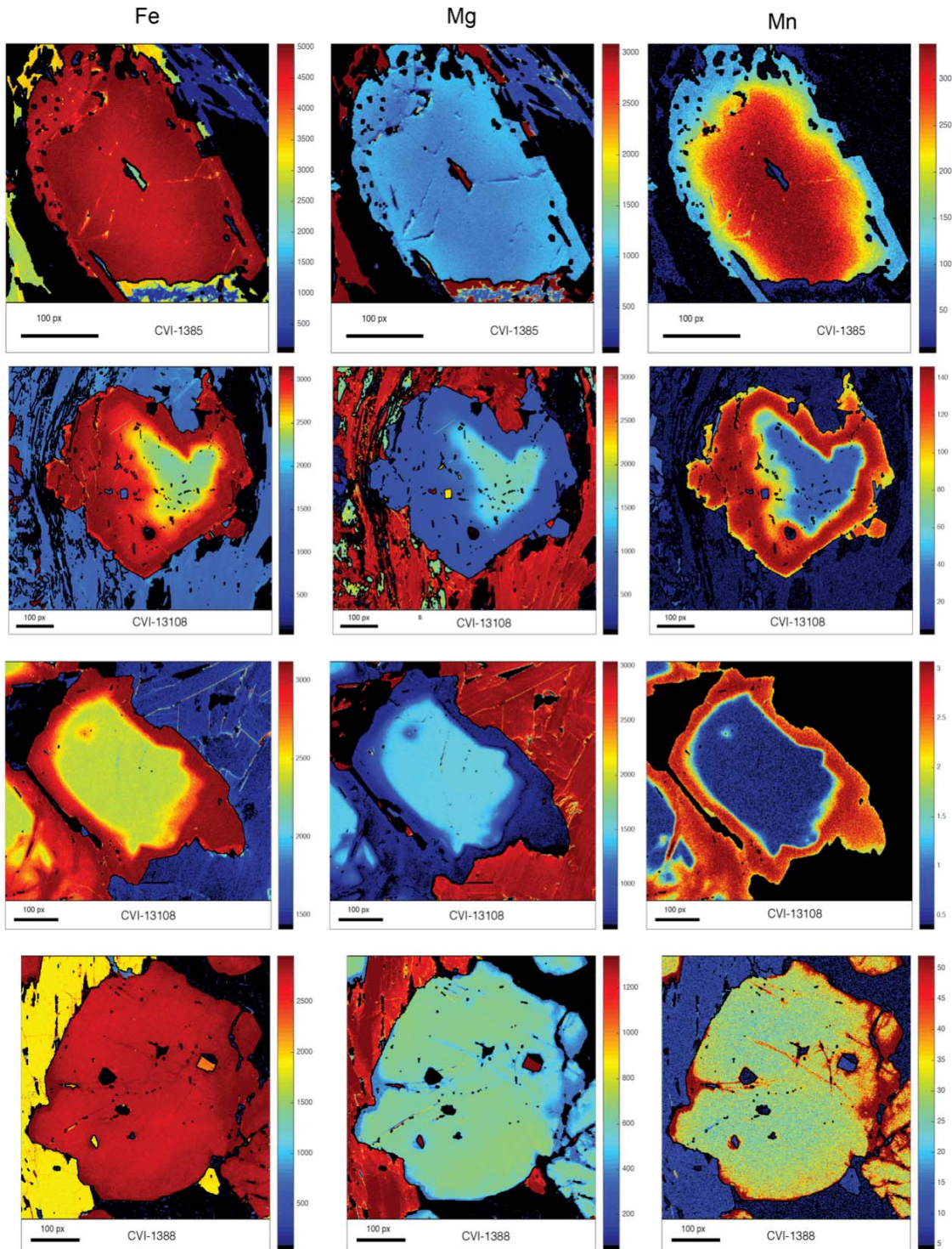


Figure 18. Mn, Mg and Fe X-ray maps from Hbl+Grt+Pl amphibolite (CVI-1388) (bottom), Bt+Grt+Pl schist (CVI-13108) (center) and Bt+Ms+Grt+Qtz+Pl+Chl CVI-1385 (top), collected adjacent to the Aguja fault. Observe compositional zoning varying, being more intense at the Bt+Grt Schists (CVI-13108), with higher values of Fe and Mn at rims, and a decrease of Mg in the rim, this strong zoning is related to retrograde metamorphism.

In contrast Bt+Ms+Grt+Qtz+Pl+Chl schist shows a prograde path in the biotite schist field (Fig. 19d), and a decrease of Mn from core to rim, as it would be expected for a continuous fractionation growth process (Hollister, 1966; Tracy et al., 1976).

4.7 Thermobarometry

Thermobarometry of the Lower–Middle Triassic La Secreta Mylonites has been performed on three samples combining different approaches of quantifying raw x-ray electron microprobe data using internal standards, and including empirical and semi-empirical geothermobarometers from the literature, using XMapTools, a MATLAB based GUI program. We used for samples CVI-13108 and CVI-1385 garnet and biotite geothermometers (Bhattacharya et al., 1992), and for sample CVI-1388 the Hbl+Grt geothermometers (Ravna, 2000) and Hbl+Pl+Qtz geothermobarometers (Lanari et al., 2014). This first approach was complemented by calculations on pressure and temperature using the Garnet- Al_2SiO_5 -Plagioclase (GASP) geobarometer for metapelites (Ghent, 1976; Holdaway, 2001). Results of both methods are compared for estimating pressure-temperature conditions of mineral crystallization in this metamorphic suite.

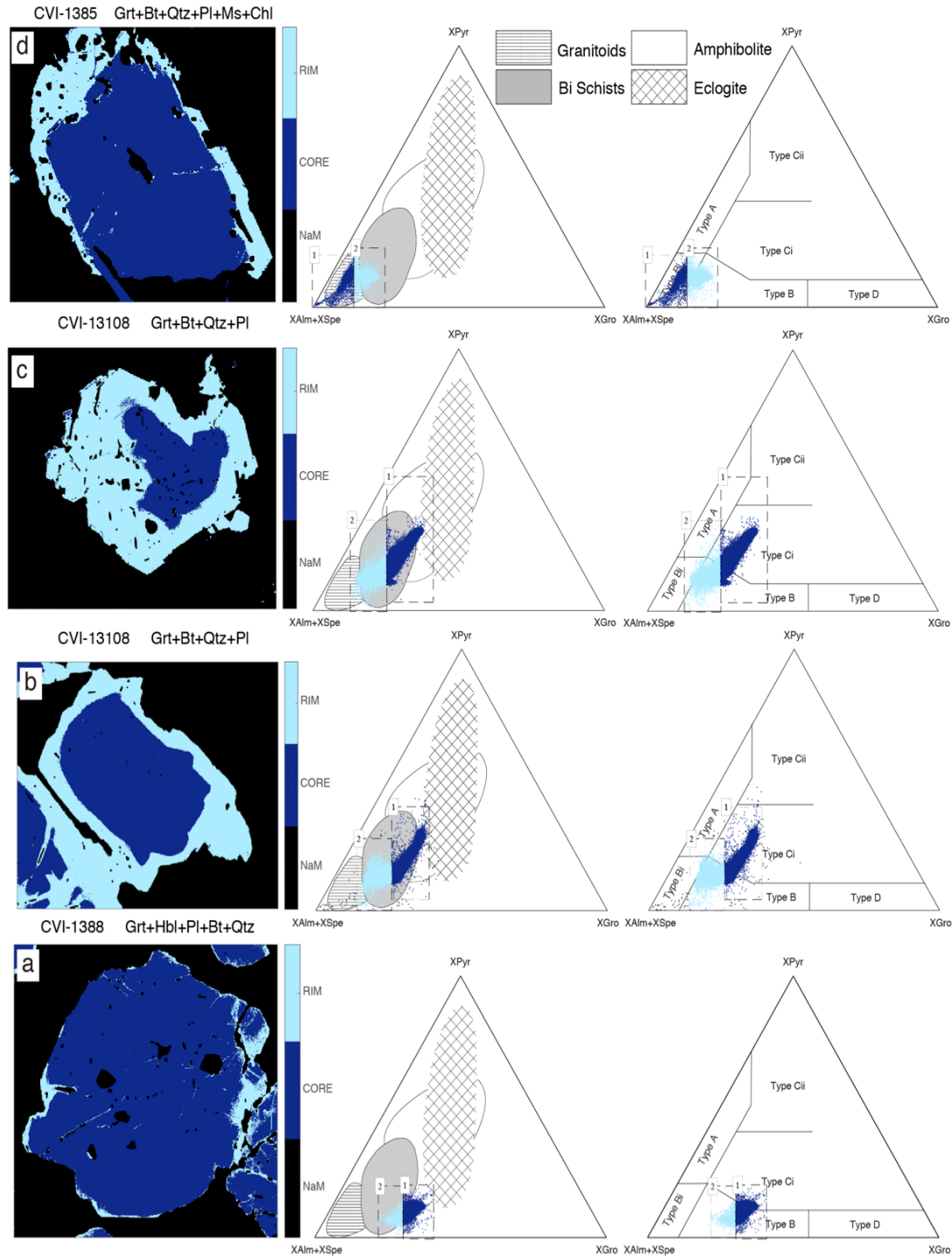


Figure 19. Retrograde zonation is evidenced in lower metamorphic degree towards the rims in samples CVI-1388 a), CVI-13108 b & c), in contrast sample CVI-1385 d), shows a prograde zoning. Samples reside within the amphibolite and biotite schists fields.

Garnet composition discrimination diagrams (left) image, extract from the Grt maskfile generated in XMapTools® (Lanari et al., 2014) showing core and rim selected areas for each crystal, each pixel from the maskfile is plotted in the triangular diagrams according end-member composition, tones indicate

compositional ranges pertaining to the crystal core and rim. (Center) Ternary discrimination diagram after (Wright, 1938), legend of the fields in the top center of the figure; (right) ternary discrimination diagram after (Mange and Morton, 2007). A – mainly from high-grade granulite-facies metasediments or charnockites and intermediate felsic igneous rocks, B – amphibolite-facies metasedimentary rocks, Bi – intermediate to felsic igneous rocks, Ci – mainly from high-grade mafic rocks, Cii – ultramafics with high Mg (pyroxenites and peridotites), D – metasomatic rocks, very low-grade metamafic rocks and ultrahigh temperature metamorphosed calc-silicate granulites.

4.7.1 La Secreta Mylonites - Garnet Amphibolites

Metamorphic temperatures of garnet amphibolite sample CVI-1388 were estimated using the Hbl+Grt geothermometer by Ravana (2000), with values ranging from 530-570 °C for the rims and 771-835 °C for the cores. In both cases, we had chosen 20 spots from core to rim in an Hbl+Grt assemblage.

Our second calculation involved the geothermobarometer Hbl+Pl+Qz (Lanari et al., 2014) resulting in values of 660 °C- 750 °C for Hbl crystal cores and at crystal borders respectively, in the vicinity of the Grt+Hbl interaction. Pressure for this assemblage was calculated with 10.2 - 11 kb for crystal cores and 11.6 - 12.1 Kb at crystal borders of Hbl+Grt and Hbl+Plg mineral pairs. Peak conditions for hornblende crystallization were 750 °C and 12.5 kb (Fig. 20) based on 11000 data points measured by x-ray map standardization on 15 amphibole spot measurements and seven feldspar spot measurements (Table A3). Thermobarometry results evidence contemporaneous Gr+Hbl growth during prograde metamorphism as temperature and pressure values increase from hornblende cores to borders (Fig 20). Garnet crystals show a similar behavior with increasing T values towards the border of the crystals. Retrograde rims in garnet are discrete (Fig 21b), indicating that garnet amphibolites experienced a less intense retrograde metamorphism.

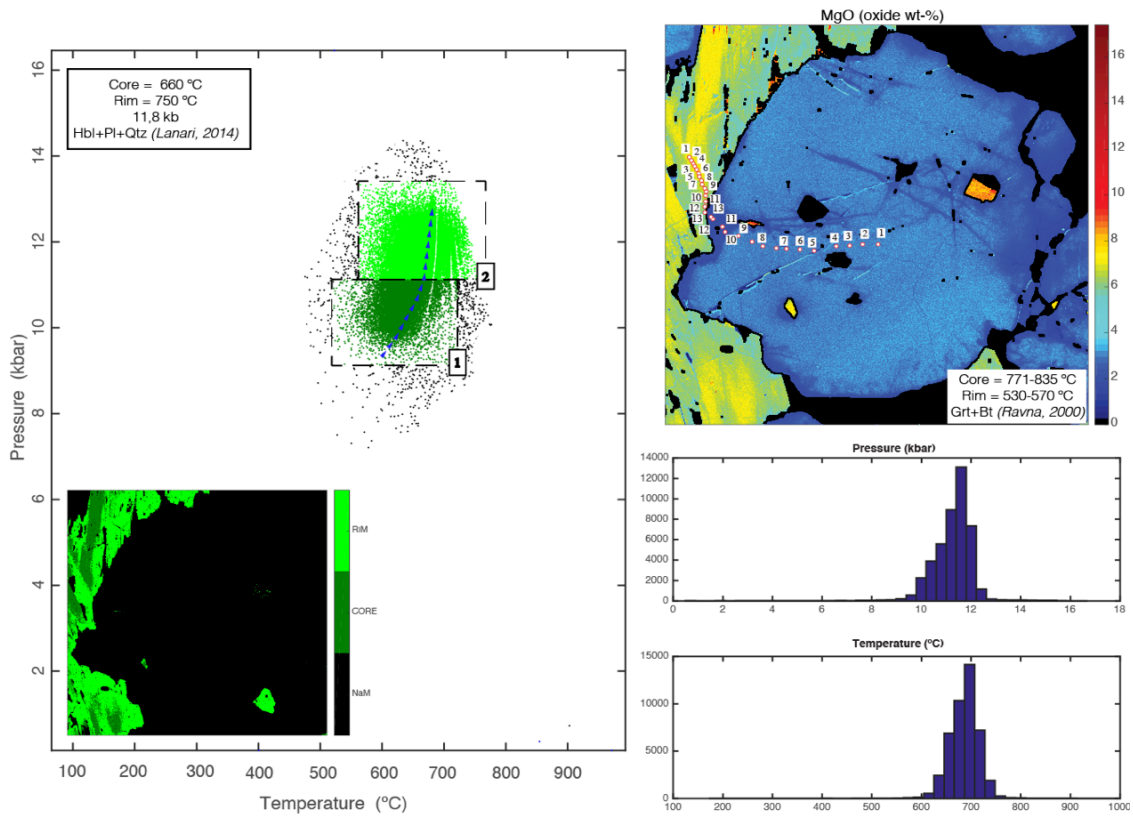
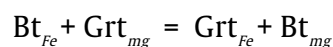


Figure 20. P-T paths for sample CVI-1388 showing (left) calculation for the Hbl+Pl+Qtz and pairs, green shaded Hbl crystal evidences a prograde growth with increasing conditions of P and T from core to rim, peak at = 750°C and 11.8 kb. T and P probability density plots. (top right) the Grt+Hbl thermometer was calculated using 13 paired spots measured from core to rim, peak T conditions for Grt are 771-835 °C, note incipient zoning with lower temperature range between 530-570°C, describing a retrograde path.

4.7.2 La Secreta Mylonites - Biotite-Garnet schists

Up section mylonites with Grt+Bt+Pl+Qz paragenesis geothermometry was calculated using the calibration by Bhattacharya et al. (1992) based on the exchange equilibria with relative large entropy change considering the reaction:



Measurements were made in spot couplets of Grt+Bt from crystal centers to borders. The results show temperatures from 512°C to 830°C. With a first prograde growth from the center of the crystal from 781-832°C, and after this peak with temperature decreasing progressively towards the rim where the measured values are between 513-536°C (Fig 20). This change depends on the compositional zoning occurring as a gentle transition or a sharp boundary. In either case, the temperature response is a gradual temperature decrease or a 200°C gap between core and rim respectively.

An additional approach is Ti-in-biotite thermometer of Henry (2005), which allows estimating an average 600°C ±12°C within a range of 510-630°C ± 12°C. However, this result should be addressed carefully because this geothermometer is recommended to be applied to ilmenite or rutile-bearing, graphitic, peraluminous metapelites, and for this sample graphite is absent, which could lead to underestimating the peak temperature by about 80°C.

Thermometry in sample CVI-1385, with a Bt+Ms+Grt+Pl+Qtz paragenesis, showed lower *T* conditions with values between 300-475°C ± 12°C with the Ti-in-Biotite thermometer, with the highest values in Bt inclusions inside Grt. The results of this thermometer are again underestimated by 50°C ± 12°C, because of the absence of graphite and ilmenite.

Garnet thermometry showed homogeneous values from core to rim ranging from 497-560 °C, with a *T* increase towards the rim in contrast to the samples described above. Coexistence of muscovite and light red to transparent pleochroic biotite support this assumption.

For samples CVI-13108 and CVI-1385, we determined pressure values using the GASP geobarometer (Ghent, 1976; Holdaway, 2001), calculating molar fraction from the

major oxide composition obtained in core to rim spot couplets measured in x-ray maps and considering sillimanite as aluminosilicate results are summarized in (Fig. 21).

4.7.3 *Ti-in zircon thermometry*

Trace element measurements in zircon provide Ti concentrations for each spot ablated. The Ti-in zircon thermometer is a powerful petrological tool. Recent studies have demonstrated that most zircons cluster strongly at $\approx 700^{\circ}\text{C}$ (Watson, 2005), in plutonic rocks however this apparently fixed temperature can decrease depending on the presence of near-water saturated conditions (Harrison et al., 2007). Experimental and natural results define a linear dependence of equilibrium Ti content (expressed in ppm by weight) upon reciprocal temperature (Watson et al., 2006). This is expressed by the relations:

$$1) \quad \text{Log} (\text{Ti}_{\text{zircon}}) = (6.01 \pm 0.03) - (5080 \pm 30) / T(\text{K})$$

$$2) \quad T (^{\circ}\text{C})_{\text{zircon}} = (5080 \pm 30) / ((6.01 \pm 0.03) - \text{Log}(\text{Ti}))$$

Results on La Secreta Mylonites samples showed an average Triassic zircon T of 730-800 \pm 10 $^{\circ}\text{C}$, with an average of 745 $^{\circ}\text{C}$ which is within the range of T predicted by the thermobarometric systems Bt+Grt, Hbl+Grt. Inherited Grenvillian zircons showed values between 820-1110 $^{\circ}\text{C}$. (Table A2).

El Encanto Orthogneiss yielded slightly lower temperatures between 600-850 $^{\circ}\text{C}$, with an average value of $\approx 685^{\circ}\text{C}$. Eocene intrusive rocks show a temperature range similar

to an average zircon crystallization temperature of 683°C. Precambrian cores preserve the original thermal signature of 820°C. Rocks from the Precambrian basement show values between 850-1200°C, attesting for a higher thermal regime during their crystallization.

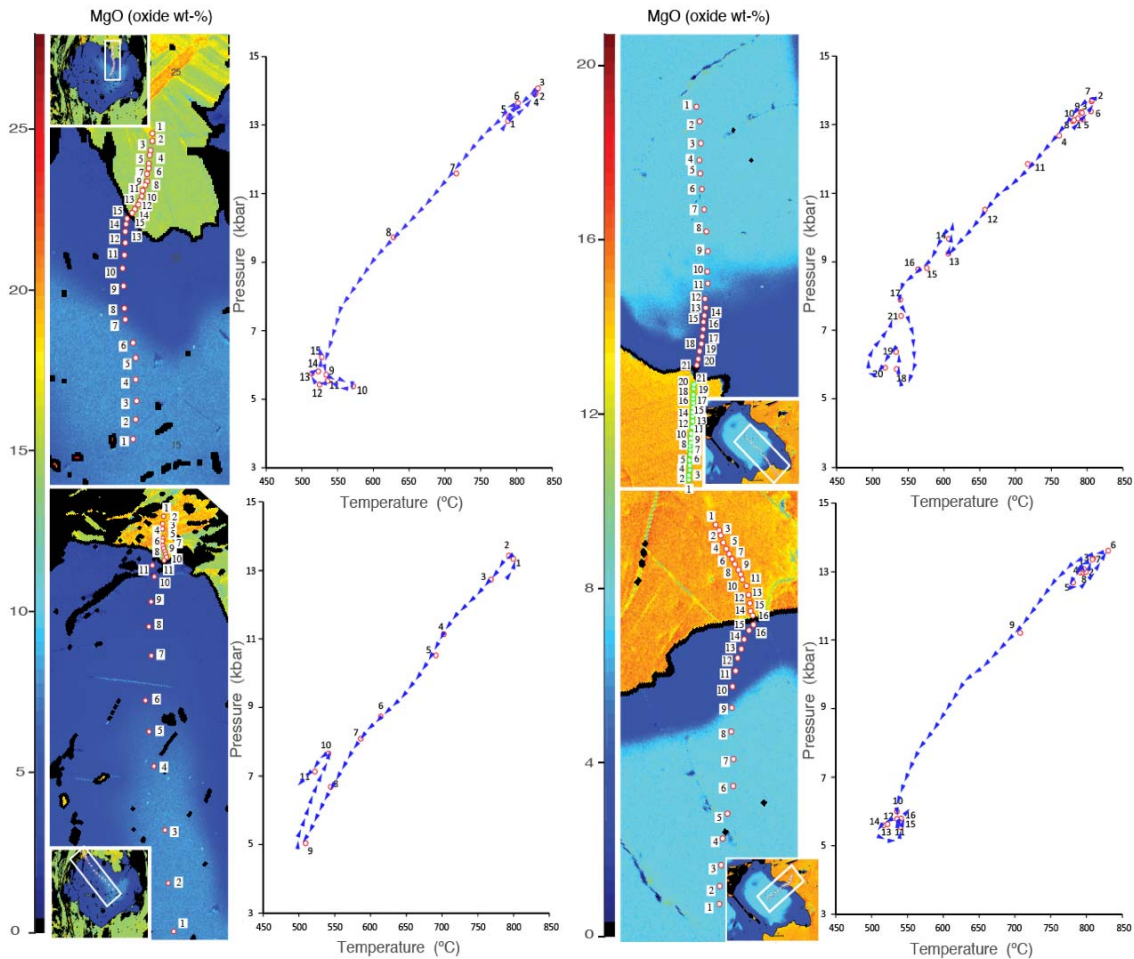


Figure 21. P-T paths calculated for sample CVI-13108 in four Grt+Bt core to rim profiles within two different crystals, each profile has between 15 to 20 measured spot couplets. Extracted major oxide compositions are used for calculating P-T conditions using the GASP geobarometer and the thermometry calibration by (Bhattacharya et al., 1992), results are plotted in the P Vs T diagrams at the right of each measurement, an initial prograde phase reaches metamorphic peak around 830°C and 14kb, and is followed by cooling and decompression down to 530°C and 5.5 kb.

5. DISCUSSION

5.1 Geochronological implications for the SNSM

Originally the SNSM has been subdivided into different tectonic provinces: 1) the Sierra Nevada Province SNP, 2) the Sevilla Metamorphic Belt or SB, 3) the Inner Santa Marta Metamorphic Belt IB, and 4) the Outer Santa Marta Metamorphic Belt OB (Tschanz et al., 1974). Our results help to redefine the extent of these provinces in a litho-chronostratigraphic context (Fig. 22).

Zircons recovered from the samples can be classified either as magmatic or metamorphic. Most of the samples exhibit a mixture between these two types of crystals. The rounded and larger crystals with subrounded shapes and thick outer rims are associated with inheritance from older sources.

Zircons commonly present overgrowths that may be related to metamorphic episodes. The first metamorphic episode corresponds to early Neoproterozoic ages and is found in all Precambrian samples 900-790 Ma. The second episode of middle Permian age was identified in the Gaira and San Lorenzo schists as a magmatic reworking of crystals ca. 270-260 Ma. A third metamorphic episode was identified in La Secreta Mylonites, related to crystallization of zircon under high-amphibolite facies, with very low inheritance and two distinctive populations at 250-220 Ma (Fig. 22).

We found that there is a significant age variation from the Sierra Nevada Province and the SB. Zircon U-Pb data show that the presumed Paleozoic Buritaca and Muchachitos cataclastic gneisses to the north of the Sevilla Lineament present Neoproterozoic age

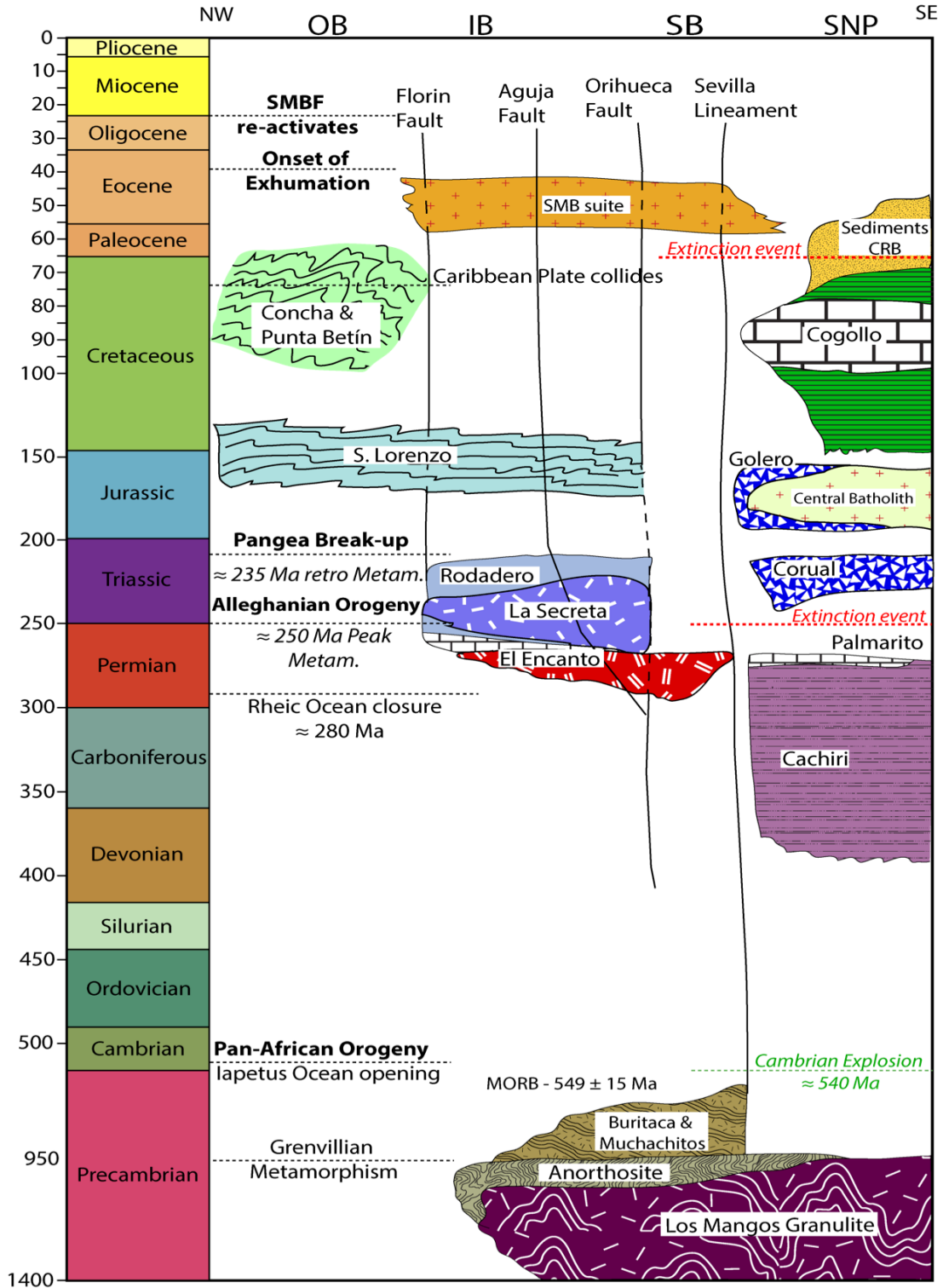


Figure 22. Stratigraphic summary of the Sierra Nevada Province, Sevilla Belt, Inner Santa Marta Metamorphic Belt, and Outer Santa Marta Metamorphic Belt according to the field relationships and stratigraphy section. Major tectonic events, geochronological significant ages, and massive extinctions are indicated.

populations with a spectrum from 790-1350 Ma recrystallized at ca. 980 Ma. Including a late generation of MORB oceanic crust aged 549 ± 15 Ma at late Neoproterozoic-Cambrian, evidencing that these gneisses are probably remnants of the opening of the Iapetus Ocean after Rodinia break up at this part of the Proto-Andean margin (Fig. 23), although slightly younger these ages, are closer to the Dibulla Gneiss and Los Mangos Granulite ages (Cardona et al., 2010a; Cordani et al., 2005; Ordóñez et al., 2002; Restrepo-Pace et al., 1997) than they are to the previously suspected Paleozoic age (Tschanz et al., 1974). Late Neoproterozoic ages found in these units are correlative with autochthonous rocks from the Proto-Andean margin of South America related to anorogenic magmatism within the southern Appalachians, during the initiation of the Neoproterozoic active margin (Chew et al., 2007, 2008; Mišković et al., 2009). Previous U-Pb zircon ages (Cardona et al., 2006; Cordani et al., 2005) (Neoproterozoic cores, Permian rims) as well as Jurassic ca. 147 Ma K/Ar ages of the SB (Tschanz et al., 1974), evidence at least two major thermal events. The first thermal event can be linked to the intrusion of Permian metaluminous granitoids in the Gondwana basement. The second thermal event may relate to Jurassic pluton emplacement at around 147 ± 6 Ma (Tschanz et al., 1974). Later thermal disturbances associated with the Caribbean arc accretion are recorded in the AFT and ZFT systems (Villagómez et al., 2011b).

Regarding our age data we propose that the Orihueca fault is the limit between the SB and IB, separating the Neoproterozoic gneisses from the Permian-Triassic and Jurassic schists that post-date, in turn, the intrusion of the El Encanto orthogneiss, which proves a continental-arc derived affinity during the late Paleozoic. This faulted contact is sharp and is marked by the juxtaposition of amethyst bearing amphibole gneisses with a zircon U-Pb age of ca. 902 Ma against quartzites with a Middle Jurassic

maximum depositional age of 174.2 ± 6 Ma. These meta-sediments overlie the La Secreta Mylonites (Figs. 6 & 22).

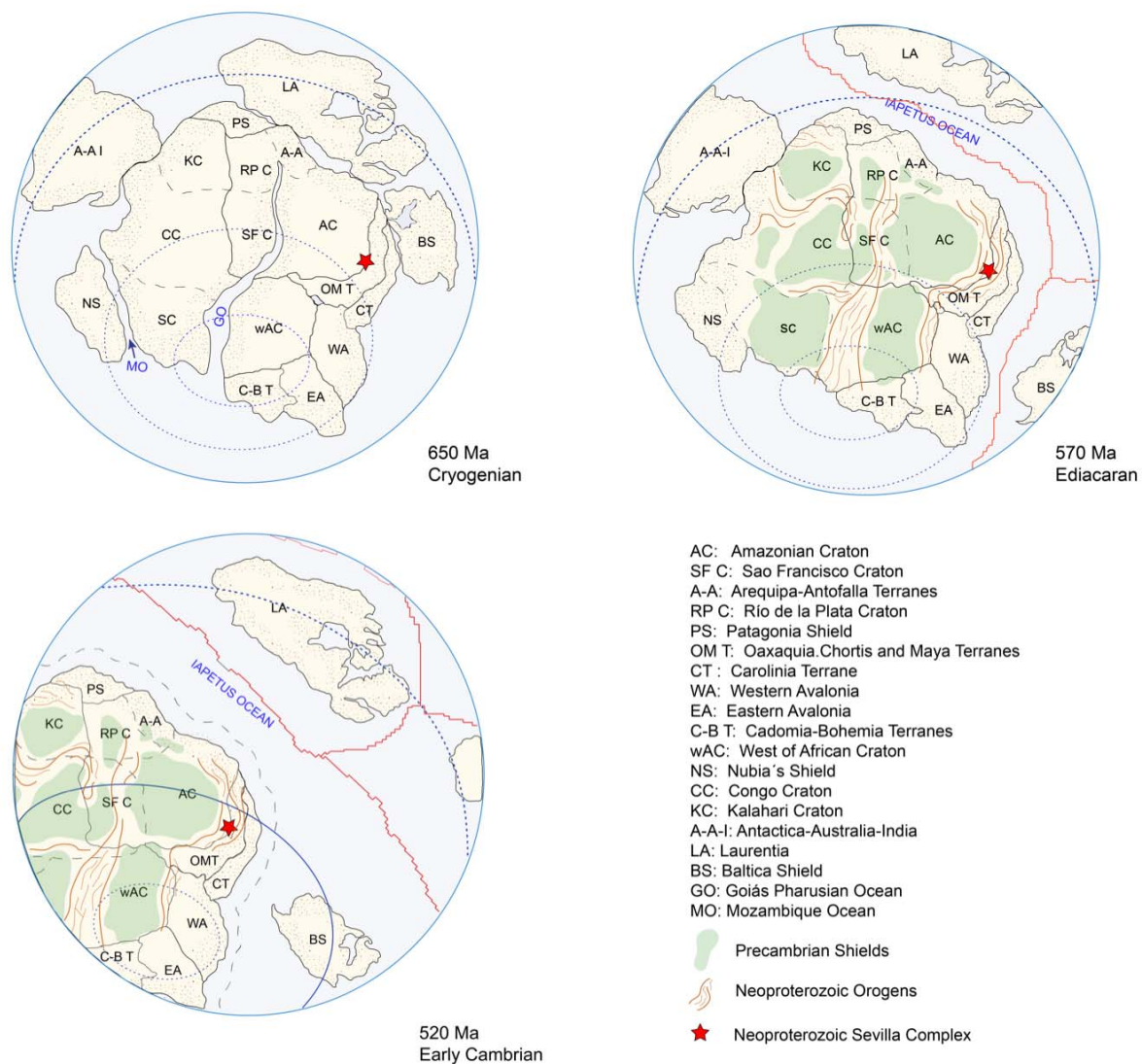


Figure 23. Palaeogeographic reconstruction at the Neoproterozoic- Cambrian transition $\approx 650 - 520$ Ma. The closure of the Goiás Pharusian and Mozambique Oceans led to Gondwana assembly, and the subsequent Brazilian- Pan African Orogeny, zircons at the Sevilla Metamorphic Belt attest a Grenvillian and Pan African sources (see Fig. 7), this reconstruction is modified after (Chicangana and Kammer, 2016a, 2016b; Cordani et al., 2013; Ganade de Araujo et al., 2014; Nance et al., 2014).

The El Encanto Orthogneiss can be correlated with the mylonitic orthogneisses of the Valencia creek with zircon U-Pb ages of 264.9 ± 4 Ma, $276.5 \text{ Ma} \pm 1.8$ and 288.1 ± 4.5 Ma (Cardona et al., 2010c).

El Encanto Orthogneiss of Permian age underlying the Gaira and San Lorenzo Schists is in contradiction to the widely accepted view that the IB is of Caribbean origin (Cardona et al., 2010b; MacDonald et al., 1971; Tschanz et al., 1969; Zuluaga and Stowell, 2012). The geochronologic evidence in this case favors the affinity of the IB to Gondwana and precludes a Caribbean origin of this province. The Permian granitoids intrude the Precambrian units at the Sierra Nevada Province at the Perijá Range and Mérida Andes (Macdonald and Hurley, 1969, van der Lelij et al., 2016).

Zircons recovered from the Gaira Schists showed maximum depositional ages of ca. 283.67 ± 6.1 (sample GLV-11) and 261.46 ± 2.6 (sample MG-063). Both meta-sedimentary rocks with lower Permian maximum depositional ages are interbedded with marble layers and amphibolites. We presume that it is possible that the age of the marbles included into the Gaira schists is Permian. For the San Lorenzo Schists, maximum depositional ages range between 188.3 ± 7.3 in the footwall of the Aguja fault, and to 174.2 ± 6 Ma adjacent to the Orihueca Fault.

These ages do not allow constraining a stratigraphic age but a range between Late Triassic and Middle Jurassic for the deposition of the protolith. Towards the base of the San Lorenzo Schists the maximum depositional ages are slightly older (220.33 ± 4.6 Ma), and the Precambrian inheritance is minimal as Permian and Ordovician sources are the most relevant sources for these sediments. These extensive age components suggest that deeper levels of the circum-Pangaeian orogen were already exhumed and eroded. We assume that sedimentation took place during the Jurassic rifting phase due

to the fact that meta-sedimentary units rest in a non-intrusive contact over Permian orthogneisses. In this scenario, the Aguja fault controlled sedimentation (Figs. 24 & 25).

In the OB we identified meta-sediments represented by the clastic Cinto Formation (CP-055) with a maximum depositional age of ca. 213.33 ± 4.7 Ma. The affinity of this unit to the continental margin is evident by the prolific content in zircon populations that includes a complete Triassic record, and with Permian, Carboniferous and early Paleozoic crystals attesting to continuous erosion from a continental source.

The widely recognized upper Cretaceous Caribbean arc inception led to accretion of oceanic island arc terranes recognized in the OB with an overall age of ca. 82 Ma (Cardona et al., 2010b). The accretion of this arc is documented on the continental margin at least since ca. 92 Ma by dikes that intruded the Jurassic meta-sediments of the IB (Cardona et al., 2011).

All samples that yield this upper cretaceous signature correspond to the Concha and Punta Betín formations. These low-grade metamorphic units are related to the Caribbean arc. Coeval with arc collision this metamorphic suite underwent metamorphism during the upper Cretaceous. In the IB during metamorphism hornblende schists retrogressed to greenschist facies rocks (Bustamante et al., 2009; Tschanz et al., 1974; Zuluaga and Stowell, 2012).

The magmatism affected extensively the Precambrian basement of the SB: the Latal, Buritaca and Sevilla plutons yield crystallization ages of 55.5 ± 0.65 Ma, 65.1 ± 0.9 Ma, and 51 ± 0.26 Ma respectively. Dioritic dikes intrude the OB, and yield a Paleocene age of ca. 62 Ma, followed by the emplacement of the SMB between 58-50 Ma (Duque, 2009), whereas in the northern greenschist from Concha Formation porphyritic

andesites crystallized at 55.4 ± 2.6 My and record volcanic arc. This age span defines the duration of the Caribbean arc plutonism whose thermal overprint annealed ZFT and AFT systems (Villagómez et al., 2011b). The SMB and coeval plutons (Latal, Buritaca and Sevilla stocks) I-type plutons resulted from collisional tectonics between the Caribbean plate and NW South America (Cardona et al., 2011). Eocene plutonism thermal overprint caused migmatization of the remobilized Permo-Triassic rocks of the IB, the leucosomes associated to this episode yield Lower Eocene ages ca. 50.8 ± 0.4 Ma. Migmatization represents the culmination of the post-collisional magmatic activity marked by cooling and middle Eocene exhumation of the SNSM massif that had persisted until the present day.

5.2 Concluding remarks on zircon trace elements and mineral chemistry

We found relationships between variations in REE content, sample age and structural position in many samples. In several cases, it is common to observe HREE depletion in crystal rims. Precambrian samples indicate a metamorphic crystallization event in a sub-solidus medium with a clear HREE depleting towards 980 Ma, either in cores or in rims. The compositional change occurs in a clockwise rotation from a steep to a flat HREE profile (Fig. 12).

Negative Eu anomaly in Precambrian zircons indicates feldspars grew coeval with zircon and during metamorphism. Trace element content along with textural and field observations within the IB defines the magmatic origin for a Permian arc located on the margin of the SNSM massif. Permian aged rims in zircon commonly cover

Precambrian cores, attesting for a metamorphic recrystallization induced by partial melting of the basement. Eocene granitoids intruded the Precambrian crust causing zircon recrystallization; the melt provided a high amount of HREE to the newly formed zircons and rims around older cores.

Eocene migmatites yield high HREE values, and plutonic rocks of the same age from the Sevilla Stock showed HREE concentrations 10 times lower. The decrease in concentration of HREE elements from core to rim in the Guachaca Migmatites is related to garnet crystallization, whereas garnet in the Sevilla Stock rocks is absent.

The Th/U ratio in zircons has been traditionally used for discriminating between magmatic (Th/U >0.01) and metamorphic (Th/U <0.01) zircons. Our results however show that at the SNSM, Eocene magmatic zircons commonly have Th/U ratios <0.01 (Table A1). Instead many zircons from metapelites and mylonites yield Th/U ratios >0.01 and evidence that Th/U ratios by themselves are not useful for discriminating if a zircon has a magmatic or metamorphic origin. This paradox can be explained for solid-state recrystallized zircons typically inherit the Th/U characteristics of its parents (Möller, 2004; Möller et al., 2003). Combining of Th/U ratios, the Lu_n/Sm_n ratio and Y content and core-rim variations allows better to discriminate between zircons that grew in the presence of a melt or otherwise in a sub-solidus environment with depleted HREE. The Y content in magmatic zircons is an order of magnitude higher than in metamorphic zircons (Table A2). There is not a specific marker for a zircon to be magmatic or metamorphic based on trace element composition, because several other factors, such as the coeval growth of phases retaining Th, U or REE influence the final REE composition in zircon.

The chemical profiles of samples CVI-13108 and CVI-1388 (Fig. 20) show the characteristics of garnets that had undergone homogenization and retrograde zoning, with a core-to-rim increase in Mn and decrease in Mg (Tuccillo et al., 1990). Mn increase in the rims in Bt+Grt schists and Hbl+Grt+Bt+Pl+Qtz amphibolites corresponds to a retrograde diffusional effect driven by cooling during interaction with an aqueous phase (Tracy et al., 1976). In normal prograde zoned garnets, Mn profiles are typically bell-shaped, Mn depletion is related to the higher Mn fractionation factor in contrast to Fe, and Mg, as Mn in the rest of the rock is depleted during a continuous growth process (Hollister, 1966; Tracy et al., 1976).

Thermometry of the Grt+Hbl and Grt+Bt pairs show similar *P-T* estimates for amphibolites and mica schists for both cores and rims, which gives a reliable *P-T* path for the thermal evolution of this unit. A two-phase metamorphic event was identified in the La Secreta Mylonites in the hanging wall of the Aguja fault. The metamorphic paragenesis varies from Grt+Hbl+Pl amphibolites at the base to Grt+Bt+Ms+Pl+Qtz to the upper part of the suite, passing from the sillimanite zone to the Grt+Bt zone. Temperature and pressure estimates derived from zoned garnet in amphibolites and mica schists indicate a two-stage evolution with pro-grade peak temperatures of 770-830°C and pressures of 11.66-14 kb, and retrograde temperatures between 510-550°C and pressures between 5-7 kb (Fig 20 & 21).

Hbl+Pl+Qtz thermobarometry of garnet amphibolites indicates a temperature of 80°C and a pressure of 1.2 kb below the calculated peak metamorphic conditions shown above. This difference can be related to disequilibrium between the phases due to transient thermal perturbation or a late-stage re-equilibration of the phases at conditions different from maximum *P-T* conditions, thus reflecting final equilibration

during a retrograde stage. Under the described HP-HT range zircons are prone to recrystallize in a sub-solid phase as result of the metamorphism (Peucat et al., 1985; Rubatto, 2002; Sawyer and Barnes, 1988). Garnet zoning is therefore the product of continuous reactions in the presence of an aqueous fluid, which can be related either to an external source or to prograde dehydration fluids unable to escape (Tracy et al., 1976).

Therefore, zircon crystallization between 250-220 Ma is coeval with garnet growth, the two age peaks date metamorphic ages and can be related to the two metamorphic phases identified in garnet crystals. This process occurred under a sub-solidus remobilization product of dynamic metamorphism during crustal thickening of the first phase (11.7 kb), and subsequently by crustal attenuation and sub-solidus melting under lower P - T conditions (500°C -5.6 kb).

The high temperatures of 770-830°C and pressures of 12 kb reached by these rocks are consequence of crustal thickening and subsequent thermal relaxation and shearing, which is not compatible with an arc-subduction setting, as it has been postulated before (Cardona et al., 2010c). The absence of HP-LT assembly also precludes a subduction context, because P - T trajectories evidence an isothermal decompression and not an isobaric cooling, as it would be expected from subduction. Examples of subduction HP-LT assemblages due to underplated mafic crust that validate a subduction context are further to the south in the Central Cordillera of Colombia and the El Oro Metamorphic Complex in Ecuador (Bustamante et al., 2012; Riel et al., 2013). Former P - T paths obtained from IB rocks were interpreted in a subduction context during the Caribbean accretion (Zuluaga and Stowell, 2012). Such P - T paths

should be reconsidered in the light of the new geochronological data presented in this study that preclude a Caribbean origin for this particular area.

5.3 Pangaea Assembly during Alleghanian Orogeny

Based on our results we aim to discuss in a regional scope an evolutionary scenario for this part of the Proto-Andean margin from the late Paleozoic, describing the tectonic phases related with episodes of metamorphic and magmatic zircon crystallization. Boundary conditions were constrained from several geothermobarometers and evaluated along with the geochronological data.

290-250 Ma

The Permo-Triassic Tahamí (Vinasco et al. 2006) terrane was accreted onto the NW margin of Gondwana during collision. The convergent zone changed into a dextral shear zone that culminated with the Pangaea assembly and the closure of the Rheic Ocean. Following the Rheic Ocean closure during the Laurentia collision with NW Gondwana at ca. 310 Ma (Cardona et al., 2016), crustal thickening reached its maximum over a time span of 10-15 Myr, causing the formation of collisional anatectic melts in the continental crust during the early Permian at around 290-280 Ma (Figs. 24 & 25).

The Cisuralian orthogneisses and mylonites in the SNSM, at around 278 ± 0.5 Ma are only present in the NW metamorphic belt. Rocks of an equivalent age have been recognized in the Central Cordillera of Colombia as crustal anatectites originating from the Laurentia-Gondwana collision (Cochrane et al., 2014a; Vinasco et al., 2006). Permian granitoids that present inliers of metasedimentary rocks of equivalent age

have been recognized as far as the southern Maya block in the Chiapas Massif (Weber et al., 2007). Mexican and Colombian orthogneisses share a compressional origin that differs from other Permian granitoids found further to the south of the Andes interpreted as being the result of lithospheric thinning (Mišković et al., 2009; Sempere et al., 2002).

Zircon age and trace element core-rim relations prove that the principal Permian arc plutons intruded and recycled Gondwana crust as suggested by Cochrane et al. (2014a). In the SNSM the post collisional crystallization occurred at ca. 278 ± 0.5 Ma, and later high-grade metamorphism at ca. 250 Ma. Inheritance of zircons with Grenvillian U-Pb ages evidences a continuous magmatic arc through the Central Cordillera, the Perijá Range and the Merida Andes existed with a highly similar tectonic history to coeval granitoids northeastern Mexico (Weber et al., 2007).

Furthermore, during Permian times calcareous reefs fringed the arc and shallow marine sediments deposited in a nascent back-arc basin (Fig. 24 & 25). Remnants of these basin deposits are found on the E flank of the SNSM, the Perija Range, the Santander Massif, and the Merida Andes (Cardona et al., 2016; Langenheim, 1959; Laya and Tucker, 2012; Patarroyo, 2001; Villarroel and Mojica, 1987).

Geochronology and detailed mapping shows that an extensive area mapped homogeneously as mylonitic orthogneisses in the Valencia creek (Cardona et al., 2010c) is in fact a sequence consisting of meta-sediments, metabasites and meta-igneous rocks evidenced by the occurrence of marble seams and layered amphibolites. We propose to include within the Gaira Schists the undifferentiated amphibole schists that crop out in the northern part of the SNSM, where marble beds are frequent as well as interbedded with amphibolites and biotite schists.

In this context, the SNSM displays the record of a Permian arc and basin, which underwent a period of collisional magmatism between 300-280 Ma during Laurentia-Gondwana collision in the Ouachita-Alleghanian orogenic episode. The accretion of the so-called “Tahamí” terrane in the NW of Gondwana constituted the basement of the Central Cordillera (Vinasco et al., 2006), this interaction was responsible for an orogeny whose climax was the Pangaea amalgamation followed by Barrovian metamorphism. The final stage of this orogenic phase was marked by the onset of Pacific subduction during the Early Triassic beneath the recently formed Pangaea supercontinent (Figs. 24 & 25).

5.4 Post-Orogenic Collapse and metamorphism during Pangaea Break-Up

250-225 Ma

Pangaea agglutination at this part of the proto-Andean margin is marked by post-collisional anatectic melting at ca. 278 ± 0.5 Ma. This magmatism was followed 30 Myr later by *P-T* peak metamorphic conditions at ca. 250 Ma. This implies that the early Triassic peak metamorphism occurred several tens of million years after the deformation phases that were responsible for crustal thickening.

Between 240-230 Myr isothermal decompression was triggered by crustal extension that caused zircon recrystallization (Figs. 10, 13, 14, 24 & 25). The mechanism for extension can be explained by delamination of the slab of the overriding plate that causes plate decoupling, trench retreat and decrease on the convergence rates (Gorczyk et al., 2007). In this context, crustal delamination influenced a transition from compression to extension, given that crustal delamination triggers extension in

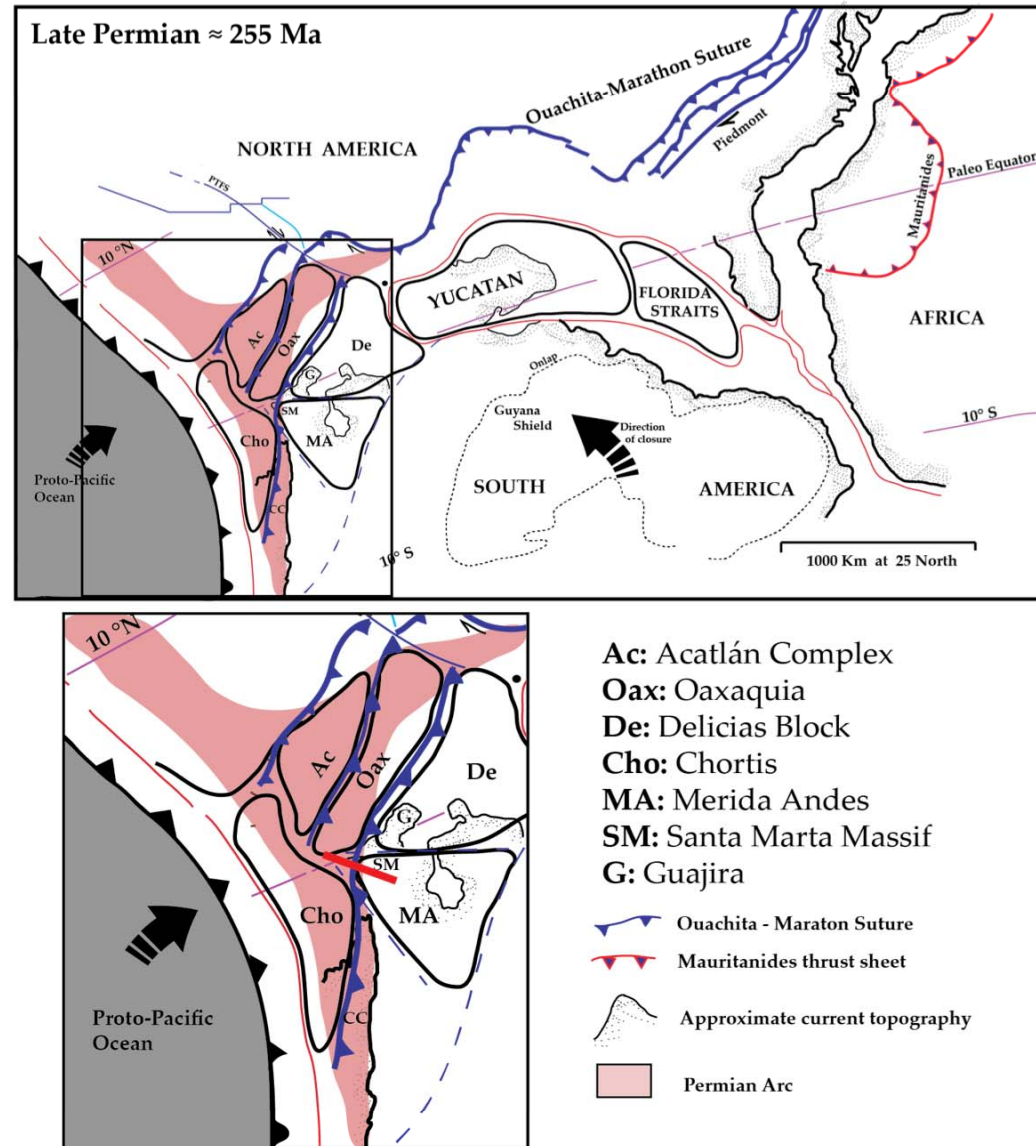
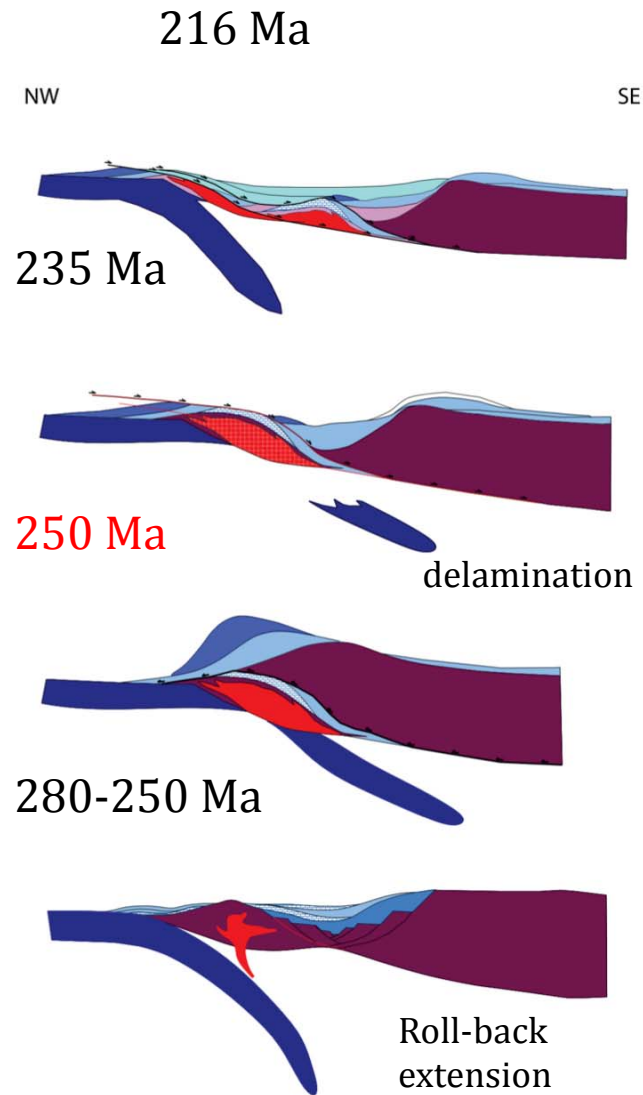


Figure 24. a) Schematic multi-stage evolution for the conjugate NW Gondwana margin and related arcs during Laurentia-Gondwana interaction. b) Palaeogeographic reconstruction of Pangaea late Permian ≈ 250 Ma, Gondwana and Laurentia are assembled entraining the Mexican- Central American Blocks in which a Permian arc is emplaced, modified after (Pindell, 1985; Van Der Lelij, 2013; Weber et al., 2007)

convergent orogens (Houseman and Molnar, 2001). It has been shown that continued extension causes partial melting of the crust and remobilization that can last for a range of 10 My (Ashworth and Brown, 1990; England and Richardson, 1977; Stüwe, 2007). Extension as well enhances the geothermal gradient, for this particular case documented during retrograde HT-LP metamorphism by an isothermal decompression that lasted between 240-225 Myr. Coeval with thermal weakening and extension occurred the formation back-arc marginal basins in half-graben systems filled with shallow marine carbonates, volcanic deposits and clastic sediments from the erosion of exhumed Precambrian basement (Fig. 24a). The geochemical composition of the La Secreta Mylonites allowed identifying volcanoclastic and mafic protoliths for this unit, which is coherent with Triassic mafic magmatism associated with an attenuated crust. For instance, based on matching 3591 analyses from the Ancomp database, the Bt+Grt+Pl schist sample CVI-13108 is comparable with the chemical composition of the middle to upper Triassic volcanoclastic Los Indios and Corual formations that crops out on the southeastern flank of the SNSM (Tschanz et al., 1969) (Table A5).

In contrast, the garnet amphibolite samples from La Secreta Mylonites (CVI-1388, LRW-21) shows a very dominant gabbroid-basaltic trend in the field of ultrabasic rocks (Bas et al., 1986). Further calibration with the Ancomp database hints at olivine-gabbro, gabbro, norite, diabas or a mafic volcanoclastic rock as protolith. The protolith database evaluation allows defining spilite or pillow lavas as protolith. Furthermore, the common occurrence of ultramafic pockets due to hyperextension within La Secreta Mylonites provides additional evidence for a tectonic extensional regime.

The continuous extensional process that permitted the onset of mafic magmatism induced an increase in the thermal gradient along with decompression. Isothermal

decompression was the main cause for retrograde metamorphism that affected the lower crustal sequences of the Alleghenides orogen as observed in the garnet retrograde zoning in metapelites, which for the proposed model, would fit a thermal re-equilibration of a thinned lithosphere. Ti-in-zircon thermometry demonstrated that metamorphic zircon grew within the P-T conditions estimated from the mineral paragenesis Grt+Bt, Hbl+Grt+Pl of the previously described samples from La Secreta Mylonites. These results bracket the period from 250-220 Ma for the Triassic metamorphism followed by the onset of magmatism during the upper Triassic that reworked the basement of the IB.

Kinematic indicators attest to a dextral shear (Fig. 7a, b) and in this case crustal extrusion more likely occurred by dextral strike slip tectonics during the Appalachian-Ouachitas Orogeny between about 315-273 Ma (Sacks and Secor, 1990). This tectonic segregation is coeval with the dynamic metamorphism (Sawyer and Barnes, 1988), and in consequence produced the mylonitic fabric c.a. \approx 240 Ma. Crustal thickening lasted at least for 20 Myr until metamorphic peak conditions and thermal re-equilibration were reached. After the metamorphic peak, orogenic collapse led to crustal extension, and continued Pacific subduction caused the development of a continental arc from the Early Triassic to the Late Jurassic. Remnants of this arc are well preserved in the SE flank of the SNSM, where the rift had its major extension and thick volcanoclastic sequences developed in a back-arc basin, which were later affected by intra-plate arc magmatism (Figs. 24 & 25). *P-T* trajectories indicate that deformation and metamorphism of the second metamorphic phase occurred simultaneously.

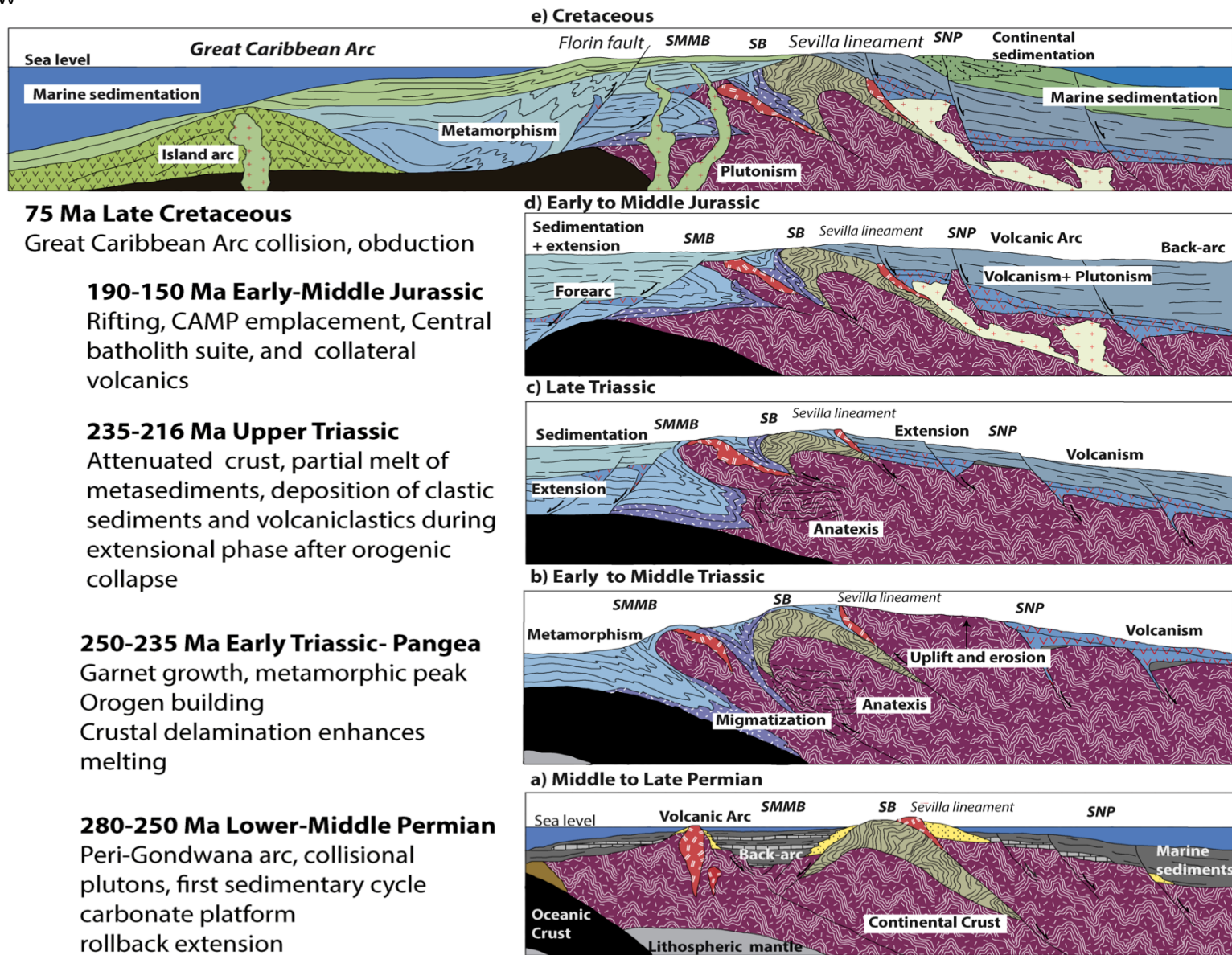


Figure 25. a) Schematic not-to-scale multi-stage tectono-metamorphic evolution for the SNSM massif since Late Permian to Late Cretaceous times, involving metamorphism, anatexia, and sedimentation cycles.

During the Triassic onset of rifting, magmatism evolved from intra plate to oceanic during Pangaea break-up. This process advanced in a rift that gave origin to the Atlantic Ocean (Cochrane et al., 2014a). Emerged Paleozoic and Precambrian massifs acted as source areas for the sediments that filled the growing rift basins. This is clearly evidenced by zircon populations that show a regional episode of arc plutonism during the Early Jurassic. This plutonic episode is distributed along the Andes and defines the drifting between the Central Cordillera to the south and the Chortis-Mexico block in North and Central America (Pindell and Kennan, 2009) (Fig 24b).

6. CONCLUSIONS

1) The gneissic units of the Sevilla Metamorphic Belt showed late Neoproterozoic ages, precluding an allocthonous origin for the SB and a Permian origin for the Sevilla lineament (Tschanz et al., 1974). Units correlatable with the Central Cordillera basement (Tahamí Terrane) are found instead in the metamorphic suites of La Secreta Mylonites and El Encanto Orthogneiss at the IB that are the continuation of the Permo-Triassic basement of the Lower Magdalena Valley. The Sevilla Metamorphic Belt units are a remnant of a late Neoproterozoic arc related to an ocean that existed between Laurentia and Amazonia during early Rodinia break-up and the later Brazilian/Pan-African orogeny that led to Gondwana assembly. This is evidenced in the ages of the gneisses of the Sevilla Metamorphic Belt of 890.83 ± 12 , 796 ± 18 that are too young to be related with the Grenvillian orogeny. The Sevilla lineament therefore is a suture that suffered normal reactivation as suspected by (Tschanz et al., 1974) during Triassic-

Jurassic rifting and facilitated pluton emplacement, this major structural feature played as well a fundamental role during the Cenozoic evolution of the SNSM.

2) We document collisional plutonism that caused partial melting of the continental crust at ca. 278 ± 0.5 Ma and our prograde metamorphic event (D1) at ca. 250 Ma that may reflect the delayed thermal response of a crustal thickening. This first tectonic cycle was linked to the closure of the Rheic Ocean that juxtaposed the Permian arc to the Neoproterozoic Pan-African units of the SB at the NW margin of Gondwana. Collisional Permian granitoids with Triassic metamorphic overprint are found in Oaxaquia, Acatlán and Maya blocks, and the Central cordillera of Colombia. These plutons confirm the existence of a late-Carboniferous-Permian magmatism that connected the Ouachita fold belt with the Oaxaquia terrane at the Central Pangaea Mountains, magmatism extended from southern North America through Central America and northern South America.

3) During collision prograde Hbl growth in an anti-clockwise P-T-t path, defines the Alleghanian-Ouachitan belt in NW Gondwana as an example of an evolving orogen that experienced metamorphic peak after crustal thickening. This is documented by an early high-pressure event (11 Kb \approx 35km depth Hbl+Pl+Qtz), and the influence of a non-constrained tectonic overpressure due to dextral shear.

4) The rocks reached their thermal metamorphic peak at 830°C and 13 kb at ca. 250 Ma, at this moment zircon metamorphic growth took place. We interpret this phase as result of collisional processes of the Central Pangaeian Orogen or Alleghenides.

5) Metamorphism at ca. 250 Ma has so far been regarded as independent of the Alleghanian-Ouachitan orogeny, because its age was not linked to Pangaea amalgamation and closure of the Rheic Ocean at around 300-280 Ma for this part of

Gondwana but rather to terrane accretion or a different orogenic event. However, this assumption underestimates the fact that the timing of crustal thickening and thermal re-equilibration can post-date collision, and that metamorphic peak conditions are reached several tens of millions of years after crustal thickening occurred (Stüwe, 2007). For this reason we postulate a continued single metamorphic event that reached its metamorphic peak ca. 40 Myr after the Laurentia-Gondwana collision, with at least 10 Myr for crustal thickening and 30 Myr until thermal re-equilibration occurred at peak *P-T* conditions. Understanding this equilibration as the moment at which mineral paragenesis crystallize. This long lasting coupling between Laurentia and Gondwana is related as well to oblique top to NE dextral shear acting during Pangaea amalgamation.

6) After the thermal peak conditions were reached the thickened crust created enough potential energy to hamper the convergence rate between the Pacific plate and the Pangaea supercontinent, which led to slab retreat and later delamination of the lower crust. As recently documented the removal of a lithospheric root causes a rapid increase in the topographic relief and onset of erosion of a previously thickened crust (Duretz and Gerya, 2013; Stüwe, 2007). The erosional cycle lasted for several million years as recorded by the Permian-Triassic clastic back-arc basin deposits, which register maximum depositional ages of ca. 260 Ma in the SNSM. The metamorphic age of the Chiapas massif at ca. 257 Ma (Weber et al., 2007) provides a temporal limit for the duration of the Pangaea formation.

7) A retrograde phase of isothermal decompression is recorded on garnet rims enriched in Mn with an episode of zircon recrystallization at ca. 240 Ma evidencing thinning and thermal weakening of the crust. The geochemical signature of garnet

amphibolites, which points to ultrabasic protoliths, evidences the onset of mafic magmatism following the pronounced crustal attenuation with deformation occurring coeval with metamorphism as evidenced by sheared zircon crystals of this age.

8) Zircon recrystallization persisted between 240–225 Ma at HT-LP conditions indicating a progressive thinning of continental crust during rifting and break-up of western Pangaea. During this process, mafic magma generation persisted at least until ca. 220 Ma. By about 216 Ma the removal of the conjugate margin suppressed most of the Permian arc remnants from NW Gondwana isolating remnants of this arc in the Mexican and Central American terranes from the continental margin. The continuation of Pacific plate subduction reactivated by slab break-off and isostatic rebound caused the arc to migrate continent-ward, and the former back-arc Permian basin occupied a fore-arc position fed by an emerging Jurassic arc. The basin fill of back-arc volcanoclastic deposits is recorded in the Gaira Schists and the volcanoclastic San Lorenzo Schists. This sedimentary cycle remained active during the Jurassic.

7. REFERENCES

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CHAPTER 2

Cenozoic exhumation of the Sierra Nevada de Santa Marta: From a collisional margin to a transpressive boundary

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Abstract

New low-temperature thermochronological data from the Precambrian, Permian-to Jurassic, Late Cretaceous and Paleocene-Eocene rocks that make up the northern metamorphic belt of the Sierra Nevada de Santa Marta (SNSM) of the Northern Andes of Colombia reveal exhumation trends in this part of the continental margin during Paleocene to Miocene times. Our analysis involves apatite and zircon fission track, as well as apatite (U-Th)/He dating, which allowed to constrain the thermal histories (~240 °C – ~60°C) of an inverted crustal sequence that exposes greenschist to granulite facies metamorphic rocks. Temperature-time paths evidence that cooling occurred diachronously in the massif through the activation of NW verging thrusts with the onset of rapid erosional exhumation after collision and TTG series pluton emplacement under an initially high geothermal gradient of 40°C/km. In this regime the massif exhumed at elevated rates ~0.9 km/Myr during 50–45 Ma. These rates incremented to 2 km/Myr for the interval 45–40 Ma with a decreasing geothermal gradient of 30°C/km, reaching the highest rates of 2.7km during 37–35 Myr. We interpret the exhumation of this inverted lower crustal sequence because of coupling of ductile extrusion of a thickened continental crust and surface denudation. Post-Eocene exhumation events show subsequent cooling under a decreased exhumation rates of ~0.2 km/Myr at ca. 25 Ma, related to the start of a composite transpressive-transtensive regime in the southern Caribbean plate boundary as a result of changes in convergence between the Caribbean and South American plates. This reorganization led to the activation of the NW trending Santa Marta - Bucaramanga (SMBF) and E-W trending Oca (OF) faults. These faults contributed to the partitioning of a subsequent oblique convergence and transferred strain from the Caribbean Plate into the Santander Massif, the Eastern Cordillera of Colombia, the Perijá Range, and the Merida Andes until the late Miocene. We interpret that this new tectonic regime inhibited further NW thrust propagation and caused a NE tilting of the SNSM massif due to extensional activity of the SMBF coupled with the

tectonic drag caused by the transit of the Caribbean plate. These processes resulted in attenuated exhumation rates of ~ 0.09 km/Myr that persisted until the late Miocene during the SNSM thermal re-equilibration.

Key words: Northern Andean block, low temperature thermochronology, collisional plutonism, lower crustal channel

1. INTRODUCTION

Bedrock thermochronological studies play a major role in elucidating evolutionary trends in mountain building around the globe. Recent thermochronological studies in the Northern Andes revealed the influence of tectonic process coupled with climatic influence during the growth of this mountain belt (Bermúdez et al., 2010, 2011, 2013; Michalak et al., 2016; Mora et al., 2008, 2010, Parra et al., 2009, 2010; Restrepo et al., 2009; Spikings and Simpson, 2014; van der Lelij et al., 2010, 2016; Villagómez et al., 2011b; Villagómez and Spikings, 2013; Whipple, 2009). Nevertheless, it is still matter of debate if exhumation is strongly dependent on climate-induced erosion or if it is mostly controlled by tectonics (Molnar and England, 1990; Montgomery et al., 2001; Raymo and Ruddiman, 1992; Reiners et al., 2003; Whipple, 2009). This problem has intrigued geologists for decades, and it seems that every specific geological problem must be addressed under the light of sedimentological and petrographic evidences and a well-constrained structural context.

At the current location of the Northern-Andean block the Caribbean-South American boundary presents a transpressive situation (Cardona et al., 2011a; Kennan and Pindell, 2009; Taboada et al., 2000). A major question remain on the evolution of the Sierra Nevada de Santa Marta massif (SNSM): 1) What is the mechanism or

mechanisms that operated during the uplifting phases experienced by SNSM during the Cenozoic?

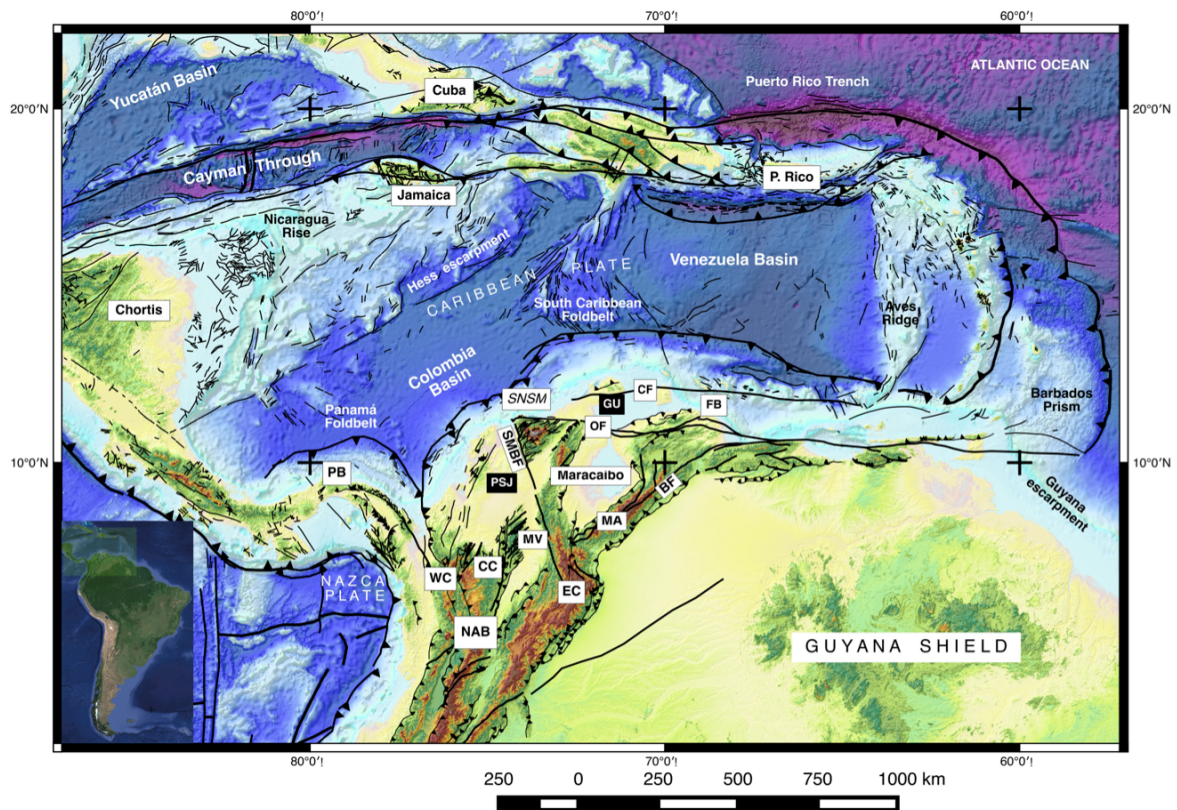


Figure 1. Regional tectonic map of the Caribbean realm showing the most relevant geological features intervening in the plate boundary configuration. after (Veloza et al., 2012). Elements of the North Andean Block (NAB): Maracaibo block, SNSM: Sierra Nevada de Santa Marta, OF: Oca Fault, CF: Cuisa Fault, SMBF: Santa Marta-Bucaramanga Fault, GU: Guajira Basin, PSJ: Plato-San Jorge Basin, MA: Mérida Andes, BF: Boconó Fault, FB: Falcon Basin, WC: Western Cordillera, CC: Central Cordillera, EC: Eastern Cordillera, MV: Magdalena Valley, PB: Panamá Block.

In this paper we present a new low-temperature thermochronology dataset for the SNSM massif, which we combine with previously published apatite fission-track (AFT) data of Villagómez et al., 2011b, and apatite (U-Th)/He by Cardona et al., 2011a. Our aims are 1) to constrain erosion rates during a first phase of rapid post-magmatic cooling 2) to identify the major role played by thrust faults during lower crustal extrusion, 3) to evaluate the magnitude of exhumation of three different tectonic

blocks separated by master faults that are inherited reactivated sutures, 4) to evaluate new and pre-existent thermochronological data in order to assess a tectonic model for Eocene-Miocene exhumation of the metamorphic complexes, and 5) to establish regional correlations between the SNSM, the Merida Andes and the Eastern Cordillera of Colombia.

2. TECTONIC SETTING

The interaction between the Caribbean and South American plates can be traced to Campanian times, and is directly related with the Caribbean Large Igneous Province (CLIP) accretion into the western margin of South America (Kerr et al., 1997; Spikings et al., 2015; Villagómez et al., 2011a). The orogenic evolution of the Northern Andes during the Cenozoic is the result of a complex interaction between the Caribbean, Nazca, Cocos and South American plates (Fig. 1). The transit of the Caribbean plate from the Pacific around the NW corner of South America constituted a major event that is recorded at least since Late Cretaceous (Campanian) times, when the old crust of the Pacific plate was incorporated in a juvenile crust of the nascent Caribbean plate. Between 139-69 Myr the CLIP, an oceanic plateau, formed above the Galapagos Plume by the interaction of multiple oceanic intraplate igneous structures (Hoernle et al., 2004; Kerr et al., 1997). The origin of this plateau has been attributed to the rise of an asthenospheric plume that induced subduction at its margins, which later evolved into the Great Caribbean Arc and the Central American Volcanic arc systems during the Late Cretaceous (Whattam and Stern, 2015). The CLIP approached NW South America

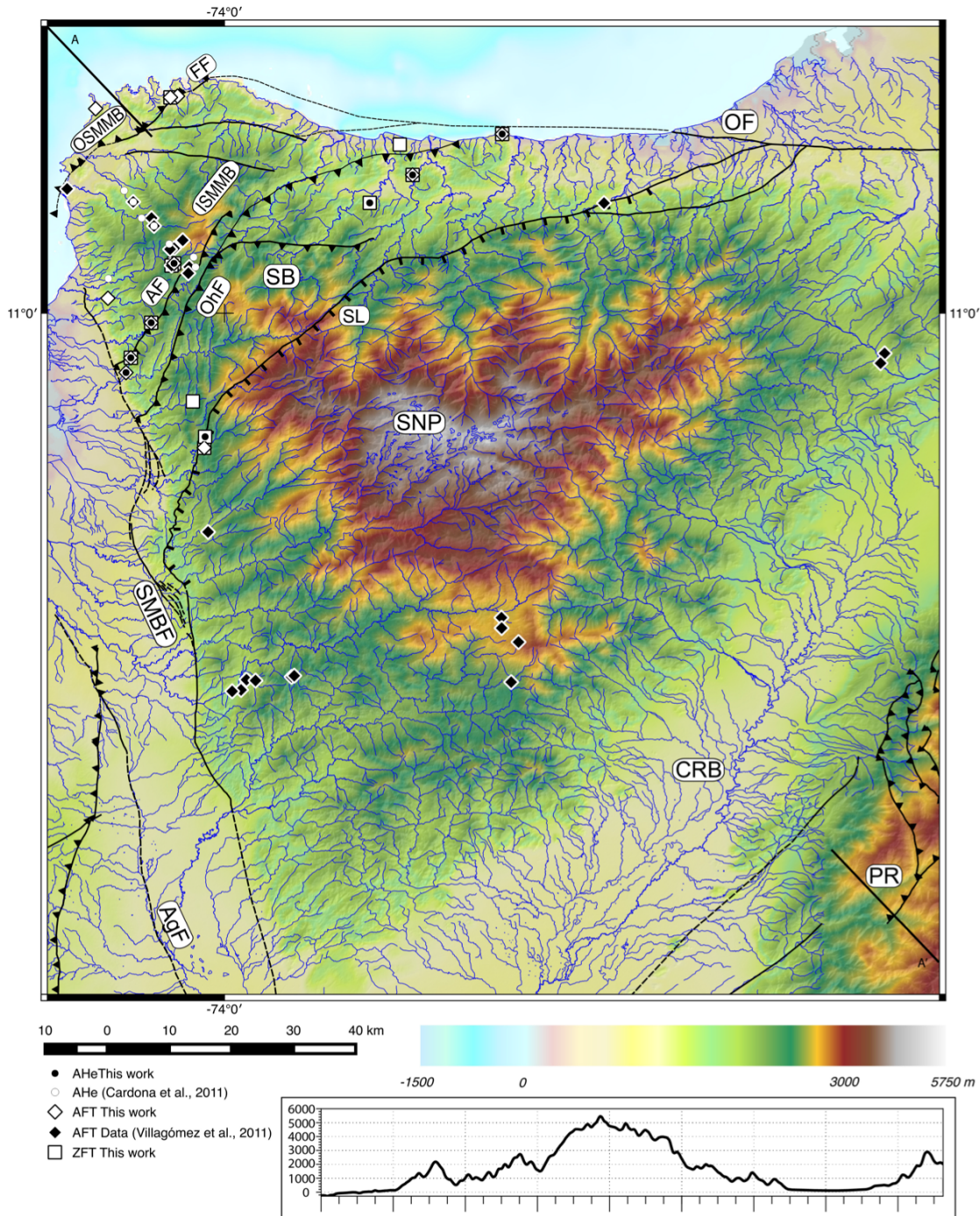


Figure 2. Digital elevation map (DEM) of the Sierra Nevada de Santa Marta Massif, including sample localities from previous studies and the new acquired samples. Major structural elements are indicated: OB: Outer Santa Marta Metamorphic Belt, IB: Inner Santa Marta Metamorphic Belt, SB: Sevilla Metamorphic Belt, SL: Sevilla Lineament, SNP: Sierra Nevada Province, SMBF: Santa Marta Bucaramanga Fault, OF: Oca Fault, AF: Aguja Fault, OhF: Orihueca Fault, FF: Florin Fault, CRB: Cesar Ranchería Basin, PR: Perija Range, AgF: Algarrobo Fault. the black thick lines at the corners indicate the trace of the topographic profile at the bottom.

and its marginal arcs collided against the continental margin during Late Cretaceous-Paleogene times. Remnants of this collision can be traced along Ecuador the Western Cordillera of Colombia, the SNSM massif, the Guajira Peninsula, and the Leeward Antilles (Cardona et al., 2010; Duque-Caro, 1979; Jaillard et al., 2009; Spikings et al., 2015; Vallejo et al., 2009; van der Lelij et al., 2010; Villagómez et al., 2011a) The collision of this massive arc involved obduction and accretion of the oceanic arc into the continental crust. Interaction of the Caribbean plate with the South American plate caused fragmentation of the continental margin at the northern Andes as plate coupling increased with the advance of the Caribbean plate towards the E. The initiation of a low-angle SE dipping subduction zone resulted in the development of a magmatic arc along NW South America during Paleocene-Eocene times (Kennan and Pindell, 2009). Magmatism manifests in felsic plutons that were emplaced in the oceanic plate, as the Parashi Stock and the Mandé Batholith in the Guajira Peninsula and the Western Cordillera respectively (Cardona et al., 2011a, 2014; Duque, 2009; Villagómez et al., 2011a; Weber et al., 2010). In the SNSM massif, the Eocene plutons were emplaced in a transitional crust, and within the Precambrian basement producing wide migmatite zones.

The SNSM massif, with 5775 m elevation is the highest mountain in the Caribbean realm (Fig. 2), is an isolated tectonic block. Mainly an assemblage of Precambrian to Late Paleozoic metamorphic rocks, Jurassic-Cretaceous granitoids, and remnants of Caribbean plate metabasites constitutes this remarkable topographic element. These arc-units were obducted onto tilted Permo-Triassic lower crustal rocks. The different units have been grouped in four major morphotectonic provinces: 1) the Sierra Nevada Province (SNP) that corresponds to Precambrian granulites, anorthosites and gneisses

related to the Grenvillian-Sunsas-Putumayo orogenic belt with nested Jurassic to Cretaceous plutons (Cardona et al., 2006; Cordani et al., 2005; Ibanez-Mejia et al., 2011; Kroonenberg, 1982; Ordóñez et al., 2002; Tschanz et al., 1974), 2) the Sevilla Metamorphic Belt (SB) composed by Neoproterozoic gneisses probably associated with the Pan-African episode (Chapter 1), 3) The Inner Santa Marta Metamorphic belt (IB) composed of, Permian orthogneisses, Triassic mylonites and Jurassic metapelites (Chapter 1), 4) The Outer Santa Marta Metamorphic belt (OB) at the northwestern corner of the SNSM composed of accreted Caribbean metavolcanic rocks as supported by all geochronological data available (Cardona et al., 2010); (Chapter 1).

Internally the relations between the provinces that conform the SNSM massif are very complex and involve a Grenvillian basement of the Sierra Nevada Province (SNP) amalgamated to the late Neoproterozoic complex of the SB through an elusive structure known as the Sevilla lineament (SL). The SB thrusts over the Permo-Triassic (IB) through a major shear zone denominated the Orihueca fault (OhF). Internally to the IB a thrust fault exposes a mylonite belt of Triassic age and is named as the Aguja fault. At the northernmost tip of the massif, an accreted Caribbean oceanic sliver (OB) is thrust along the Florin fault (FF) onto the continental margin. Paleocene-Eocene plutonism affected all provinces, and great volumes of granite were emplaced in the Santa Marta Batholith (SMB), the Buritaca Pluton (BP), the Latal Stock (LS), and the Sevilla Stock (SS) (Fig. 3).

In this context, the Paleocene-Eocene plutonic episode is extremely relevant because its thermal overprint is recorded in low- and high-temperature thermochronological systems. This thermal event masked most of the pre-Caribbean thermal events that

took place on the continental margin. During the late Eocene cessation of plutonism was directly linked to Caribbean slab flattening (Bernal-Olaya et al., 2015).

During the Eocene- Miocene, the SNSM was dragged along the continental margin to its current position as a consequence of the tectonic force exerted by the continuous convergence between the Caribbean plate and South America (Ayala et al., 2012; Bayona et al., 2010). The clockwise rotation of the SNSM led to the opening of the Lower- Magdalena valley basin until Oligocene times, when convergence was partitioned. Then the current boundaries of the SNSM were established, the Santa Marta-Bucaramanga fault on the western flank and the Oca fault on the northern flank of the massif (Bernal-Olaya et al., 2015; Flinch, 2003; Flinch and Castillo, 2015; Montes et al., 2010).

Accretion of the Great Caribbean arc against northwestern South America highly influenced the end of subduction (Vence, 2008). To the east of the Guajira Peninsula, the Leeward Antilles underwent rotation during their transit around the northwestern corner of South America at ca. 45 Ma. Contemporaneously the basement was exhumed next to the continental margin (van der Lelij et al., 2010), and the Great Caribbean arc was segmented along major extensional faults, creating basins that received up to 3000 m of sediments. This process marked the final stages of the accretion of the Great Caribbean arc (Gorney et al., 2007).

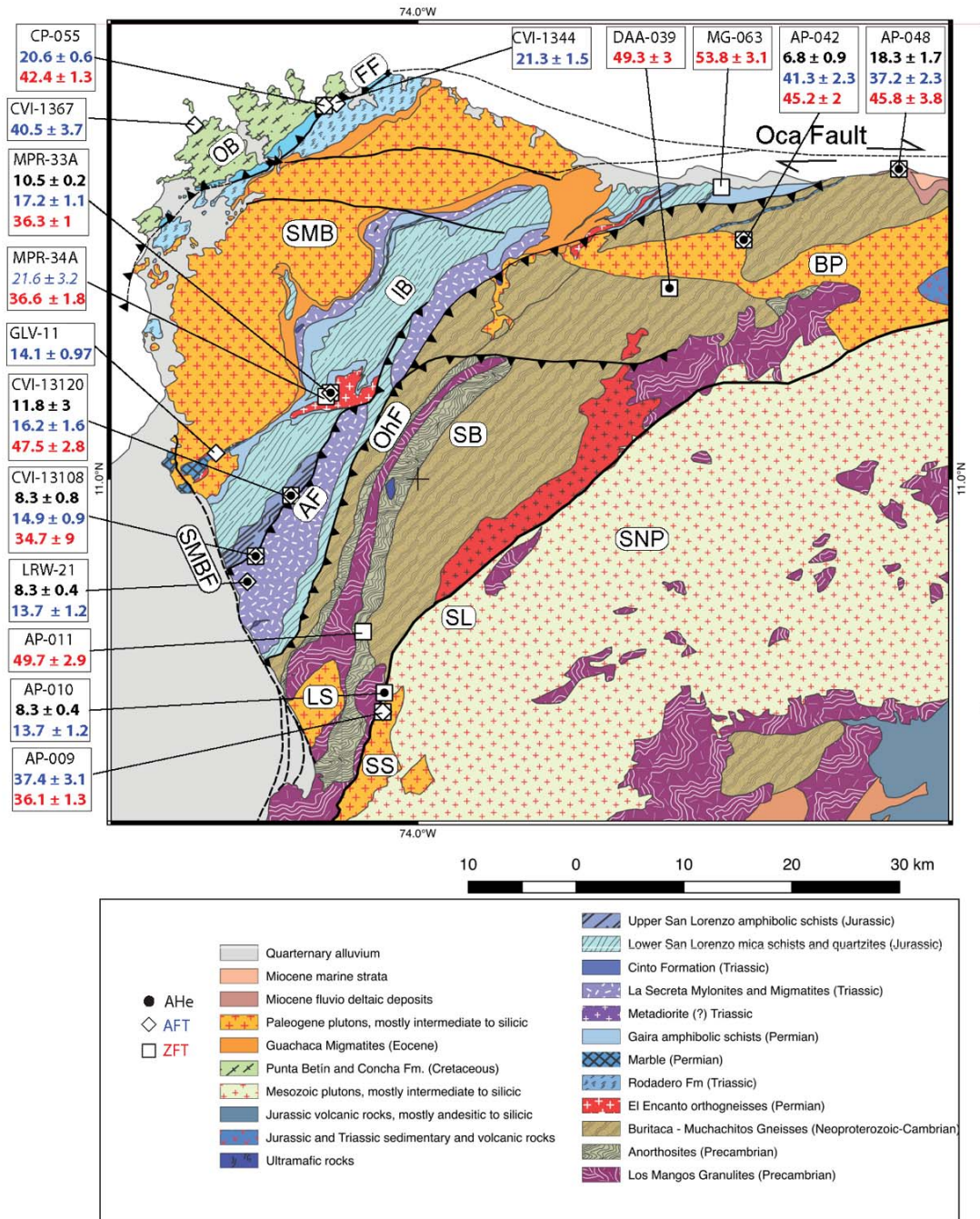


Figure 3. Geological Map of the NW metamorphic belt of the Sierra Nevada de Santa Marta Massif (SNSM) with location of samples analyzed in this study, color labels represent each thermochronometric system. OB: Outer Santa Marta Metamorphic Belt, SMB: Santa Marta Batholith, IB: Inner Santa Marta Metamorphic Belt, SB: Sevilla Metamorphic Belt, SNP: Sierra Nevada Province, BP: Buritaca Pluton, SS: Sevilla Stock, SL: Sevilla Lineament, LS: Latal Stock, SMBF: Santa Marta Bucaramanga Fault, AF: Aguja Fault, OhF: Orihueca Fault, FF: Florin Fault.

The Santa Marta-Bucaramanga fault (SMBF) separates the crystalline basement of the SNSM from the Oligocene-Miocene Lower Magdalena basin to the west (Bernal-Olaya et al., 2015; Montes et al., 2010; Tschanz et al., 1974; Villagómez et al., 2011b). The SMBF is a long-lived major discontinuity that extends more than 500 km, and connects the Southern Caribbean plate boundary with the Santander Massif in the Eastern Cordillera of Colombia (Fig. 1). The Santa Marta Bucaramanga fault was active during the Early Jurassic rifting episode, accommodating an important amount of right lateral displacement, and was reactivated during the Neogene (Kammer and Sánchez, 2006). During reactivation it accommodated strain from the Caribbean plate into the North-Andean block driving the surface uplift and exhumation of the Eastern Cordillera since the Miocene, with a major pulse during the Pliocene (Cortés et al., 2005; Parra et al., 2009). Implying that the SMBF is perhaps the most relevant structure that controlled the evolution of the Northern Andes from the Early Jurassic to the late Pliocene. Previous thermochronological studies on the SNSM had shown six time intervals in which the SNSM exhumed at elevated rates 0.2 -0.7 km/Myr, in response to the collision of the Caribbean plateau with northwestern South America, subsequent underthrusting of the Caribbean plate, and late propagation of NW verging thrust faults (Villagómez et al., 2011b). In the following sections, we present our thermochronological data, combined with data from previous studies, in the context of the redefined tectonic provinces that make-up the SNSM, and its exhumation histories constrained from middle Eocene to late Miocene times.

3. METHODS

3.1 Bedrock fission-track thermochronology

3.1.1 Zircon fission track

Zircons of the 125-250 μm fraction, were mounted in Teflon® sheets, with at least two mounts per sample. Zircons were polished to expose internal grain surfaces and etched in a NaOH-KOH solution at 228°C. Etching time varied between 20-60 hours in order to reveal countable tracks in at least 100 grains per sample. Samples were irradiated together with Fish Canyon Tuff and Buluk Tuff zircon age standards and IRMM541 dosimeter glasses.

3.1.2 Apatite fission track

Apatite crystals were mounted in epoxy, polished and etched in 5.5 M HNO_3 at 21°C for 20 s and covered with muscovite detectors. The samples were irradiated together with Fish Canyon Tuff and Durango Tuff age standards and IRMM540R dosimeter glasses. Both apatite and zircon samples were irradiated at the FRM II research reactor at München, Germany.

3.1.3 (U-Th)/He in Apatites

Crystals separated from the metamorphic and igneous rocks of the SNSM were prepared for (U-Th)/He analysis at the Institute of Geochemistry and Petrology at ETH Zürich. The objective was to obtain euhedral, inclusion free, colorless crystals without fractures. Selected apatite crystals were photographed, measured and wrapped in Pt foil. Apatites were first analyzed for ^4He by degassing in a static vacuum using a diode laser heating to temperatures around 1000°C, and using a sector field mass

spectrometer equipped with a Baur-Signer ion source. Gases were pre-cleaned with liquid nitrogen traps and zirconium-vanadium-iron getters (SAES, St707). Net signal intensities were interpolated to the inlet time of the gas into the mass spectrometer, and then compared to the corresponding mean signal from reference gas aliquots of known absolute amounts analyzed by the same procedure. The degassed apatites were retrieved, dissolved in HNO₃ at ~90°C for 1 hour, spiked with ²³³U, ²³⁰Th, and ¹⁴⁹Sm, and analyzed on an ElementXR ICP-Mass spectrometer for ²³⁸U, ²³²Th, and ¹⁴⁷Sm.

4. RESULTS

In this section, results from ZFT, AFT and (U-Th)/He analyses obtained from samples of the distinctive lithologies of the northwestern corner of the SNSM are presented. Our aim was to perform these analyses on all the same samples analyzed with the U-Pb LA-ICP-MS method for Chapter-1. Nevertheless, because of several different factors, such as U content and/or grain quality for example not all of these samples could be analyzed. In some cases, zircons were scarce and only served for U-Pb dating. Apatites in most of the metamorphic rocks were severely damaged and fractured, and presented low U content. Regardless of these inherent complications, we were able to obtain 31 thermochronological ages (Fig. 3, supplementary Tables A1, A2, and A3). The results are presented here according to structural position of the sampled rock within the different lithotectonic units of the SNSM massif. Two-sigma analytical errors are reported in this study.

4.2 Sevilla Metamorphic Belt (SB)

Five samples from the high-grade metamorphic gneisses of the Buritaca and Muchachitos gneiss suite and one sample of the Eocene Sevilla Stock at the SB in the hanging wall of the Orihueca fault were selected for obtaining zircon and apatite ages (Fig. 3). Fission track ages are quite similar along the SB in both AFT and ZFT systems. Apatite yield was low, so only two samples were triple dated by ZFT, AFT and AHe; two samples were double dated by ZFT and AHe and one sample was double dated by ZFT, AFT. (Fig. 3, 4 & 5).

4.2.1 ZFT

Quartz-feldspathic gneiss sample AP-048 from the northwestern most part of the unit at sea level next to the Oca fault has an age of 45.8 ± 3.8 Ma, whereas amphibole gneiss sample DAA-039 yields an age of 49.3 ± 3 Ma. For the Buritaca Pluton a cooling age of 45.2 ± 2 was obtained. Amphibole gneiss sample AP-011, from the western flank of the SNSM has a ZFT cooling of 49.7 ± 2.9 Ma, and 45.3 ± 1.9 Ma for mafic gneiss sample AP-010. The Sevilla Stock sample AP-009 presented a slightly younger cooling age of 36.1 ± 1.3 Ma.

4.2.2 AFT

Sample AP-048 has a pooled age of 37.2 ± 2.3 Ma, and was the only sample of Precambrian rock that could be analyzed with the apatite fission-track method, as crystals from the remaining samples were highly damaged, or presented to many flaws for counting. The Sevilla Stock sample AP-009 has an age of 37.4 ± 3.1 , which is

almost identical to its ZFT age. The Buritaca Pluton tonalite has a cooling age of 41.3 ± 2.3 Ma.

4.2.3 Apatite (U-Th)/He

Sample AP-048 has an average age of 18.3 ± 1.7 Ma replicated from five grains, whereas sample AP-010 from the western flank of the SNSM yields an age of 13.6 ± 0.9 Ma. at one grain. The Eocene Buritaca Pluton yields a younger age of 6.8 ± 0.89 Ma replicated from five grains.

4.3 Inner Santa Marta Metamorphic Belt (IB)

Zircons and apatites were separated from seven samples (Fig 4 & 5) from the meta-igneous and metasedimentary rocks of the IB, either in the hanging wall or the footwall of the Aguja fault (Fig. 3). From these samples, two belong to the La Secreta Mylonites unit, one to the El Encanto Orthogneiss, three to the Gaira Schists, and one to the Upper San Lorenzo Schists.

4.3.1 ZFT

Sample MPR-33A from the El Encanto Orthogneiss yield a cooling age of 36.3 ± 1 Ma. Sample MPR-34A from the Gaira Schists yields a cooling age of 36.6 ± 1.8 Ma. Sample CVI-13120 of the Upper San Lorenzo Schists in the footwall of the Aguja fault yields an age of 47.5 ± 2.8 Ma. Sample CVI-13108 Bt+Grt+Pl schist from the hanging wall of the Aguja fault yields a cooling age of 34.7 ± 1.9 Ma.. Sample MG-063 from the Gaira Schists from the northern flank of the SNSM yields a cooling age of 53.8 ± 3.1 Ma.

4.3.2 AFT

El Encanto Orthogneiss sample MPR-33A yields a cooling age of 17.2 ± 1.1 Ma, whereas sample GLV-11 from the Gaira Schists yields an age of 14.1 ± 0.97 Ma. Sample CVI-13120 from Upper San Lorenzo Schists shows a cooling age of 16.2 ± 1.6 Ma. In the hanging wall of the Aguja fault at La Secreta Mylonites sample CVI-13108 shows a cooling age of 14.96 ± 0.9 Ma, and sample LRW-21 yields a similar age of 13.7 ± 1.2 Ma.

4.2.3 Apatite (U-Th)/He

From the IB apatites were recovered from four samples. Sample MPR-33A from the El Encanto Orthogneiss has an age of 10.5 ± 0.2 Ma replicated from five grains. Sample CVI-13120 from Upper San Lorenzo Schists yields a cooling age of 11.8 ± 3 Ma replicated from three grains. Samples LRW-21 and CVI-13108 from La Secreta Mylonites yield identical cooling ages of 8.3 ± 0.4 Ma and 8.3 ± 0.8 Ma respectively replicated from five grains.

4.4 Outer Santa Marta Metamorphic Belt (OB)

Zircon and apatite grains were separated from three samples of low-grade metamorphic rocks from the OB. Two of the samples are from the Cinto Formation and one sample from to the Concha Formation (Fig. 3). In the units of the OB apatite and zircon content is very scarce. For this reason there were not enough apatite crystals for (U-Th)/He dating, and only one of the samples yielded enough zircons for fission track analysis (Fig. 4 &5).

4.4.1 ZFT

Sample CP-055, from the Cinto Formation has a cooling age of 42.4 ± 1.3 Ma.

4.3.2 AFT

Sample CP-055 from the Cinto Formation yields a cooling age of 20.6 ± 0.66 Ma, and sample CVI-1344 from the most external part of the SNSM massif yields a cooling age of 40.5 ± 3.7 Ma,

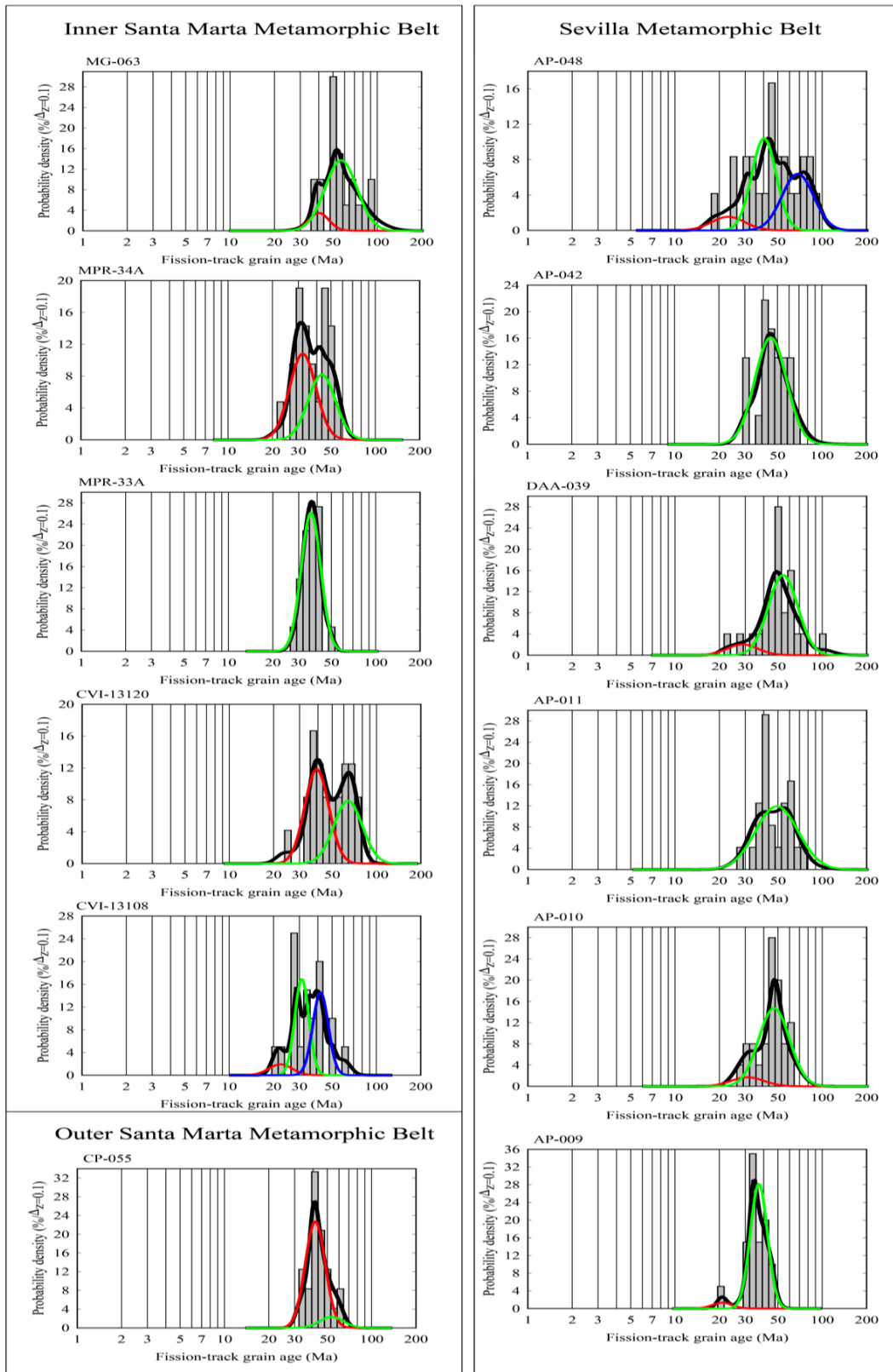


Figure 4. Bedrock ZFT data from the SNSM, probability density plots elaborated with Binomfit (Ehlers et al. 2005; Stewart & Brandon 2004).

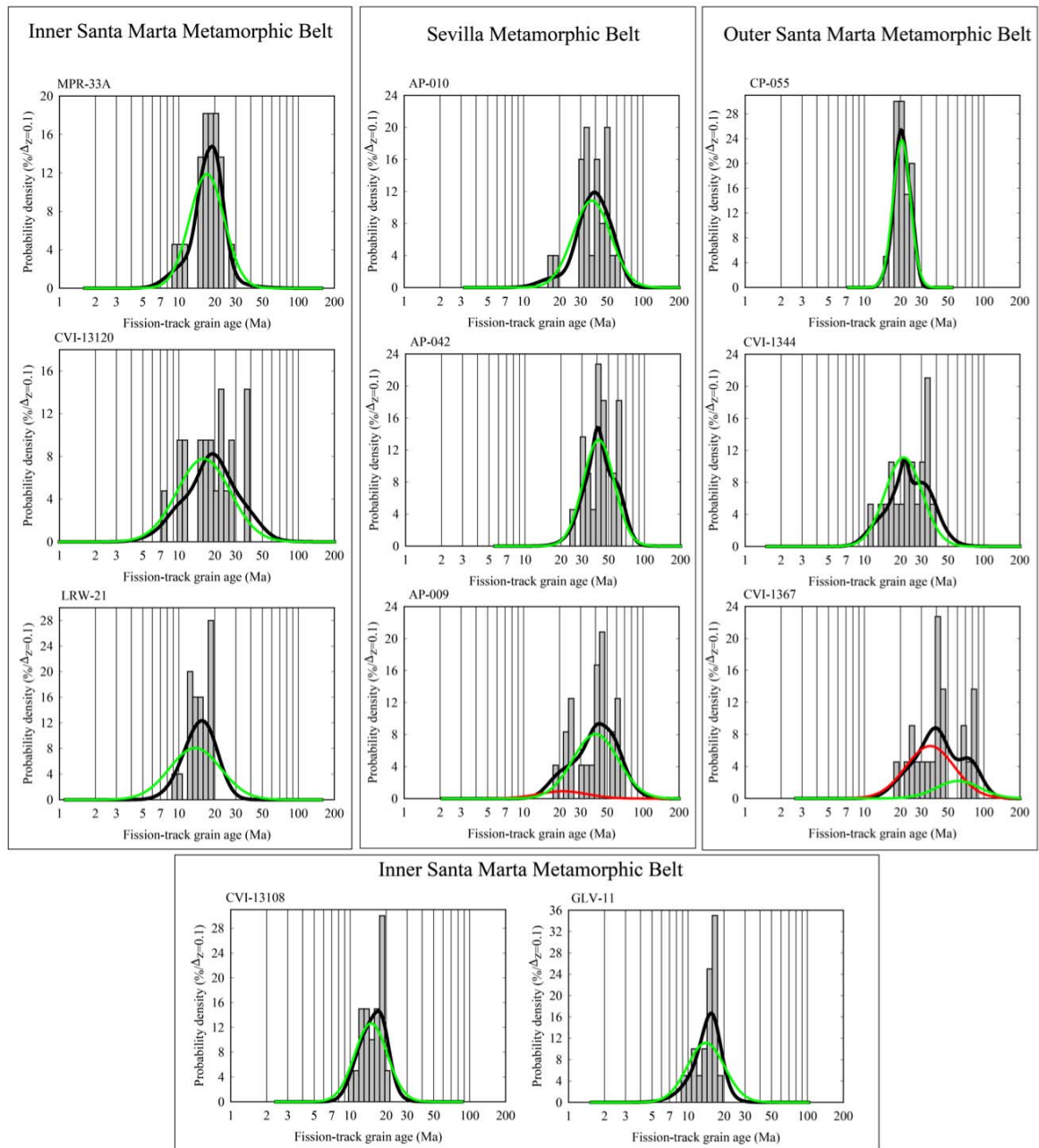


Figure 5. Bedrock AFT data from data from the SNSM, probability density plots elaborated with Binomfit (Ehlers et al. 2005; Stewart & Brandon 2004).

5. DISCUSSION

5.1 Accretion of Caribbean Terranes and onset of magmatism

The beginning of exhumation in the SNSM corresponds the later stages of diachronic accretion of the Caribbean Plateau to the Pacific margin of the South American plate between 65 and 58 Ma (Villagómez et al., 2011b). Zircon U-Pb ages of 69.3 ± 0.5 Ma for the first dikes intrusions in metamorphic rocks of the OB and the youngest ages ca. 55.4 ± 2.6 Ma for andesitic arc volcanism (Chapter 1) along with crystallization ages for the SMB and associated plutons between 65-49 Ma (Duque, 2009). It shows that arc accretion and pluton emplacement occurred contemporaneously in a time span of ≈ 15 Ma.

Paleogene plutonism affected the OB, the IB, the SB and the SNP. For this reason the AHe, AFT and ZFT systems were regionally reset and annealed in the SNSM in the vicinity of the Paleocene-Eocene plutons

During the late stages of Caribbean and South American plate convergence, the intrusion of the Santa Marta Batholith occurred within the upper plate and transitional crust. Pluton emplacement continued until about 50 Ma and rocks cooled at fast rates of $97\text{-}80$ °C/Myr until about 48 Ma (Duque, 2009). The thermal overprint produced by the Eocene plutons caused migmatization and skarns in the metapelitic Permian-Jurassic rocks. The Precambrian basement of the Sevilla Belt and Sierra Nevada Provinces was affected by mafic magmatism producing hornblendite clusters around plutons. P-T conditions for the pluton emplacement estimated from Al-in-amphibole content, yield values of 4.9 ± 0.6 kb and 6.4 ± 0.6 kb (Cardona et al., 2011a), indicating at

least 16.6 -21.8 km depth with an ΔP of 0.2493 b/Km assuming an average crustal density of 3.0 g/cm³ given the geological context of the SNSM massif. For the Buritaca Pluton, thermobarometry analyses performed on sample AP-042 (Mateus, 2014) with the Qtz+Hbl+Bt+Pl calibration yield P-T conditions of 3.18 kb and 762°C for an approximated depth of 10.8 km (Chapter 2).

5.2 Erosion rates

Thermal histories had been modeled recently for the SNSM (Villagómez et al., 2011b) using apatite fission track length using the HeFTy program of Ketchum (2005). We constructed thermal paths for our samples (Fig. 6 & 7) for three thermochronological systems (ZFT, AFT, AHe). Erosion rates, closure depths and cooling rates were calculated with the Age2Edot program of Brandon (see Ehlers et al., 2005; Willett and Brandon, 2013), using a geothermal gradient, the current surface temperature and cooling ages. We defined closure temperature depending on cooling rates for the different low-temperature thermochronometers (Ehlers, 2005).

Recent estimates on the geothermal gradient for the SNSM obtained from the geothermal map of Colombia and from curie depth values (CDP) revealed a thermal gradient in the range of 12-18 °C/km (Vargas et al., 2009, 2015). Nonetheless, thermal gradients during the Eocene should have been higher in consequence of the Paleocene-Eocene island to continental arc. Given its tonalite-trondhjemite-granodiorite (TTG) series character, the SMB required a high thermal gradient during emplacement \approx 40°C/km (Duque, 2009). Furthermore, the different plutons (SMB, SS, and BP) intruded at different crustal depths. In this context the Cenozoic paleo-thermal gradients estimated for the SNSM, which have been estimated between 20-

40°C (Cardona et al., 2011a; Ceron-Abril, 2008; Mendoza and Ojeda, 2006; Villagómez et al., 2011b), should be analyzed carefully for calculating cooling rates regarding the timing and depth of magmatism according to temperature paths (Fig. 6). During pluton emplacement a gradient of 62°C/km is assumed, whereas during post magmatic cooling the geothermal gradient decreased progressively to about 20°C/km. It is very important to acknowledge that cooling occurred diachronously across the SNSM. First at the Buritaca Pluton in the north and subsequently in the plutons of the SMB with later cooling of the Sevilla Stock in the south-western flank of the SNSM (Fig. 6), this cooling trend is the consequence of plutons of the same age intruded at different depths with even 8 km of difference in intrusion depth between the northern and western flanks of the SNSM.

For this reason, geothermal gradients used for calculating erosion rates take into account the magmatic activity and its influence at least until ca. 45 Ma at the northern flank of the SNSM (Buritaca Pluton), and 37 Ma at the western flank of the SNSM (Sevilla Stock) as constrained by ZFT.

After assigning a geothermal gradient according to the chronological and structural context, we estimated erosion rates as a linear function on the change in temperature ΔT vs change in time Δt . For a given sample in this way defining a cooling rate for a particular interval of time, this is for example the amount of time in Ma for a particle to travel from the isotherm of closure temperature for zircon U-Pb crystallization $\approx 700^\circ\text{C}$ calculated from Ti-in-Zr thermometry (Chapter 1), until the ZFT closure temperature $\approx 350^\circ\text{C}$, such elevated temperature corresponds to fully reset zircons with zero radiation damage and is applied to the Eocene plutons samples.

Thermal paths Sevilla Belt (SB)

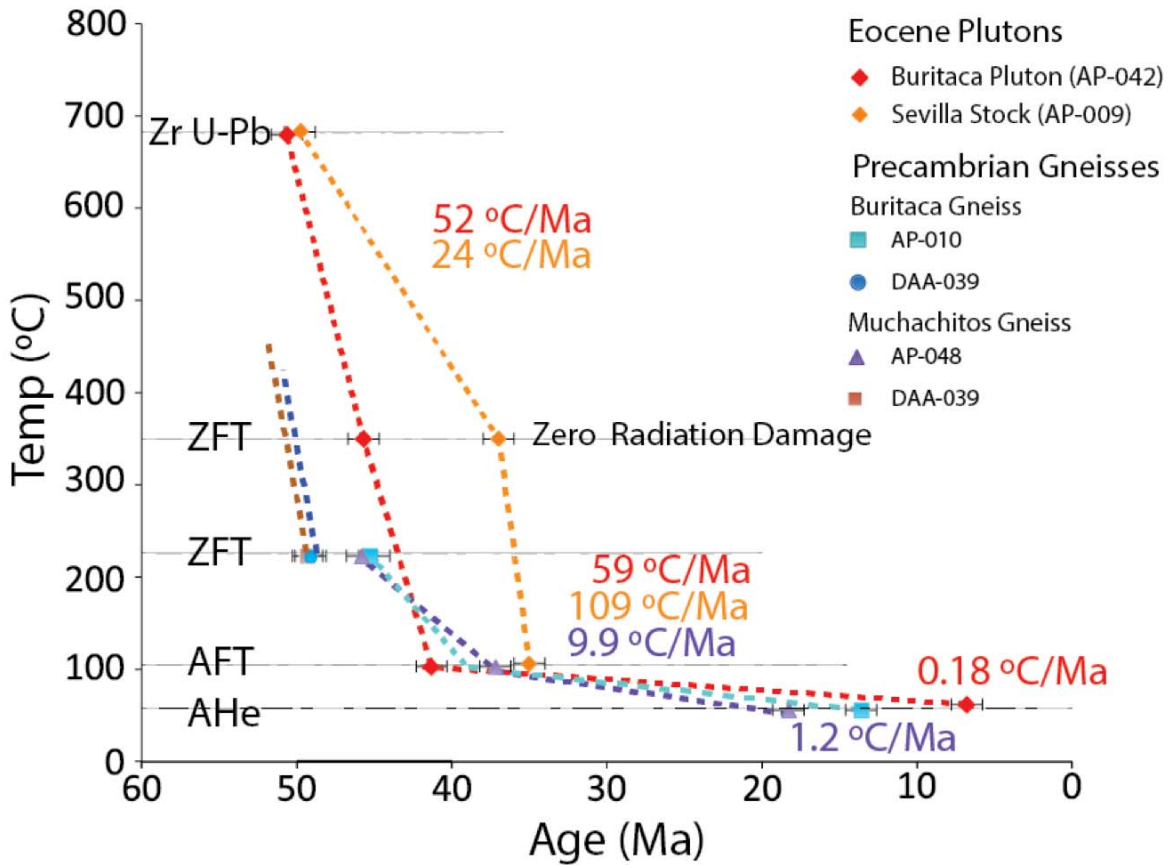


Figure 6. Temperature-time Paths of the Sevilla Metamorphic Belt (SB), constructed with ZFT, AFT and (U-Th)/He cooling ages.

Post magmatic cooling at the Sevilla Belt occurred, between 50–45 Ma at rates of ≈ 0.8 km/Myr, and cooling rates of about $52^{\circ}\text{C}/\text{Myr}$ documented for the Buritaca Pluton estimated from $\Delta T/\Delta t$.

Thermobarometry estimates that the pluton intruded at 10.8 km depth (Mateus, 2014) at 50.7 Ma and cooled at elevated rates of $52^{\circ}\text{C}/\text{Ma}$ until reaching closure temperature of 342°C at ≈ 5.52 km which is the calculated closure depth for the ZFT system for a T_c of 342°C , considering zero radiation damage zircon occurring at 44.9 Ma. This accounts for an exhumation rate of 0.9 km/Myr estimated from $\Delta Z/\Delta T$. These very

similar results indicate that initial post magmatic cooling was highly influenced by exhumation. Erosion rates increased to 2 km/Myr between 45 - 40 Myr, during the establishment of a decreasing geothermal gradient ($\approx 30^{\circ}\text{C}/\text{km}$) as evidenced by the increase of cooling rates to $59^{\circ}\text{C}/\text{Ma}$. The last cooling phase is recorded between 40-5 Ma estimated on AHe data with a stable geothermal gradient of $20^{\circ}\text{C}/\text{km}$ and low erosion rates $\approx 0,01$ km/Myr (Fig. 6).

On the western flank of the SNSM the Sevilla Stock shows fast cooling interval between 49.8- 37 Ma with values of 0.4 km/Myr and cooling rates of $24^{\circ}\text{C}/\text{Myr}$, related to a latest stages of Eocene plutonism. These values are coherent with rates of 0.73 km/Myr estimated for the SMB between ca. 50-41 Ma (Cardona et al., 2011a). The Sevilla Stock cooled at a very high rate of $109^{\circ}\text{C}/\text{km}$, $\approx 12\text{Ma}$ after crystallization, we interpret that this rapid cooling was driven by erosional exhumation between 37-35 Myr at an increasing rate of 2.7 km/Myr. Differently from the other units of the Sevilla Metamorphic Belt (Fig. 6), this rate evidences the onset of intense erosive process at this part of the SNSM. The Precambrian Muchachitos gneiss hosting the Buritaca Pluton presented cooling rates of $10^{\circ}\text{C}/\text{Myr}$ between 46-37 Ma, which can be transferred into erosion rates about 0.32 km/Myr during that period. The Buritaca Pluton shares a similar thermal history, showing a significant decrease in cooling rate after this epoch and reaching erosion rates of 0.03 km/Myr (Fig. 6). The Buritaca Gneiss at the western flank of the SNSM shares a similar thermal path with respect to the previously described Muchachitos Gneiss and with erosion rates of 0.14 km/Myr at ca. 45.8 Ma and 0.08 km/Myr after 45.8 Ma that had persisted until the present.

Permian, Triassic and Jurassic meta-igneous and metasedimentary rocks from the IB evidence as a general trend younger cooling ages in all systems. For these rocks we

assumed a geothermal gradient decreasing from 30 to 20°C/km. The exhumation of the IB postdates the middle Eocene major cooling event that led to exhumation recorded as an unconformity, this is corroborated by their time-temperature path that describes an independent history from the Eocene plutons and the Sevilla Belt rocks.

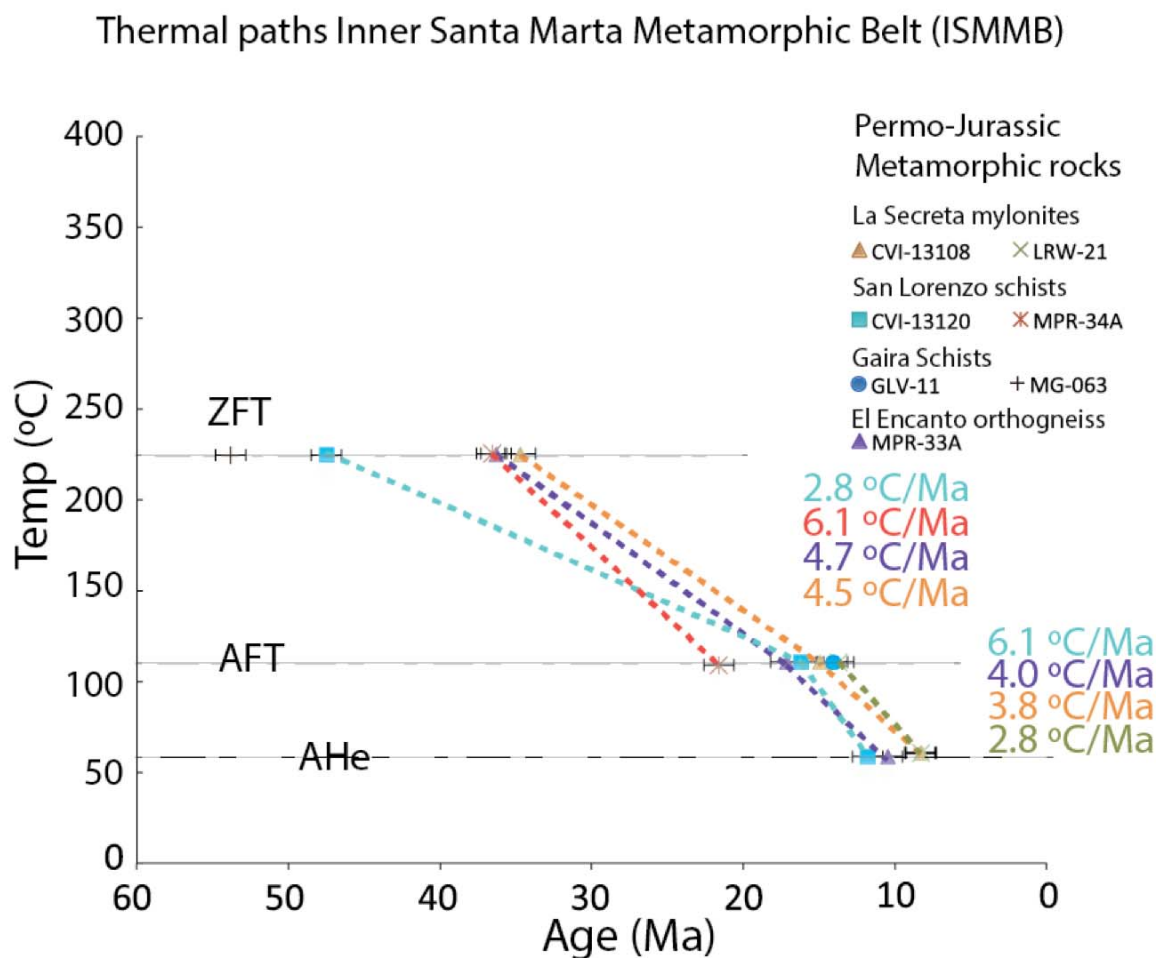


Figure 7. Temperature-time Paths of the Inner Sant Marta Metamorphic Belt (IB), constructed with ZFT, AFT and (U-Th)/He cooling ages.

Between 37-13 Ma all samples from the IB cooled at rates between 3-6°C/Myr, exhuming with erosion rates ranging from 0.09-0.2 km/Ma (Fig. 7). From this group

two samples presented older ZFT ages of 47.5 and 53.8 Ma and showed the slowest cooling rates. The last exhumation phase is recorded in the AHe system between 12-8 Ma with a minor increase in exhumation rates between 0.14-0.3 km/Myr. Meta-volcanic rocks of the OB cooled at rates of 4°C/Myr between 42-20 Ma, with exhumation rates estimated at ≈ 0.15 km/Myr

Our results show that a first phase of rapid cooling and erosion followed the pluton emplacement, this phase occurred between 50-45 Ma in the northern SNSM at rates of 0.8 km/Myr, and later in the western SNSM between 45-35 M at rates of 0.4 km/Myr (Fig. 6). A second exhumation phase evidences its erosional influence by the increase in cooling and therefore in erosion rates. At the northern SNSM with values of 2 km/Myr during 45-40 Myr whereas at the western SNSM rates reached 2.7 km/Myr during 37-35 Myr defining diachronous elevated rates for the middle Eocene. A third phase of exhumation took place between 32-19 Ma with decreased exhumation rates between 0.09-0.2 km/Myr. The last exhumation pulse occurred between 17-8 Ma as constrained by apatite (U-Th)/He data, and with slow cooling at a slower exhumation at 0.09 -0.03 km/Myr in the northern SNSM massif, in contrast with more elevated erosion rates in the western SNSM of 0.14-0.3 km/Myr (Fig. 7). In synthesis, our results based on three thermochronological systems reveal that during the first phase of exhumation of the SNSM relatively high exhumation rates dominated at the -Sevilla Belt. This phase occurred immediately after post-magmatic cooling for ~ 10 Myr in both the northern and western flanks of the SNSM diachronously with a time difference of about 8 Myr. This difference in time may be related to intrusion depths, as zircon crystallization ages of Eocene plutons are coeval throughout the SNSM. Because thermal histories are similar, we propose that the rapid cooling and

exhumation in the SNSM between 50-35 Ma should be considered as a single phase of post magmatic and exhumational cooling, rather than two independent phases.

5.3 Exhumation driven by crustal channel flow

Accelerated exhumation under high geothermal gradients may have occurred during the formation of a low-viscosity mid crustal channel, in which intruding pods at different depths were fed from a thickened continental crust (Beaumont et al., 2001, 2004; Grujic, 2006), during pluton emplacement, the flow of partially-molten rocks corresponds to lateral channel flow. Such flow may have led eventually to exhumation of migmatites and granites by extrusion along the edges of the orogenic wedge (Vanderhaeghe, 2009). The subducting buoyant Caribbean crust may have favored doming triggered by an underthrust ramp with exhumation of the dome at 1.1 km/Myr. Channel flow is favored by the current structural configuration of the SNSM in which a lower crustal section that preserves greenschist to granulite facies rocks in an inverted position occupies the NW corner of the massif (Fig. 8). Similar configurations of inverted crustal sections hosting granitoids and migmatites are a common feature of channel flow as documented in the High Himalayas (Godin et al., 2006; Jamieson et al., 2004). The “inversion” of the sequence may be syn-metamorphic or post-metamorphic and was most likely caused by ductile shear during strong plate coupling (Hubbard, 1996). As other documented inverted crustal sections, the SNSM is separated by two major shears zones: the lower shear zone with a thrust-sense kinematic is the Orihueca fault, and the upper shear zone dominated by normal-sense kinematics is the Sevilla lineament (Fig. 8).

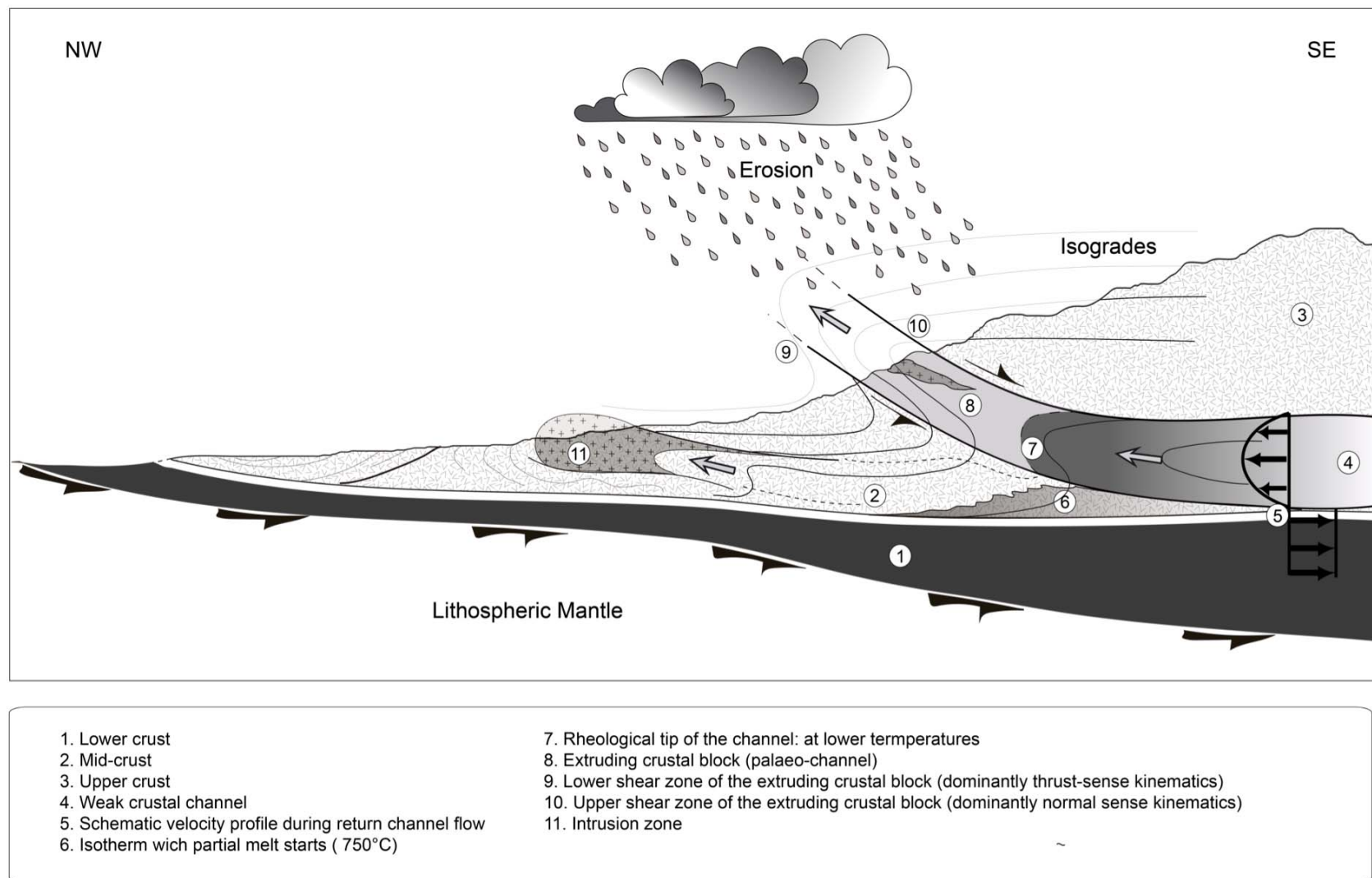


Figure 8. Schematic diagram of kinematic relationship between channel flow and extrusion of a lower crustal channel, applied to the crustal section of the SNSM. Modified from (Godin et al., 2006), in this representation the crustal channel would correspond to the Sevilla Metamorphic Belt bounded by major shear zones, granitoids hosted in the channel and migrated into the upper plate may correspond to Paleocene-Eocene plutons. Erosion enhances lower crustal extrusion.

5.4 Exhumation along major faults

Our data complements the existing SNSM thermochronological database, including a complete new dataset for the SB. The new data shows that the tectonic blocks of the northwestern metamorphic belt have different thermal histories. The boundaries between these tectonic blocks are the Orihueca fault and the Sevilla lineament, and at the tip of the SNSM the Florin fault separates the continental margin sequences from accreted Caribbean slivers (Fig. 9).

After the high-temperature phase of pluton crystallization, post magmatic cooling lasted from 48-35 Ma (Duque, 2009). This cooling age trend is observed in ZFT system in the SB, where the Precambrian rocks cooled around 45-50 Ma below about 350° C. The Sevilla Stock to the SE of the Sevilla lineament shows a much younger cooling age ca. 36.1 ± 1.3 , its cataclastic texture may indicate a syn-tectonic exhumation driven by the Sevilla Lineament.

Elevation profiles validate different thermal histories for the hanging wall and the footwall of the Orihueca fault (Fig. 10). Middle Eocene AFT cooling ages in the hanging wall are juxtaposed against early Miocene AFT cooling ages in the footwall of the Orihueca fault, evidencing that the advancing thrust front was located at this structure, this also evidences that the Miocene exhumation involved the two fault blocks that exhumed as a whole. ZFT, AFT and AHe cooling ages in the IB vary between 33-8 Myr, and are slightly younger than ZFT-AFT ages in the OB ca. 42-20 Ma. These values indicate that after the post-magmatic cooling exhumation in the NW part of the SNSM, the fault blocks were individualized at the middle Eocene.

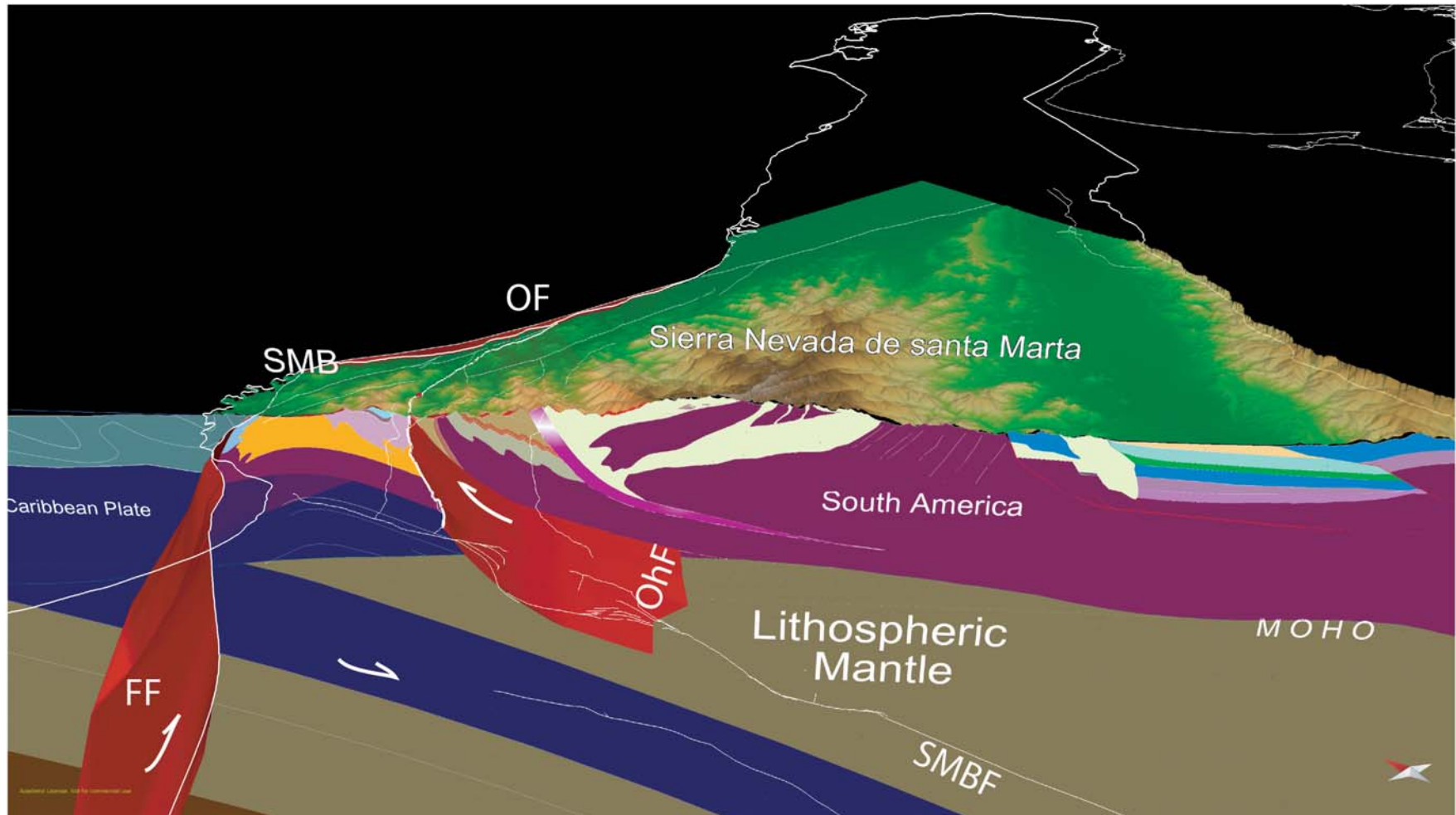


Figure 9. 3D model of the Sierra Nevada de Santa Marta massif, cross section trace corresponds to the topographic profile presented in Fig. 2. OF: Oca Fault, SMBF: Santa Marta Bucaramanga Fault, SMB: Santa Marta Batholith, FF:Florin fault and OhF: Orihueca fault geometries were constructed from serial cross sections, the Sevilla Lineament is indicated as a highlighted polygon in the cross section due the lack of structural data, Moho depth is incorporated in the model from (Camargo, 2014), north direction towards the upper left corner of the render.

In contrast exhumation at the SE flank of the SNSM was already active since upper Paleocene corresponding with coeval deposits at the Cesar-Rancheria basin (Bayona et al., 2011; Montes et al., 2010; Villagómez et al., 2011b), this means that exhumation migrated from continent interior to the active margin.

Exhumation of Caribbean crustal slivers can be traced up to 40 Ma in the hanging wall of the Florin fault in the OB according to ZFT ages. At the OB during 40-20 Myr the metavolcanic rocks obducted the continental crust, by the Florin fault with a constant exhumation rate ≈ 0.12 km/Myr, attesting for continuous accretion-obduction that lasted at least until the early Miocene. These ZFT ages ca. 40 Ma overlap the most intense period of Orihueca fault activity and showed a similar age trend as the IB, implying that Florin fault activity was coeval with IB doming and exhumation. The Florin fault activity postdates the activity of the Orihueca fault and the Sevilla lineament, and predates the activity of the Aguja fault.

Relief increased as consequence of slip through NW verging thrust faults, active since the late Paleocene this process led to erosional exhumation. The major structure involved during this process was the Orihueca fault, which is responsible for at least 11 km of vertical rock exhumation towards the Caribbean since the early Eocene (ca. 50.7 Ma), from the calculated depth of intrusion of sample AP-042 (Mateus, 2014). In the SNP the Sevilla lineament may have acted as a normal shear zone along which at least 6 km of rock were exhumed since the middle Eocene at a rate of about 1.17 km/Myr, evidenced by almost equal cooling ages in the ZFT and AFT systems. Most of the rock was eroded and transported into the Baja Guajira basin, which was displaced by the Oca fault 65 km to the west (Tschanz et al., 1974). Some sediment may have reached the Caribbean Sea, where it was recycled back into the accretionary wedge,

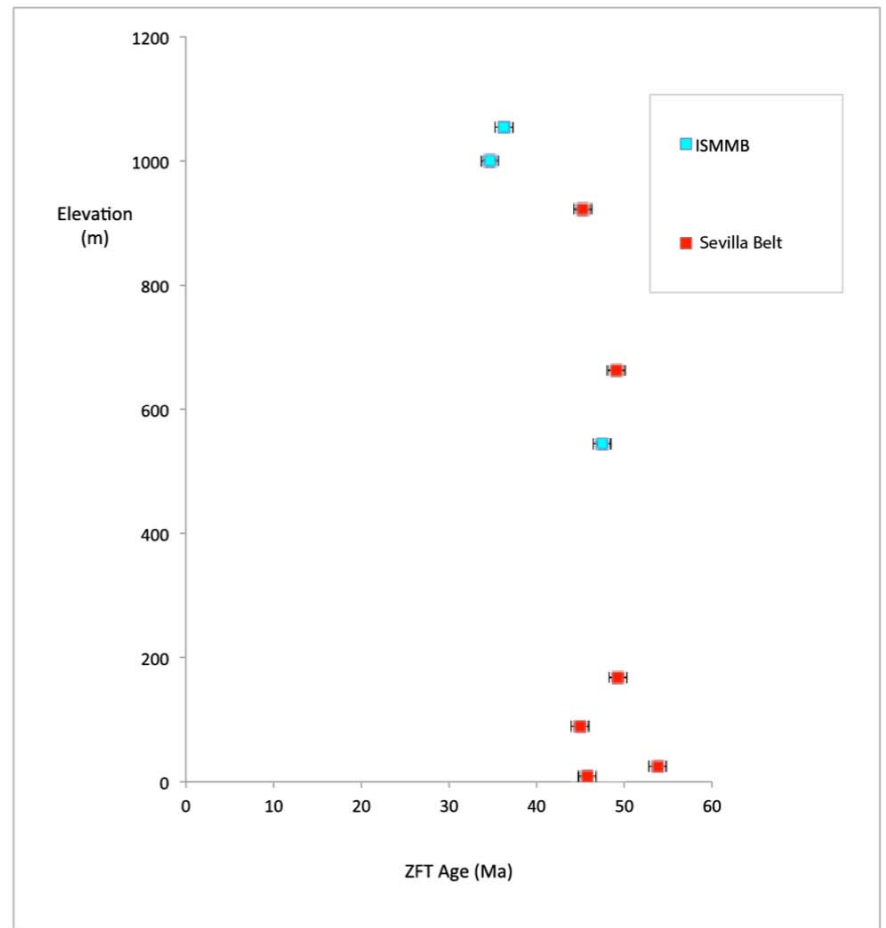
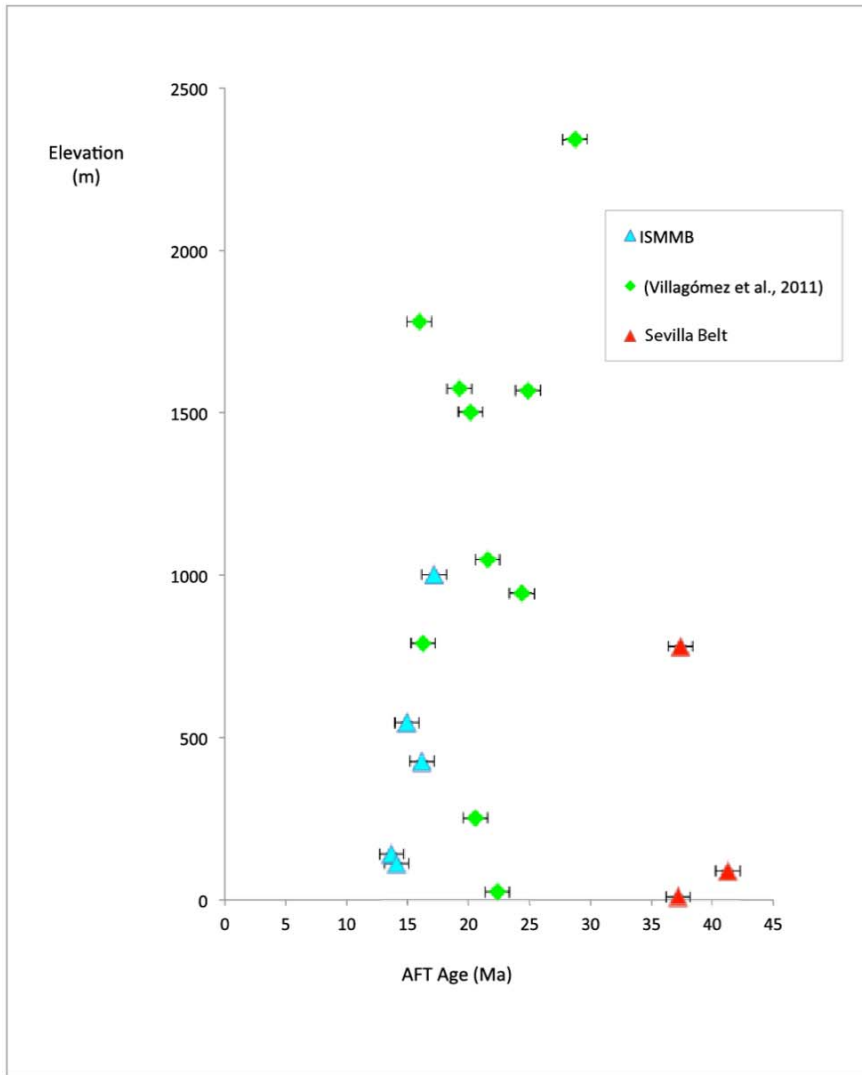


Figure 10. AFT and ZFT elevation profiles from samples of the IB and the Sevilla Metamorphic Belt.

either by tectonic accretion or by sediment accumulation on the shelf and basin slope, contributing to increasing the relief of the oceanic plate. This process worked until a significant decrease in slab angle led to the Oligocene activation of the Oca and Santa Marta - Bucaramanga faults. The SMBF is a transfer fault that accommodate strain differences between the Lower Magdalena and the SNSM massifs. The SMBF contributed to an enhanced exhumation on the western flank of the SNSM, as evidenced by younger AFT ages towards the fault trace. The youngest age is 13.7 ± 1.2 Ma. The regional pattern of AFT, and ZFT cooling ages suggests a regional tilting of the SNSM towards the NE as a consequence of increased subsidence at the SMBF hanging wall (Fig. 11).

At the same time the onset of sinistral displacement on the SMBF during the late Oligocene was coeval with a decrease in thrust advance towards the Caribbean. The increased amount of slip assumed by the reactivated Santa Marta-Bucaramanga and Oca faults was recorded in the adjacent clastic basins of Aracataca and Palomino in thick clastic sequences that evidence unroofing of the Precambrian basement since the late Oligocene (Piraquive et al., 2016, Chapter 3). From late Eocene-Oligocene times on exhumation rates decreased progressively to ~ 0.2 km/Myr to a steady-state equilibrium, which culminated with the transference of slip from the Orihueca fault to the Aguja fault, which acted as a splay on the advance of the deformation front. In the late stages of deformation, the obducting Caribbean plate impeded the advance of the thrust front.

5.5 Implications on the evolution of the Nor-Andean Block

Since the Paleocene the southeastern flank of the SNSM was exhumed, sourcing sediments to the newly formed Cesar-Rancheria basin. At the start of exhumation, the cover of the SNSM was the first material eroded. In a regional, basin-wide drainage system fed by the Cretaceous sedimentary cover exposed in low-amplitude topography and developed a mixed siliciclastic-carbonate platform in the shelf areas, with a major sediment contribution from the craton, as evidenced by a dominant population of Paleoproterozoic zircons (Ayala et al., 2012). During the late Paleocene, unroofing of the Permo-Triassic metamorphic basement from the Central Cordillera and Perija Range started to contribute to the clastic sediments.

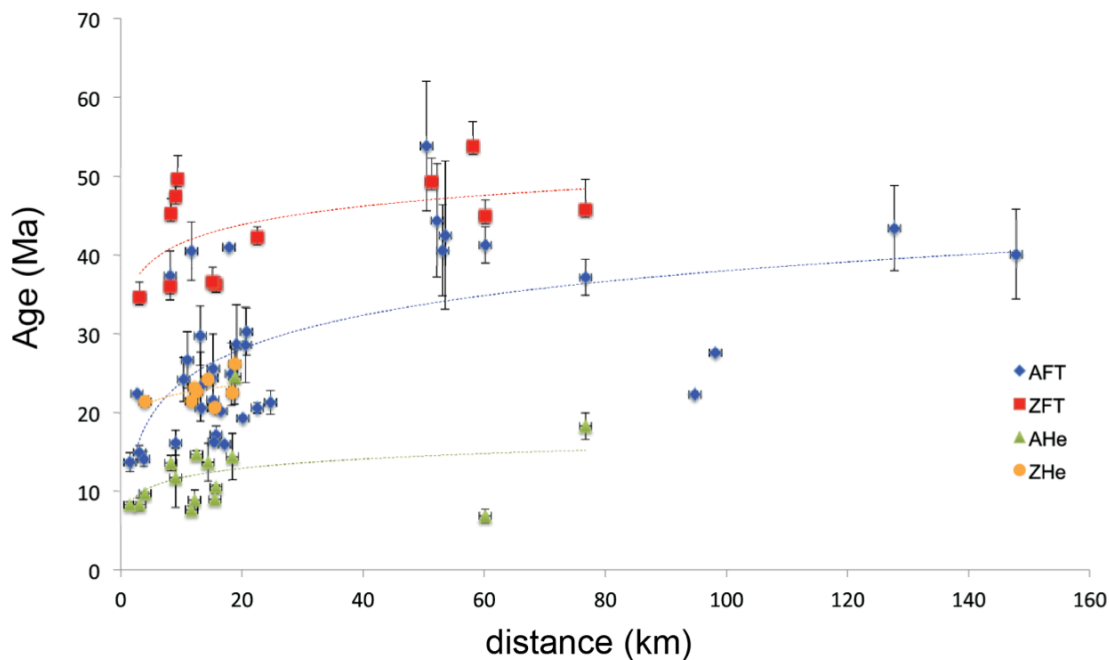


Figure 11. AFT, ZFT, AHe, ZHe thermochronometric ages obtained in this study and compiled with previous data (Cardona et al., 2011a; Villagómez et al., 2011), n= 64. Data are plotted relative to the linear distance from the Santa Marta Bucaramanga fault (SMBF), at the Sierra Nevada de Santa Marta massif (SNSM). All distances were measured as normal lines to the mapped SMBF, uncertainties are 2σ errors, and the dashed lines show the age trend towards the SMBF.

Evidence of Paleogene sedimentation is preserved as well on the western margin of the SNSM in the Lower Magdalena basin, recording arc-continent collision and basin opening. In this area, the main sediment sources areas were the upper Cretaceous volcanic arc and a continental margin composed upper Paleozoic to Triassic metamorphic rocks (Cardona et al., 2012). Minor contributions of Permo-Triassic zircons found in the Catatumbo, Cesar-Rancheria and NW Maracaibo basins since the Mid-Paleocene are related to sources in the Perijá Range and Santander Massif (Ayala et al., 2012). An important shift in provenance occurred during the middle Eocene, when the Precambrian basement of the SNSM, the Perija Range, the Merida Andes, and the Santander Massif exhumed and supplied Grenvillian zircons to the Cesar Rancheria, Catatumbo and Maracaibo basins. These mountain belts became the main sediment source areas, diminishing cratonic input because of the topographic configuration that individualized these basins (Bayona et al., 2007; Montes et al., 2010; van der Lelij et al., 2016; Villagómez et al., 2011b).

High middle Eocene exhumation rates (1.1 km/Myr) in the SNSM caused the removal of at least 9 km of upper crustal rocks that contributed to the filling of the adjacent sedimentary basins. The onset of exhumation was probably a consequence of Caribbean slab flattening and strong plate coupling that led to surface uplift of the Eastern Cordillera of Colombia (Caballero et al., 2010). The magnitude of regional exhumation during the late Eocene- late Oligocene is recorded in the Cenozoic strata of the Middle Magdalena basin and the Eastern Cordillera (Horton et al., 2010a, 2010b; Mora et al., 2010; Nie et al., 2010). The mechanism that triggered this accelerated exhumation during the middle Eocene is closely related to strain transfer between the

Caribbean plate and the Eastern Cordillera, through reactivation of the SMBF, which accommodated significant subsidence during the Oligocene favoring carbonate platform development at marginal highs (Arjona, El Cicuco, El Difícil) (Duque-Caro, 1979; Montes et al., 2010).

The reorganization of the Nazca, Caribbean and South American plates with a change in convergence direction from an E-W during the Paleocene to NW-SE during the Oligocene was responsible for the strain transfer (Cortés et al., 2005). The Paleocene-late Eocene shift in the regional stress field can be observed in fracture patterns, kinematic markers and structural styles in the Eastern Cordillera of Colombia, coeval with the middle Eocene unconformity. The mid-Eocene unconformity coincides with the episode of the highest rates of exhumation (0.8 - 1.2 km/Myr) in the SNSM.

The early indentation of the Panama block at ca. ~ 23-25 Ma against western South America (Farris et al., 2011; Montes et al., 2012) induced a pulse of surface uplift and exhumation. This led to doming and exhumation of the SMB in the IB. The last pulse of exhumation recorded in the IB, about 2 km to the NE of the SMBF, is related to the activity of the Aguja fault. (U-Th)/He ages of 8.3 ± 0.8 for the IB attest that this pulse of exhumation continued at least until the late Miocene with erosion rates of 0.14-0.18 km/Myr. The onset of exhumation of the Merida Andes at ca. 14 Ma was due to Boconó fault activation (Bermúdez et al., 2010, 2011; Javadi et al., 2011). This exhumation persisted at high rates of about 1.5 km/Myr from 14 -4 Ma and at 0.4 km/Myr from 4 Ma to the present for the Sierra Nevada block. Whereas the El Carmen block, to the north of the Bocono fault was exhumed at 1.5 km/Myr since about 2 Ma until the present (Kon et al. 1984; Bermudez et al. 2011). Simultaneously elevated erosion rates on the eastern flank of the Eastern Cordillera started during the late Pliocene (Mora et

al., 2008, 2010; Parra et al., 2009). Evidence of late Pliocene exhumation in the SNSM is absent, and its elevated topography may be the consequence of a very recent surface uplift triggered by mantle upwelling (Ceron-Abril, 2008; Villagómez et al., 2011b). We favor this hypothesis considering the orogenic float model (Monod et al., 2010), given that inversion of gravity and magnetic data constrain crustal thickness to only 25 km and doming under the SNSM (Camargo, 2014). Rapid mantle upwelling beneath the SNSM could be a consequence of late Pliocene convective removal of the SNSM underthrust Caribbean slab causing isostatic rebound, which may explain the high topography. Convective removal can occur much more rapidly than mountain building, over 1-10 Myr (Houseman and Molnar, 2001). The removal of the dense root probably caused 1-3 km of surface uplift (England and Houseman, 1989; Platt and England, 1994).

7. CONCLUSIONS

1) The history of exhumation of the SNSM has been complemented with new low-temperature thermochronological data, which has allowed to calculate cooling rates and determine exhumation rates for the Sevilla Metamorphic Belt (SB). Our data show that this province presents an independent cooling history from the thermal histories presented in the SNP and IB (Cardona et al., 2011a; Villagómez et al., 2011b).

2) Paleocene-Eocene plutonism was produced by collision between the South American and Caribbean plates. Although, previous studies tried to define a geochemical character for these plutons, the results were ambiguous (Cardona et al., 2011a; Duque, 2009; Tschanz et al., 1974). Field relationships as well as thermal

history models support a thickened continental crust origin. Intrusion depth of the SMB batholith has been constrained between 15-20 km. Emplacement occurred during subduction failure, which was a consequence of an abnormally thick oceanic crust of the Caribbean plate (Stern, 2002). During convergence and plate coupling part of the Caribbean plate was obducted onto the South American plate margin through the Florin fault. The onset of extrusion of crustal slivers is calculated after ca. 40 Ma by ZFT.

3) Widespread Paleocene-Eocene plutonism affected the Precambrian basement of the Sevilla Metamorphic Belt. In this tectonic block, exhumation occurred diachronously, with ZFT cooling ages 10 Myr older at the Buritaca pluton with respect to the Sevilla Stock. Exhumation of the SB occurred at rates that increased from 0.8 – to 2 km/Myr on the northern flank of the SNSM, between 50-42 Ma. On the western flank of the SNSM, exhumation took place between 49-35 Ma, with rates that increased from 0.4 to 2.7 km/Myr.

4) The SNSM exhibits an inverted metamorphic sequence, in which the SB occupies a position between the Sierra Nevada Province and the IB. The SB is separated from the adjacent blocks by two major shear zones: the Orihueca fault and the Sevilla lineament, with thrust and normal components respectively. This structural configuration and the elevated exhumation rates make the SB a possible example of crustal channel flow during the middle-Eocene. In this model the SB was extruded as a weakened viscous crustal channel between major faults, with focused exhumation in a narrow zone. In such channel partial melting occurs commonly in the core of the channel and in the marginal thrust zones. Therefore, we interpret that the Paleocene plutons of the Latal Stock (LS) and Sevilla Stock (SS) are a consequence of lower crustal

extrusion. This model satisfies the conditions in which an extruded narrow zone of lower crustal rocks bounded by faults, incorporating deformed granitoids of a more recent age, overlies younger rocks of a lower metamorphic grade.

5) Elevated exhumation rates during the middle Eocene are related to collision of the Caribbean plate against South America. This collision induced I-type plutonism because of crustal thickening of radiogenic pelites and subsequent exhumation of the lower crust under a high geothermal gradient $\approx 62^{\circ}\text{C}/\text{km}$. Subduction failure due to strong plate coupling led to reactivation of mid crustal structures of the continental margin and strain transfer from the Caribbean plate to the South American plate along the SMBF causing surface uplifting and exhumation of the SNSM, the Perijá Range and the Eastern Cordillera of Colombia.

6) The Santa Marta-Bucaramanga fault caused tilting of the SNSM towards the NE during the Miocene as evidenced by younger thermochronological ages close to the fault trace. This tilting may be related to subsidence of the SMBF hanging wall coupled with the tectonic drag produced by the Caribbean plate during its eastward motion.

7) An overall change in convergence during the Oligocene caused a decrease in exhumation rates. Slip was transferred from the Sevilla fault to the Santa Marta-Bucaramanga fault, and exhumation rates decreased to $0.09 \text{ km}/\text{My}$, at the same time plutonism ended and geothermal gradients decreased. With this scenario of a partitioned convergence strike slip faults started because of the eastward motion of the Caribbean plate, generating subsidence along the Santa Marta-Bucaramanga and Oca faults.

8) The last pulse of surface uplift triggered by the indentation of the Panama Block at ca. 23 Ma influenced renewed thrusting that persisted on the Aguja fault until about 8

Ma. In the already established transpressive regime the SNSM may have reached its maximum thickness. The current lack of isostatic compensation and absence of younger cooling ages at the surface may be related to a very recent episode of convective removal of the crustal root that resulted in accelerated exhumation of the massif during the late Pliocene.

8. REFERENCES

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CHAPTER 3

Early Neogene unroofing of the Sierra Nevada de Santa Marta determined from detrital geothermochronology and the petrology of clastic basin sediments

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Abstract

Geo-thermochronologic data from sedimentary rocks, coupled with stratigraphic analyses, unravel an early Neogene exhumation of the Sierra Nevada de Santa Marta massif (SNSM). This fault-bounded triangular block in the northern Andes of Colombia exposes Precambrian to Paleozoic metamorphic rocks and suites of a Triassic-Jurassic magmatic arc. We compare the Neogene basin fills of two marginal basins, a western one located along the Santa Marta-Bucaramanga fault (SMBF) and the northern one occupying the southern margin of the Oca fault. The western sequence consists of gravitational conglomeratic deposits ≈ 1200 m that define a progradational Gilbert-type delta emanating from a scarp coinciding with the SMBF. By its stratigraphic and structural position this sequence compares to the northern transgressive sequence. Provenance analyses of representative sections of these two sequences, evidence a sourcing from the underlying basement. This finding is in concordance with zircon U-Pb age spectra that lack pre-Grenvillian and Ordovician signals of coeval sediments of the Magdalena basin, demonstrating, that they were deposited in a local basins disconnected to the trunk system of the proto-Magdalena river. In order to reconcile the different organization of the two marginal basins we propose a regional-scale tilting of the SNSM towards the northeast during the early Neogene. This model takes account of the coeval progradational sedimentary cycle at the fault-bounded western flank and the transgressive onlapping of the northern conglomerates on a tilted surface during an overall exhumation of the massif.

Key words: Santa Marta Massif, Marginal basins, Neogene tectonics, Gilbert-type delta, Santa Marta-Bucaramanga fault

1. INTRODUCTION

The northwest corner of South America is divided into the Northern Andean block which, on its turn, is subdivided into a number of micro-blocks as a result of the continued interaction between the South American, Caribbean and Nazca plates that record various Cretaceous to Cenozoic contractional events. Although, fault-bounded morphotectonic provinces (Fig. 1) readily define the assemblage of the individual micro-blocks, controversy exists about tracing kinematic limits between the North Andean block (or continental South America) and adjacent oceanic plates. Velocity gradients deduced from Regional Global Positioning of the Central and South America Project (CASA) (Trenkamp et al., 2002) argue for diffuse limits that emerged as a result of a transpressive stress regime between the Caribbean plate and South America (Fig. 1). Inherited sutures and present plate boundaries are partly concealed by the mountain ranges of the North Andean block, that comprise the Eastern, Central and Western Cordilleras, the Sinú folded belt, the Sierra Nevada de Santa Marta (SNSM) and the Guajira massifs in Colombia, and the Perijá Range and Merida Andes in Venezuela and the Neogene basin fill of the LMV (Fig.1).

During the Maastrichtian-Paleocene a juvenile Caribbean plateau with associated marginal island arcs collided against the NW corner of South America, (Ayala et al., 2012; Bayona et al., 2011; Cardona et al., 2011a, 2012; Higgs, 2009; Pindell et al., 2005; Pindell and Kennan, 2009; Spikings et al., 2015; Villagómez et al., 2011), juxtaposing oceanic crustal slivers against the continental margin. This accretion was followed by a post-collisional Eocene plutonic event, marking the closure of a Cretaceous-Paleogene subduction cycle. During the Oligocene-Miocene the onset of a transpressive regime

led to the opening of the Plato-San Jorge basin to the west and the separation of the Sierra Nevada de Santa Marta massif from the Central Cordillera, as the continental margin became fragmented along the Santa Marta - Bucaramanga fault (Duque-Caro Mora-Bohórquez et al., 2017, Montes et al., 2010) and the drainage of the Magdalena River originally communicating with into the Maracaibo basin became diverted to its present site

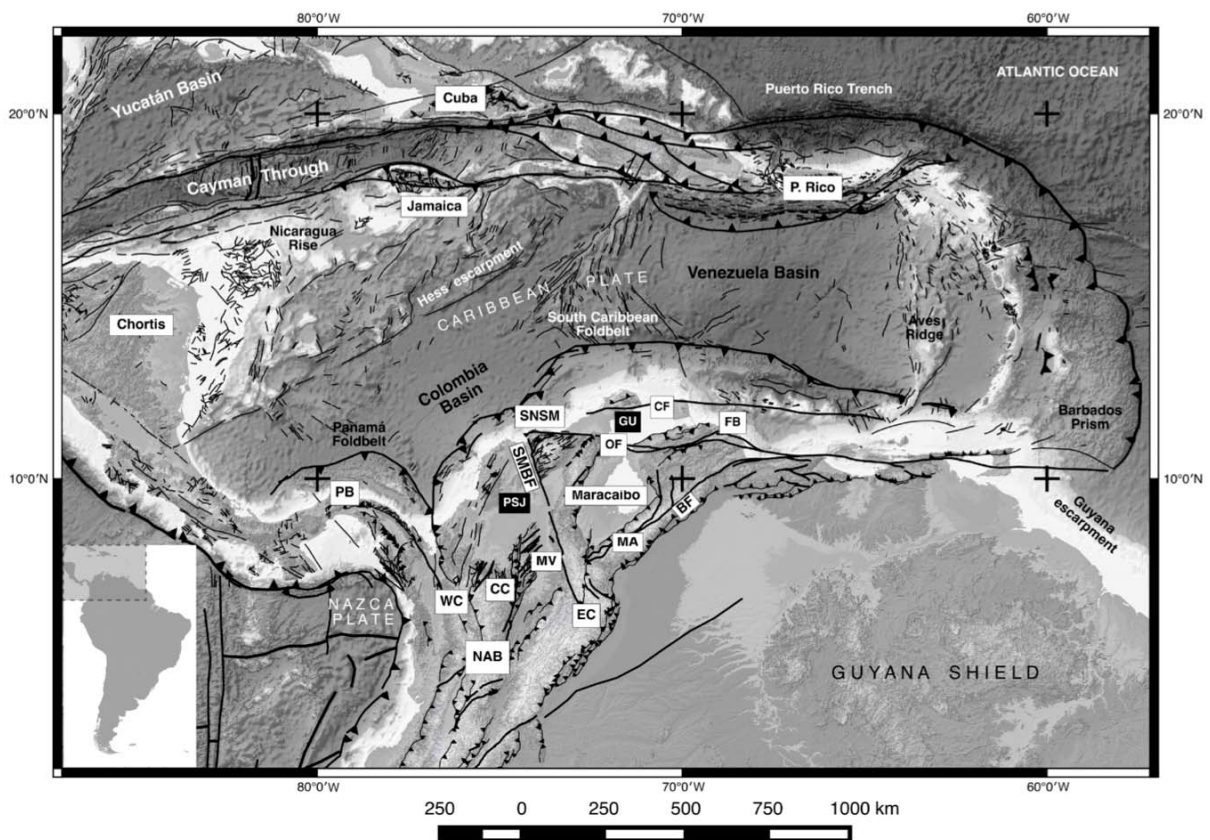


Figure 1. Regional tectonic map of the Caribbean realm showing the most relevant geological features intervening in the plate boundary configuration. after (Veloza et al. 2012). Elements of the North Andean Block (NAB): Maracaibo block, SNSM: Sierra Nevada de Santa Marta, OF: Oca Fault, CF: Cuisa Fault, SMBF: Santa Marta-Bucaramanga Fault, GU: Guajira Basin, PSJ: Plato-San Jorge Basin MA: Mérida Andes, BF: Boconó Fault, FB: Falcon Basin, WC: Western Cordillera, CC: Central Cordillera, EC: Eastern Cordillera, MV: Magdalena Valley, PB: Panamá Block.

Recent bedrock thermochronological studies demonstrated an early Eocene to early Miocene exhumation of the SNSM, with youngest cooling ages (25-16 Ma) located to the NW and older ages (50-40 Ma) located to the SE of the SNSM (Villagómez et al., 2011). This asymmetric age distribution may be attributed to an enhanced exhumation at the Santa Marta – Bucaramanga fault in the course of a separation of the SNSM from its co-terminal western Central Cordillera, as the collisional impact of the Caribbean plate induced a clockwise rotation of the continental margin (Montes et al., 2010). A coeval accumulation of a Cenozoic marine sedimentary pile of ≈ 8 km in the Plato San Jorge basin supports this scenario (Duque-Caro, 1979). This predominantly extensional regime gave rise to fault bounded marginal basins and persisted up to the middle Miocene, when the stress regime became transpressive and the Oca fault assumed its strike-slip displacement (Audemard and Audemard, 2002). The late stages of the tectonic evolution of the SNSM are recorded in the basins surrounding the massif.

This study focuses on the Oligocene-Miocene sedimentary fill of the fault-bounded sub-basins located in the footwall block of the Santa Marta-Bucaramanga fault associated with the Plato-San Jorge basin and the hanging wall block of the Oca fault at the Guajira basin. We defined these marginal fault related basins as the Aracataca basin at the western foothills of the SNSM, and the Palomino basin along the northern foothills of the SNSM respectively.

We analysed the late exhumation history of the SNSM using the most proximal Neogene sediments found in the Aracataca and Palomino basins. The studied Cenozoic clastic sedimentary successions in both basins consist mainly of conglomeratic fan delta deposits that are interbedded with marine to transitional sandstones at their

base and at their top. Compositional changes and age data record relevant stages of a Paleogene exhumation of the SNSM.

Our methodological approach is based on the determination of sediment provenance, relating compositional data with bedrock units of a mainly Jurassic arc assemblage of the SNSM. This compositional analysis is performed by heavy mineral counts in sandstones and pebble variations in the conglomeratic sequence. These results are combined with U-Pb ages measured in detrital zircon, and fission-track cooling ages from apatites and zircons.

The integration of stratigraphic and geochronological data helps constraining the timing of deformation phases that led to the exhumation of basement rocks that form the footwall of the Santa Marta-Bucaramanga fault and NW-verging thrust slices of the interior of the massif.

2. TECTONIC SETTING

The Caribbean continental margin of South America was shaped/constituted itself during a Cretaceous to Paleogene subduction cycle and a subsequent collisional event (Audemard and Audemard, 2002; Cardona et al., 2011b, 2012; van der Lelij et al., 2010; Villagómez et al., 2011). During Cretaceous subduction of the Farallones plate, the convergence was not partitioned, as strike-slip displacements are absent until the Paleogene (Acosta et al., 2007; Cortés et al., 2005; Kennan and Pindell, 2009; Vence, 2008). The Paleocene-Oligocene collisional phase, however, entailed a strike-slip reactivation of margin-parallel faults, as the North-Andean block rotated into its present position (Nova et al., 2013; Vence, 2008), culminating in the late Miocene to

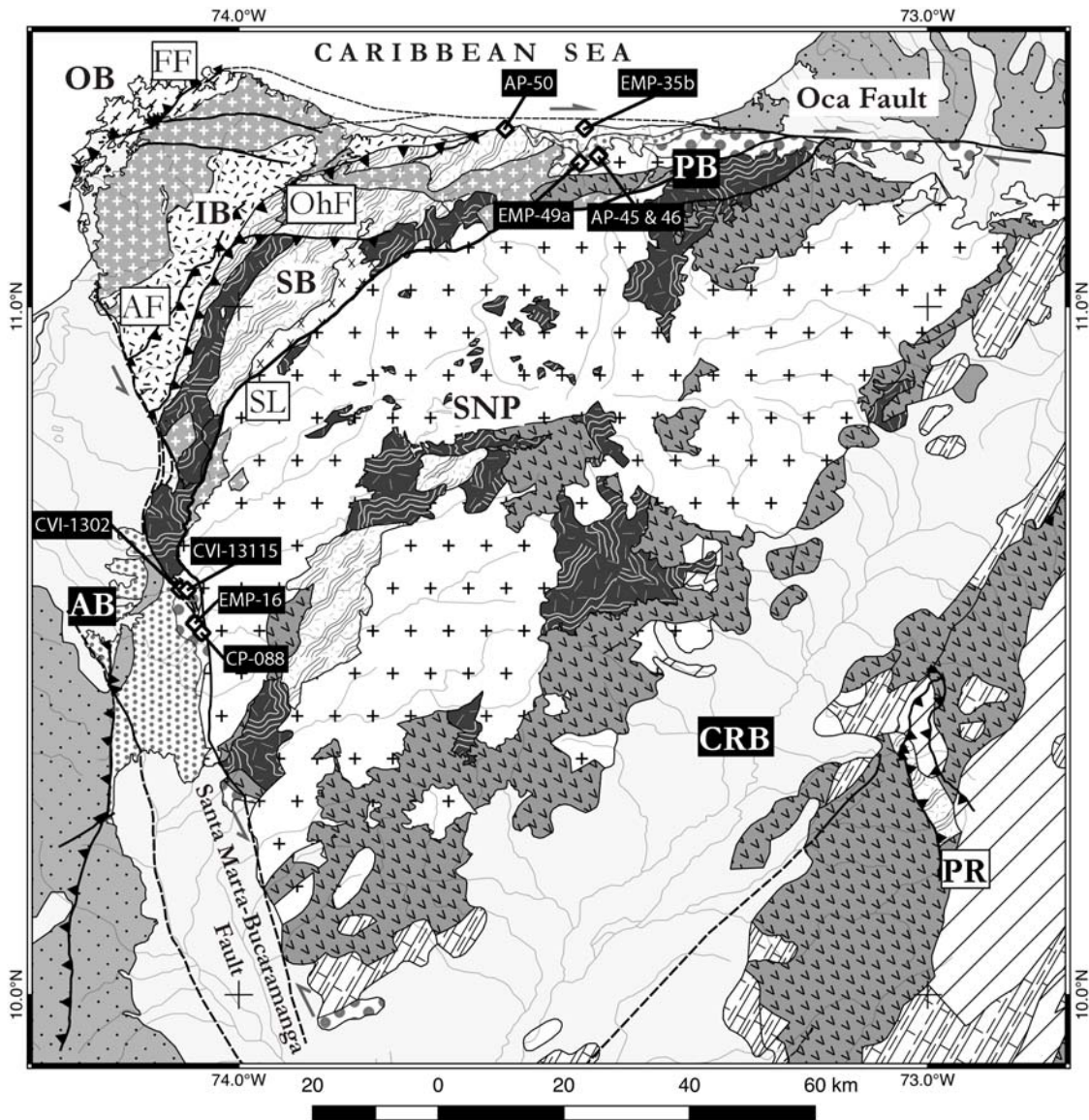
Pliocene with the activation of the Boconó fault of the Merida Andes (Bermúdez et al., 2010; Egbue and Kellogg, 2010; Javadi et al., 2011). In this tectonic environment, upper crustal heterogeneities caused the formation of minor tectonic blocks during Paleogene-Neogene transpression. Contrary to the presence of directional faults that form sharp boundaries (Motagua fault, Cayman trough, Puerto Rico trench) on the northern edge of the Caribbean plate (Fig. 1), deformation on the southern edge of the Caribbean plate is widely distributed along crustal heterogeneities of the different mountain chains (Pennington, 1981; Taboada et al., 2000).

The northern edge of the South American plate in the Colombian Andes is segmented by the Santa Marta-Bucaramanga fault, between the N-S striking Andean Cordilleras and the Maracaibo block (Fig. 1). The Santa Marta-Bucaramanga fault originated during the early Mesozoic continental break-up (Kammer and Sánchez, 2006), and subsequently decoupled the structural evolution of the Central Cordillera and the Maracaibo block (Figs. 1, 2, 3.). The Western Andean domain still preserves the tectonic frame of an arrested Pre-Campanian subduction setting with the Romeral suture separating continental and oceanic terrains. However, in the Eastern block of the Santa Marta-Bucaramanga fault, the triangular SNSM massif and the isolated Perijá Range formed as response to the early subduction of the Caribbean plate during the Paleogene (Ayala et al., 2012; Bayona et al., 2011; Cardona et al., 2011a; Monod et al., 2010).

2.1 Sierra Nevada de Santa Marta massif

The SNSM constitutes a remarkably elevated massif along the South-Caribbean margin with the Bolivar peak reaching an altitude of 5775 m. This triangular block displays the structural array of a monocline, with its northwestern corner exposing an imbricated lower crustal section, capped by nested Jurassic plutons and correlative volcanic sequences forming its southwestern flank (Tschanz et al., 1974). On its western side it is limited by the normal-sinistral Santa Marta-Bucaramanga fault, which produced a structural relief of >7000 m between the basement of the SNSM and the basement of the Lower Magdalena basin (LMB), and accumulated as much as 45 km of left lateral displacement since the Oligocene (Montes et al., 2010) (Fig. 2 & 3). At its northwestern corner, an obducted contact of the Caribbean plate can be inferred (Doolan, 1970). On its northern side the contact of the Caribbean and South American plates is cut by the right-lateral Oca fault (Fig. 2).

The juxtaposition of the South American and Caribbean plates sustains the idea of a delaminated crustal section “floating” on the Caribbean plate (Audemard and Audemard, 2002) this exceptionally thick crustal section forms the upper crust of a continent-ward dipping main suture that is underlain by strongly sheared platform sediments and transitional basement rocks of the lower crust. The lower crustal section incorporates low to high-grade metamorphic units, the Sierra Nevada Province and the Sevilla Metamorphic Belt composed of Precambrian granulites, anorthosites and gneisses (Cardona et al., 2010a; Cordani et al., 2005; Kroonenberg, 1982; Ordóñez et al., 2002; Restrepo-Pace et al., 1997), the Inner Santa Marta Metamorphic Belt that is composed of a Permian mylonitic belt (Cardona et al., 2010c) of orthogneisses and



Legend

Quaternary alluvium	Outer Santa Marta Metamorphic Belt
Pliocene Marine and continental strata	Granitoids (Jurassic-Cretaceous)
Miocene marine strata	Volcanic arc deposits (Triassic-Jurassic)
Miocene fluvio deltaic deposits	Inner Santa Marta Metamorphic Belt (Permian-Jurassic)
Lower Miocene Fan Delta	Metadiorite (Permian)
Palaeogene marine and continental strata	Paleozoic strata
Granodiorites (Palaeogene)	Buritaca-Muchachitos gneisses (Neoproterozoic- Cambrian)
Upper Cretaceous marine strata	Los Mangos Granulite (Precambrian)

Figure 2. Geological Map of the Sierra Nevada de Santa Marta Massif (SNSM) and sampling sites. SNP: Sierra Nevada Province, SB: Sevilla Belt, IB: Inner Santa Marta Metamorphic Belt, OB: Outer Santa Marta Metamorphic Belt, AB: Aracataca Basin, PB: Palomino Basin, CRB: Cesar Rancheria Basin, AF: Aguja Fault, OhF: Orihueca Fault, FF: Florin Fault, SL: Sevilla Lineament, PR: Perijá Range.

meta-sedimentary rocks, a Triassic to Lower Cretaceous sequence of high amphibolite facies meta-sedimentary rocks and S-type granites (Zuluaga and Stowell, 2012), and the Outer Santa Marta Metamorphic Belt, a phyllite belt that is related to oblique accretion of oceanic crust composed of low-grade metamorphic schists, meta-conglomerates, and Upper Cretaceous-Paleogene meta-basites (Cardona et al., 2010b; Doolan, 1970; MacDonald et al., 1971). These units were subject to surface uplift and erosion as a consequence of collisional tectonics along the continental margin since the Paleogene.

3. STRATIGRAPHY OF THE ARACATACA AND PALOMINO MARGINAL BASINS

3.1 Aracataca Marginal basin (Western SNSM)

The Cenozoic Aracataca basin, forms a narrow depositional fill along the western flank of the SNSM (Fig. 2), and is reported to record marine to lagoonal and continental depositional conditions (Tschanz et al., 1969). Considering this setting, the Santa Marta-Bucaramanga is likely to have controlled its subsidence, as its elevated eastern block contains the Precambrian Mangos granulite and the Jurassic acid to intermediate plutonic and volcanic arc assemblage, which essentially sourced the basement fill (see below). A compilation of the stratigraphy is given in (Table. 1).

We divide the stratigraphic profile of the Aracataca basin, into a lower shaley to conglomeratic unit with a marine fossils, informally designated as Macaraquilla

conglomerate (Hernandez et al., 2003) and an upper unit of sandstones and conglomerates formerly attributed to a fluvial to transitional molasses facies of a Miocene age (Tschanz et al., 1969). This eminently conglomeratic sequence is succeeded by interbedded sandstones, marls and mudstones of the Zambrano Formation which yielded pollen of a lower Pliocene age (Hernandez et al., 2003). Upper Pliocene deposits are defined as the informal Guamachito conglomerate unit, which towards the west of Aracataca decreases in grain size and give rise to the “Unidad Arenosa de Fundación” (Table. 1).

3.2 Palomino Marginal Basin (Northern SNSM)

In the southern footwall of the Oca fault a fining-upward sequence rests unconformably on Pre-Cretaceous basement rocks. In the Palomino basin along the road from Santa Marta to Riohacha, the outcrops preserve a Miocene conglomeratic succession defined as molasse facies (Tschanz et al., 1969). The conglomeratic facies rests directly over the crystalline basement with no remnants of pebbly mudstones and associated marine deposits. The sedimentary succession of pebble conglomerates ends against the Oca fault, fine grained Upper Miocene – Pliocene sequences are preserved north of the Oca fault. (Tschanz et al., 1969).

TABLE 1. STRATIGRAPHIC NOMENCLATURE FOR THE CENOZOIC SEDIMENTARY ROCKS AT THE ARACATACA BASIN

	AGE	TSCHANZ et al. (1969)	HERNANDEZ, (2003)	COLMENARES, (2007)	THIS WORK
Q.	PLEISTOCENE	Upper Tertiary sedimentary rocks	Unidad arenosa de Fundación		Unidad arenosa de Fundación
	PLIOCENE		Guamachito Conglomerates	Guamachito Conglomerates	Guamachito Conglomerates
			Zambrano Formation	Zambrano Formation	Zambrano Formation
NEOGENE	MIOCENE	Miocene sedimentary rocks		Miocene sedimentary rocks	Aracataca Conglomerates
	OLIGOCENE				
PALEOGENE	EOCENE	Eocene Sedimentary rocks			
	PALEOCENE		Macaraquilla Conglomerates	Macaraquilla Conglomerates	

An appraisal of the two sedimentary basins clearly needs the precision of a biostratigraphic frame, a task still to be performed. Our argument for an Oligocene to early Miocene stratigraphic age of the Aracataca conglomerates is based on the regional framework of the Lower Magdalena basin and the Paleogene evolution of its sub-basins, as documented by their subsidence histories (Bernal-Olaya et al., 2015; Flinch and Castillo, 2015). These data are fit the frame of exhumation ages (Villagómez et al., 2011) of the Santa Marta Massif reasonably well.

A well-constrained subsidence history has been elaborated recently for the Plato sub-basin (Bernal-Olaya, 2015), which constitutes one of the well-evolved depocenters to the W of the study area. Burial plots document two phases of subsidence. A first late Paleogene/early Neogene depositional phase was initiated along faults oriented perpendicularly to the continental margin (Bernal-Olaya, 2015; Montes et al., 2010).

Subsidence plots constrain this rift re-related basin evolution to the middle Oligocene until the late Miocene age. Among the graben-bounding faults, the Santa Marta-Bucaramanga fault and subsidiary strands form a western basin margin and control the NNW-striking depocenter of the Arjona sub-basin (Bernal-Olaya et al. 2015), which connects further north to Aracataca sub-basin.

A second subsidence phase spans the middle Miocene to Pliocene and concludes with the deposition of deep-water shales of the Tubará Formation (Flinch, 2003). Subsidence was accommodated without rifting.

A late Paleogene exhumation phase of the western border of the SNSM massif has been documented by fission track-data by Villagómez et al. (2011), who evidenced a principal exhumation interval of 29 to 26 Ma estimating exhumation rates of 0.6 to 0.7 km/Ma. Considering the proximal depositional setting for the Aracataca conglomerates with respect to a mountain front located at the SMBF, the time lag between denudation, transport and deposition could be below our very approximate time frame and amount to just a few million years, as alluded for a gravelly-dominated depositional settings of a similar tectonic environment (Jones et al., 2014). Such a largely synchronous interplay between exhumation and deposition may be presumed for the late Oligocene exhumation of the western flank of the SNSM and the largely coeval rifting of the Lower Magdalena Valley (LMV), which most likely triggered the accumulation of the Aracataca conglomerates.

4. METHODS

4.1 Clast Counting

Along the two stratigraphic sections conglomerate clast counts were performed at 13 locations shown in Fig. 3 & 4. Clast counting was intended at least to have one representative count for each conglomeratic facies associations. An average of 100 clasts (range between 60 up to 200 clasts) with sizes varying from very coarse pebbles (>3 cm) to boulders were counted using the ribbon method (Howard, 1993). Conglomerates are commonly matrix-supported poorly sorted and highly immature. Clasts were classified into plutonic, volcanic, metamorphic, sedimentary and minor abundances of gabbroic and ultramafic related clasts. Within the discrimination twelve specific litho-types were defined according to macroscopic compositions, results are summarized in Fig. 6.

4.2 LA-ICP-MS U-Pb Geochronology

Zircon U-Pb ages were obtained from 6 samples, using LA-ICP-MS analyses. Zircons were extracted from samples using conventional mineral separation technique including rock crushing, sieving (250-60 μm fraction), and concentration with heavy liquid and magnetic separation techniques. Detrital zircons (50 to 200 grains per sample) were randomly picked, selecting clean, crack and inclusion-free grains. Subsequently polished grain epoxy mounts were coated with graphite and imaged at

the Geosciences Department of the Université de Lausanne using a Vega ©Tescan MV2300VP SEM, to create a detailed set of panchromatic cathodoluminescence images. Laser ablation spots were selected, in both grain cores and rims.

LA-ICP-MS analyses and data processing were carried out at the Institute of Geochemistry and Petrology at ETH Zürich, Switzerland. Zircons were ablated with a NewWave UP-193nm Ar-F excimer ablation system coupled to a PerkinElmer PE SCIEX Elan 6100 ICP-MS to measure Pb/U and Pb isotopic ratios. The following parameters were applied during this process: 30 µm diameter beam size, 10 Hz repetition rate, 30–45 second signal and a beam intensity of 2.2–2.5 J/cm². Reproducibility of U/Pb data was monitored by measurements of GEMOC GJ-1 CA-ID-TIMS ²⁰⁶Pb–²³⁸U age of 608.5 ± 0.4 Ma; (Jackson et al., 2004) used as a primary standard. The external reference standards Plešovice 337.13 ± 0.37 Ma (Sláma et al., 2008), and TEMORA 1 416.75±0.24 Ma, (Black et al., 2003), were used to calibrate and monitor fractionation and consistency in the measured U–Pb dates. Ages were calculated using LAMTRACE (Jackson, 2008). Additional data reduction details are given by Ulianov et al, (2012). Statistical analyses of zircon data were performed using Isoplot 3.71 (Ludwig, 2003). All discordant (> 1–3%) analyses of magmatic zircons were discarded. Only zircons with concordance greater than 90% were accepted and plotted. Statistical interpretation of the results, regarding to discordance, maximum depositional ages and selection of the best age were done considering a threshold of 1.5 Ga given the change in chronometric power, the details on these proceedings are explained in Spencer et al. (2016).

4.3 Detrital fission-track thermochronology

Zircons of the 125-250 μm fraction from eight samples, mounted in Teflon® sheets, with at least two mounts per sample. Zircons were polished to expose an internal surface and etched in a NaOH-KOH solution at 228°C. Etching time varied between 20-60 hours in order to reveal countable tracks in at least 100 grains per sample. Samples were irradiated together with Fish Canyon Tuff and Buluk Tuff zircon age standards and IRMM541 dosimeter glasses.

Apatite crystals of 4 samples were mounted in epoxy, polished and etched in 5.5 M HNO_3 at 21°C for 20 s and covered with muscovite detectors. The samples were irradiated together with Fish Canyon Tuff and Durango Tuff age standards and IRMM540R dosimeter glasses. Both apatite and zircon samples were irradiated at the FRM II research reactor at Munich, Germany. A summary of the dated samples and analytical details as zeta calibration, grain number and U content are presented in Table A2, Probability density plots for selected examples are shown on figures 9 & 10.

5. RESULTS

5.1 Stratigraphy

In this section, we present a summary of the lithofacies of two representative sections of the Aracataca and Palomino basins, which we measured along the Rio Aracataca and the Rio Negro respectively (Fig. 3 & 4).

The analysed deltaic deposits are mainly conglomeratic, commonly with presence of calcareous cement, gravels that can be monomict but in general, a polymict composition is dominant especially towards the top. The detailed analysis and classification of the sedimentary rocks (Boggs, 2009; Folk, 1980; Pettijohn, 1980) led to the definition of 19 different facies, associated in four different assemblages. For the facies definition, we rely on the fluvial system classification defined by (Miall, 1978, 2006).

This stratigraphic scheme allowed defining the presence of four facies assemblages in the Aracataca basin and three facies assemblages in the Palomino basin. Depositional processes attest for major differences between two delta systems that are coeval but respond to different genetic process.

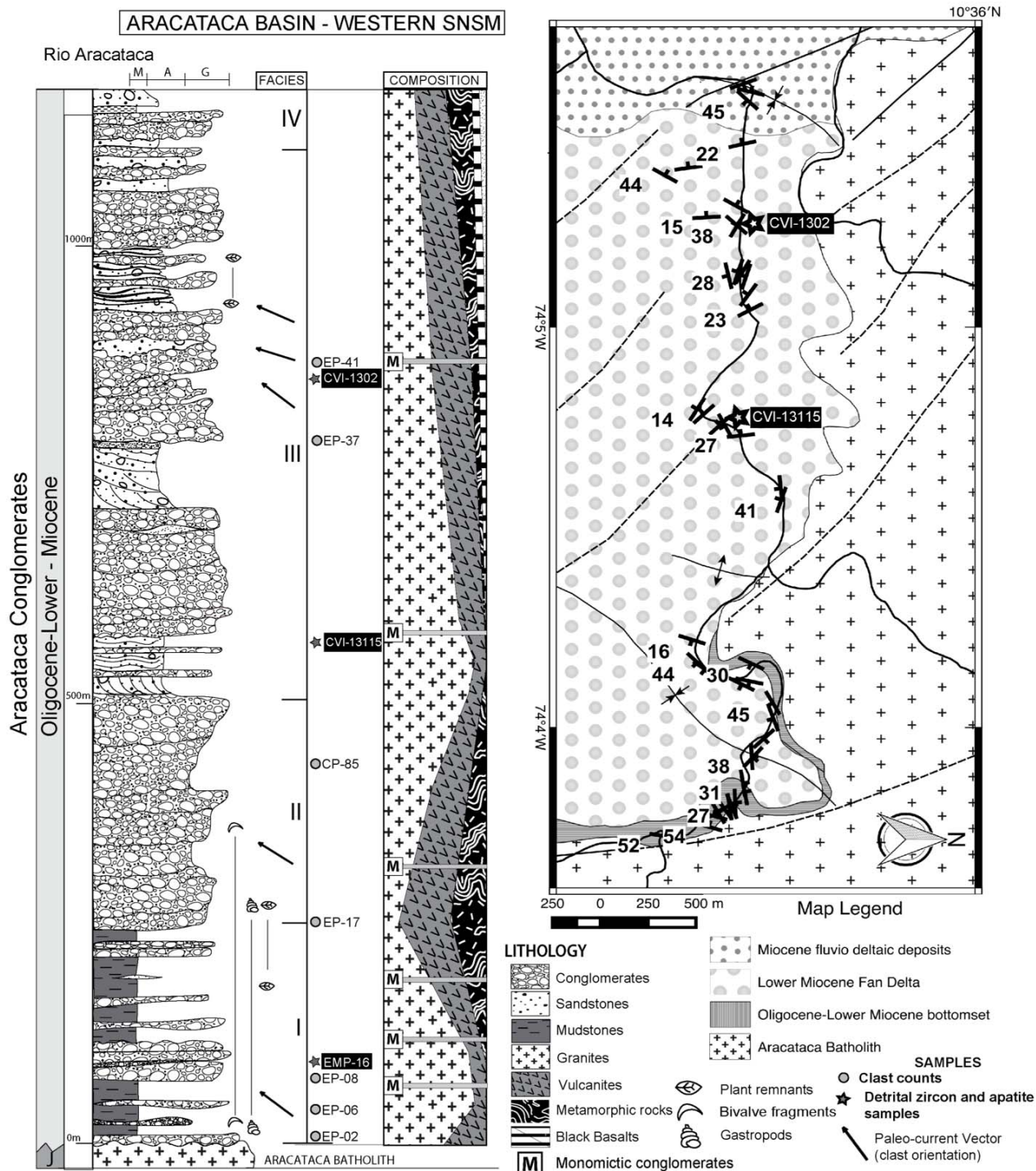


Figure 3. Stratigraphic column of the sedimentary successions at the Rio Aracataca with compositional variations of clast counts. Right: location map of the Rio Aracataca section. Facies associations are described in detail in Table 2.

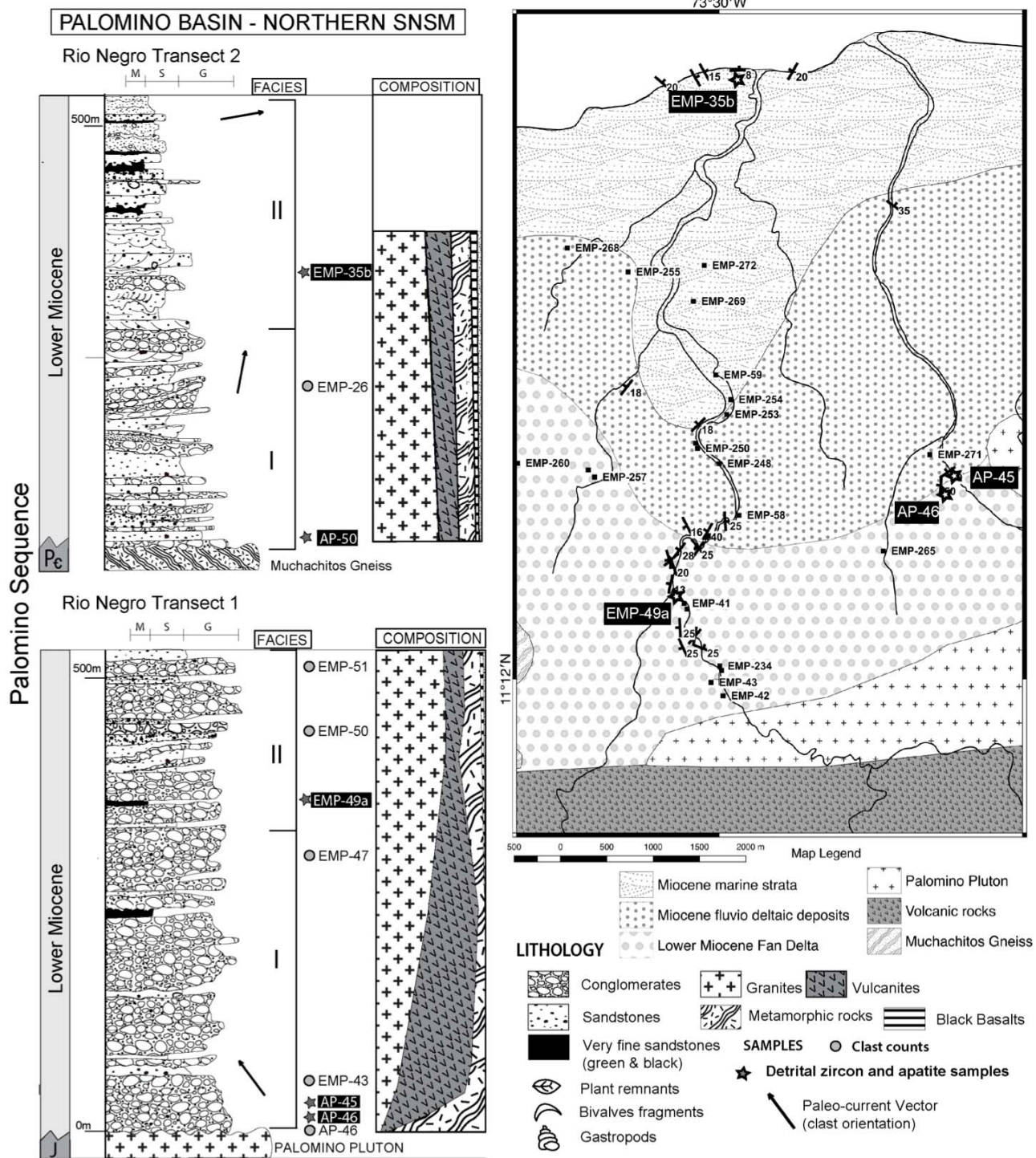


Figure 4. Stratigraphic column of the sedimentary succession at the Rio Negro and at the Palomino basin with compositional variations from clast counts. Right: location map of the Rio Negro section. Facies associations are described in detail in Table. 3.

5.1.1 Aracataca Conglomerates

Along the Aracataca River we measured a composite section of 1200 m of a mainly conglomeratic succession, which include fossil-bearing mudstones at its base. Based on lithological criteria and facies analysis we divide this profile into four sequences, numbered from base to top and labeled by interpretative names for ease of description (Fig. 3, Table 2). Sequence I consists of mud- to siltstones, alternating with discrete conglomeratic lenses. Sequences II and III are characterized by a predominance of tabular conglomeratic beds, which form amalgamated in sequence II and separated by sandy layers in sequence III. Sequence IV displays conglomeratic beds in discontinuous lenses and channel-like bedforms, which interfinger with sandy units and are separated by muddy beds. We base our facies analyses on the classification Miall (1978) proposed for fluvial deposits.

Our provenance study depicts trends in compositional variation of the gravel fraction. These variations may be divided into a two-fold scheme, according to the contribution of the local substratum, the Aracataca Batholith that belongs to the Central Batholith suite, and a wider bedrock spectrum of the southwestern flank of the SNSM. Thus, monomict conglomerates exclusively consist of granitic components of the Aracataca batholith, while polymict conglomerates include acid and basic volcanic clasts, supplemented by granulitic rock fragments of the Los Mangos unit.

Sequence I starts with a basal polymict pebble-boulder conglomerate, 20 m thick, composed of stacked beds, each distinguished by a characteristic clast size (facies *Ia*). Up-section, muddy to silty beds with calcareous intercalations denotes the extent of this unit. They alternate with lenses and mound-like accumulations of debris flows

with boulders of granite and micritic limestone (Figure 5-a) and enclose solitary boulders and monomict lenses of angular granitic clasts (facies *A-IIb*), which are virtually devoid of matrix, but locally intermingled with shell fragments. Overlying laminated silt- and mudstones smooth out rugged surfaces of these breccia-like conglomerates and mark a steady background sedimentation.

The base of sequence II is traced by the appearance of a first thick tabular conglomerate bed that overlies a basal mud layer in sharp contact (facies *A-IIa*). Large cobble-sized clasts show a bed-parallel orientation and a tendency of a reverse grading in the lower part of this bed (Figure 5-b), displaying thereby characteristics of a non-cohesive debris flow (Nemec & Steel, 1984) or a concentrated particle flow (Mulder & Alexander, 2001). Other similar beds contain solitary granitic boulders embedded along their base or floating in cobble conglomerates (facies *A-IIb*). Where defining closely packed lenses, they display an *a(p)a(i)* imbrication with their long axes dipping in a southeastern upflow direction. Other interfaces of debrites contain pockets of polymict cobbles, which are angular and highly disorganized. Occasionally, these conglomeratic layers are superseded by laminated medium-grained and coarse, massive sandstone (facies *A-IIIa*). Throughout this sequence, conglomeratic beds are intermittently separated by thin beds of silt- and mudstones, which contain lenses of monomict breccias, as described for unit I. Their presence defines the extent of sequence II.

Amalgamated debrites prevail in the lower part of sequence III and may be separated into distinct m-scaled flow units, according to particular clast size and sandy matrix content (Figure 5-c). Further up-section, flow units are capped by thin to medium-sized sandstone beds. These sandy layers are laminar to massive (facies *A-IIIa*) and

display water-escape structures, as evidenced by sheared sandy flames intruded into overlying conglomerates, patchy granular pockets within the sandstone and conglomeratic pillars or dikes emanating from underlying debris flows (nomenclature adopted from Postma, 1983). These debrites include in their upper part outsized granite boulders, which are aligned close to or may drape the very interface to the overlying sandy layer, displaying imbricate arrays, where closely packed (figure 5-d). Outsized clasts may also be present in the overlying sandy layers, suggesting an emplacement mechanism related to their common interface.

Within sequence III, matrix-supported polymict conglomerates compose a distinct facies, as they define channel fills with an abundant sand- to granule-sized matrix (facies *A-IIIb*).

Sequence IV is composed of discontinuous beds of pebble-cobble conglomerate, both clast- and matrix-supported, which form channel-shaped lenses up to 3m thick and display no distinct or a normal grading (facies *A-IVa*). Clasts are rounded and display an *a(t)b(i)* imbrication. They alternate and interfinger with trough-cross stratified sand beds that display conglomeratic base lags (facies *A-IVb*). This assemblage is subdivided by thin to medium beds of massive mudstone, which are variably scoured by channel-shaped bedforms (facies *A-IVc*).

Interpretation

Our facies inventory combines diagnostic characteristics of the three principal architectural elements of a Gilbert-type delta that defines, from base to top, a progradational succession. Sequence 1 constitutes a bottom- or toset, sequences 2

and 3 a lower and upper foreset assemblage of a delta slope and sequence 4 a topset of a delta plane. We base our arguments mainly on the facies distribution, as detailed in table 2.

Mud- and silt-stones of facies with embedded conglomerate lenses and outsized boulders (facies *A-Ia* and *A-IIIb*) define a bottomset within the reach of debris slides and falls of a nearby delta slope. Our measured section (Fig. 3) illustrates a decrease of muddy lithologies towards, which supports a progradational change from distal bottom- to proximal toeset. By their composition and textural maturity, debris flow units comprise two distinct classes. Monomict breccias are derived from the local bedrock of the Aracataca granite and are texturally highly immature, as they encompass granule- to boulder-sized rock fragments within a same bed. This lack of sorting and their disordered fabric (akin to an open-work fabric; Gobo et al., 2014) suggest an origin by debris fall from cliffs of a contiguous topographic scarp. Their intermingling with little fragmented marine fossil shells and boulders of calcareous beds (Fig. 5-a) points to the existence of a subaqueous platform or delta plain which they must have crossed during their avalanching. Considering an elevated sediment supply for the progradational setting of a fan delta, the formation of calcareous beds is atypical and may be restricted to minor transgressional cycles, as documented for a Paleogene clastic fan-delta of the northeastern Ebro basin (López-Blanco et al., 2000). Isolated outsized clasts embedded within mudstones attest to the presence of a pronounced delta slope, which contributed to their separation from smaller-sized clasts, according to a high momentum imparted by their fall or down-slope tumbling (Nemec, 1990).

Polymict conglomerates, on the other hand, are texturally more mature and represent hyperconcentrated flows. By their inverse grading in basal divisions, they document an incipient sorting of bigger clasts by dispersive pressure (Lowe, 1982; Mulder & Alexander, 2001). In the lower part of sequence 2 these beds represent homogeneous units enriched in a sandy matrix and lack, as a distinctive feature, scoured contacts even overlying soft muddy beds (figure 5-b). These relations point to a reduced basal drag of relatively impermeable flow units, and may be associated to an emplacement by hydroplaning (Mohrig et al., 1998). In contrast to these compact plugs, lenses of polymict cobble-boulder conglomerates, which by their disordered fabric and textural immaturity resemble debris fall deposits, may have been derived by the extraction of the sandy matrix of a “leaky” debris flow within a gravel-rich front of a debris flow (Nemec, 1990), from which gravels may have detached and formed debris fall deposits in front of a parental debris flow (Sohn, 2000).

Amalgamated conglomeratic units of segment 2 may represent multiphase debris flows deposited under different rheological regimes (Sohn, 2000). Basal divisions of inversely graded, thin-bedded pebbly sandstones and conglomerates may be attributed to a flow regime of traction carpets, while normally graded upper conglomeratic divisions may have accumulated by the fall-out of particles initially supported by a turbiditic suspension mechanism (Figure 6-c; Sohn, 1997, 2000). The segregation of conglomeratic and sandy layers becomes distinctly bipartite in segment 3 with a basal conglomeratic traction carpet (or inertia layer; Postma et al., 1988) separated from an overlying sandy turbidite layer by a sharp interface. These debrite-turbidite couplets contain outsized clasts reaching the fraction of boulders that align close to the top of conglomeratic beds or at the very interface to overlying sandy

layers. This situation closely relates to the experimental findings of Postma et al. (1988) and their conceptual model of a laminar inertia flow capable of supporting large clasts and a segregating upper turbulent layer. The trapped oversized clasts of figure (5-d) may thus have slid close to the interface of a conglomeratic laminar and sandy turbulent flow unit. In contrast to other deltaic foreset sequences, secondary reworking of the depositional units is subordinate and can only be inferred from an isolated channel-shaped bedform with the matrix-rich conglomerate that composes facies *A-IIIb*.

Conglomeratic and sandy deposits of sequence 4 display channel-like bedforms and may be associated to a distributary delta plain. Intervening muddy layers are likely to represent floodplain deposits. The lack of pedogenic horizons supposes a subaquatic environment and suggests that sea-level rise outpaced depositional aggradation.

Summarizing, the depositional environment of a Gilbert-type delta may be constrained by its three basic architectural elements and a progradational setting. The bottomset contains outrunners of oversized clasts and debris-fall deposits, signaling the proximity of a delta front. Distal conglomeratic units intercalated with muddy layers of the bottom- to toeset consist of little differentiated, matrix-rich debris flows. These conglomeratic deposits become increasingly differentiated into composite basal conglomeratic inertia and overlying sandy turbidity flow units. The presence of oversized clasts at clear-cut interfaces attests to elevated grain-support mechanisms in the proximal foreset setting of facies assemblage 3. Delta plain deposits of sequence 4, finally, may be assigned to a topset.

Ubiquitous oversized clasts point to a steep foreset gradient and pose the question about a tectonically active catchment area. For polymict conglomerates, gross

compositional trends of sequences 1 and 2 record a gradual increase in input of the Precambrian Mangos Granulite suite and gneiss components of the southern Sevilla Metamorphic Belt. The exclusion of metasedimentary components of the Inner Santa Marta Metamorphic Belt restricts the catchment area to the central and southern part of the SNSM. Compositional variations of the gravel spectrum are minor, presupposing that the drainage system maintained a similar structural position since its onset, as might be expected for the gradual denudation of a topographic scarp at an active fault. This relatively uniform contribution is punctuated by the monomict input of the of the Aracataca batholith, which must have constituted both fault scarp and hillslopes of an incised valley. These findings support a localized uplift at the western flank of the SNSM, for which Villagómez et al. (2011) modeled a time span of 29 Ma to 25 Ma, based on AFT data.

TABLE 2. SUMMARY OF FACIES AND FACIES ASSEMBLAGES FOR THE ARACATACA BASIN

ARACATACA BASIN - WESTERN NSM				
Facies assemblage	Code*	Description	Inferred depositional process	Interpreted environment
A-IVc	Fm	Massive mud- to siltstone layer with sand dikes	Overbank deposit	Gravelly to sandy braid-plain constituting the topset of a Gilbert-type delta
A-IVb	Sm-t	Medium- to coarse-grained sandstone, filling scours or forming planar beds; scoured surfaces are draped by pebbly lag.	Isolated channel fill or crevasse splay on flood plain	
A-IVa	Gh-t	Pebble-cobble conglomerate in lenses, 0.8m to 3m thick. Components are well-sorted, rounded and sand- to clast-supported. Beds tend to be normally graded, display a scoured base (Gt) or are interdigitized with sandstone (Gh). Cobbles display an <i>a(t)b(i)</i> imbrication.	Channelized hyperconcentrated mass flow	
A-IIIb	Gmt	Pebble-cobble conglomerate in lenses with an erosional base, uniformly dispersed rounded components or angular fragments, sand-supported.	Channel-fill of a cohesive (?) debris-slide associated to an erosional scar (Postma, 1984; Nemeč 1990)	Slide-scar in the foreset domain of a Gilbert-type delta
A-IIIa	Sh-m	Laminated medium-grained sandstone grading upwards into massive coarse-grained sandstone with interspersed pebbles and cobbles.	River-generated hyperpycnal flow descending from a delta brink, <i>or</i> : grain flow associated to the turbulent layer of a descending laminar inertia-flow of a debris avalanche (Postma et al., 1988)	Foreset domain of a Gilbert-type delta
A-IIb	Gmb	Lenses, pockets or horizons of outsized boulders with diameters up to 2m, gravel-supported, composed of granite.	Boulders derived from a delta brink zone or from canyon walls of an incised valley by debris-fall avalanches	Boulder accumulation in the fore to bottomset domain of a Gilbert-type delta
A-IIa	Gms	Pebble-cobble conglomerate in tabular beds, 0.8m to 1.5m thick. Clasts are subangular to rounded, sand- to gravel-supported, poorly to regularly sorted. Beds display a coarse-tail inverse grading in their lower part. Cobbles are oriented parallel to bedding or imbricated, exhibiting an <i>a(p)</i> or <i>a(p)a(i)</i> fabric.	Cohesionless debris flow with sheared basal part. Inverse coarse-tail grading due to the forced upward movement of larger clasts (Lowe, 1982).	Foreset of a Gilbert-type delta
A-Ia	Fsm	Laminated mud- and siltstone with calcareous lenses; contain shells and shell fragments of molluscs shells and gasteropods. Associated with lenses of monomict breccia conglomerates and outsized granite boulders (facies IIb).	Suspension fall-out of turbidity flow <i>or</i> of an homopycnal river outflow (Colella, 1987)	Toe- to bottomset of a Gilbert-type delta

*Facies code modified from (Miall,1978)

5.1.2 Palomino Conglomerates

This conglomeratic to sandy sequence forms an isolated strip to the south of the Oca fault. In an internal or southern position it involves facies associations that comprise hyperconcentrated mass flows of a braid plain that grade, northwards, into sandy sequences associated to flood plain and siliciclastic shoreface deposits. It overlies the Muchachitos Gneiss of the Sevilla Metamorphic Belt (Tschanz et al., 1974), a quartz monzonite of a probable Jurassic age, which has been mapped as the Palomino Stock (Tschanz et al., 1974) and a Late Jurassic volcanic suite. Further toward the hinterland, the basement exposes the Precambrian Mangos Granulite and igneous to volcanic suites of the Sierra Nevada Province constitute the present exposure level (Tschanz et al., 1974).

We discuss this sequence by means of a southern conglomeratic and a northern sandy section, each one resting unconformably on the Jurassic igneous or Late Paleozoic metamorphic basement of the Sevilla Metamorphic Belt (Palomino stock and Muchachitos Gneiss; Fig. 4). Their comparable structural position above a pre-Cretaceous basement suggests that their particular facies associations belong to an equivalent systems tract, which comprise three laterally equivalent sequences and are associated to (1) a gravelly braidplain, (2) a sandy floodplain and (3) a shoreface environment. These sequences are mapped as facies belts (figure 4), according to the predominance of conglomeratic, sandy and sandy to muddy alternations.

We label facies according to the classification of Miall (1978c, complemented by Miall 2006), combining some gravelly and sandy facies of similar bedforms into one category for simplification.

Sequence 1 is composed of pebble-cobble conglomerates and includes subordinately intercalations of sandy and muddy beds (Table 3). Pebble-cobble conglomerates form units up to 4 m thick (facies *P-Ia*, Table 3). They display a sand- to gravel-supported fabric with an ill to moderately sorted clast assemblage. Scoured bases and internal reactivation surfaces may be draped by cobbles and outsized boulders (Fig. 5-e). Discontinuous sand lenses subdivide conglomeratic beds into individual flow units. Major bedforms with erosive bases grade upwards into pebbly sandstones, which display trough-shaped laminations of channel fills (Fig. 5-e). Toward the internal part of this braidplain belt (or sequence 1) channel-shaped bedforms comprise stacked units. At the transition to the sandy facies belt (or sequence 2), isolated conglomeratic lenses increasingly interfinger with sandstones and form thin intercalations within sandy beds.

Facies *P-Ib* consists of gravel- to sand-supported, well-sorted and rounded pebble-cobble conglomerates, with oriented pebbles displaying an *a(t)b(i)* imbrication. These m-sized units are subdivided by thin sandstone beds with sharp contacts.

Facies *P-Ic* is made up of pebbly sandstones and pebble conglomerates with angular components arranged in foresets that abut against underlying beds at a sharp, non-tangential contact.

In sequence 2 conglomerates are still present, though subordinate. Facies *P-IIa* consists of ill- to well-sorted and angular to rounded pebble-cobble conglomerates, with a sand-supported texture. Bedforms are tabular and display a scoured base. In pebbly stringers and thin beds granule and pebble-sized clasts are disorganized. Facies *P-IIb* contains laminar coarse-grained sandstones, overlain by massive sandstone. Facies *P-IIc* consists of massive sandstones making up channel fills. Facies *P-IId* is represented

by normally graded silt- to mudstone beds, the latter displaying a bluish-grey tone and evolving from the silty substrate by a normal gradation.

Sequence 3 is composed of sandy units up to 4 m thick embedded within massive bluish-grey mudstones of variable thickness (facies *P-IIIa* and *P-IIIb*). Basal contacts of the sandy units are sharp and planar or shaped by open scours. Fluidization is evident by flame structures that, in an advanced stage, may disrupt even thick beds. Internally, sandstone units display amalgamated sets of cross-stratification with bi- or multi-directional flow directions. Interwoven laminae between coarse- and medium-grained cross-sets argue for aggradation under bi-polar currents. Muddy beds, on their turn, are massive and devoid of fossil remnants and trace fossils. By their gently inclined laminae and bi- (or multi- polar) cross-sets the low-angle cross stratification of sandy units bears a resemblance to a hummocky cross stratification (HCS). Where basal laminae abut against underlying mudstone, these bedforms may derive from migrating dunes.

Interpretation

Conglomeratic channel and sandy flood plain are the principal components of the *gravelly braidplain facies association* of sequence 1. Stacked channels are aligned in a southern continuation of an embayment delineated by the facies belts of sequences 2 and 3 (Fig. 4) that may designate the axis of a fluvial trunk system. Within this environment, bedforms of facies *P-Ib* are interpreted as longitudinal bars and those of facies *P-Ic* as straight-crested dunes, considering a deposition of relatively well-sorted gravel by avalanching along foresets. Occasionally, conglomeratic beds associated to

longitudinal bars and transverse bedforms overly dark muddy layers, which we group into the sandy flood plain facies *P-IIId*. By their dark color and their lack of pedogenetic features, these muddy layers may indicate a subaqueous depositional environment.

In the *sandy flood plain facies association* of sequence 2 matrix-rich conglomerates of facies *P-IIa* are interpreted as cohesionless debris flows, on account of their high and variable proportion of sandy matrix and their strongly disorganized pebbles, whose occasional unstable position may indicate a cohesive, in-situ freezing of a mass flow (Lowe, 1982; Rasmussen, 2000). The laminar sandstones of facies *P-IIb* may record flashes of minor sheet floods. The bluish-grey color of the tabular mudstones of facies *P-IIId* suggests again a subaqueous sedimentation by suspension fall-out. This interpretation is corroborated by their lateral transition into mottled sand- to siltstones, which overly the Palomino Pluton at the eastern margin of the mapped sandplain belt (AP-46, Fig. 4). These reddish sand- and silt-stones include pedogenic structures, like mudcracks, coated nodules and carbonized rootlets, advocating a subaerial exposure.

Coarse-grained sandstones with HCS and large-scale wave ripples or dunes characterize the *shoreface facies association* of sequence 3. Gradual changes between these structures have been attributed to textural changes from fine- to coarse-grained sandstones (Bourgeois & Leithold, 1984; Leckie, 1988) and characterize oscillating to unidirectional flow conditions of a lower shoreface (Leckie, 1988; Duke et al., 1991). The lack of fossil remains points to a high-energy, near-shore environment (Bourgeois & Leithold, 1984). In this facies belt well-sorted conglomerates with a shape-preferred selection diagnostic of a beachface facies (Bluck, 2010) are absent, a circumstance that could be attributed to a low preservation potential of this type of deposit in a wave- to

storm-dominated siliciclastic coast (Bourgeois & Leithold, 1984). The key features of these depositional environments should be based on a low-gradient braidplain to coastal setting dominated by basinal storm- and wave-driven processes. If our interpretation of an embayment of the coastline at a major river trunk is correct, we should discard the existence of a delta and refer to this coastal configuration as a small-scaled estuary of a transgressive setting (Dalrymple & Choi, 2007). A transgressive frame is indicated, further, by the superposition of sandplain over braidplain and shoreface over sandplain facies associations, as documented by the two logs presented in figure 4.

TABLE 3. SUMMARY OF FACIES AND FACIES ASSEMBLAGES FOR THE PALOMINO BASIN

PALOMINO BASIN - NORTHERN SNSM				
Facies assemblage	Code*	Description	Inferred depositional process	Interpreted environment
P-IIIb	Fm	Massive mudstone layers	Suspension sedimentation	Middle to lower shoreface, below fair-weather wave base
P-IIIa	St/Sp	Well-sorted coarse sand, grouped into low-angle sets of trough or planar cross stratification constituting units 1 m to 4 m thick. Major bedforms and the base of the units are erosive and contain pebbly lags. Foresets indicate bi- or multidirectional currents. Sand units affected by fluidization.	Wave-worked tractional transportation of sand at middle to lower shoreface conditions and migration of predominantly sinuous bedforms or formation of rip channels	
P-IId	Fm	Massive silt- to mudstone in tabular beds	Suspension sedimentation	Sandy flood plains with channels representing cut-and-fill events and
P-IIc	Sm	Medium- to coarse-grained sand in isolated channel fills	Sands emplaced by cut-and-fill processes during flood	
P-IIb	Sh, Sm	Coarse laminar sandstone, grading into massive	Flashy overbank floods	
P-IIa	Gh, Gm	Pebble conglomerate in tabular beds with scoured bases or thin non-erosive conglomerate layers, clast or sand-supported, grading into pebbly sandstone. Disk-shaped clasts tend to be oriented parallel or at high angles to bedding.	Hyperconcentrated gravel or cohesionless debris flow emplaced on flood plain. Clasts oriented at a high angle to bedding attest to rapid frictional freezing.	
P-Ic	Gp, Sp	Pebbly sandstone with moderately inclined foresets (<30°) that form sharp angles with an unscoured base. Foreset laminae are graded.	Straight-crested dunes, associated to transverse bedforms	Gravelly channel fills of a sub-aqueous(?) braidplain. - Major units may be composed of stacked channel units along a structural depression or less confined flow conduits, that interfinger with a sandy flood plain of facies association II.
P-Ib	Gh, Sm	Pebble-cobble conglomerates in tabular, ungraded units, with rounded, relatively well-sorted clasts, which are sand- to gravel-supported. Disk-shaped clasts are oriented parallel to bedding or display, where closely packed, an a(t)b(i) imbrication. Beds have an unscoured base and are overlain by sandstone with sharp contacts.	Longitudinal bar of braided stream formed by bedload transport	
P-Ia	Gh-t, Sm-t	Pebble-cobble conglomerates with rounded, relatively well-sorted clasts, sand- to gravel-supported, forming lenses and channel fills 0.8 m to 4 m thick. Their scoured base may be draped with outsized clasts. Units are stacked or form individual channel fills. Minor channels fills are ungraded, major channel fills fine upward, grading into gravelly sandstone. Isolated conglomerate lenses interfinger with sandstone. Closely packed cobbles display an a(t)b(i) fabric.	Channelized hyperconcentrated gravel flows defining stacked patterns or relatively isolated shallow channels, when alternating with sandy beds.	

*Facies code modified from Miall (1978c, 2006)

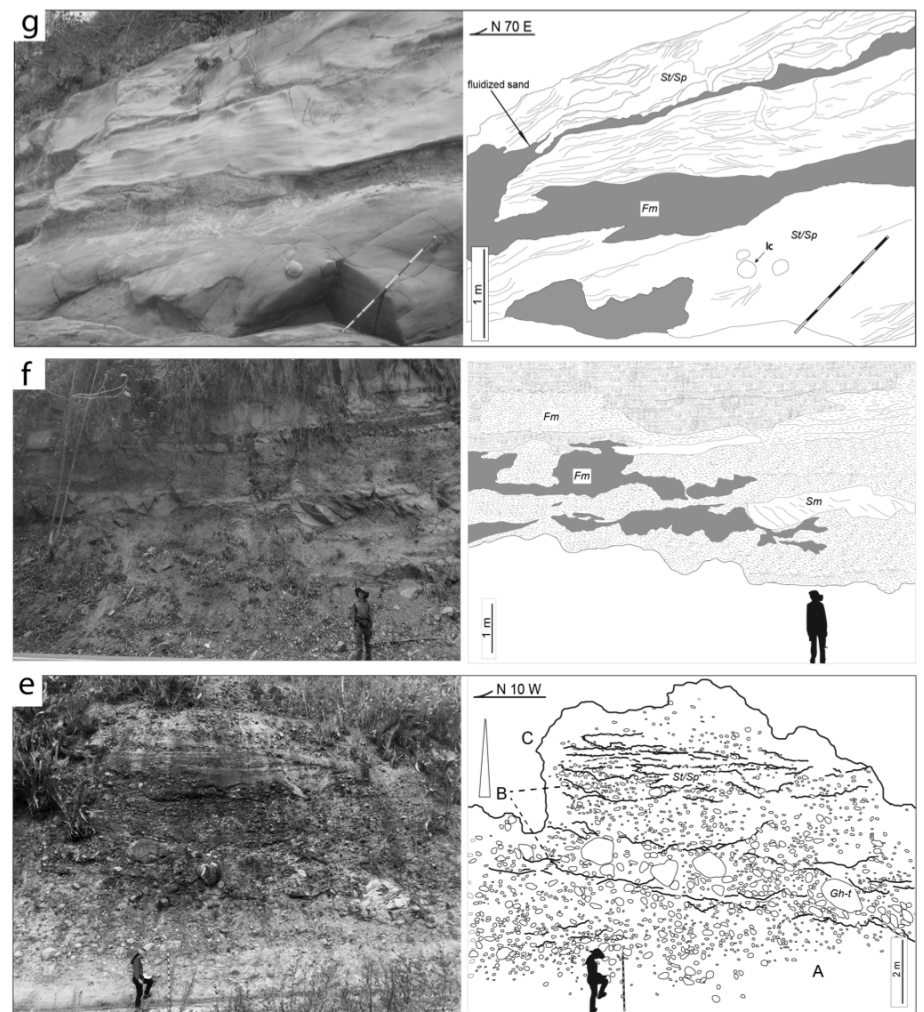
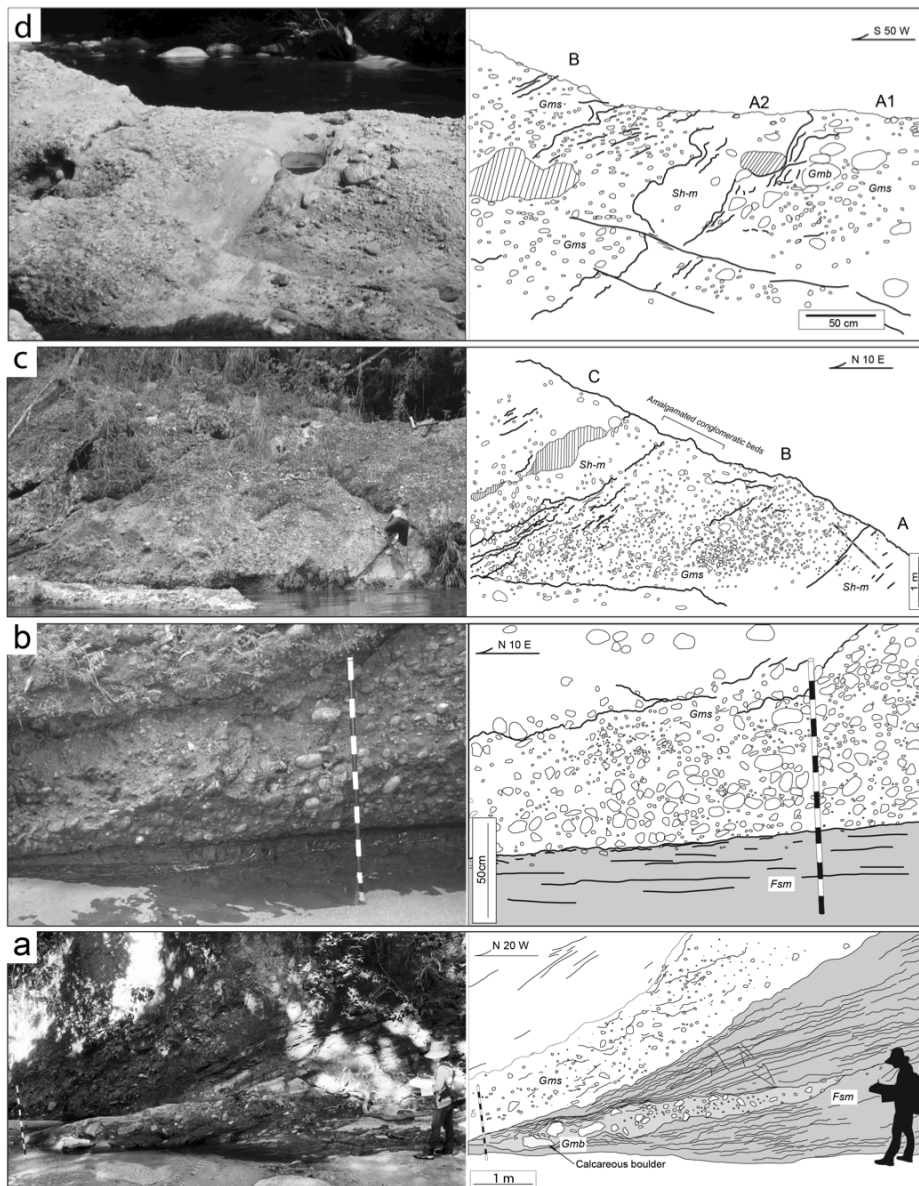


Figure 5.

Left:

Figures from the Aracataca section (facies refer to Table 2.):

a) Facies *A-Ia* and *A-IIb*, bottomset of a Gilbert-type delta: Lens of debris flow embedded within mud- and siltstone. Boulders are concentrated at the head of this mound-like accumulation (termination to the left) and consist of granite and micritic limestone. The boulders are variously tilted with respect to bedding and, according to the muddy matrix, point to a cohesive flow rheology. For scale, refer to the Jacob staff with dm-scaled divisions.

b) Facies *A-IIa*, lower foreset of a Gilbert-like delta: Polymict pebble-cobble conglomerate with sub-rounded clasts, displaying a tendency of a coarse-tail inverse grading at its base and a normal grading in its upper half. Contact to underlying mudstone is planar. Imbricated and closely packed cobbles may display an *a(p)a(i)* fabric (basal domain to the right of the Jacob staff). Flow direction is to the right.

c) Facies *A-IIa*, lower to upper foreset of a Gilbert-like delta: Three amalgamated conglomerate beds may be differentiated by textural criteria, such as clast size and orientation that vary from disordered to bed-parallel and may be imbricated. Despite of patchy clast-size variations parallel to bedding, units A to C display some contrasting trends of an internal organization: unit B sets in with a sand-supported pebble conglomerate and grades into a coarser, clast-supported cobble-pebble conglomerate. Pebbly sandstones form a top-most division and are loosely stratified, attesting to frictional aggradation. A similar stratified pebbly sandstone forms the top of unit A. The cobble conglomerate of unit C sets in with boulder-sized clasts and displays a highly disordered fabric. The scale is given by the Jacob staff (1.4 m). Label units A, B, C; indicate tendency for normal/reverse grading; flow direction.

d) Facies *A-IIa* and *A-IIIa*, upper foreset of a Gilbert-like delta: A thick unit of pebble-cobble conglomeratic (unit A1) is capped by a pebbly sandstone (unit A2) (center of figure). Outsized angular granite cobbles and boulders drape or are closely aligned to this lithological boundary. In the central upper part, this interface steps up to a stratigraphically higher horizon and marks the termination of a row of imbricated boulders (center-right of figure). Flow direction is to the left.

Right:

Figures from the Río Negro section (facies refer to Table 3.):

e) Facies *P-Ia* and *P-IIIa*, channel fill of braidplain: A crudely stratified conglomeratic unit A is cut by a densely packed unit B that, on its turn, grades into a stratified pebbly sandstone (unit C). Units B and C are interpreted to define a channel fill. Outsized boulders mark the erosive base of unit A.

f) Facies *P-IIc* and *P-IId*, sandy floodplain: Scoured sandy beds alternate with tabular strata of siltstone and gray mudstone. To the right-hand side of the center, a sandstone forms asymmetric channel fills within a muddy substrate.

g) Facies *P-IIIa* and *P-IIIf*, middle to lower shoreface: Sandy cross-stratified units alternate with thin beds of gray mudstone. The coarse- to medium-sized sandstone displays low-angle, tangential foresets, which may abut against underlying muddy beds. Foresets indicate bi- or multidirectional currents. Sandy units are strongly distorted and partially homogenized by fluidization.

5.2 Clast Counting

Thirteen clast counts were performed in Rio Aracataca and Rio Negro sections, at the western and northern foothills of the SNSM respectively; results are summarized in Figures 3, 4 & 6 and location of clast counts in Table 4. The conglomerates are mainly matrix to clast supported, with clast sizes ranging from granules to boulders. Many of the pebbles are sub-rounded, and clast variety increases up-section in both of the basins with twelve different litho-types identified.

Plutonic clasts: From granites through sienogranites and granodiorites, with high contents of plagioclase, quartz, biotite, and hornblende and alkaline feldspar is present in a lower proportion. Granodiorites are in fact the most abundant population of clasts in both basins, being at 100% of clast counts at the base of the sequence of the Rio Aracataca section where this litho-type decreases towards to a proportion between 30%-50%. In the Rio Negro section, granites and sienogranites are present at 20%-50% of the counts throughout the whole section.

Volcanic clasts: Are the second major component represented by a suite of multi-coloured basalts (red, green, black) and spilites that vary widely in texture (aphanitic, porphyritic, amigdular). In the Rio Negro section, the presence of andesites and rhyolites is common; in the Rio Aracataca, basalts are dominant over more acid volcanic rocks. Volcanic clasts constitute between 20%-40% of the composition, and its contribution increases to the top in both basins.

TABLE 4. SUMMARY OF CLAST COUNT ANALYSIS, AND LOCATIONS

Clast/station	EP-02	EP-06	EP-08	EP-17	CP-085	EP-37	EP-41	AP-46	EMP-43	EMP-47	EMP-50	EMP-51	EMP-26
Lat (°N)	10°35'06"	10°35'14"	10°35'15"	10°35'21"	10°35'19"	10°35'20"	10°35'18"	11°13'07"	11°11'58"	11°12'14"	11°12'39"	11°12'41"	11°15'44"
Long (°W)	74°03'16"	74°03'18"	74°03'21"	74°03'47"	74°04'28"	74°04'53"	74°05'06"	73°28'41"	73°30'02"	73°30'08"	73°30'16"	73°30'17"	73°35'25"
Granite	118	119	70	13	45	83	21	N.D.	21	46	42	60	16
Rhyolite	N.D.	N.D.	N.D.	N.D.	N.D.	N.D.	N.D.	N.D.	17	19	4	6	N.D.
Andesite	N.D.	N.D.	N.D.	N.D.	N.D.	N.D.	N.D.	N.D.	10	3	5	N.D.	7
Red Basalt	31	N.D.	1	9	5	26	14	1	2	1	N.D.	N.D.	8
Green Basalt	N.D.	N.D.	N.D.	16	8	32	42	N.D.	2	8	3	5	N.D.
Black Basalt	51	N.D.	2	22	20	23	27	6	4	9	4	6	24
Gabbro	N.D.	N.D.	N.D.	2	N.D.	2	5	N.D.	1	1	3	3	N.D.
Volcanic Breccia	N.D.	N.D.	N.D.	N.D.	N.D.	N.D.	N.D.	N.D.	N.D.	2	N.D.	8	N.D.
Gneiss	N.D.	N.D.	N.D.	5	13	N.D.	13	122	N.D.	2	13	5	7
Quartzite	N.D.	N.D.	N.D.	10	N.D.	N.D.	10	N.D.	1	N.D.	1	N.D.	N.D.
Sandstones	N.D.	N.D.	N.D.	1	N.D.	N.D.	2	N.D.	N.D.	N.D.	N.D.	3	N.D.
Mudstones	N.D.	N.D.	N.D.	N.D.	N.D.	N.D.	N.D.	N.D.	N.D.	N.D.	N.D.	1	2

Metamorphic clasts: Are characterized by amphibolite gneisses, leucocratic gneisses bearing blue quartz, anorthosites, granulites, and a minor contribution of quartzites. In the Rio Aracataca section, these litho-types appear after sequence I (Fig. 3) and account for up to 30% of the clasts. In contrast in the Rio Negro section the base of the sequence consists of oligomictic conglomerates from facies *P-Ia* with leucocratic gneisses as the main component (90%), decreasing to the top to 20% and quartzites (5%).

Sedimentary clasts: Mainly reworked sandstones and rare mudstones. Carbonate clasts are common in sequence I at the Aracataca River, facies *Fsm* from facies assemblage Ia and often contain macrofossils. The presence of 1%-4% of sedimentary clasts coincides with the occurrence of metamorphic clasts.

Ultramafic clasts: These clasts appear at the same locations as the metamorphic and sedimentary clasts. This litho-type (3%-5%) is

represented mainly by dark green gabbros and volcanic breccias, which locally may represent up to 10% of the counts.

5.3 Heavy Mineral Analysis

The heavy mineral spectra found in the SNSM Neogene sedimentary rocks includes garnet, chlorite, amphibole, pyroxene, rutile, muscovite, biotite, andalusite, sphene, tourmaline, epidote, and opaque minerals such as ilmenite and pyrite (Table. 5). The Aracataca conglomerates samples CVI1302, CVI13115 contain rutile, pyroxenes, and garnet. Green augite-aegirine clinopyroxenes are present in all samples. Amphiboles are present in the form of green and brown hornblende. Garnet is also present, but is much more abundant in the Palomino basin in samples AP-45 and AP-46, whereas in the Aracataca basin garnet occurs as a minor component. Sphene and rutile are found in both basins and rutile content is observed to increase in samples towards the top of the Rio Aracataca section.

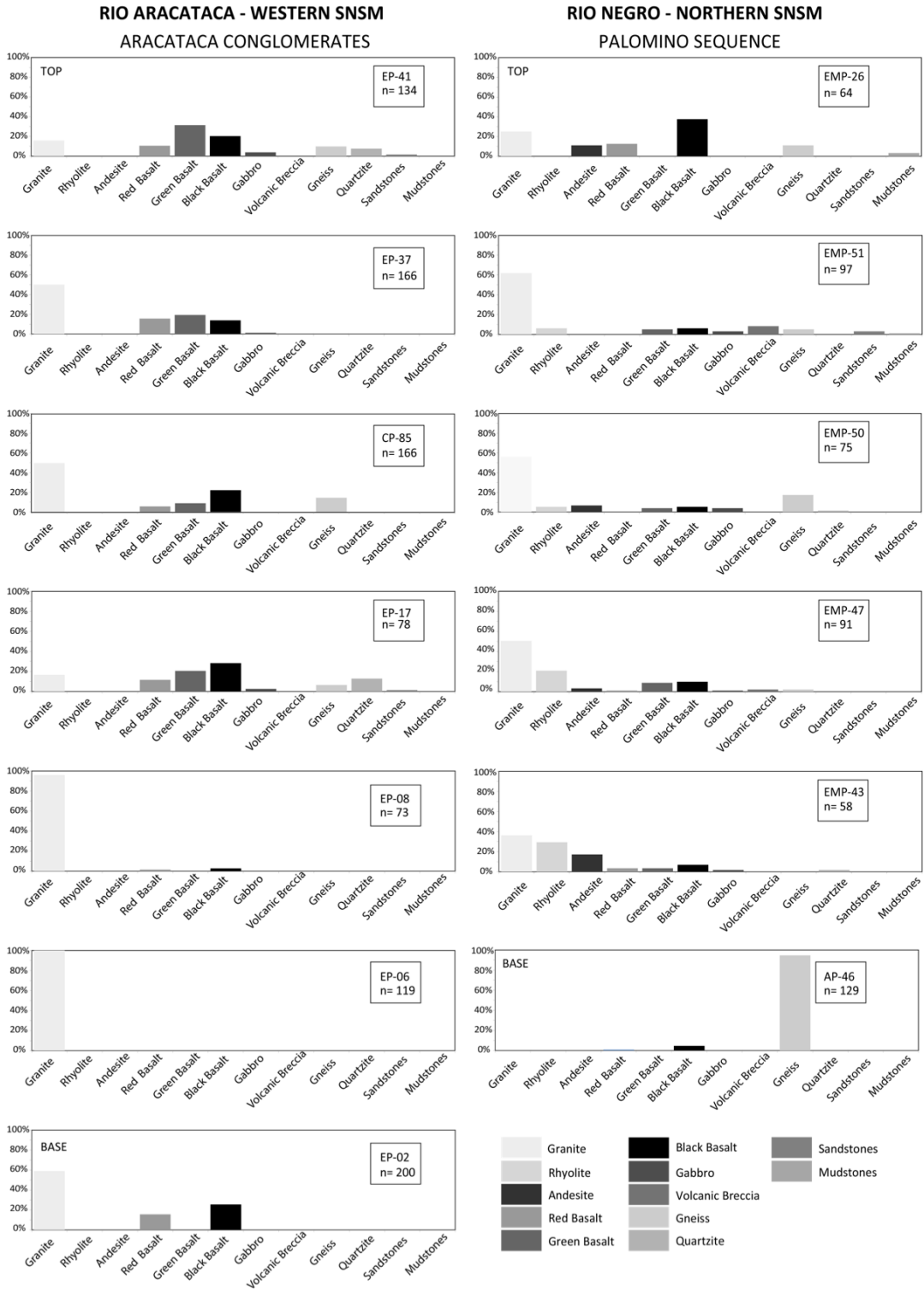


Figure 6. Clast counts from The SNSM Cenozoic Aracataca and Palomino basins at the Negro and Aracataca rivers, respectively, counts were done following the ribbon method (Howard, 1993).

TABLE 5. LOCATION, DESCRIPTION, DIAGNOSTIC MINERALS, ZIRCON MORPHOLOGY, AND AGES OF SAMPLES ANALYZED BY U-Pb LA-ICP-MS

Sample Code	Coordinates		Description	Diagnostic Minerals*	Zircon Morphology and CL Zoning	Age Core	Age Rim
	Lat (°N)	Long (°W)					
<i>Post-Oligocene Sediments</i>							
CVI1302	10°35'18"	74°05'06"	Coarse grained sandstone matrix, of medium pebble polymictic conglomerates with abundant clasts of igneous and metamorphic rocks, rounded to subrounded	Zr, Rt, Ap, Chl, Aeg, Hbn, Bio, Plag	Angular to subangular, subhedral and tabular prismatic, purple brown translucent, irregular edges, with inclusions, showing xenocrystic cores in metamorphic grains and rim growth by recrystallization,	803 -1434 Ma	889-1295 Ma
CVI13115	10°35'15"	74°04'30"	Silica cemented grey conglomeratic sandstone, well sorted, subrounded grains	Rt, Zr, Hbn, Chl, Str, Aeg, Gr,	Angular to subangular, subhedral and tabular prismatic and xenomorphic, purple brown translucent, irregular edges, with inclusions, showing xenocrystic cores in metamorphic grains and rim growth by recrystallization; magmatic grains are translucent, euhedral and tabular prismatic, with regular zoning and convolute zoning	986-1182 Ma 116.5-135.6 Ma	881 -1324 Ma 126.1-193.3 Ma
EMP-16	10°32'21"	74°03'50"	Conglomeratic sandstone	Zr, Gr, Ep, Ap, Str, Aeg, Hbn	Subangular to subrounded, subhedral and tabular prismatic and xenomorphic, purple brown translucent, irregular edges, with inclusions, showing xenocrystic cores in metamorphic grains and rim growth by recrystallization; magmatic grains are translucent, euhedral and tabular prismatic, with regular zoning and convolute zoning	944-1507 Ma 592 Ma 179.5-205.8 Ma	955-1365 Ma 181.5-185.5 Ma
CP088	10°31'33"	74°03'13"	White conglomeratic quartzsandstones, angular-subangular, regular sorted with lithics, very friable highly porous, probably from fluvial coastal deposits	Zr, Rt, Chl, Aeg, Hbn, Bio	mainly subrounded and some subangular, subhedral prismatic, purple and brownish, translucent, irregular edges, with inclusions, showing xenocrystic cores in metamorphic grains and rim growth by recrystallization, magmatic grains present are translucent, euhedral and tabular prismatic and showed no recrystallization	997-1359 Ma 132.7-189.5 Ma	904-1398 Ma 127.9-197.6 Ma
AP-045	11°13'12"	73°28'39"	medium to coarse grained sandstone matrix from a coarse pebble polymictic conglomerate, with abundant gneissic and granitic clasts and also a minor mafic fraction	Zr, Gr, Hbn, Chl, Ap, Aeg, , Ep, Str, Rt	Rounded to subangular, subhedral fewer prismatic, purple, translucent and yellowish, seldom inclusions, some grains exhibit xenocrystic cores and rim growth by recrystallization, magmatic grains are translucent and more prismatic, euhedral with oscillatory zoning	897-1300 Ma 189.2 - 258.6 Ma	888 - 1241 Ma 191.2 -265.8 Ma
EMP-49A	11°12'33"	73°30'17"	Sub-lithic green coarse sandstone	Gr, Chl, Hbn, Str, Aeg,	Subangular to subrounded, subhedral and tabular prismatic and xenomorphic, purple brown translucent, irregular edges, with inclusions, showing xenocrystic cores in metamorphic grains and rim growth by recrystallization; magmatic grains are translucent, euhedral and tabular prismatic, with regular zoning and convolute zoning	850-1259 Ma 178.2-192.7 Ma	945-1132 Ma 180.2-198.9 Ma

*Mineral abbreviations Aeg = Aegirine, Ap= Apatite, Bio= Biotite, Chl= Chlorite, Gr=Garnet, Hbn= Hornblende, Rt= Rutile, Str= Staurolite, Tre= Tremolite, Zr= Zircon

5.4 Detrital Zircon U-Pb Dating Results of the Conglomerate Matrix

Four samples were collected in the western foothills of the SNSM from outcrops along the Rio Aracataca, and outcrops at the adjacent Río Piedras and the Macaraquilla Creek. The analysed samples are located into the Aracataca stratigraphic section (Fig. 3). Additionally, two more samples were collected from outcrops in the northern foothills of the SNSM in the Negro and Ancho rivers at the Palomino basin that are located in the Rio Negro stratigraphic section (Fig. 4). We collected the sandy matrix of the conglomeratic facies. The detrital samples contain a major amount of euhedral and tabular prismatic magmatic zircons with classical oscillatory zoning patterns. Metamorphic zircons are mostly purple, brownish or translucent, and show more complex growth zoning with local intermediate reabsorption, sector and patchy zoning (Fig. 7). U-Pb ages range from 1450 to 116 Ma, with two main peaks at 1000 Ma and 200 Ma (Fig. 8). The main age population corresponds to the 1000 Ma peak, which is present in all samples, independently of stratigraphic position in the western and northern basins (Fig. 8). The Lower Jurassic to Lower Cretaceous ages between 200 and 116 Ma are the second major population of zircons of the conglomerates. This population significantly decreases up-section and Precambrian ages become dominant. In the Aracataca basin, samples CP088 and CVI13115 are similar, with core ages mainly from 950 to 1360 Ma and a minor population of zircons with Mesozoic (190 to 130 Ma) crystallization ages. One zircon with a discordant Meso-Archean core age of 2970 Ma was also detected.

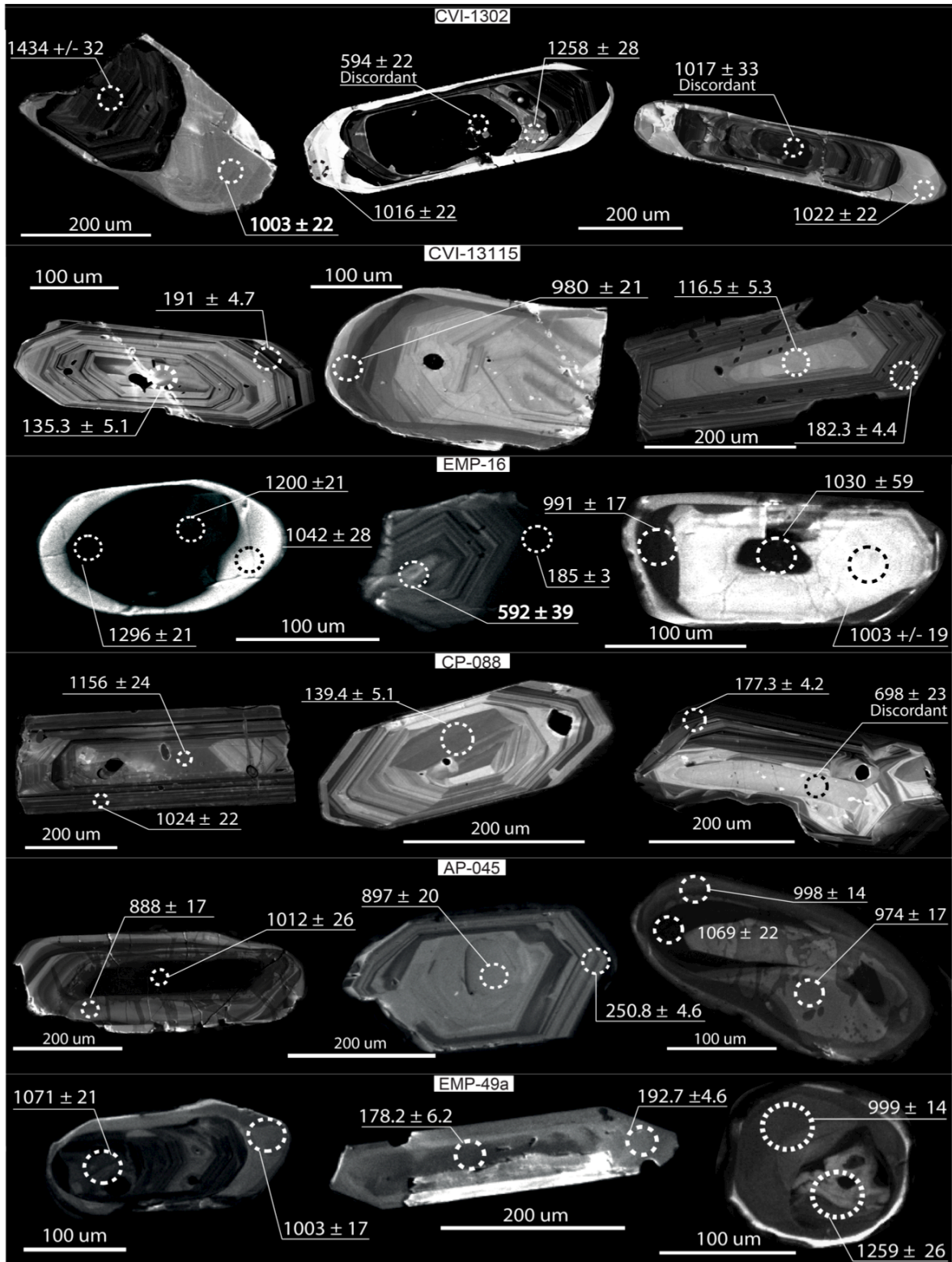


Figure 7. Panchromatic cathodoluminescence imagery of zircons analysed in this study, there are representative crystals for each sample labelled at the top with the name of sample. A high variability of crystals is observed with different age distributions showing two main distinctive populations. Circles correspond to the ablation spot.

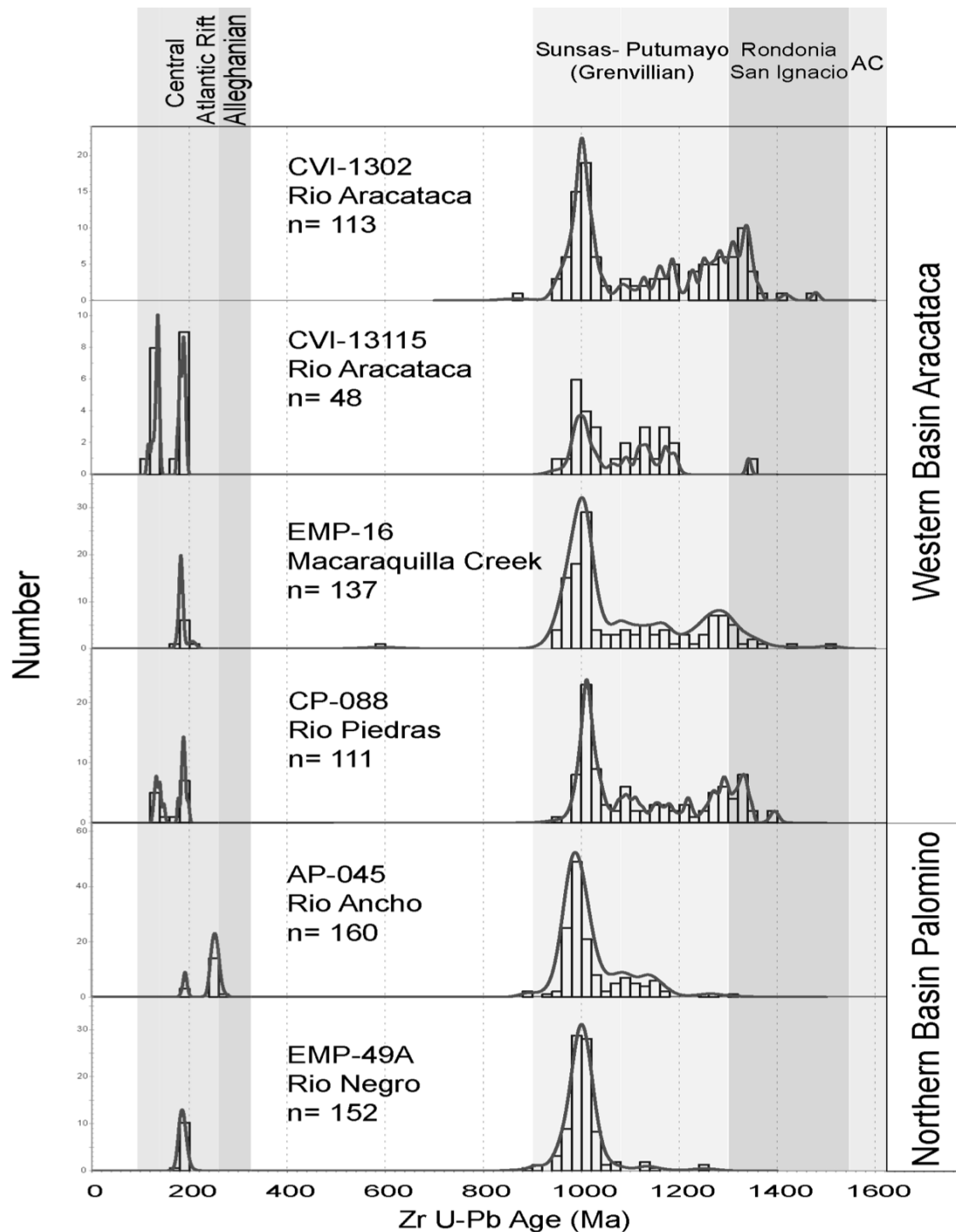


Figure 8. Detrital zircon U-Pb LA-ICP-MS probability density plots obtained from the Post-Eocene sediments of the western and northern foothills of the Sierra Nevada de Santa (Aracataca and Palomino basins), data are organized from base to top for each basin, there is still not a precise stratigraphic correlation between basins due the lack of fauna in the coarser deposits, however zircon populations are intimately related implying the same sources supplied with sediments both basins, with an older signature in the Aracataca basin at the western border of the SNSM massif.

Zircons from sample CVI1302 have ages older than 800 Ma, with peaks at 1100 Ma, 1200 Ma, and 1320 Ma, similar to sample EMP-16. Samples AP-45 and EMP-49a from the northern basin show nearly equal age trends with the most significant 1000 Ma peak in both samples and lesser populations of Jurassic and Triassic peaks in sample AP-45. The only significant difference between age populations from samples of the Aracataca and Palomino basins is the presence of a 1300 Ma age peak in the Aracataca basin (Fig. 8).

5.5 Detrital Fission Track Thermochronology

5.5.1 Zircon fission-track results

Zircon fission-track ages were determined for 751 crystals in eight samples from the western and northern foothills of the SNSM. In the Aracataca basin, samples CVI-1302 and CP-088, show a fixed major peak age of 103 ± 5.5 Ma, sample EMP-16 shows a peak of 107 ± 7.2 Ma; with a secondary age population of 52-74 Ma. A minor younger peak at 29-33 Ma is present in samples CVI-1302 and CP-088, but is absent in sample EMP-16, which was collected closer to the base of the stratigraphic section. Sample EMP-16 yields as well an older age component of 158 ± 27 Ma. In the Palomino basin, the fission-track age distributions are characterized by a major moving peak between 94-132 Ma present in all samples that rejuvenates towards. A second moving peak population ranging from 55-80 Ma is also present (Fig. 9). The pre-orogenic signature is defined by ages between 155-170 Ma, and an older population of 250 ± 30 Ma peak.

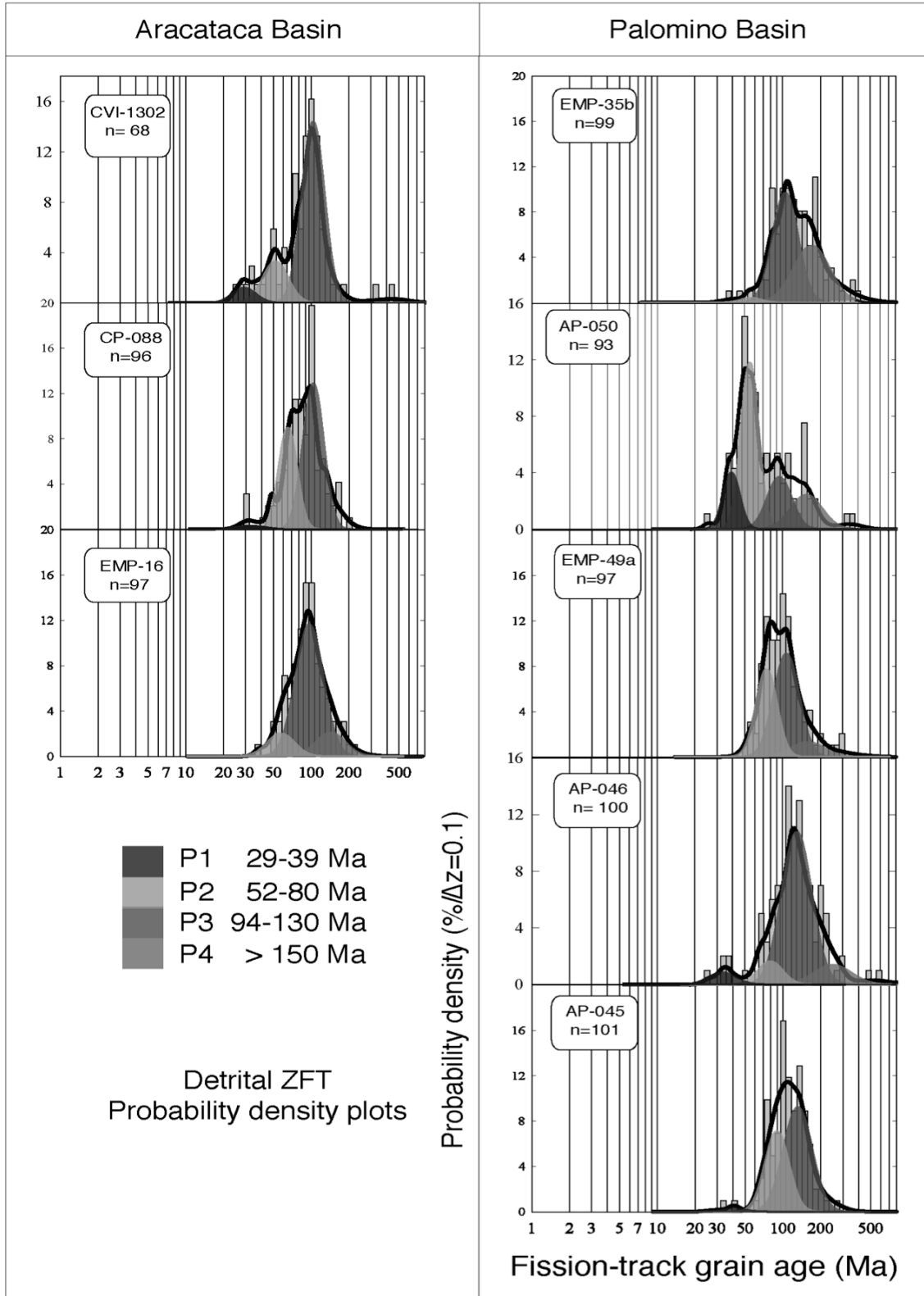


Figure 9. Detrital ZFT data from Aracataca and Palomino basins, probability density plots elaborated with Binomfit (Ehlers et al. 2005; Stewart & Brandon 2004).

The youngest age population was found at samples AP-45, AP-46 and AP-50 and consists of a 33-49 Ma peak. In synthesis, the main age populations for the Aracataca and Palomino basins are defined as an older P4 at 150 Ma, P3 at 94-130 Ma, P2 at 52-80 Ma and P1, the youngest at 29-39 Ma. Probability density plots of all samples are presented in Fig. 9.

5.5.2 Apatite fission-track results

Apatite fission-track ages were obtained for 153 crystals in four samples; CVI1302 and EMP-16 from the Aracataca basin, and from samples AP-45 and EMP-49 from the Palomino basin. In general, other processed samples contained few apatites and in several cases, samples were discarded due to the low quality of crystals, fractures, and inclusions. Three age populations were identified in both basins, and these populations are discriminated as follows with P1 at 19-25 Ma, P2 at 25-35 Ma and P3 at 42-60 Ma. Similarly, to the ZFT peaks, the AFT peaks are becoming younger up-section (Fig. 10).

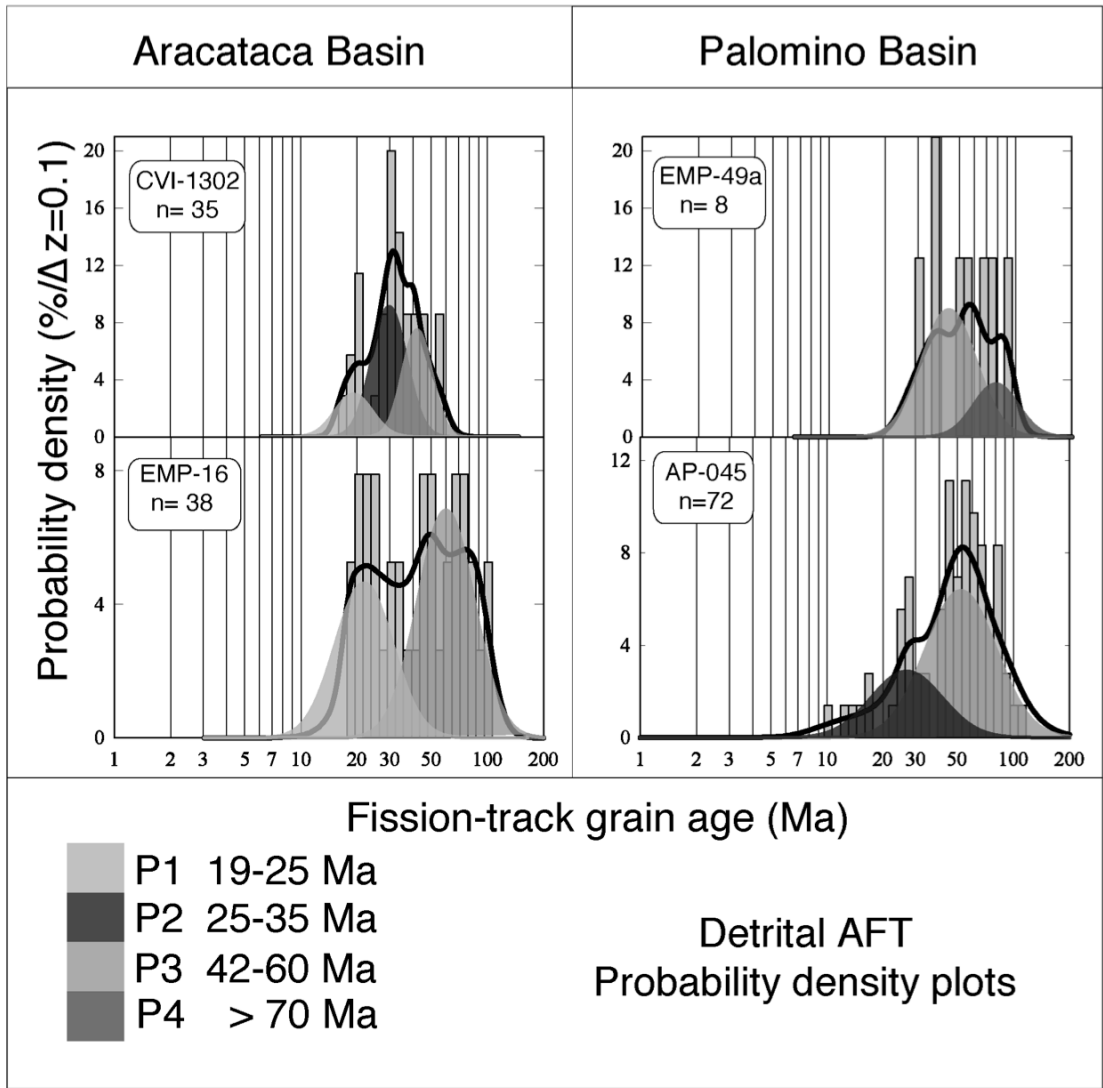


Figure 10. Detrital AFT data from data from Aracataca and Palomino basins, probability density plots elaborated with Binomfit (Ehlers et al. 2005; Stewart & Brandon 2004).

6. DISCUSSION

6.1 Sediment Provenance

There are two main zircon U-Pb age populations in the detrital zircons from the Neogene sedimentary rocks of the Aracataca and Palomino basins. Proterozoic ages

correspond to a widespread population around 850 to 1400 Ma which is found in the SNSM and the Alta Guajira massifs high-grade metamorphic rocks (Weber et al., 2010), which were reworked in Neogene sediments. This population can be related to the Grenvillian belt in the northern Andes (Cardona et al., 2010a; Restrepo-Pace et al., 1997).

Zircon populations from the conglomerates matrix are strictly dominated by the Precambrian U-Pb signal, even if the granitic and volcanic clasts are more abundant. The main sources for the Neogene sediments in the Aracataca and Palomino basins, were the Los Mangos Granulite from the SNP, and the Muchachitos and Buritaca gneisses from the SB, as the detrital zircon U-Pb age spectra are clearly related to the reworking of the Precambrian basement older than \approx 850 Ma. The main zircon U-Pb age population between 900-1150 Ma, which can be found in all samples, corresponds to inherited Grenvillian zircons (Cordani et al., 2005), from the Sunsas-Putumayo province (Ibanez-Mejia et al., 2011).

The older populations of 1300-1600 Ma point to a source in the Rondonia San Ignacio Province in the Amazon Craton. This composite Proterozoic signature has been detected in the Precambrian inliers of the Northern Andes (Cardona et al., 2010a, 2006; Cordani et al., 2005) as well as in the post-Permian metamorphic rocks of the SNSM, the Alta Guajira, and in Cenozoic sediments of the Leeward Antilles (Cardona et al., 2010c; Weber et al., 2010; Zapata et al., 2014). Jurassic and Cretaceous zircons are related to volcanoclastic deposits and arc granodiorites and sienogranites. These zircons indicate the constant erosion of the Central Batholith of the SNSM as well as the associated Jurassic plutonic and volcanic units. In the top-most samples, the Precambrian ages became the dominant population.

The deposition in a Gilbert-type delta system, which traversed the Santa Marta – Bucaramanga fault in the western basin, developed since the late Oligocene. The bottomset segment with an input of mainly granodiorites is related to the Aracataca Batholith, part of the Central Batholith suite. The foresets shows a mixture of Precambrian and Upper Jurassic to Lower Cretaceous ages.

Zircons found at the Neogene sedimentary rocks of the Aracataca and Palomino basins are subrounded to subangular and the presence of unstable heavy minerals indicates a relatively short sediment transport distance from the source area. Rutile presence is also characteristic of high-grade regionally metamorphosed terrains (Force, 1980), which is supported by the fact that conglomerates are highly immature and contain at least ten different litho-types that include high-grade metamorphic and ultramafic rocks.

The provenance information derived from the zircon fission-track maximum relative ages suggests that the main sedimentary source area was the core of the SNSM massif. ZFT and AFT ages are younger in the Aracataca basin in comparison to the Palomino basin, suggesting that the sediments of the western limb of the SNSM were derived from a more deeply exhumed lower crustal source (Fig. 11).

Provenance on the Aracataca and Palomino basins shows that during the Oligocene - Miocene, sediments were sourced from the SNP, and evidence that the Santa Marta Batholith, and Buritaca pluton, although exhumed during the Miocene did not contributed as a sources for the Oligocene-Miocene sediments, probably due to the development of a drainage system with a NE-SW and N-S trend, controlled by the main structures (i.e. Orihueca fault, Aguja Fault, Sevilla Lineament). This would explain

the Miocene exhumation of the Eocene batholiths and the absence of Paleocene-Eocene zircons in both basins.

The Buritaca Pluton yields an early Miocene AFT age for this unit of 22.3 ± 3.1 Ma (Villagómez et al., 2011), and is cut by N-S drainages, but in the sedimentary rocks of Palomino basin, there is no record of Paleogene zircons, therefore is possible to constrain the deposition of the sediments in this basin maximum the middle Miocene (≈ 16 Ma), with a lag time of 5 Myr constrained from an exhumation rate of 0.8km/Myr (Villagómez et al., 2011) and a surface temperature of 30°C. Sediments deposited after the unroofing of the Buritaca pluton containing Eocene zircons are probably bypassed offshore in the late Miocene-Pliocene deposits (Vence, 2008).

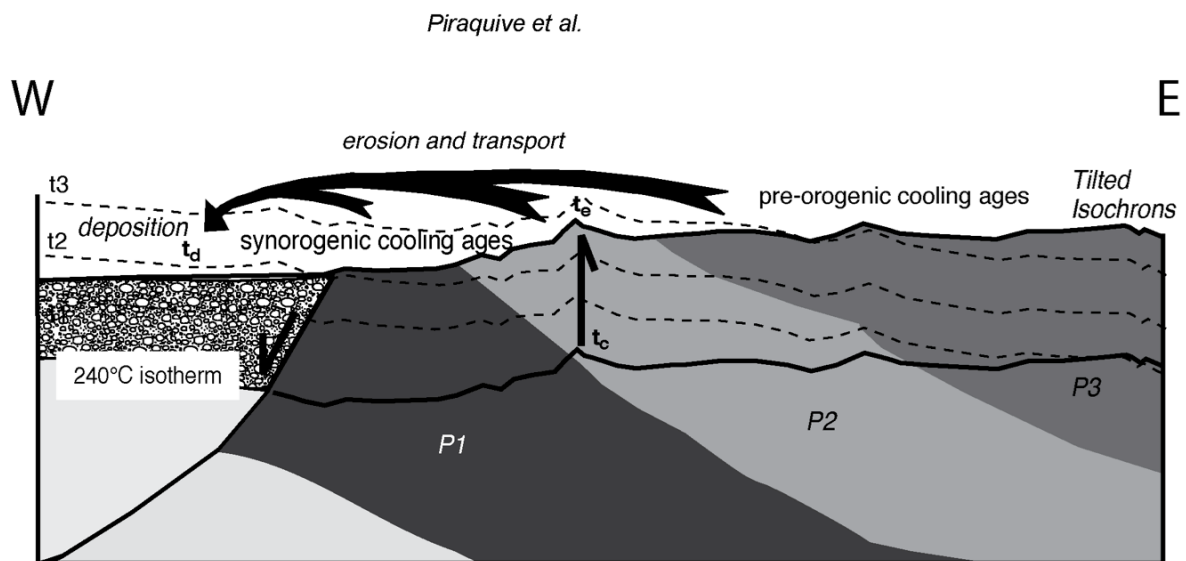


Figure 11. Schematic cross section showing the sources for the ZFT age populations found at the Aracataca Basin, The Santa Marta-Bucaramanga fault activity controls exhumation and isochrone tilting. Modified from (Bernet et al. 2006).

6.2 Regional Stratigraphic Correlations

Lower Miocene sedimentary rocks found either in the Aracataca and the Palomino basins correspond to proximal fan delta systems that developed during the first pulse of surface uplift of the SNSM during the late Oligocene. The clockwise rotation of the SNSM towards the east led to the development of an extensional phase defining the Lower Magdalena basin as a rotational basin linked to the advance of the Caribbean plate (Bayona et al., 2010; Montes et al., 2010). The record of the first tectonic pulse of the Santa Marta-Bucaramanga fault, are the basal deposits of the Aracataca border basin in the proximity of the SNSM massif. At the western edge of the Plato-San Jorge basin, the Romeral suture separated the continental basement of South America from the Caribbean intra-oceanic crust that hosted volcanic arcs and accreted during the Paleocene. The accretion of terranes evolved into I-type plutonism during the Eocene, which gave a characteristically mixed oceanic arc-basement zircon U-Pb age signal in the Paleogene sediments from the San Jacinto belt (Cardona et al., 2012). Towards the SNSM the detrital zircon U-Pb ages became dominated strictly by the continental crust basement signature, with the start of sedimentation in the late Oligocene, sediments in the Aracataca and Palomino basins record a late Oligocene activation of the Santa Marta- Bucaramanga and Oca faults, as the North-Andean block indented against the Caribbean plate. Coarse sediments derived from the SNSM massif, were deposited as growth strata along the active faults. The coarse facies are typical of hyperconcentrate flows, being the product of rapid slope instabilities. During the Miocene, in the Plato-San Jorge basin, the Lower Magdalena basin, and the San Jacinto basin, distal deposits of mudstones, siltstones, and carbonate facies were deposited in a low-energy

environment, coeval with the conglomerates of the Aracataca and Palomino basins (Fig. 12).

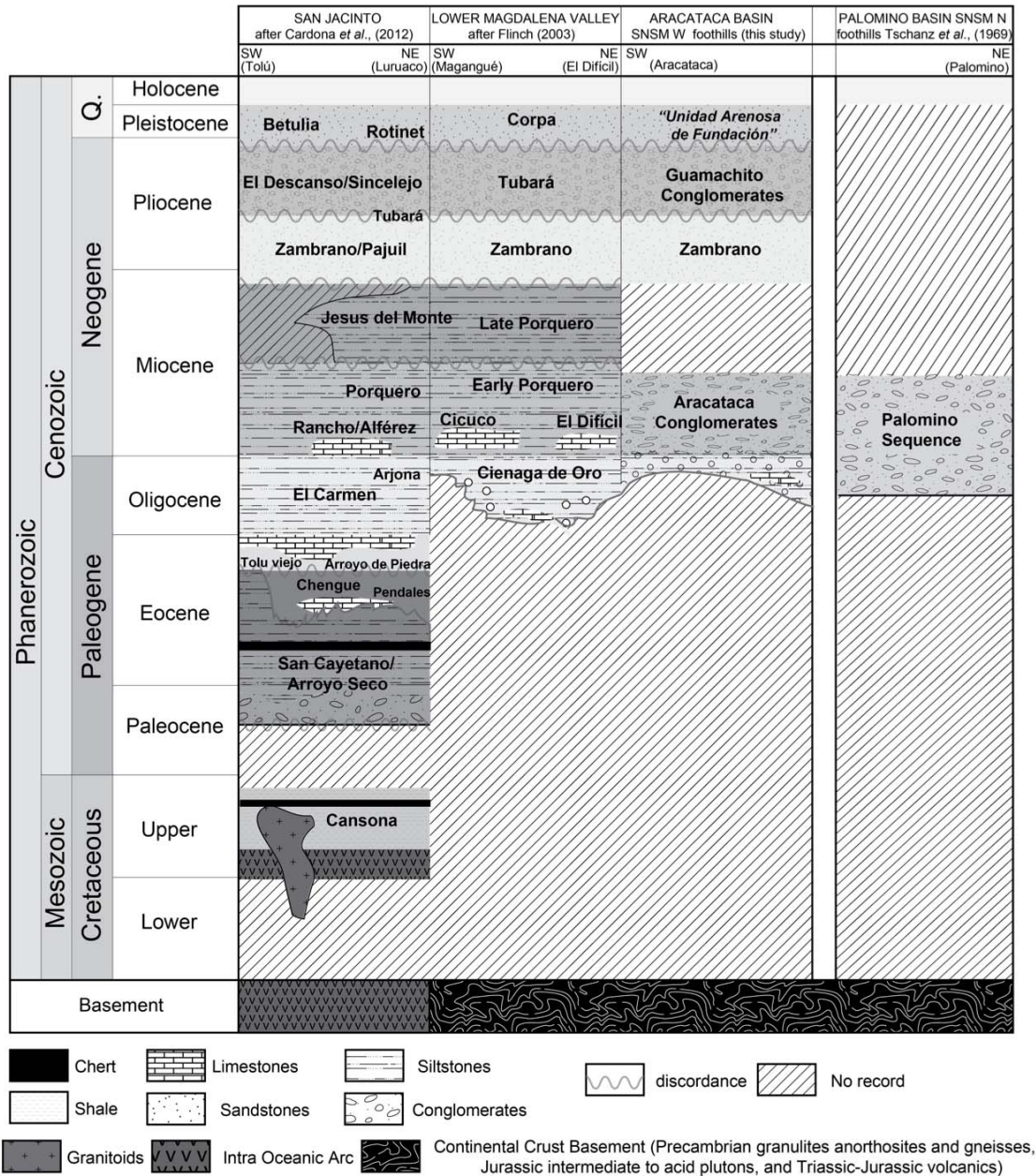


Figure 12. Stratigraphic correlation chart from the Plato-San Jorge basin, the adjacent Lower Magdalena Valley and San Jacinto belt; as well as the Guajira basin in the N foothills of the SNSM. Modified from (Cardona *et al.*, 2012; Duque-Caro, 1979; Flinch, 2003; Tschanz *et al.*, 1969).

The Santa Marta Batholith, the Palomino Pluton, Sevilla, Latal, and Buritaca Stocks (Fig. 2) share Paleocene-Eocene zircon U-Pb ages and are outcropping closely to both the Aracataca, and Palomino basins. Paleogene zircon U-Pb ages are not found within the Neogene sedimentary rocks of Aracataca and Palomino basins, suggesting that by Oligocene-Miocene times, the Paleogene intrusive rocks were not yet exhumed to the surface, and therefore did not act as sources. Although, Paleocene deposits of the Cesar-Ranchería basin (Fig. 2) contain a distinctive population of Paleocene zircons (Bayona et al., 2011), we discard a direct supply of these crystals from the plutons in the NW part of the SNSM, given that thermochronological data of Villagómez et al. (2011) and Duque (2009) evidence exhumation of the Eocene Buritaca stock and Santa Marta Batholith plutons during the Miocene. Accordingly, the Palomino conglomerates lack of corresponding quartzdioritic and tonalitic components of this Paleogene intrusive suite. The sediments from the Paleocene Cerrejón and the Eocene Tabaco Formation in the CRB are derived from Precambrian and Permo-Triassic source rocks (290-220 Ma) with a younger Cretaceous-Paleocene zircon U-Pb age signature (90-65Ma), nevertheless they do not necessarily support exhumation of the NW corner of the SNSM during the Paleocene (Bayona et al., 2011; Montes et al., 2005), neither erosion of the plutons of the Paleogene arc in the NW corner of the Massif at the Santa Marta Metamorphic belts (Cardona et al., 2011), instead thermochronological AFT ages record Paleocene-Eocene exhumation at the SE flank of the SNSM in the Sierra Nevada Province, adjacent to the Cesar-Ranchería basin (Villagómez et al., 2011). In consequence, the older exhumation of the SE flank of the SNSM would have created topography that may have acted both as a barrier for sediment derived from the NW part of the SNSM, and as a source for the CRB.

Zircon U-Pb ages from the lower Eocene Misoa Formation, and Cerrejón and Tabaco Formations from the Cesar-Ranchería basin, indeed overlap with zircon U-Pb ages of plutons in the Santa Marta Metamorphic Belts at the NW corner of the massif, but erosion of these source rocks during the Paleocene seems unlikely of the Miocene cooling histories documented for these igneous rocks (Villagómez et al., 2011). In this context the Late Cretaceous-Paleogene zircon U-Pb populations in the Cesar-Ranchería basin (56-70 Ma) are most likely related to a volcanic source found further to the east, as evidenced by tuffs within the Misoa Formation of the Maracaibo basin, coeval with the Eocene Tabaco Formation (Bayona et al., 2011). Paleocurrent directions measured in the Cerrejón Formation show an E to SE trend (Bayona et al., 2011), given that the SNSM has rotated clockwise by about 30° after the Eocene according to Montes et al. (2010), paleocurrent directions result in E to NE flow directions for the Cerrejón Formation, which in turn are coherent with an E-NE sediment dispersal direction during the early-mid Paleocene 65-58 Ma as documented by Ayala et al. (2012). In this epoch the Precambrian, Permian-Triassic and Cretaceous zircons found at the Cerrejón and Tabaco Formations from the Cesar-Ranchería basin may not be exclusively derived from the SNSM and could also be supplied by the rocks of the Proto-Perijá Range that yield as well Permian to Jurassic granitoids (Martin, 1968), and the Central Cordillera (Mora-Bohórquez et al., 2017), regarding that at Paleocene times the SNSM, Perijá ranges and Central Cordillera were exhumed (Ayala et al., 2012; Shagam et al., 1984; Villagómez et al., 2011; Villagómez and Spikings, 2013), with the major exhumation of the Perijá Range during the Oligocene (Shagam et al., 1984). Other sources probably supplied Paleocene zircons to the Cesar-Ranchería basin. Such sources can be related to Paleocene volcanism, and plutons of the Guajira Peninsula (Cardona et al., 2014), the

Central Cordillera and the Eastern Cordillera. Alternatively a plutonic source of Paleocene zircons found at the Cerrejón and Tabaco formations could be related to unmapped and undated Paleocene plutons (i.e. Atánquez Laccolith, Tschanz et al., 1969) located within the Sierra Nevada Province. This alternative would agree with Paleocene-Eocene exhumation of this province and would correlate with the unstable metamorphic minerals found at the Tabaco and Cerrejón formations indicative of short sediment transport (Bayona et al., 2007). Adjacent to the Cesar-Ranchería basin the onset of a NE directed drainage system occurred during the Paleogene incising the newly formed topography and precluding a connection with the drainage systems, in the NW of the SNSM. The interpretation of a disconnected NW-SE drainage of the SNSM is validated by the zircon U-Pb ages found in the Sinú San Jacinto Belt (Fig. 13), in which Precambrian and Permian-Triassic source rocks are dominant along with the youngest arc derived material from the Late Cretaceous, 71.3 ± 1.3 Ma and 68.7 ± 1.4 Ma (Cardona et al., 2012), without Paleocene-Eocene zircons record in the sediments. Instead of a Paleocene tilting towards the SE which would have generated an inclined topography towards the Cesar-Ranchería basin and an extensive NW-SE drainage system feeding a clastic wedge, the distribution of U-Pb and detrital thermochronologic ages in Oligocene-Miocene sedimentary rocks in the Aracataca and Palomino basins indicates that the SNSM was tilted during Oligocene-Miocene exhumation towards the NE, with an increasing relief in the western limb of the SNSM controlled by the Santa Marta Bucaramanga fault, whose activation was contemporaneous with the dextral Oca fault and the onset of transpressive tectonics. The Oligocene-Miocene progradational deposits of the Aracataca conglomerates coeval with the transgressive deposits of the Palomino conglomerates during exhumation can

be explained by tilting of a block bounded by normal faults (Leeder and Gawthorpe, 1987). For the NE tilting, we presume a fulcrum that separates elevated from subsided segments of a ramp. In this block, two basins are opposed by a drainage divide correspondent to the uplifted crest of the domino-block (Fig. 13): a fault scarp prefiguring a steep slope (drainage basin of Aracataca conglomerates) and ramp of a gentle inclination (Palomino sub-basin). The latter is not fault-controlled: This mechanism provides a simple kinematic explanation for a contemporaneous uplift and subsidence; only valid for a thick crustal block. This block can be delimited by the Sevilla Lineament and involves the Sierra Nevada Province.

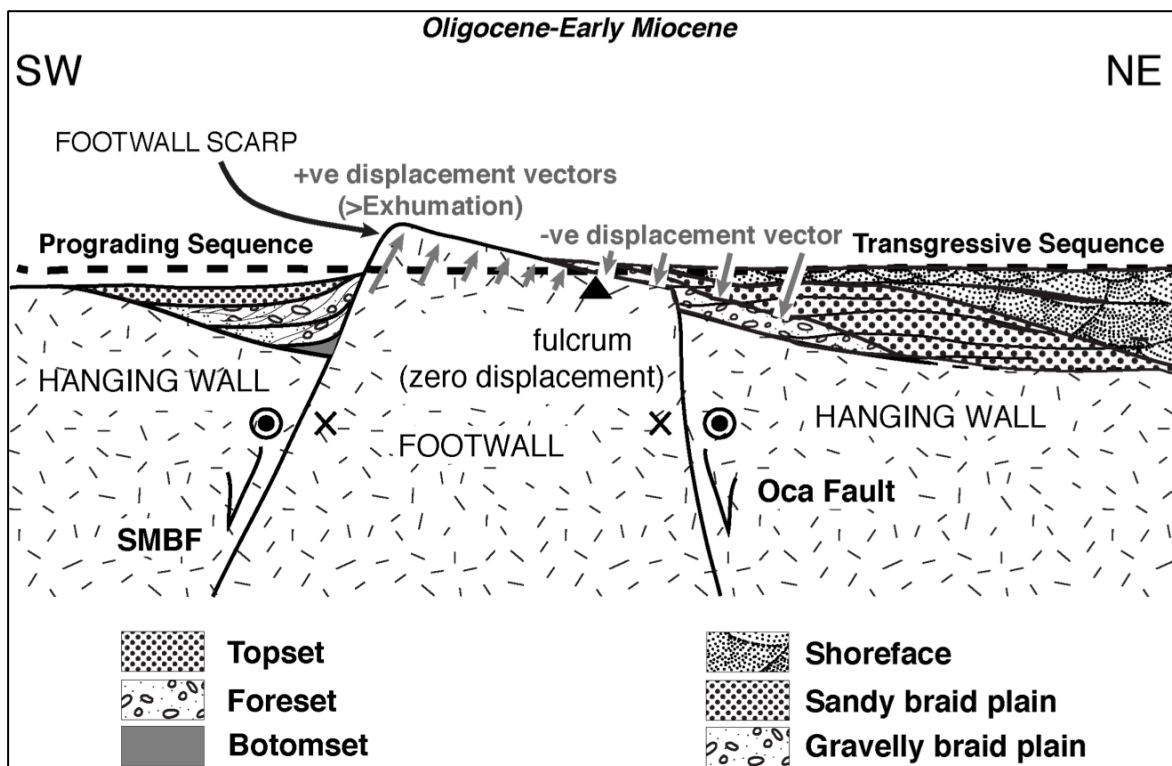


Figure. 13 Conceptual model not to scale, depicting tectonic slopes associated with a simple tilt block/half-graben, involving a fault-bounded syn-extensional clastic wedge, as the case of the Aracataca basin gravel rich fan-delta, at the Santa Marta-Bucaramanga fault, and its counterpart deposited on a gently inclined slope dominated by estuarine deposits of the Palomino basin, modified after (Henstra et al., 2016; Leeder and Gawthorpe, 1987; Ravnås and Steel, 1998).

6.3 Tectonic Implications for the Caribbean Realm

Transition in the Caribbean plate subduction angle from 30° to a shallower dipping subduction between 4°-8° occurred between the late Eocene and middle Miocene (Bernal-Olaya et al., 2015). Since the late Eocene a relevant change in the subduction regime is evidenced by a gradual cessation of arc-plutonism and late cooling, which during the Oligocene-Miocene is defined by a gently inclined slab dipping less than 30°. The establishment of a shallower dipping slab 4°-8° since the middle Miocene is related to the subsequent underthrusting of a thickened Caribbean crust under NW South America responsible for the Santa Marta-Bucaramanga fault reactivation (Kammer and Sánchez, 2006; Villagómez et al., 2011) and accumulation of left lateral slip, as the Caribbean plate moved to the East during the late Miocene. In response to the rigid coupling of oceanic and continental lithospheres, the Santa Marta-Bucaramanga fault slip was prevented and in the interior of the SNSM massif, Early Mesozoic sutures became reactivated as NW verging thrusts (Villagómez et al., 2011) (Fig. 14).

The sedimentary rocks of the SNSM basins attest basement exhumation, which was controlled mainly by block tectonics at the NW corner of the SNSM massif, driven by the normal-sinistral slip of the Santa Marta - Bucaramanga fault (Villagómez et al., 2011).

Zircon provenance evidences that the fan-delta systems and estuaries sourced by the Sierra Nevada de Santa Marta massif remained isolated from the main trunk system of the proto- Magdalena paleochannel (Fig. 15). The proto-delta was controlled by the

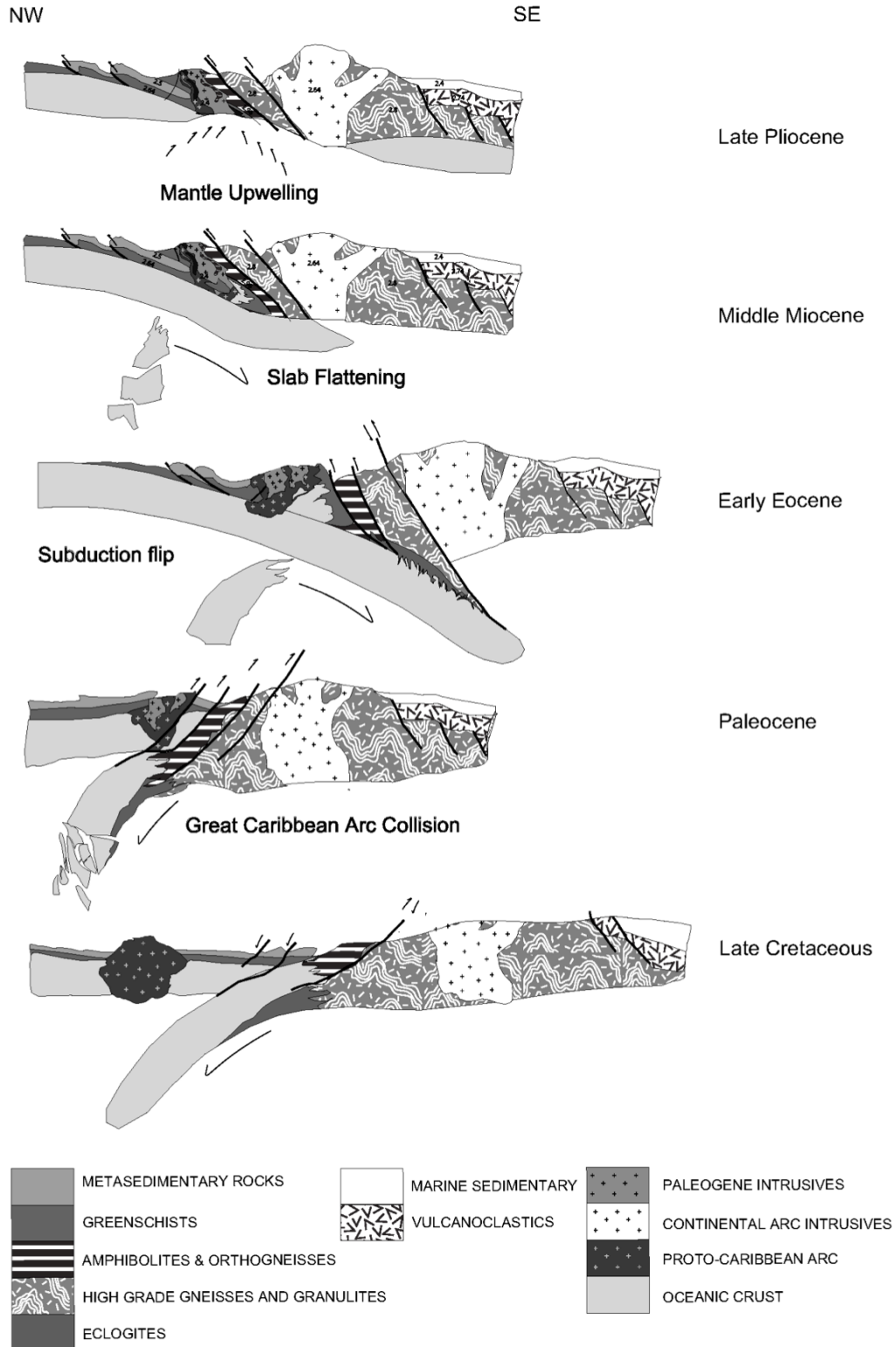


Figure 14. Schematic multi-stage evolution for the Caribbean-South American Margin, in the sense of cross section A-A' Fig 3. During first stages Proto-Caribbean arc approaches to South America; in the late Eocene slab flattening is responsible for the reactivation of crustal structures as the Santa Marta-Bucaramanga Fault.

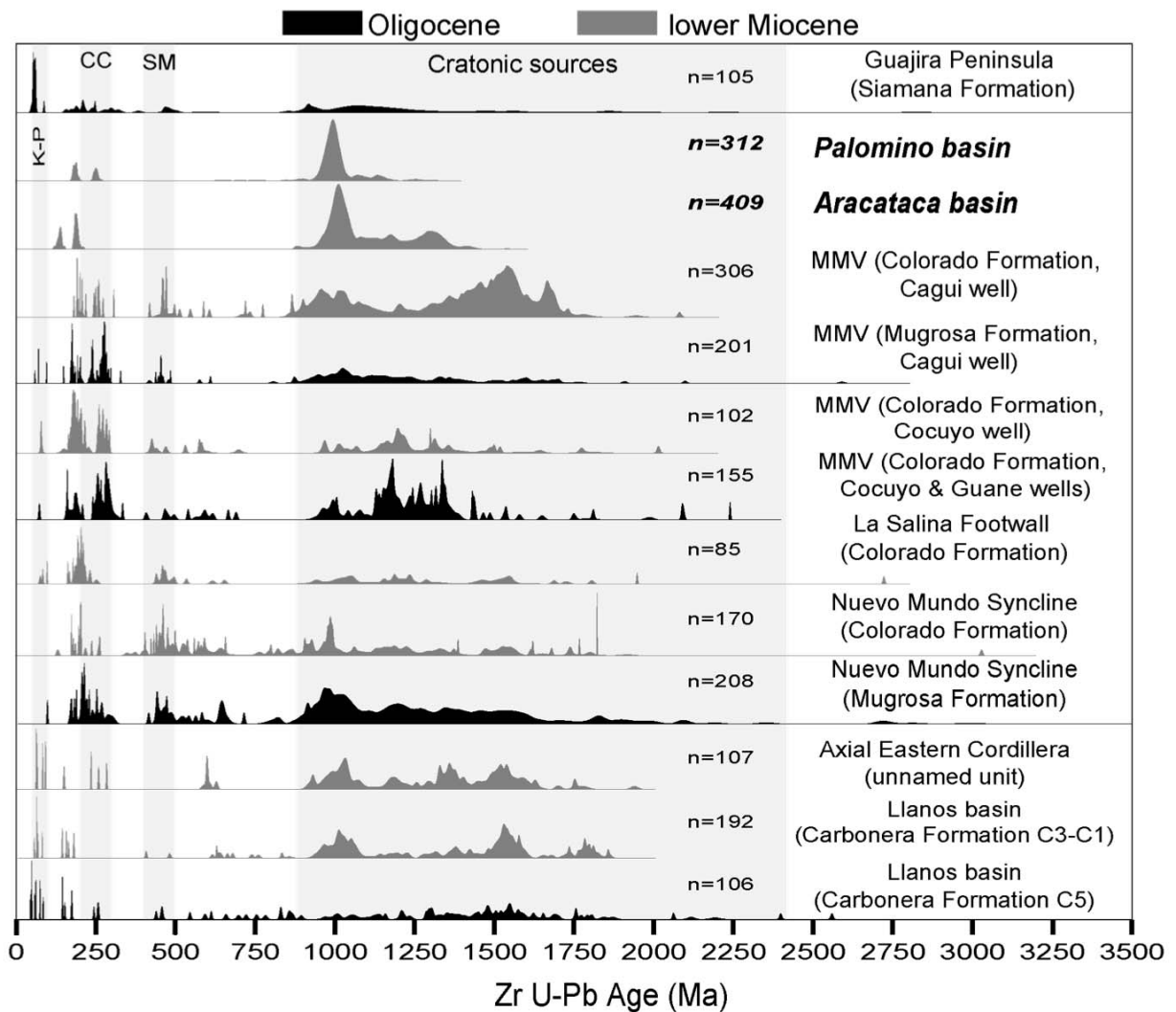


Figure 15. Comparative plots of detrital zircon U-Pb age spectra for Oligocene - lower Miocene sedimentary rocks from the Guajira Peninsula (Zapata et al., 2010), the Middle Magdalena Valley (Caballero et al., 2013; Horton et al., 2015), the axial Eastern Cordillera and Llanos basin (Horton et al., 2010b, 2010a), and the Aracataca and Palomino basins from this study. Diagnostic age populations are delimited by grey shaded zones, SM: Santander Massif, CC: Central Cordillera, K-P: Cretaceous-Paleogene volcanic-plutonic arc.

Santa Marta Bucaramanga fault activity at the southern boundary of the Lower Magdalena basin (Bernal-Olaya et al., 2015). These new findings improve greatly paleogeographic reconstructions by this epoch and allow to precise timing between

fault activity and the onset of sedimentation in the Lower Magdalena basin and the sourcing of sediments from the Eastern Cordillera, the Central Cordillera, the Santander Massif, the Perijá Range, and the SNSM (Ayala et al., 2012; Caballero et al., 2013; Horton et al., 201b, 2015; Zapata et al., 2010).

The Miocene activity of the Orihueca fault marks a tectonic pulse driven by a significant variation in convergence rate between the Caribbean and South-American plates, and the shift from collision to transpression produced the increased NW verging thrusting in the interior of the SNSM (Fig. 14), culminating in the displacement shift from normal to dextral-slip of the Oca fault, that segmented the Guajira basin from the North-Andean margin in the Pliocene and displaced it to its current location about 55 km to the east (Kellogg, 1984; Tschanz et al., 1974).

The sediments of the SNSM basins attest basement exhumation, which was controlled mainly by block tectonics at the NW corner of the SNSM massif, driven by the normal-sinistral slip of the Santa Marta - Bucaramanga fault (Villagómez et al., 2011). The Miocene activity of the Orihueca fault marks the onset of a tectonic pulse driven by a significant variation in convergence rate between the Caribbean and South-American plates, and the shift from collision to transpression that produced the increased NW verging thrusting in the interior of the SNSM, culminating in the displacement shift from normal to dextral-slip of the Oca fault.

Other evidence of this episode can be observed in the Leeward Antilles (van der Lelij et al., 2010; Zapata et al., 2014) where oblique displacements occurred diachronously, and also in the Venezuelan Andes when transpression was accommodated by vertical tectonics and caused a rapid exhumation through the Boconó fault in the Pliocene (Bermúdez et al., 2010).

7. CONCLUSIONS

1) Post-Oligocene sedimentary rocks found in the Aracataca basin show a major input from Precambrian source rocks from the SNP. This major input is the consequence of the unroofing of the basement that was accelerated since the deposition of the Aracataca conglomerates foresets. The shift in composition of the clasts means that the Aracataca conglomerates are growth strata that evidence both extension tectonics along the Santa Marta - Bucaramanga fault since the Upper Oligocene and exhumation of the massif through the Orihueca fault that occurred at least since the Lower Miocene supported by AFT and ZFT cooling ages.

2) The abundance of the Precambrian source is related to an exhumation due to extensional tectonics at the Santa Marta - Bucaramanga fault and a subsequent activation of NW verging thrusts at the interior of the SNSM.

3) The Precambrian sources are derived from inherited Grenvillian terranes that passed through the apatite partial annealing zone approximately since 25-20 Ma, and were lately exhumed and affected by surface erosion.

4) The sudden change from muddy oligomictic conglomerates and siltstones to a coarser facies in the sediments attests to a steepening slope in the source areas, which led to debris flow into Gilbert-type delta environment during the transition from a shallow marine to a more fluvial dominated environment. The facies change is as well related to an increase in erosion and therefore probably also in exhumation rates.

5) Provenance analysis of the basin sediments clearly shows that the clastic material was derived by erosion of the SNSM Precambrian inliers, and that Permo-Triassic, and Late-Cretaceous accreted terrains at the NW corner of the SNSM did not contributed to

sediments, as these tectonic blocks were not yet exhumed during Oligocene-Miocene times.

6) The major shift in provenance between the bottomset and foresets of the Aracataca conglomerates is related to exhumation of the massif during increasing displacement along the Santa Marta – Bucaramanga fault, which occurred between the late Oligocene and early Miocene. The increase in the variety of source rock lithologies registered in the stratigraphic record is evidence for unroofing of the Precambrian basement and the Central Batholith.

7) Detrital thermochronologic ages from the Aracataca basin are evidence for the Santa Marta-Bucaramanga fault being responsible for the increase of relief in the western flank of the SNSM massif, as it acted coeval with the NW verging Orihueca fault and Oca fault. In the Palomino basin extensional faulting started since the Late Eocene-Oligocene with the onset of a dextral strike-slip since the Miocene in consequence of the coupling of the South American crust and the Caribbean crust. This caused faster exhumation of the western flank of the SNSM massif, exhuming lower crustal levels that acted as sources for the Miocene sequences that yield the oldest zircon U-Pb ages and younger detrital ZFT and AFT ages.

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OVERALL CONCLUSIONS

At The Sierra Nevada de Santa Marta massif I had found evidences that this part of the Andean margin has acted as an accretionary margin throughout its history, and that the rocks that form the SNSM preserve the record of at least three orogenic cycles that involve processes of crustal thickening, orogenic wedge growth, thermal equilibration, crustal weakening and late orogenic collapse due plate kinematic reconfiguration (Fig. 1). In this dissertation, I integrated several geochronological and thermochronological techniques to construct a scenario on P-T-t-d paths during the evolution of the SNSM; during the Phanerozoic.

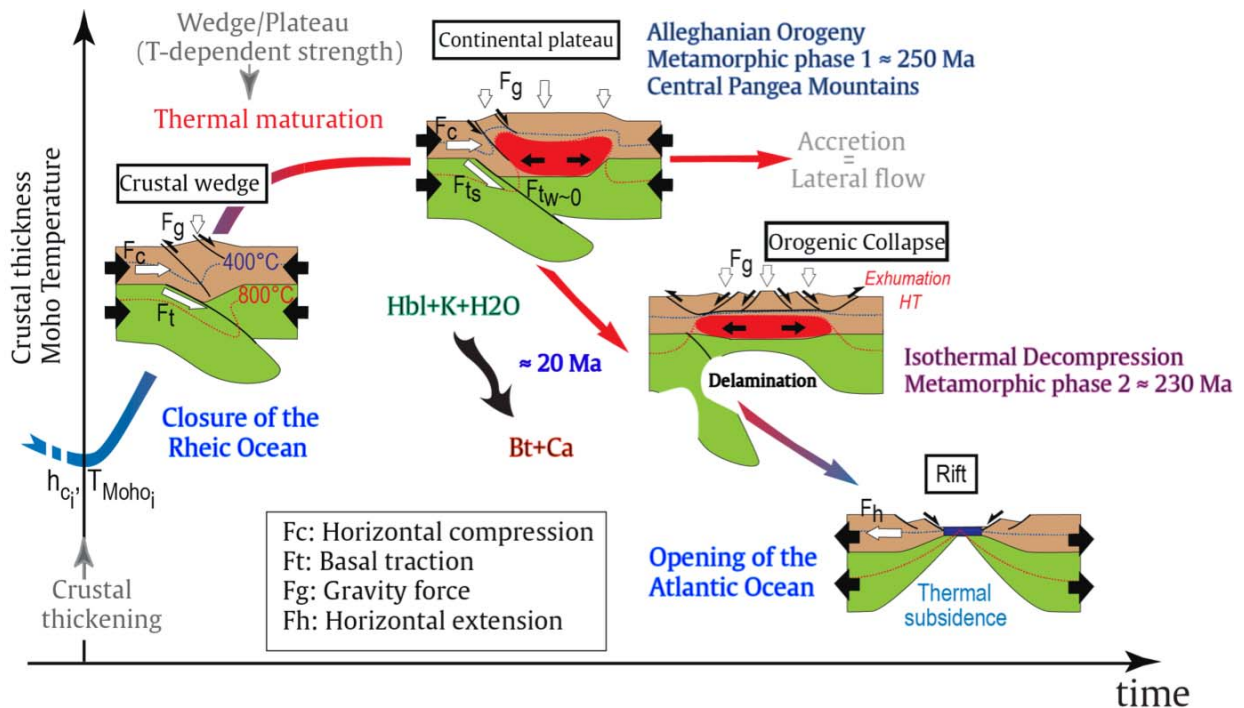


Figure 1. Sketch of a Thermal-Mechanical evolutive scenario of the crust at plate boundaries and the crustal tectonic cycle. Illustrated by the evolution of crustal thickness and Moho temperature as a function of time. The first stage of the thermal-mechanical evolution of the crust at convergent plate boundaries is characterized by the development of a crustal wedge under a low-geothermal gradient dominated by subduction. Thermal maturation of the crustal wedge is responsible for the weakening of the lower crust, which leads to the

wedge/plateau transition. At this stage, the basal traction force decomposes into F_{tw} beneath the plateau and F_{ts} along its edges. A steady state is reached if accretion is balanced by gravity-driven lateral flow of the weakened crust. A modification of the lithospheric-scale dynamics boundary condition causes the switch from crustal thickening to thinning and the mode of crustal extension ranges from metamorphic core complex to rift as a function of the rheologic state of the crust. Crustal thinning is associated with isothermal decompression and is followed by thermal subsidence. Taken from (Vanderhaeghe, 2012).

My interpretation is that at the Sierra Nevada de Santa Marta this evolutionary trend occurred during three orogenic cycles, at the end of Neoproterozoic, In the middle Permian - Upper Triassic, and during the Late Cretaceous - Miocene. During these interactions, lower crustal delamination and slab retreat played a major role influencing the modification of lithospheric-scale dynamics boundary conditions i.e. Oligocene plate reorganization, caused a transition from lower crustal extrusion to extension through reactivated faults.

In Chapter 1, I presented a tectonic history of the SNSM massif that includes its evolution during the supercontinental cycle of Pangaea amalgamation and disassembly.

A secondary topic of interest discussed briefly in Chapter 1, derived from our results is related to the geochronological implications of previous interactions of the continental margin related to the opening of the Iapetus Ocean during Rodinia disassembly and the subsequent Brazilian/Pan-African orogeny. Although the new evidences found in this study at the high metamorphic grade units of the Sevilla Metamorphic Belt suggest a palaeogeographic context for the Iapetus opening at this part of the Proto-Andean margin. It, is necessary to continue with deeper and extended analyses on this particular topic since most of the mapping and sample collection on the late Neoproterozoic to early Paleozoic units at the SNSM remains to be done (Fig. 2).

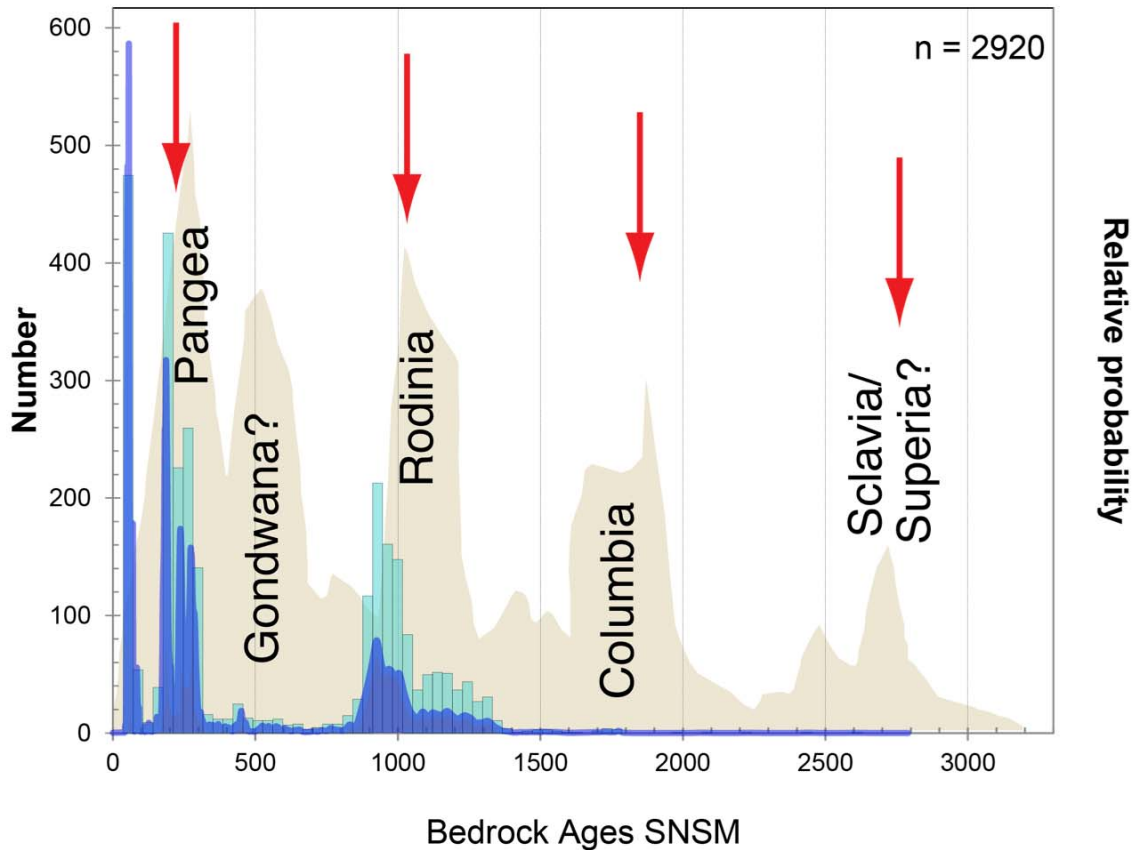


Figure 2. Comparison of all U-Pb data previously published and from this work of bedrock samples from the SNSM massif represented by the blue curve and bins, the red arrows indicate the orogenic peaks recognized by Runcorn, (1962) brown shade in the background corresponds to the spectra of U-Pb ages compiled by (Hawkesworth et al., 2010). Although minor, a Neoproterozoic-Paleozoic population is present and may indicate sampling bias or either orogenic quiescence. There are two major currents on the interpretation of age peaks: one is related with crustal growth driven by mantle convection and the second to preservation of the continental crust due supercontinent assembly, within this discussion U-Pb peaks may correspond to periods of accelerated crustal growth (Arndt, 2013), and detrital zircon populations may act as a measure of the intensity of continental arc magmatism (McKenzie et al., 2016). Modified after (Nance et al., 2014).

My results in Chapter 1 redefine the morphotectonic provinces of the SNSM (Tschanz et al., 1974) based in geochronology and petrology. This new considerations have a high impact in the palaeogeographic models discussed for Pangæa assembly (Fig. 3) in

the Proto-Andean margin, and geodynamics of the South American and Pacific plates, in which I had reevaluated the role played by the Pacific plate in the arc inception and the subduction of a juvenile oceanic crust of the Rheic Ocean during Pangæa Assembly. Interaction of Laurentia and Gondwana resulting from a dextral oblique collision between their irregular margins caused the closure of the Rheic Ocean, and led to continental arc magmatism that extended from the current Peruvian Andes to the Southern North America Maya block including the Colombian Andes and the Sierra Nevada de Santa Marta Massif.

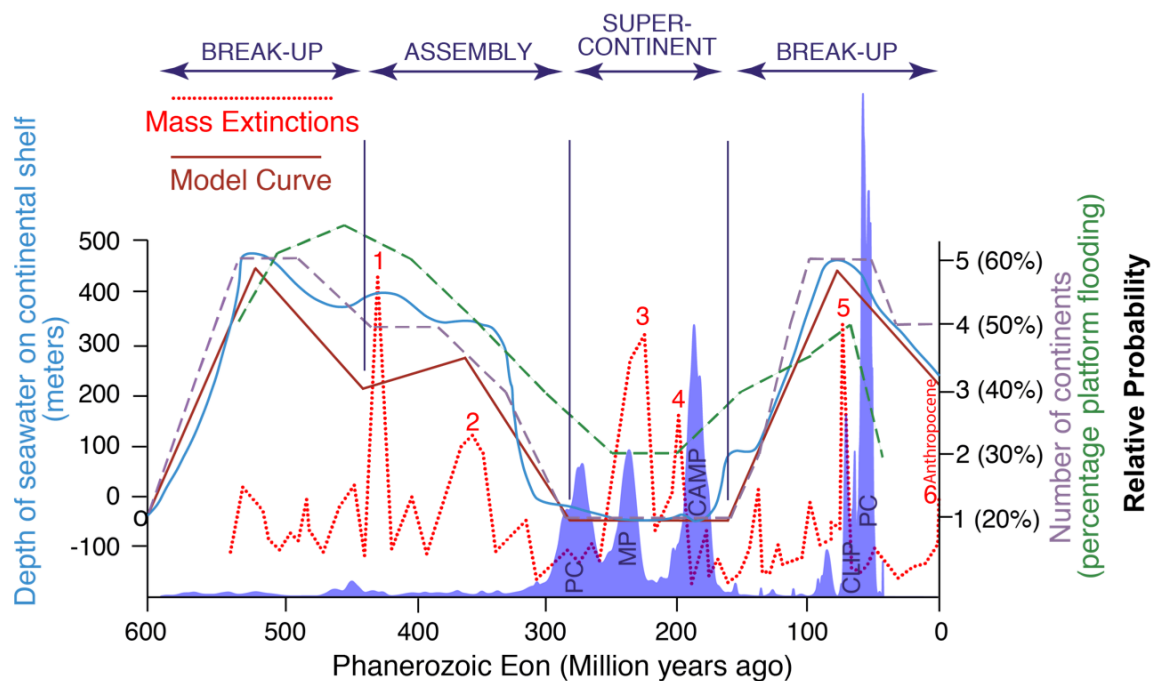


Figure 3. Spectra of U-Pb data from the SNSM massif illustrates periods of crustal growth associated to tectonic events PC: Post collisional I-type magmatism, MP: Metamorphic peak, CAMP: Central Atlantic Magmatic Province, CLIP: Caribbean Large Igneous Province. Blue curve represents variation in sea level, purple dashed line represents the number of continents and green dashed curve depicts amount of flooding of shallow shelves by seawater for the last 600 Ma. Stages of supercontinent are indicated in this case for Pangæa. Modified after (Condie, 2011; Hawkesworth et al., 2010; Nance et al., 2014). Massive extinction curve in red reproduced from the University of California Museum of Paleontology's Understanding Evolution (<http://evolution.berkeley.edu>).

The youngest arc magmatism of this episode is early Permian 290-280 Ma in western Pangæa and is associated with collision tectonics (Weber et al., 2007), and found further to the south of the Andes interpreted as being the result of lithospheric thinning (Mišković et al., 2009; Sempere et al., 2002). In East Pangæa Permian magmatism is associated to a Post-Variscan extensional regime (Beltrando et al., 2010; Beltrán-Triviño et al., 2016; Franke, 2006; Friedl et al., 2004; Letsch et al., 2015; Nance et al., 2012, 2010). This diachronous closure was controlled by a north oblique convergence in Europe in the early Carboniferous as Gondwana collided with southern Baltica, and a late Carboniferous-early Permian dextral transpressive convergence between Gondwana and Laurentia that originated the Appalachian-Ouachitan orogen. In Western Pangæa approximately 30 Myr after the ca. 280 Ma magmatism, peak conditions of Barrovian metamorphism were reached along the assembled margin (ca. 250 Ma), 10 Myr later orogenic collapse of the Appalachian-Ouachitan orogen led to crustal thinning and isothermal decompression during the Triassic. This was the beginning of Western Pangæa disassembly and is recorded at this part of the Proto-Andean margin in zoned garnet and metamorphic zircon and in-situ rim overgrowths. We also point out the relevance of the late Paleozoic tectonics during Pangæa amalgamation and the regional correlations of the Permo-Triassic terranes in the Colombian Andes, with the Inner Santa Marta Metamorphic belt were rocks of this epoch have been recognized and studied by geochronology and thermobarometry. It had revealed the metamorphic conditions (P-T) during a time span of 60 Ma that shifted between two phases, that constrain Pangæa amalgamation in a first collisional phase with a prograde path and a second phase of isothermal decompression. It attests

for crustal thinning and orogenic collapse prior Pangæa disassembly by rifting and subsequent drifting during the Late Triassic ca. 216 Ma.

Finally, the correlation between the depletion of trace element concentrations in zircons and coeval crystallization of garnet provides a reliable estimate of the age of metamorphism that rather is defined as a time span.

Chapter 2 explored the geodynamic constraints that operated during the Late Cretaceous Caribbean Plate interaction against the South American plate. In this scenario, we had discussed the mechanisms for terrane accretion, collisional plutonism and high-elevated geothermal gradients and the later events that correspond to major plate reorganization during the Oligocene in Chapter 3.

In Chapter 2, we animate this discussion in the context of the Caribbean plate origin, which is debated, between a pacific realm or a segmentation and incorporation between Cocos and Nazca plates, against cratonic South America. The main boundary of this scene can be reconstructed up to Late Cretaceous ages, when is widely recognized, an active subduction process related to Thethys closure, is evidenced from the marine sequences with a important basaltic input acting until early Paleocene. Such marine deposits and its arc related deep basaltic sequences are most known as the CLIP.

The thermal overprint produced by the CLIP accretion annealed the fission track systems preserving exclusively the last orogenic event since the Eocene.

Models for Caribbean plate evolution involve the development of this proto Caribbean oceanic floor and its accretion against a continental South America. The main interest in Chapter 2 resides in the tectonic relationship between units that belong originally to the domain of the Caribbean plate (Santa Marta schists) and the continental basement

of the Inner Santa Marta Metamorphic Belt. For the contact between these two provinces, a suture has been discovered in the field and confirmed by geochronological data. Our results showed that this inherited discontinuities favor renewed thrusting and shear as the case of the Orihueca fault and the Sevilla lineament, and even can conduct crustal channeling exhuming the lower crust.

A major breakthrough of this work is the complete redefinition of the extent of the Caribbean accreted oceanic crust at the SNSM massif, which is confined only to the NW corner of the massif, as proven by geochronological data and field observations, metamorphic degree, deformations, and paragenesis. The metapelitic and metavolcanic sequences formerly associated to the Caribbean arc are underlain by ca. 280 Ma orthogneisses, implying a clear affinity with the continental margin in contradiction with the widely accepted Caribbean origin for these units. This new concept has major consequences in the magnitude of the subducted Caribbean slab; allowing understanding this process at the continental margin of the SNSM as a collision governed tectonics.

Collision and subduction failure are highly influenced by the consumption of a thickened oceanic crust composed by the plateau basaltic rocks and a thick sedimentary sequence of at least 6000 m, which is to be found at the Guajira Peninsula and the Western Cordillera of Colombia, but that was eroded by subduction at the portion of the continental margin that corresponds to the SNSM massif.

Collision favored a high temperature regime for the SNSM during the Cenozoic, and lower crustal weakening as well as increase on relief and hence accelerated exhumation rates at least until thermal maturity and crustal thickening.

It is likely that the SNSM had experienced these processes during its orogenic history with periods of crustal thickening and orogenic building followed by orogenic collapse and extension, and crustal delamination, as documented in this work this may have occurred during Pangaea amalgamation and during CLIP accretion.

In the most recent part of its history described at Chapter 3, strong plate coupling reactivated mid-crustal structures that controlled exhumation during the transient stages of the Caribbean plate towards the East. The Santa Marta Bucaramanga fault, a major fault inherited from Jurassic rift, suffered a normal sinistral strike slip activation due to the establishment of a partitioned convergence between the Caribbean Plate and the Nor-Andean block. This process created enough subsidence for accommodating sediments and redistributed strain onto the continental margin causing the uplift and late exhumation of the Eastern Cordillera, the Santander Massif and the Merida Andes. The perspectives for a future work on the SNSM involve a major expedition for acquiring samples from several elusive areas as the peaks of the glacial zones that may hold keys for the Neoproterozoic-Paleozoic history of the massif. It is clear that most of the research has been done in the vicinity to the most populated areas, and although it seems that the remnants of an Ordovician arc present further to the SE in the Merida Andes, Santander Massif, and Central Cordillera (Anaconda terrane) is absent in the SNSM. This assumption has yet to be demonstrated, the relevance of these investigations is related with the regional correlations of the proto-Andean margin. In this same workflow it is fundamental to construct a more precise time frame for the metamorphic episodes, for which is necessary to perform geochronological dating in minerals as garnet, monazite, and mica using techniques as Lu-Hf and oxygen isotopes in mica by TIMS.

The author recommends as well investigating the active sediments integrating thermo-geochronology, further research needs a detailed hydrographic analysis and sampling on the river catchments for studying the erosive processes acting during the Neogene, this research may be aided by cosmogenic nuclide dating.

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I had worked with Professor Andreas Kammer during my bachelor and master’s at the Eastern Cordillera of Colombia, and suddenly one day we were redirecting our efforts to a new adventure: the Sierra Nevada de Santa Marta massif, the pioneers in this endeavor were Prof. Jean Francois Stephan director from the Institut national des Sciences de l’Universe (INSU - CNRS) and Dr Thomas Maurin from Geoazur at the Université Sophia Antípolis, together we enjoyed excellent journeys on field and shared good times in Santa Marta either at the Aracataca river, or the Tayrona park,

this first field trip ultimately induced me to start my PhD. These first explorations revealed an incredibly complex geology for us in that time and from the beginning, I felt that I seriously wanted to pursue a research project in this area.

I want to thank Prof. Kammer for always guiding me along all these years, thanks for all the adventures and experiences, in the jungle, deserts and glaciers in all these places where we took a look at the rocks, sometimes with great efforts, but every time with great rewards!!

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the material we gathered constitutes an archive for future geologists, and all of us contributed to build a new Geological map.

I want to dedicate this work to my friends Sebastián and Tomás, your departure in 2016 was a very difficult moment, and it grieves me so to write this particular line, but you must know that you belong in my memories, and I'll remember you.

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APPENDIX CHAPTER 1

TABLE A1. U, Pb AND Th ISOTOPIC DATA OBTAINED USING THE LA-ICP-MS METHOD

I.D.	Isotopic Ratios				Isotopic Ages				Parts per million (ppm)	
	$^{206}\text{Pb}/^{238}\text{U}$	$\pm 2\sigma$	$^{207}\text{Pb}/^{235}\text{U}$	$\pm 2\sigma$	$^{206}\text{Pb}/^{238}\text{U}$	$\pm 2\sigma$	$^{207}\text{Pb}/^{235}\text{U}$	$\pm 2\sigma$	$^{232}\text{Th}/^{238}\text{U}$	$\pm 2\sigma$
AP-01	1.433	0.036	0.1389	0.023	0.0756	0.0019	0.0416	0.0041	902	15
AP-02	1.791	0.038	0.1762	0.003	0.0738	0.0011	0.0475	0.0038	1042	14
AP-03	0.0461	0.005	0.0061	0.001	0.0061	0.0001	0.0024	0.0009	514	4
AP-04	0.201	0.034	0.0274	0.001	0.0274	0.0002	0.0162	0.0012	186	5
AP-05	1.611	0.037	0.1659	0.003	0.0705	0.0029	0.0479	0.0047	972	23
AP-06	1.718	0.043	0.1733	0.007	0.0733	0.0043	0.0488	0.0048	1024	18
AP-07	0.0521	0.007	0.0058	0.001	0.0058	0.0002	0.0037	0.0007	515	2
AP-08	1.557	0.034	0.1557	0.004	0.0704	0.0024	0.0464	0.0024	903	16
AP-09	0.0522	0.007	0.00787	0.0005	0.00787	0.0002	0.0026	0.0002	516	2
AP-10	0.0521	0.0025	0.00787	0.0002	0.00787	0.0001	0.0024	0.0002	515	2
AP-11	0.0528	0.0028	0.00818	0.0004	0.00818	0.0001	0.0025	0.0001	512	2
AP-12	0.0545	0.0039	0.00915	0.0006	0.00915	0.0003	0.0038	0.0003	518	3
AP-13	1.2	0.051	0.1251	0.009	0.0633	0.0012	0.0376	0.0047	1182.2	13
AP-14	0.0521	0.0039	0.0083	0.0002	0.0083	0.0001	0.0038	0.0004	515	4
AP-15	1.352	0.058	0.1384	0.029	0.077	0.0025	0.0492	0.0025	886	25
AP-16	0.0539	0.007	0.0069	0.0008	0.0069	0.0004	0.0024	0.0008	513	6.8
AP-17	0.057	0.0052	0.00663	0.0004	0.00663	0.0002	0.0026	0.0002	509	4.8
AP-18	1.851	0.059	0.1798	0.009	0.0751	0.0021	0.0236	0.0025	1065	21
AP-19	0.1978	0.0071	0.03989	0.0007	0.03989	0.0001	0.015	0.0011	183.3	6
AP-20	1.514	0.025	0.1561	0.007	0.07063	0.0006	0.0453	0.0008	936	18
AP-21	2.472	0.061	0.2152	0.003	0.08139	0.0013	0.062	0.0053	1263	18
AP-22	0.0524	0.0048	0.00789	0.0008	0.00789	0.0007	0.00272	0.00033	518	4.7
AP-23	1.91	0.038	0.1836	0.002	0.0733	0.0013	0.0429	0.0008	1086	12
AP-24	2.057	0.064	0.2275	0.006	0.0838	0.0017	0.0582	0.0018	1115	18
AP-25	0.0501	0.0005	0.00605	0.0002	0.00605	0.0002	0.0024	0.0002	505	5.1
AP-26	0.0571	0.0041	0.00629	0.0006	0.00629	0.0002	0.0025	0.0005	56.3	3.9
AP-27	1.737	0.029	0.1738	0.0029	0.0743	0.0013	0.0443	0.0013	1014.3	11
AP-28	0.0524	0.0034	0.00611	0.0004	0.00611	0.0002	0.00257	0.00043	519	2.9
AP-29	0.0499	0.0049	0.00605	0.0009	0.00605	0.0009	0.0024	0.0002	50.2	4.6
AP-30	1.373	0.038	0.1216	0.003	0.0631	0.001	0.035	0.0012	912.2	13
AP-31	0.0548	0.0058	0.00626	0.0005	0.00626	0.0003	0.0026	0.0002	50.2	4.9
AP-32	1.465	0.059	0.1521	0.004	0.0700	0.001	0.045	0.002	915	12
AP-33	2.86	0.11	0.2495	0.007	0.0879	0.0026	0.0533	0.0026	1189	38
AP-34	0.0521	0.0026	0.00605	0.0002	0.00605	0.0002	0.0024	0.0002	515	2.5
AP-35	0.0571	0.0047	0.00602	0.0002	0.00602	0.0002	0.0024	0.0002	515	2.5
AP-36	0.054	0.0048	0.00616	0.0002	0.00616	0.0002	0.0025	0.0002	50.7	4.6
AP-37	0.0521	0.0021	0.00605	0.0001	0.00605	0.0001	0.0024	0.0001	515	2.5
AP-38	1.55	0.037	0.1518	0.007	0.0745	0.0025	0.045	0.0025	92.7	11.1
AP-39	0.0523	0.0029	0.00629	0.0003	0.00629	0.0003	0.0025	0.0003	50.7	4.6
AP-40	0.0523	0.0037	0.00611	0.0004	0.00611	0.0004	0.0025	0.0004	51.9	3.9
AP-41	0.0523	0.0031	0.00621	0.0003	0.00621	0.0003	0.0025	0.0003	50.7	4.6
AP-42	0.242	0.041	0.20932	0.006	0.0904	0.0015	0.051	0.0015	118.3	12
AP-43	2.207	0.049	0.2021	0.004	0.0792	0.0014	0.0486	0.0017	1183	24
AP-44	0.0594	0.0046	0.00645	0.0003	0.00645	0.0003	0.0028	0.0003	56.8	5.6
AP-45	1.758	0.065	0.1741	0.004	0.072	0.0025	0.047	0.0025	1028	24
AP-46	0.0569	0.0037	0.00609	0.0002	0.00609	0.0002	0.0024	0.0002	50.6	4.9
AP-47	0.0587	0.0022	0.00676	0.0002	0.00676	0.0002	0.00236	0.00025	57.9	2.1
AP-48	0.266	0.034	0.2396	0.001	0.0499	0.0008	0.0312	0.0008	238	27
AP-49	1.599	0.052	0.1599	0.002	0.0726	0.0012	0.0453	0.0012	914	18.2
AP-50	0.121	0.017	0.10181	0.001	0.078	0.001	0.049	0.0005	114	6.3
AP-51	0.235	0.024	0.2024	0.002	0.0665	0.002	0.04	0.002	714	49.6
AP-52	0.0543	0.0028	0.007997	0.0004	0.007997	0.0004	0.00258	0.0002	51.6	2.7
AP-53	0.22	0.021	0.20072	0.0005	0.0517	0.0009	0.039	0.0005	201	18
AP-54	0.0543	0.0047	0.00608	0.0007	0.00608	0.0007	0.00254	0.0002	50.2	4.9
AP-55	0.0591	0.0051	0.00608	0.0002	0.00608	0.0002	0.0024	0.0002	51.6	4.5
AP-56	1.673	0.029	0.1679	0.003	0.0731	0.0011	0.0437	0.0037	998.3	11
AP-57	1.493	0.041	0.1493	0.002	0.0713	0.002	0.0437	0.002	913.7	10
AP-58	0.0551	0.003	0.00606	0.0005	0.00606	0.0005	0.0024	0.0003	54.5	2.9
AP-59	16.23	0.01	0.1628	0.002	0.07137	0.0009	0.045	0.0009	979.1	7.8
AP-60	0.056	0.0031	0.00608	0.0002	0.00608	0.0002	0.0024	0.0002	51.6	4.5
AP-61	1.303	0.042	0.1303	0.004	0.0742	0.0003	0.0463	0.0003	946	21
AP-62	0.0535	0.0039	0.00624	0.0005	0.00624	0.0005	0.0024	0.0004	52.9	1.8
AP-63	1.236	0.021	0.1236	0.005	0.0617	0.0017	0.0376	0.0017	812	16
AP-64	0.051	0.002	0.051	0.001	0.051	0.001	0.026	0.001	50.6	2.7
AP-65	0.1492	0.009	0.1492	0.0009	0.00707	0.0012	0.046	0.0012	141.1	8.3
AP-66	0.0569	0.0022	0.00612	0.0002	0.00612	0.0002	0.0024	0.0002	51.6	2.8
AP-67	1.689	0.039	0.1726	0.004	0.0709	0.003	0.054	0.003	1025	20
AP-68	1.428	0.049	0.1428	0.004	0.0717	0.004	0.052	0.004	1027	19
AP-69	1.527	0.026	0.1563	0.003	0.0711	0.002	0.0468	0.0008	943	9.8
AP-70	0.072	0.008	0.0688	0.0008	0.0644	0.0003	0.0375	0.0003	69.2	7.9
AP-71	0.24	0.031	0.2161	0.003	0.0691	0.0027	0.041	0.0027	186.2	13
AP-72	2.355	0.029	0.2398	0.004	0.0813	0.0034	0.0549	0.0034	1229	12
AP-73	0.132	0.017	0.1017	0.001	0.061	0.001	0.035	0.001	393	37
AP-74	0.0578	0.0039	0.00783	0.0004	0.00783	0.0004	0.0026	0.0004	2.3	2.3
AP-75	0.0569	0.0045	0.00608	0.0005	0.00608	0.0005	0.0024	0.0005	54	5.4
AP-76	1.409	0.033	0.1474	0.003	0.0694	0.0015	0.0461	0.0014	892	14
AP-77	0.0533	0.0023	0.00623	0.0002	0.00623	0.0002	0.0024	0.0002	52.7	2.1
AP-78	1.086	0.034	0.1086	0.006	0.0751	0.0011	0.0467	0.0011	1098	12
AP-79	0.414	0.024	0.414	0.001	0.221	0.001	0.035	0.001	304	11
AP-80	1.287	0.034	0.1287	0.006	0.062	0.002	0.036	0.002	1099	11
AP-81	1.711	0.04	0.1662	0.003	0.0738	0.0018	0.0458	0.0012	1012	15
AP-82	0.0566	0.0026	0.00608	0.0002	0.00608	0.0002	0.0024	0.0002	51.6	2.9
AP-83	0.1488	0.004	0.1488	0.0004	0.00778	0.0005	0.0329	0.0004	141.6	8.3
AP-84	1.626	0.026	0.1646	0.002	0.0719	0.0019	0.0449	0.0019	979	17
AP-85	0.0523	0.0029	0.00608	0.0002	0.00608	0.0002	0.0024	0.0002	51.6	2.9
AP-86	0.284	0.012	0.284	0.002	0.1646	0.0004	0.0468	0.0004	141.6	8.3
AP-87	0.052	0.002	0.052	0.0002	0.00599	0.0002	0.0024	0.0002	51.6	2.9
AP-88	0.0511	0.0018	0.00584	0.0001	0.00584	0.0001	0.0024	0.0001	51.6	2.9
AP-89	0.0536	0.0026	0.0062	0.0002	0.0062	0.0002	0.0024	0.0002	53	2.5
AP-90	1.5	0.045	0.15	0.007	0.074	0.002	0.047	0.002	979	17
AP-91	1.777	0.041	0.1735	0.006	0.0749	0.002	0.0478	0.002	1041	14
AP-92	1.565	0.044	0.1565	0.004	0.0735	0.0038	0.046	0.0038	1068	19
AP-93	0.0707	0.0077	0.00615	0.0002	0.00627					

AP00-42	1.07	0.15	0.1112	0.0051	0.0097	0.0008	0.045	0.0002	730	70	680	30	890	180	850	270	18.73	3.06	0.4834	
AP00-43	1.066	0.14	0.0847	0.0052	0.0047	0.0027	0.033	0.0007	1103	1103	1075	25	870	150	820	240	19.2	2.6	0.5943	
AP00-44	1.29	0.11	0.139	0.0044	0.0678	0.0058	0.057	0.013	836	50	830	20	800	190	760	167	2.3	0.1377		
AP00-45	1.072	0.14	0.087	0.0052	0.0074	0.0045	0.036	0.0006	1009	1009	1009	1009	1009	1009	1009	1009	1009	1009	1009	
AP00-46	1.49	0.12	0.1538	0.005	0.0707	0.0068	0.04	0.012	924	51	922	28	790	200	740	200	25.48	4.0	0.1193	
AP00-47	3.61	0.14	0.2463	0.0062	0.1063	0.002	0.0724	0.0055	1549	31	1519	32	1412	100	1372	34	3.94	146	0.7076	
AP00-48	1.72	0.14	0.1637	0.0051	0.0607	0.0045	0.036	0.0007	1007	1007	1007	1007	1007	1007	1007	1007	1007	1007	1007	
AP00-49	1.542	0.072	0.1531	0.0026	0.0722	0.0027	0.0506	0.0065	945	28	940	28	940	130	1003	85	5.14	11.75	0.2286	
AP00-50	1.626	0.058	0.1622	0.0008	0.0962	0.0008	0.054	0.0007	1007	1007	1007	1007	1007	1007	1007	1007	1007	1007	1007	
AP00-51	1.665	0.047	0.1589	0.001	0.076	0.001	0.0509	0.0055	995	18	991	17	1002	100	1009	130	1.38	9.48	0.0280	
AP00-52	1.3	0.14	0.0865	0.0043	0.061	0.001	0.0479	0.0036	1006	1006	1006	1006	1006	1006	1006	1006	1006	1006	1006	
AP00-53	1.699	0.043	0.1531	0.0054	0.054	0.0011	0.0479	0.0042	1007.4	1007.4	1007.4	1007.4	1007.4	1007.4	1007.4	1007.4	1007.4	1007.4	1007.4	
AP00-54	1.5	0.062	0.1553	0.0041	0.0702	0.0024	0.048	0.0065	925	34	930	20	940	130	925	94	2.09	3.47	0.1756	
AP00-55	1.16	0.12	0.0847	0.0052	0.0406	0.0041	0.0278	0.01	841	41	840	140	840	190	840	190	15.88	12.82	0.1468	
AP00-56	2.05	0.066	0.1674	0.0049	0.0794	0.0024	0.0512	0.0042	1132	22	1137	20	1130	100	1130	95	14.5	15.4	0.4435	
AP00-57	1.46	0.14	0.1648	0.005	0.074	0.004	0.048	0.004	1006	1006	1006	1006	1006	1006	1006	1006	1006	1006	1006	
AP00-58	1.386	0.083	0.143	0.006	0.0703	0.0033	0.0407	0.0049	861	35	861	26	861	95	861	95	9.3	30.1	0.3152	
AP00-59	0.228	0.032	0.0321	0.0029	0.0521	0.0034	0.0153	0.0014	207	26	204	18	202	18	200	20	30.4	87	0.2862	
AP00-60	1.443	0.041	0.1774	0.0025	0.1074	0.0022	0.0465	0.002	1006	1006	1006	1006	1006	1006	1006	1006	1006	1006	1006	
AP00-61	2.025	0.095	0.1852	0.0054	0.0795	0.0031	0.0533	0.005	1122	32	1095	29	1095	100	1100	81	5.47	37.4	0.8837	
AP00-62	1.448	0.046	0.145	0.0075	0.0756	0.0083	0.0441	0.0055	900	68	971	15	971	100	970	100	14.7	44.1	0.0320	
AP00-63	2.14	0.15	0.1893	0.0054	0.0826	0.0061	0.0542	0.0085	1150	50	1117	29	1150	160	1110	140	11.31	3.53	0.3121	
AP00-64	1.7	0.12	0.1672	0.0059	0.0495	0.0059	0.0595	0.0056	999	29	996	23	990	180	990	120	16.4	1.64	0.1454	
AP00-65	1.707	0.078	0.1677	0.0051	0.0748	0.0023	0.0529	0.0039	1014	27	1007	28	1030	100	1050	63	28.6	2.6	0.9909	
AP00-66	2.02	0.068	0.1838	0.0057	0.0799	0.0031	0.0563	0.0042	1122	21	1089	21	1088	81	1093	28	389	166.1	0.4270	
AP00-67	1.6	0.13	0.123	0.0074	0.0807	0.005	0.011	0.068	208	51	971	20	960	190	960	100	14.50	1.60	0.1450	
AP00-68	1.581	0.05	0.1517	0.0033	0.0747	0.0013	0.0462	0.0048	962	20	922	18	911	92	1057	34	479	18.4	0.3384	
AP00-69	1.93	0.13	0.123	0.0074	0.0727	0.0089	0.054	0.015	851	69	850	100	850	100	850	250	18.9	5.79	0.2910	
AP00-70	2.091	0.064	0.1922	0.0043	0.0749	0.0013	0.0707	0.0117	1145	21	1133	23	1180	140	1169	33	130	36.9	0.1118	
AP00-71	0.216	0.038	0.0315	0.0016	0.0501	0.0036	0.018	0.011	108	15	109.8	94	120	100	100	100	102.7	2.46	0.0240	
AP00-72	1.701	0.057	0.1664	0.0049	0.0844	0.0046	0.0484	0.0068	1008	21	960	100	1027	100	1027	100	7.86	11.11	0.1617	
AP00-73	1.58	0.13	0.1662	0.0056	0.0716	0.006	0.016	0.106	52	963	31	1280	300	800	170	842	1.087	0.1154		
AP00-74	1.983	0.046	0.1656	0.0064	0.0621	0.0054	0.048	0.009	1048	100	1007	31	1000	100	1000	100	100	100	100	
AP00-75	1.65	0.13	0.1619	0.0048	0.0745	0.006	0.057	0.012	990	47	947	27	1110	100	1010	100	9.61	1.065	0.1108	
AP00-76	1.609	0.11	0.1609	0.0047	0.0734	0.0056	0.048	0.0077	1000	1000	1000	1000	1000	1000	1000	1000	1000	1000	1000	
AP00-77	1.457	0.088	0.1511	0.004	0.0734	0.0041	0.0445	0.0062	948	35	947	22	880	120	1020	110	10.7	43.7	0.2617	
AP00-78	1.446	0.096	0.1478	0.0042	0.0715	0.0047	0.049	0.012	901	40	880	20	910	140	100	140	11.8	14.3	0.1324	
AP00-79	1.48	0.1	0.1492	0.0051	0.0705	0.0057	0.013	0.017	910	36	896	110	910	110	910	110	11.9	11.9	0.119	
AP00-80	1.809	0.039	0.1662	0.004	0.0787	0.004	0.0564	0.0074	1026	37	1000	24	1110	100	1080	110	20.78	3.88	0.1867	
AP00-81	1.172	0.064	0.1524	0.0054	0.0673	0.0053	0.043	0.003	907	907	902	701	902	701	902	701	902	701	902	
AP00-82	1.574	0.086	0.1588	0.0043	0.0725	0.0042	0.0477	0.0063	955	34	940	24	940	130	950	130	10.7	4.5	0.2542	
AP00-83	1.171	0.136	0.1536	0.0059	0.0709	0.0049	0.046	0.0079	940	44	940	24	940	130	940	140	14.1	10.7	0.1670	
AP00-84	1.55	0.11	0.1536	0.0059	0.0709	0.0049	0.046	0.0079	941	44	954	34	1040	180	800	140	17.0	2.15	0.1692	
AP00-85	1.597	0.043	0.156	0.0031	0.0741	0.0099	0.047	0.0041	968.4	17	935	27	928	78	1050	100	7.79	23.9	0.3037	
AP00-86	1.996	0.096	0.1816	0.004	0.0683	0.0043	0.056	0.006	935	737	940	100	940	100	940	100	10.8	10.8	0.108	
AP00-87	2.143	0.063	0.2	0.0047	0.0515	0.0047	0.0515	0.0047	1168	20	1125	17	1125	110	1125	38	149.1	15.46	0.1037	
AP00-88	1.188	0.108	0.0778	0.0048	0.0778	0.0048	0.0778	0.0048	1007	1007	1007	1007	1007	1007	1007	1007	1007	1007	1007	
AP00-89	1.6	0.13	0.1649	0.0056	0.0718	0.0069	0.068	0.014	957	53	983	31	1020	270	860	200	11.16	1.34	0.1201	
AP00-90	1.388	0.094	0.1601	0.004	0.0711	0.0044	0.047	0.011	876	41	850	19	930	190	930	140	16.2	1.87	0.1154	
AP00-91	1.809	0.091	0.1711	0.0091	0.0941	0.0067	0.0504	0.0051	1004	1004	1018	25	1018	110	1107	100	9.0	4.0	0.2888	
AP00-92	1.405	0.073	0.1488	0.0042	0.0731	0.0036	0.0411	0.0059	888	32	866	24	810	150	940	110	9.75	9.75	0.2899	
AP00-93	2.132	0.057	0.2084	0.0047	0.0843	0.0079	0.051	0.0039	1230	230	1230	230	1230	230	1230	230	1230	230	1230	
AP00-94	0.657	0.064	0.0654	0.0021	0.0731	0.0038	0.04	0.011	1057	38	408	13	780	220	90	140	2.04	2.04	0.1000	
AP00-95	1.96	0.056	0.1825	0.0042	0.0742	0.0045	0.0515	0.0045	1021	39	1021	39	1021	39	1021	39	1021	39	1021	
AP00-96	1.55	0.17	0.174	0.0011	0.0654	0.0079	0.053	0.018	945	72	1035	60	1040	370	780	280	13.2	2.08	0.1576	
AP00-97	1.65	0.13	0.1623	0.0051	0.0735	0.0046	0.0463	0.0038	977	50	990	30	910	170	970	140	13.2	1.84	0.1625	
AP00-98	1.739	0.082	0.2163	0.0039	0.0823	0.0039	0.0628	0.0042	821	847	781	847	781	847	781	847	781	847	781	
AP00-99	1.759	0.079	0.1735	0.004	0.0737	0.0025	0.0515	0.0075	1027	29	1031	22	1000	100	1016	69	5.36	3.64	0.0679	
AP00-100	1.56	0.11	0.1524	0.0043	0.0744	0.0029	0.044	0.005	945	823	1200	1200	1200	1200	1200	1200	1200	1200	1200	
AP00-101	1.85	0.14	0.1765	0.0063	0.0762	0.0049	0.063	0.011	1051	49	1047	49	1047	49	1047	49	1047	49	1047	
AP00-102	1.079	0.054	0.1535	0.0042	0.0731	0.0047	0.048	0.0077	933	38	940	27	940	100	940	100	940	100	940	
AP00-103	1.666	0.11	0.1631	0.0041	0.0744	0.0042	0.053	0.013	1061	35	1030	26	1030	140	1060	100	10.6	10.6	0.106	
AP00-104	1.84	0.11	0.1844	0.0051	0.0726	0.0036	0.0539	0.0039	1053	39	990	28	1061	82	970	100	24	61.7	2.708	
AP00-105	1.346	0.068	0.1608	0.0049	0.0719	0.0043	0.054	0.004	884	833	848	833	848	833	848	833	848	833	848	

AP_011_107	3.15	0.21	0.182	0.039	0.1258	0.0085	0.1187	0.013	0.1443	52	1078	22	2260	340	2050	120	9.47	4.17	0.4923	
AP_011_108	1.24	0.1525	0.003	0.0788	0.0853	0.940	0.028	0.185	12	116	185	10	1168	1050	1156	25	11.67	1.747	0.1747	
AP_011_109	1.854	0.642	0.1753	0.032	0.0768	0.012	0.0226	0.043	0.1654	15	1041	18	1108	83	1119	32	6.32	1.15	0.1820	
AP_011_110	1.22	0.152	0.0015	0.0244	0.0026	0.0107	0.0041	0.016	0.1844	10	1162	10	1168	1200	1200	10	10.48	1.148	0.1848	
AP_011_111	1.519	0.019	0.1562	0.0024	0.0708	0.00885	0.046	0.038	0.937	7.6	95.3	13	110	73	95	25	1.088	110.9	0.9813	
AP_011_112	1.61	0.11	0.161	0.0041	0.0719	0.0057	0.0338	0.007	0.980	39	926	30	1100	130	1100	140	12.94	5.99	0.4629	
AP_011_113	1.545	0.1577	0.0034	0.0172	0.0045	0.0012	0.0045	0.048	0.948	94	944	94	1167	972	972	84	4.16	0.668	0.1688	
AP_011_114	0.087	0.088	0.0028	0.0787	0.0032	0.0442	0.0048	0.044	2.141	29	1325	23	1168	82	44	39.4	0.895	0.895	0.895	
AP_011_115	1.43	0.053	0.0173	0.0033	0.0422	0.0044	0.003	0.044	0.83	0.4	95.7	10	1167	1023	1023	81.2	15.38	0.4511	0.4511	
AP_011_116	1.9	0.15	0.1725	0.0057	0.0993	0.0069	0.071	0.0087	1073	53	1026	31	1120	170	1130	150	12.74	3.28	0.2975	
AP_011_117	1.648	0.0225	0.1676	0.0025	0.0422	0.0071	0.041	0.041	0.91	19	1167	19	1167	100	1167	100	10.88	14.1	0.8151	
AP_011_118	1.46	0.126	0.1849	0.0049	0.0819	0.0279	0.0702	0.01	1270	83	1039	16	1130	180	1140	180	11.33	10.78	0.8275	
AP_011_119	1.942	0.049	0.1878	0.0041	0.0745	0.0018	0.0146	0.047	1.094	17	1147	22	1073	80	1063	49	4.58	59.6	1.2013	
AP_011_120	1.14	0.215	0.154	0.0075	0.0404	0.0026	0.0044	0.279	0.172	0.2	1162	24	1120	287	1120	287	11.78	21.8	0.4102	
AP_011_121	1.565	0.077	0.1873	0.0043	0.072	0.0039	0.0412	0.0044	957	31	947	24	111	85	86	110	11.5	11.09	0.8215	
AP_011_122	1.16	0.09	0.1884	0.0028	0.0748	0.0023	0.042	0.044	0.169	0.1	1167	10	1167	1160	1160	103.2	15.82	0.1523		
AP_011_123	1.094	0.038	0.1634	0.0027	0.0701	0.0027	0.0504	0.044	967	15	975.5	15	994	85	967	46	6.82	40.9	0.5997	
AP_011_124	1.058	0.048	0.1638	0.0029	0.0711	0.0022	0.0488	0.0399	958	19	978	16	104	75	966	60	6.29	37.7	0.5994	
AP_011_125	1.01	0.149	0.1519	0.0073	0.0551	0.0052	0.0322	0.044	1122	11	985.5	14	988	78	964	51	8.81	18.51	0.2325	
AP_011_126	1.661	0.077	0.1678	0.0057	0.0712	0.0019	0.0533	0.005	999	29	999	32	1048	96	954	50	23.7	0.2833	0.2833	
AP_011_127	1.794	0.052	0.1676	0.0027	0.0749	0.0022	0.0461	0.0041	1000	96.8	954	15	966	1058	966	79	10.6	12.6	1.2887	
AP_011_128	1.527	0.089	0.1526	0.0024	0.0723	0.0032	0.0448	0.0042	946	27	945.5	14	885	81	885	91	29.79	13.54	0.4545	
AP_011_129	1.578	0.075	0.1595	0.0024	0.0698	0.0033	0.0467	0.0046	957	19	954	19	960	80	968	150	11.86	0.542	0.542	
AP_011_130	1.582	0.06	0.1629	0.0031	0.0691	0.0025	0.0517	0.005	960	24	975	18	1019	95	895	79	21.2	11.5	0.5425	
AP_011_131	1.67	0.068	0.1638	0.0031	0.0734	0.0028	0.0496	0.0054	994	26	978	17	977	100	1004	80	19.8	63.9	0.3227	
AP_011_132	1.581	0.083	0.1602	0.0032	0.0705	0.0024	0.0479	0.0041	964	94	974	16	966	964	964	63	44.7	2.1887	2.1887	
AP_011_133	1.801	0.026	0.1775	0.0027	0.0734	0.0011	0.0462	0.0399	1045.6	9.4	1053.3	15	913	75	1022	30	5.48	61.5	0.1122	
AP_011_134	1.523	0.036	0.1544	0.0026	0.0719	0.0018	0.0475	0.0045	940	15	955.5	14	938	88	944	51	8.81	18.51	0.2325	
AP_011_135	1.6	0.044	0.1624	0.0034	0.0716	0.0017	0.0468	0.004	969	17	970	19	924	77	968	47	10.9	69.9	0.6413	
AP_011_136	1.72	0.052	0.1728	0.0034	0.0731	0.0021	0.0527	0.0027	1021	20	1027	19	1017	104	97	102	21.2	31.19	1.1487	
AP_011_137	1.662	0.045	0.1683	0.0029	0.0724	0.0021	0.0587	0.0058	1114	11	1121	11	1010	112	1010	112	40.3	26.1	0.9295	
AP_011_138	1.62	0.043	0.1683	0.0028	0.0715	0.002	0.0479	0.0046	990	17	990	16	945	88	944	57	22.84	33.02	0.4095	
AP_011_139	1.47	0.064	0.1682	0.0038	0.062	0.0022	0.0402	0.004	991	16	991	16	960	100	1100	110	11.02	31.6	0.2955	
AP_011_140	1.8	0.16	0.144	0.0051	0.0799	0.0037	0.0508	0.0087	1052	55	974	29	1160	130	1190	170	14.91	4.58	0.3072	
AP_011_141	1.507	0.057	0.1545	0.0025	0.0733	0.0025	0.0473	0.0041	1017	11	1017	11	950	12	950	12	10.88	11.86	0.542	
AP_011_142	1.608	0.036	0.1581	0.0031	0.0711	0.0025	0.0487	0.0051	970	24	974	18	916	98	956	75	25.9	14.06	0.5429	
AP_011_143	2.086	0.047	0.1593	0.0038	0.0721	0.0032	0.0527	0.0057	1136	16	1138	21	1128	110	1120	111	30.0	14.67	0.4899	
AP_011_144	1.945	0.1563	0.1493	0.0039	0.0713	0.0031	0.0456	0.0041	946	94	946	94	968	968	968	64.6	1.945	1.945		
AP_011_145	2.182	0.038	0.1941	0.0032	0.08178	0.0031	0.058	0.0407	1177	11	1143	19	1139	27	497	153	0.3078	0.3078	0.3078	
AP_011_146	1.52	0.054	0.1624	0.0027	0.0688	0.0027	0.0469	0.0027	104	19	933	23	922	80	924	28	16.58	16.58	16.58	
AP_011_147	1.646	0.077	0.1606	0.0031	0.0751	0.0039	0.0466	0.0047	984	30	940	17	920	92	1030	110	11.72	8.43	0.6627	
AP_011_148	1.42	0.1627	0.1647	0.0037	0.0687	0.0037	0.0502	0.0048	968	23	967	23	971	100	1000	110	11.87	57.6	0.542	
AP_011_149	1.079	0.1948	0.1948	0.0044	0.0755	0.0021	0.0556	0.0048	1141	15	1147	13	1092	92	1127	61.8	41.5	0.9061	0.9061	
AP_011_150	1.636	0.059	0.1644	0.0039	0.0727	0.0027	0.0502	0.0048	968	23	967	23	989	91	1000	74	30.1	13.61	0.4522	
AP_011_151	1.043	0.063	0.168	0.0037	0.068	0.0037	0.0493	0.004	966	24	972	19	960	79	960	83	49.7	0.542	0.542	
AP_011_152	1.096	0.039	0.171	0.0032	0.0723	0.0014	0.0544	0.0048	968	15	1017	18	955	78	988	78	7.33	169.9	2.3179	
AP_011_153	1.659	0.086	0.1659	0.0027	0.068	0.0027	0.0469	0.004	961	24	970	19	989	100	1000	112.3	63.2	0.3078	0.3078	
AP_011_154	1.502	0.054	0.1589	0.0028	0.0691	0.0027	0.0509	0.0059	1005	22	981	16	1003	87	891	81	27.7	18.8	0.8787	
AP_011_155	1.639	0.076	0.1645	0.0037	0.0705	0.0028	0.0509	0.0055	981	29	981	21	1003	110	971	81	13.65	9.22	0.6705	
AP_011_156	1.145	0.2004	0.2004	0.0048	0.0704	0.0044	0.0564	0.0048	1164	10	1177	19	1166	1166	1166	12.7	11.404	11.404	11.404	
AP_011_157	1.673	0.056	0.1703	0.0031	0.0723	0.0026	0.0496	0.0049	1005	24	1013	17	978	91	1001	72	17.58	7.87	0.4081	
AP_011_158	1.04	0.1527	0.1527	0.0037	0.0717	0.0035	0.0497	0.004	911	0.4	916	0.6	916	916	916	916	916	916	916	916
AP_011_159	1.55	0.11	0.1646	0.0042	0.0751	0.0046	0.0493	0.0045	945	40	944	24	970	140	910	140	15.9	5.48	0.3447	
AP_011_160	1.584	0.056	0.1613	0.0037	0.0717	0.0037	0.0469	0.004	961	24	970	19	989	100	1000	112.3	63.2	0.3078	0.3078	
AP_011_161	1.54	0.11	0.1639	0.0044	0.0696	0.0049	0.0556	0.0087	918	40	978	25	1000	170	850	140	10.11	3.01	0.2977	
AP_011_162	1.649	0.051	0.1648	0.0034	0.0726	0.002	0.0512	0.0062	990	19	998	19	1008	90	993	57	17.3	18.44	0.4944	
AP_011_163	1.724	0.046	0.1711	0.0034	0.0724	0.0032	0.0493	0.0043	1016	0.3	1018	0.8	1018	1018	1018	1018	1018	1018	1018	1018
AP_011_164	1.894	0.078	0.1867	0.0041	0.0731	0.0028	0.0459	0.0056	1080	28	1109	22	1078	110	1023	85	18.04	7.67	0.4522	
AP_011_165	1.625	0.056	0.1646	0.0026	0.0696	0.0028	0.0484	0.0028	1014	24	1014	24	1014	1014	1014	1014	1014	1014	1014	1014
AP_011_166	1.77	0.11	0.1689	0.0041	0.0745	0.0049	0.0491	0.0047	1044	38	1046	22	1060	150	1140	140	10.52	2.77	0.2833	
AP_011_167	1.502	0.052	0.1623	0.0039	0.0709	0.0039	0.0463	0.0043	962.9	39	969	34	914	100	914	100	94.1	100.1	0.2833	
AP_011_168	1.6	0.1578	0.1658	0.0036	0.0726	0.003	0.048	0.006	80	82	940	82	101	181	99	181	18.66	1.0558	1.0558	
AP_011_169	1.555	0.04	0.1608	0.0027	0.0702	0.0036	0.0482	0.004	951	16	951	15	951	77	929	47	97.			

AP-048-97	1.615	0.068	0.1654	0.0038	0.0707	0.0025	0.0554	0.0067	972	26	987	21	1287	130	933	74	49	4.87	0.0994	
AP-048-98	1.746	0.034	0.0037	0.0864	0.0051	0.0043	0.0041	0.0043	1342	13	1447	23	1366	1366	1481	23	44	60.4	0.078	
AP-048-99	2.229	0.041	0.1987	0.0037	0.08905	0.0012	0.0587	0.0048	1190	13	1168	20	1152	92	1219	29	80	174.3	0.1752	
AP-048-100	1.02	0.002	0.0013	0.07708	0.0047	0.0013	0.0047	0.0047	1024	13	1024	13	1024	1125	1125	27	101	1.601	0.1466	
AP-048-101	2.684	0.053	0.2111	0.0063	0.0845	0.0016	0.0628	0.0061	1324	15	1230	27	1320	1320	1314	41	651	48	0.2137	
AP-048-102	1.822	0.096	0.0749	0.0057	0.0739	0.002	0.0648	0.0054	1058	33	1058	31	1077	1030	1030	57	55.1	9.3	0.1797	
AP-048-103	2.231	0.002	0.2032	0.0033	0.0799	0.0013	0.0698	0.0046	1131	17	1131	17	1131	1192	1192	30	14	260	0.1466	
AP-048-104	1.486	0.037	0.1495	0.0032	0.0719	0.002	0.0449	0.0055	924	15	893	18	887	150	79	83	73	10.19	0.1285	
AP-048-105	2.295	0.043	0.0623	0.005	0.0813	0.0015	0.069	0.011	1196	13	1196	13	1196	1196	1255	15	499	199	0.2689	
AP-048-106	2.188	0.07	0.1959	0.0049	0.0808	0.0021	0.0593	0.0054	1177	22	1122	22	1122	100	1214	51	297	117	0.3939	
AP-048-107	2.076	0.034	0.1889	0.0042	0.0795	0.0015	0.0664	0.0054	1140	18	1133	29	1174	118	118	108	48	171	48	0.2838
AP-048-108	1.441	0.004	0.1641	0.0045	0.08128	0.0011	0.0664	0.0054	1124	11	1106	11	1106	1299	1291	131	245	84	0.2461	
AP-048-109	1.584	0.038	0.1529	0.0036	0.0761	0.0013	0.0444	0.0077	963	26	917	26	1070	150	1115	79	517	85.9	0.1662	
AP-048-110	2.461	0.042	0.0013	0.08203	0.0026	0.0013	0.08203	0.0026	1303	12	1303	12	1303	1266	1266	10	281	112.9	0.2787	
AP-048-111	2.058	0.068	0.1048	0.0011	0.0786	0.0013	0.0772	0.005	1114.8	10	1125	17	1125	95	1116	33	103.3	41.8	0.1253	
AP-048-112	1.155	0.029	0.1555	0.0029	0.0766	0.0012	0.064	0.0054	1124	15	1124	15	1124	1124	1124	17	44.2	0.1188	0.242	
AP-048-113	1.61	0.029	0.1649	0.0026	0.07215	0.0011	0.0478	0.004	974	11	987.2	15	944	77	984	11	192	52.8	0.2750	
AP-048-114	1.756	0.066	0.0745	0.0054	0.0725	0.0022	0.0626	0.006	1026	25	1026	25	1026	100	989	62	34	611	0.1797	
AP-048-115	1.644	0.068	0.0935	0.007	0.0711	0.0027	0.071	0.06	990	19	990	19	990	190	990	31.8	28	356	0.2822	
AP-048-116	2.048	0.081	0.19	0.0033	0.0789	0.0027	0.0513	0.0079	1133	25	1131	29	1010	1189	1189	62	27.2	2.266	0.0829	
AP-048-117	2.269	0.066	0.209	0.0076	0.0641	0.0021	0.0641	0.0057	124	14	1228	22	1228	125	125	110	102	86.2	0.149	
AP-048-118	2.08	0.051	0.1958	0.0034	0.0763	0.0022	0.055	0.056	1141	17	1153	18	1082	150	1093	58	437	9.39	0.2149	
AP-048-119	2.079	0.057	0.1898	0.0052	0.0779	0.0021	0.0604	0.0067	1140	19	1147	20	1180	160	1145	55	341	3.28	0.0962	
AP-048-120	1.592	0.051	0.1624	0.0031	0.0707	0.0024	0.0465	0.0062	965	20	975	17	915	120	933	70	50.1	5.61	0.1120	
AP-048-121	2.129	0.086	0.0986	0.0038	0.0775	0.0041	0.061	0.0025	1182	10	1187	25	1187	1259	1131	174	38	104	0.2822	
AP-048-122	2.4	0.054	0.2161	0.0038	0.0823	0.0015	0.0617	0.0033	1256	15	1261	20	1210	1200	1250	36	159	40.1	0.2522	
AP-048-123	2.16	0.049	0.1956	0.0041	0.0799	0.0036	0.0675	0.0167	1167	16	1152	24	1152	1240	1185	60	42.5	10	0.2523	
AP-048-124	2.006	0.046	0.189	0.003	0.0767	0.0019	0.0597	0.0054	1117	16	1116	16	1112	100	1112	51	90.9	14.12	0.1575	
AP-048-125	1.463	0.053	0.1489	0.0033	0.0711	0.0027	0.06	0.0078	914	22	895	29	1000	150	974	79	32.7	3.97	0.214	
AP-048-126	1.463	0.053	0.1489	0.0033	0.0711	0.0027	0.06	0.0078	914	22	895	29	1000	150	974	79	32.7	3.97	0.214	
AP-048-127	2.11	0.065	0.1965	0.004	0.0765	0.0062	0.0582	0.0146	1164	19	1156	17	1120	89	1134	33	62	167.8	0.288	
AP-048-128	2.025	0.033	0.1891	0.0011	0.0776	0.0013	0.057	0.0066	1123.6	10	1120	10	1120	100	1120	104	45.6	10.4	0.1966	
AP-048-129	2.025	0.033	0.1891	0.0011	0.0776	0.0013	0.057	0.0066	1123.6	10	1120	10	1120	100	1120	104	45.6	10.4	0.1966	
AP-048-130	2.526	0.042	0.2176	0.0035	0.084	0.0014	0.0661	0.0054	1281	12	1249	18	1249	120	1291	33	223.4	79.6	0.3963	
AP-048-131	1.959	0.064	0.1664	0.004	0.0765	0.0011	0.0661	0.0054	1189	11	1189	11	1189	1189	1189	10	103	165.3	0.283	
AP-048-132	2.231	0.064	0.2046	0.0038	0.0791	0.0025	0.0644	0.0065	1190	20	1200	120	1164	161	1164	61	49.2	8.29	0.1685	
AP-048-133	2.192	0.04	0.2107	0.0032	0.0824	0.0013	0.0673	0.0101	1240	12	1233	19	1241	97	1258	30	303	138.9	0.4684	
AP-048-134	2.32	0.024	0.2024	0.005	0.084	0.0012	0.0814	0.0052	1248	12	1248	12	1248	1186	1248	27	317	3.7	0.1023	
AP-048-135	2.425	0.073	0.2145	0.0058	0.08183	0.0012	0.0814	0.0052	1248	12	1230	28	1244	28	1244	28	75	88	7.8	0.1173
AP-048-136	2.119	0.054	0.1974	0.0031	0.0785	0.0015	0.0696	0.0156	1164	16	1164	16	1164	1168	1164	16	1164	1168	16	0.1173
AP-048-137	1.708	0.057	0.168	0.0043	0.0741	0.0016	0.0502	0.005	1005	10	1002	17	988	96	1047	41	61.6	8.36	0.1367	
AP-048-138	2.189	0.045	0.2045	0.0043	0.0782	0.0017	0.0619	0.005	1179	17	1188	21	1179	1179	1179	14	116	10.8	0.2838	
AP-048-139	2.088	0.103	0.1939	0.0054	0.0788	0.0011	0.0652	0.0048	1147.7	15	1142	18	1148	1148	1148	16	726	86.2	0.1187	
AP-048-140	2.004	0.051	0.1903	0.0038	0.076	0.0016	0.0575	0.0054	1111.9	15	1123	21	1129	100	1098	43	55.9	13.81	0.2470	
AP-048-141	2.156	0.045	0.2061	0.0039	0.0781	0.0017	0.0562	0.0066	1175	17	1175	17	1175	1175	1175	14	116	116	0.1187	
AP-048-142	2.212	0.056	0.2052	0.0057	0.079	0.0028	0.0818	0.011	1185	13	1203	30	1200	150	1157	71	297	2.39	0.0805	
AP-048-143	1.828	0.067	0.1779	0.0059	0.0767	0.0016	0.0564	0.0055	1169	16	1169	16	1169	1169	1169	11	104	55.1	11.01	0.2838
AP-048-144	1.528	0.061	0.1617	0.0041	0.068	0.0026	0.0485	0.0074	943	23	966	23	956	90	987	41.4	10.7	0.1349	0.2838	
AP-048-145	2.578	0.053	0.2246	0.0045	0.0824	0.0012	0.0814	0.0052	1293	15	1306	24	1242	90	1242	27	610	90	0.1475	
AP-048-146	2.02	0.037	0.1911	0.0032	0.0766	0.0012	0.0516	0.0025	1125	11	1127	19	1127	129	1127	27	394	8.2	0.1430	
AP-048-147	1.525	0.063	0.1613	0.0041	0.0683	0.0025	0.0525	0.0081	998	25	993	19	1000	160	984	83	52.3	5.38	0.1029	
AP-048-148	1.072	0.064	0.1674	0.0032	0.0714	0.0011	0.0647	0.006	104	24	104	24	104	104	104	27.2	4.4	0.1142	0.2838	
AP-048-149	1.423	0.053	0.1263	0.004	0.0646	0.0016	0.0575	0.0056	1278.6	11	1262	21	1321	130	1306	36	249.5	83.7	0.3955	
AP-048-150	2.608	0.053	0.2027	0.0037	0.0809	0.0017	0.0675	0.0047	1102	15	1102	15	1102	1102	1102	15	84.6	0.1779	0.2838	
AP-048-151	2.608	0.053	0.2027	0.0037	0.0809	0.0017	0.0675	0.0047	1102	15	1102	15	1102	1102	1102	15	84.6	0.1779	0.2838	
AP-048-152	2.049	0.053	0.1893	0.0042	0.0781	0.002	0.0554	0.0049	1131	18	1134	23	1134	1134	1134	31	589	106.8	0.1913	
AP-048-153	2.154	0.036	0.2107	0.0036	0.0805	0.0012	0.0672	0.0054	1187.2	10	1180	27	1180	1207	1180	27	394	2.026	0.2838	
AP-048-154	1.5	0.035	0.1533	0.0026	0.0706	0.0016	0.0481	0.0033	930	14	931.8	14	948	100	939	46	86.5	9.17	0.1060	
AP-048-155	2.414	0.053	0.2013	0.0033	0.0771	0.0043	0.0654	0.0036	1044	14	1044	14	1044	1044	1044	14	104	104	0.1060	
AP-048-156	2.196	0.042	0.2001	0.0045	0.0789	0.0013	0.0553	0.0049	1179	13	1176	19	1176	95	1171	31	295.9	28.99	0.0980	
AP-048-157	2.414	0.053	0.2013	0.0033	0.0771	0.0043	0.0654	0.0036	1044	14	1044	14	1044	1044	1044	14	104	104	0.1060	
AP-048-158	2.154	0.036	0.2107	0.0036	0.0805	0.0012	0.0672	0.0054	1187.2	10	1180	27	1180	1207	1180	27	394	2.026	0.2838	
AP-048-159	2.241	0.055	0.2018	0.0037	0.08	0.0014	0.0603	0.005	1193	17	1185	20	1185	20						

DA-39-72	1.582	0.043	0.1609	0.0032	0.0713	0.0018	0.0491	0.0042	962	27	961	18	969	81	958	53	899	32.9	0.6093	
DA-39-73	1.526	0.0545	0.0731	0.0029	0.0469	0.0029	0.0489	0.0042	936	27	936	18	944	81	943	53	899	32.9	0.6093	
DA-39-74	1.57	0.18	0.1649	0.0053	0.0677	0.008	0.0485	0.0096	930	28	940	29	950	180	70	250	4.7	1.484	0.1517	
DA-39-75	1.433	0.12	0.1623	0.005	0.0677	0.008	0.0485	0.0096	930	28	940	29	950	180	70	250	4.7	1.484	0.1517	
DA-39-76	1.338	0.02	0.1441	0.0024	0.0677	0.008	0.0485	0.0096	861.9	9.7	868	13	862	49	861	49	126.1	199.4	1.5913	
DA-39-77	1.38	0.036	0.1429	0.0024	0.0703	0.006	0.042	0.0386	882	16	865	14	832	71	938	45	108.4	40.1	0.3699	
DA-39-78	1.386	0.1	0.1484	0.0028	0.0616	0.0012	0.0416	0.0094	882	17	882	17	882	26	882	26	882	26	882	
DA-39-79	1.6	0.12	0.1528	0.0024	0.0728	0.001	0.058	0.013	965	44	965	44	965	150	250	960	150	250	960	
DA-39-80	1.638	0.043	0.1429	0.0029	0.0601	0.0021	0.044	0.0484	964	44	964	44	964	150	250	960	150	250	960	
DA-39-81	1.463	0.039	0.1504	0.0027	0.0703	0.002	0.0454	0.0408	914	16	900	15	866	94	945	60	59.1	9.8	0.1608	
DA-39-82	1.527	0.039	0.1509	0.0028	0.0719	0.001	0.0512	0.0391	914	16	900	15	866	94	945	60	59.1	9.8	0.1608	
DA-39-83	1.444	0.035	0.1512	0.0027	0.0618	0.0011	0.044	0.046	914	16	912	15	815	61	917	77	918	54	1.615	
DA-39-84	1.437	0.067	0.1509	0.0032	0.0685	0.003	0.0487	0.0508	901	28	886	18	879	110	880	100	24.7	80.8	0.3729	
DA-39-85	1.448	0.067	0.1511	0.0033	0.0693	0.003	0.0487	0.0508	901	28	886	18	879	110	880	100	24.7	80.8	0.3729	
DA-39-86	1.399	0.038	0.1469	0.0026	0.0568	0.0018	0.0455	0.040	887	26	886	18	879	110	880	100	24.7	80.8	0.3729	
DA-39-87	1.495	0.037	0.1555	0.0023	0.0719	0.001	0.0512	0.0391	914	16	912	15	815	61	917	77	918	54	1.615	
DA-39-88	1.901	0.061	0.1504	0.0034	0.0879	0.0026	0.0542	0.0404	1079	21	932	13	1066	84	1371	57	76.7	79.9	1.0417	
DA-39-89	1.41	0.063	0.1462	0.0024	0.0693	0.0027	0.0456	0.0355	892	13	891	13	862	69	900	50	92.9	70.4	0.7578	
DA-39-90	1.481	0.063	0.1529	0.0024	0.0694	0.0028	0.0457	0.0356	892	13	891	13	862	69	900	50	92.9	70.4	0.7578	
DA-39-91	1.514	0.064	0.1536	0.0028	0.0705	0.0028	0.046	0.04	895	16	895	16	895	70	905	48	88	88	88	
DA-39-92	1.407	0.124	0.1496	0.0028	0.0695	0.0017	0.0445	0.0391	901	16	886	15	885	70	905	48	88	88	88	
DA-39-93	1.402	0.124	0.1496	0.0028	0.0695	0.0017	0.0445	0.0391	901	16	886	15	885	70	905	48	88	88	88	
DA-39-94	1.421	0.01	0.1496	0.0024	0.0694	0.0016	0.0475	0.0395	901	16	886	15	885	70	905	48	88	88	88	
DA-39-95	1.552	0.086	0.1589	0.0034	0.0708	0.0043	0.0452	0.052	951	36	940	19	863	100	940	140	17.33	5.39	0.3110	
DA-39-96	1.48	0.057	0.1518	0.0027	0.0694	0.0027	0.0456	0.0401	920	24	945	19	863	100	940	140	17.33	5.39	0.3110	
DA-39-97	1.52	0.057	0.1527	0.0028	0.0698	0.002	0.0464	0.0401	920	24	945	19	863	100	940	140	17.33	5.39	0.3110	
DA-39-98	1.45	0.075	0.1503	0.0034	0.0699	0.0036	0.0439	0.0399	907	32	907	19	869	75	910	110	41.2	38.3	0.9286	
DA-39-99	1.463	0.054	0.1547	0.0032	0.0682	0.002	0.0481	0.0356	912	22	927	18	867	110	884	75	46	24.70	0.2470	
DA-39-100	0.284	0.015	0.0882	0.0008	0.0531	0.0031	0.025	0.01	203	12	200	10	490	200	100	110	53.7	0.932	0.0172	
DA-39-101	0.247	0.0057	0.0701	0.0005	0.0501	0.0016	0.0136	0.0054	203.3	44	243.2	34	232	27	101	50	448	17.2	0.3888	
DA-39-102	0.197	0.0049	0.1504	0.0023	0.0493	0.0025	0.0499	0.0019	902	21	902	13	877	76	903	79	79	79	79	
DA-39-103	0.147	0.0049	0.1506	0.0023	0.0494	0.0026	0.0499	0.0019	902	21	902	13	877	76	903	79	79	79	79	
DA-39-104	0.146	0.013	0.1619	0.0039	0.0509	0.003	0.0708	0.003	901	20	907	10	907	100	1000	1000	1000	1000	1000	
DA-39-105	1.477	0.079	0.1525	0.0032	0.0708	0.0038	0.0441	0.0555	915	33	915	33	915	100	110	90	110	739	0.3732	
DA-39-106	1.49	0.13	0.1563	0.0063	0.0693	0.0043	0.05	0.059	915	33	915	33	915	100	110	90	110	739	0.3732	
DA-39-107	1.507	0.14	0.1553	0.0065	0.0705	0.0049	0.0463	0.042	915	16	915	15	895	81	933	57	60.4	27.8	0.4003	
DA-39-108	1.756	0.058	0.1518	0.0036	0.0731	0.0021	0.0494	0.0402	1031	20	1022	20	974	81	1004	65	43.5	43.1	0.9908	
DA-39-109	1.787	0.073	0.1713	0.0036	0.0736	0.0027	0.0511	0.033	1027	20	1027	20	1027	106	1027	106	1027	106	1027	
DA-39-110	1.48	0.1	0.1531	0.0039	0.058	0.0045	0.043	0.051	913	43	913	22	890	99	870	140	11.5	5.65	0.4913	
DA-39-111	1.448	0.054	0.1524	0.0038	0.0595	0.004	0.05	0.041	914	43	914	22	890	99	870	140	11.5	5.65	0.4913	
DA-39-112	1.393	0.023	0.1487	0.0022	0.0685	0.0012	0.0442	0.0386	885.8	5.7	877	12	873	12	873	12	881	38	169.6	0.9976
DA-39-113	1.393	0.023	0.1487	0.0022	0.0685	0.0012	0.0442	0.0386	885.8	5.7	877	12	873	12	873	12	881	38	169.6	0.9976
DA-39-114	1.393	0.023	0.1487	0.0022	0.0685	0.0012	0.0442	0.0386	885.8	5.7	877	12	873	12	873	12	881	38	169.6	0.9976
DA-39-115	1.426	0.052	0.1485	0.0028	0.0688	0.0028	0.0442	0.04	899	22	904	18	878	89	899	84	49.5	24.8	0.5010	
DA-39-116	1.456	0.052	0.1512	0.0032	0.0694	0.0029	0.0442	0.04	899	22	904	18	878	89	899	84	49.5	24.8	0.5010	
DA-39-117	1.494	0.062	0.1559	0.0033	0.0706	0.0032	0.0435	0.0395	899	25	934	18	871	100	930	91	17.7	5.2	0.2938	
DA-39-118	1.504	0.054	0.1545	0.0034	0.0706	0.0032	0.0435	0.0395	899	25	934	18	871	100	930	91	17.7	5.2	0.2938	
DA-39-119	1.43	0.058	0.1519	0.0027	0.069	0.0024	0.0466	0.0407	899	21	912	15	900	91	907	70	34	15.51	0.4562	
DA-39-120	1.445	0.058	0.1517	0.0023	0.0688	0.0023	0.0468	0.0504	894	28	912	15	900	91	907	70	34	15.51	0.4562	
DA-39-121	1.427	0.043	0.1483	0.0026	0.0694	0.0024	0.0468	0.0504	894	28	912	15	900	91	907	70	34	15.51	0.4562	
DA-39-122	1.407	0.049	0.145	0.0023	0.0709	0.0025	0.0467	0.0404	890	24	873	13	923	84	962	72	48.8	20.3	0.4160	
DA-39-123	1.481	0.037	0.1566	0.0033	0.0643	0.0027	0.0411	0.0317	896	29	917	19	897	100	900	130	13.3	13.9	0.3208	
DA-39-124	1.398	0.049	0.1521	0.0036	0.067	0.0041	0.0406	0.0373	884	34	873	20	900	140	880	130	19	5.02	0.3116	
DA-39-125	1.401	0.025	0.1463	0.0025	0.0696	0.0023	0.0437	0.0335	889.9	11	889	14	864	69	914	37	294.2	153	0.5201	
DA-39-126	1.469	0.049	0.1497	0.0037	0.0718	0.0025	0.0465	0.0402	892	30	880	21	899	21	889	21	889	21	889	
DA-39-127	1.45	0.047	0.1488	0.0037	0.0702	0.0026	0.0444	0.0402	893	30	880	21	899	21	889	21	889	21	889	
DA-39-128	1.499	0.038	0.1571	0.0028	0.0704	0.0027	0.0441	0.0401	935	16	941	16	938	79	934	70	904	50.7	0.5256	
DA-39-129	1.466	0.052	0.1522	0.0039	0.0697	0.0027	0.0441	0.0401	935	16	941	16	938	79	934	70	904	50.7	0.5256	
DA-39-130	1.503	0.048	0.1548	0.0028	0.0714	0.0024	0.0457	0.0399	933	19	943	15	948	71	964	71	144.8	22.3	0.4978	
DA-39-131	1.415	0.045	0.1488	0.0025	0.0682	0.0022	0.0435	0.0337	866	18	866	15	841	35	861	35	61.6	35.6	0.7421	
DA-39-132	1.35	0.043	0.1438	0.0027	0.0689	0.0022	0.0425	0.0337	866	18	866	15	841	35	861	35	61.6	35.6	0.7421	
DA-39-133	1.74	0.045	0.1728	0.0032	0.074	0.0018	0.0507	0.0402	1022	17	1027	18	999	81	1033	50	63.9	49.73	0.7762	
DA-39-134	1.74	0.045	0.1728	0.0032	0.074	0.0018	0.0507	0.0402	1022	17	1027	18	999	81	1033	50	63.9	49.73	0.7762	
DA-39-135	1.294	0.028	0.1428	0.0028	0.0718	0.0028	0.0421	0.043	904	18	892	12	882	68	904	140	14.3	20.76	0.4856	
DA-39-136	1.422	0.069	0.1432	0.0052	0.0724	0.0046	0.0491	0.0371	906	24	896	20	900	140	880	130	23.52			

MPP-3A_72	0.273	0.029	0.0307	0.0012	0.0548	0.0099	0.0112	0.0018	244	47	154	5848	7.5	225	37	720	120	520	680	1.2077	
MPP-3A_73	0.276	0.0378	0.0313	0.0024	0.0523	0.0014	0.0293	0.0017	247	14	14	2188	8	261	310	281	310	281	310	1.0516	
MPP-3A_74	0.305	0.018	0.0447	0.0011	0.0499	0.0017	0.0150	0.0013	269	14	280	280	6.8	30	180	110	178.3	162	102	0.9386	
MPP-3A_75	0.316	0.0203	0.0387	0.0019	0.0429	0.0011	0.0223	0.0015	267	10	267	267	6.8	20	207	207	148.6	141	148.6	0.7741	
MPP-3A_76	0.313	0.013	0.0398	0.001	0.057	0.0022	0.0150	0.0013	276	10	252.7	6.3	301	25	47	86	213	145.1	169.12	0.8812	
MPP-3A_77	0.31	0.017	0.0437	0.0008	0.052	0.0027	0.01376	0.0012	273	13	273	5.5	276	24	270	110	203	274	1.0542	0.8974	
MPP-3A_78	0.346	0.0081	0.0438	0.0013	0.0481	0.0012	0.0183	0.001	282	10	282	6.8	276	20	276	110	203	274	1.0542	0.8974	
MPP-3A_79	0.3001	0.0092	0.0438	0.0008	0.05155	0.00099	0.01257	0.00095	266.4	7.2	267.5	5.4	252.4	24	270	110	203	274	1.0542	0.8974	
MPP-3A_80	0.338	0.0094	0.0438	0.001	0.05474	0.001	0.0215	0.0015	287.5	7.2	277.2	5.6	276	20	276	110	203	274	1.0542	0.8974	
MPP-3A_81	0.316	0.017	0.0439	0.001	0.0525	0.0027	0.01448	0.0012	278	13	280	6.5	291	24	280	110	217.9	19	140	0.8377	
MPP-3A_82	0.311	0.01487	0.0437	0.001	0.0507	0.0017	0.01207	0.0011	274	12	281.1	6.4	264.4	24	276	110	216	189	114	1.0531	
MPP-3A_83	0.318	0.016	0.0452	0.0012	0.051	0.0025	0.0135	0.0015	271	12	274.4	7.6	271	24	276	110	216	189	114	1.0531	
MPP-3A_84	0.315	0.014	0.04487	0.0009	0.0523	0.002	0.014	0.0012	277.7	11	278.7	5.7	281	24	288	84	264	186.2	0.7053	0.8974	
MPP-3A_85	0.314	0.014083	0.0448	0.0008	0.0523	0.0014	0.00768	0.0011	277.2	6.4	274.1	6.8	263.3	20	270	110	203	274	1.0542	0.8974	
MPP-3A_86	0.3072	0.058	0.0465	0.0076	0.0316	0.0045	0.0038	0.0028	996	22	996	22	956	24	1025	36	254	91.1	0.3587	0.8974	
MPP-3A_87	0.3088	0.0472	0.0465	0.0076	0.0316	0.0045	0.0038	0.0028	996	22	996	22	956	24	1025	36	254	91.1	0.3587	0.8974	
MPP-3A_88	0.305	0.011	0.0451	0.001	0.051	0.0051	0.0138	0.001	270.1	8.6	270	6.2	265.6	24	265	24	1055	84	68	1.265	0.7574
MPP-3A_89	0.327	0.016	0.04607	0.0011	0.0517	0.0021	0.01353	0.0013	287	12	287.3	6.6	277	26	270	110	216	189	114	1.0531	
MPP-3A_90	0.339	0.02	0.0462	0.0012	0.0522	0.0015	0.01463	0.0011	285	11	285.8	5.8	282	24	282	84	264	186.2	0.7053	0.8974	
MPP-3A_91	0.338	0.027	0.0436	0.0013	0.0516	0.0017	0.0138	0.0018	272	21	275.3	8.2	277	35	240	190	103.5	68.2	0.6899	0.8974	
MPP-3A_92	0.338	0.027	0.0436	0.0013	0.0516	0.0017	0.0138	0.0018	272	21	275.3	8.2	277	35	240	190	103.5	68.2	0.6899	0.8974	
MPP-3A_93	0.338	0.027	0.0436	0.0013	0.0516	0.0017	0.0138	0.0018	272	21	275.3	8.2	277	35	240	190	103.5	68.2	0.6899	0.8974	
MPP-3A_94	0.338	0.027	0.0436	0.0013	0.0516	0.0017	0.0138	0.0018	272	21	275.3	8.2	277	35	240	190	103.5	68.2	0.6899	0.8974	
MPP-3A_95	0.338	0.027	0.0436	0.0013	0.0516	0.0017	0.0138	0.0018	272	21	275.3	8.2	277	35	240	190	103.5	68.2	0.6899	0.8974	
MPP-3A_96	0.338	0.027	0.0436	0.0013	0.0516	0.0017	0.0138	0.0018	272	21	275.3	8.2	277	35	240	190	103.5	68.2	0.6899	0.8974	
MPP-3A_97	0.338	0.027	0.0436	0.0013	0.0516	0.0017	0.0138	0.0018	272	21	275.3	8.2	277	35	240	190	103.5	68.2	0.6899	0.8974	
MPP-3A_98	0.338	0.027	0.0436	0.0013	0.0516	0.0017	0.0138	0.0018	272	21	275.3	8.2	277	35	240	190	103.5	68.2	0.6899	0.8974	
MPP-3A_99	0.338	0.027	0.0436	0.0013	0.0516	0.0017	0.0138	0.0018	272	21	275.3	8.2	277	35	240	190	103.5	68.2	0.6899	0.8974	
MPP-3A_100	0.303	0.014	0.04295	0.0008	0.0512	0.0019	0.01385	0.0011	268.1	11	271.1	6	274	22	249	81	364	311	0.8544	0.8974	
MPP-3A_101	0.310	0.014	0.04295	0.0008	0.0512	0.0019	0.01385	0.0011	268.1	11	271.1	6	274	22	249	81	364	311	0.8544	0.8974	
MPP-3A_102	0.310	0.014	0.04295	0.0008	0.0512	0.0019	0.01385	0.0011	268.1	11	271.1	6	274	22	249	81	364	311	0.8544	0.8974	
MPP-3A_103	0.310	0.014	0.04295	0.0008	0.0512	0.0019	0.01385	0.0011	268.1	11	271.1	6	274	22	249	81	364	311	0.8544	0.8974	
MPP-3A_104	0.310	0.014	0.04295	0.0008	0.0512	0.0019	0.01385	0.0011	268.1	11	271.1	6	274	22	249	81	364	311	0.8544	0.8974	
MPP-3A_105	0.310	0.014	0.04295	0.0008	0.0512	0.0019	0.01385	0.0011	268.1	11	271.1	6	274	22	249	81	364	311	0.8544	0.8974	
MPP-3A_106	0.310	0.014	0.04295	0.0008	0.0512	0.0019	0.01385	0.0011	268.1	11	271.1	6	274	22	249	81	364	311	0.8544	0.8974	
MPP-3A_107	0.310	0.014	0.04295	0.0008	0.0512	0.0019	0.01385	0.0011	268.1	11	271.1	6	274	22	249	81	364	311	0.8544	0.8974	
MPP-3A_108	0.310	0.014	0.04295	0.0008	0.0512	0.0019	0.01385	0.0011	268.1	11	271.1	6	274	22	249	81	364	311	0.8544	0.8974	
MPP-3A_109	0.310	0.014	0.04295	0.0008	0.0512	0.0019	0.01385	0.0011	268.1	11	271.1	6	274	22	249	81	364	311	0.8544	0.8974	
MPP-3A_110	0.310	0.014	0.04295	0.0008	0.0512	0.0019	0.01385	0.0011	268.1	11	271.1	6	274	22	249	81	364	311	0.8544	0.8974	
MPP-3A_111	0.310	0.014	0.04295	0.0008	0.0512	0.0019	0.01385	0.0011	268.1	11	271.1	6	274	22	249	81	364	311	0.8544	0.8974	
MPP-3A_112	0.310	0.014	0.04295	0.0008	0.0512	0.0019	0.01385	0.0011	268.1	11	271.1	6	274	22	249	81	364	311	0.8544	0.8974	
MPP-3A_113	0.310	0.014	0.04295	0.0008	0.0512	0.0019	0.01385	0.0011	268.1	11	271.1	6	274	22	249	81	364	311	0.8544	0.8974	
MPP-3A_114	0.310	0.014	0.04295	0.0008	0.0512	0.0019	0.01385	0.0011	268.1	11	271.1	6	274	22	249	81	364	311	0.8544	0.8974	
MPP-3A_115	0.310	0.014	0.04295	0.0008	0.0512	0.0019	0.01385	0.0011	268.1	11	271.1	6	274	22	249	81	364	311	0.8544	0.8974	
MPP-3A_116	0.310	0.014	0.04295	0.0008	0.0512	0.0019	0.01385	0.0011	268.1	11	271.1	6	274	22	249	81	364	311	0.8544	0.8974	
MPP-3A_117	0.310	0.014	0.04295	0.0008	0.0512	0.0019	0.01385	0.0011	268.1	11	271.1	6	274	22	249	81	364	311	0.8544	0.8974	
MPP-3A_118	0.310	0.014	0.04295	0.0008	0.0512	0.0019	0.01385	0.0011	268.1	11	271.1	6	274	22	249	81	364	311	0.8544	0.8974	
MPP-3A_119	0.310	0.014	0.04295	0.0008	0.0512	0.0019	0.01385	0.0011	268.1	11	271.1	6	274	22	249	81	364	311	0.8544	0.8974	
MPP-3A_120	0.310	0.014	0.04295	0.0008	0.0512	0.0019	0.01385	0.0011	268.1	11	271.1	6	274	22	249	81	364	311	0.8544	0.8974	
MPP-3A_121	0.310	0.014	0.04295	0.0008	0.0512	0.0019	0.01385	0.0011	268.1	11	271.1	6	274	22	249	81	364	311	0.8544	0.8974	
MPP-3A_122	0.310	0.014	0.04295	0.0008	0.0512	0.0019	0.01385	0.0011	268.1	11	271.1	6	274	22	249	81	364	311	0.8544	0.8974	
MPP-3A_123	0.310	0.014	0.04295	0.0008	0.0512	0.0019	0.01385	0.0011	268.1	11	271.1	6	274	22	249	81	364	311	0.8544	0.8974	
MPP-3A_124	0.310	0.014	0.04295	0.0008	0.0512	0.0019	0.01385	0.0011	268.1	11	271.1	6	274	22	249	81	364	311	0.8544	0.8974	
MPP-3A_125	0.310	0.014	0.04295	0.0008	0.0512	0.0019	0.01385	0.0011	268.1	11	271.1	6	274	22	249	81	364	311	0.8544	0.8974	
MPP-3A_126	0.310	0.014	0.04295	0.0008	0.0512	0.0019	0.01385	0.0011	268.1	11	271.1	6	274	22	249	81	364	311	0.8544	0.8974	
MPP-3A_127	0.310	0.014	0.04295	0.0008	0.0512	0.0019	0.01385	0.0011	268.1	11	271.1	6	274	22	249	81	364	311	0.8544	0.8974	
MPP-3A_128	0.310	0.014	0.04295	0.0008	0.0512	0.0019	0.01385	0.0011	268.1	11	271.1	6	274	22	249	81	364	311	0.8544	0.8974	
MPP-3A_129	0.310	0.014	0.04295	0.0008	0.0512	0.0019	0.01385	0.0011	268.1	11	271.1	6	274	22	249	81	364	311	0.8544	0.8974	
MPP-3A_130	0.310	0.014	0.04295	0.0008	0.0512	0.0019	0.01385	0.0011	268.1	11	271.1	6	274	22	249	81	364	311	0.8544	0.8974	
MPP-3A_131	0.310	0.014	0.04295	0.0008	0.0512	0.0019	0.01385	0.0011	268.1	11	271.1	6	274	22	249	81	364	311	0.8544	0.8974	

MG063-132	1.467	0.045	0.1496	0.0028	0.0713	0.0022	0.0413	0.0067	916	24	7	899	16	817	150	969	63	156.3	43.89	0.2782			
MG063-133	1.444	0.059	0.1507	0.0028	0.0713	0.0022	0.0413	0.0067	916	24	7	899	16	817	150	969	63	156.3	43.89	0.2782			
MG063-134	1.517	0.061	0.1566	0.0023	0.0712	0.0028	0.0402	0.0072	942	24	24	936	13	873	140	940	81	177.7	23.21	0.4023			
MG063-135	1.759	0.062	0.1606	0.0022	0.0713	0.0024	0.0408	0.0066	918	20	20	906	12	840	130	902	75	163.2	21.74	0.3398			
MG063-136	1.523	0.049	0.1544	0.0022	0.0713	0.0024	0.0408	0.0066	918	20	20	906	12	840	130	902	75	163.2	21.74	0.3398			
MG063-137	1.669	0.055	0.1517	0.0025	0.0794	0.0024	0.0404	0.0076	994	24	24	1010	14	937	150	1170	58	82.6	37.9	0.4588			
MG063-138	1.901	0.051	0.1521	0.0022	0.0713	0.0024	0.0408	0.0066	918	20	20	906	12	840	130	902	75	163.2	21.74	0.3398			
MG063-139	1.194	0.041	0.1459	0.0026	0.0747	0.0011	0.0487	0.0081	797	14	14	750	17	707	150	1070	28	170	96.8	0.1698			
MG063-140	1.332	0.045	0.1495	0.0026	0.0747	0.0011	0.0487	0.0081	797	14	14	750	17	707	150	1070	28	170	96.8	0.1698			
MG063-141	1.558	0.027	0.1475	0.002	0.0787	0.0013	0.0486	0.0072	993	21	21	1017	11	947	111	1150	140	1111	34	449	124.9	0.2863	
MG063-142	1.421	0.041	0.1456	0.0026	0.0747	0.0011	0.0487	0.0081	797	14	14	750	17	707	150	1070	28	170	96.8	0.1698			
MG063-143	1.815	0.0095	0.1505	0.0026	0.0747	0.0011	0.0487	0.0081	797	14	14	750	17	707	150	1070	28	170	96.8	0.1698			
MG063-144	1.631	0.079	0.1592	0.003	0.0743	0.0011	0.0489	0.0072	980	30	30	1030	17	963	170	1156	140	1039	86	175	60.9	0.3480	
MG063-145	1.446	0.075	0.1461	0.0024	0.0743	0.0011	0.0489	0.0072	980	30	30	1030	17	963	170	1156	140	1039	86	175	60.9	0.3480	
MG063-146	1.151	0.055	0.1525	0.0029	0.072	0.0025	0.0519	0.0068	941	22	22	996	16	919	130	986	69	617	13.7	1.5482	0.2166		
MG063-147	1.235	0.059	0.1539	0.0029	0.072	0.0025	0.0519	0.0068	941	22	22	996	16	919	130	986	69	617	13.7	1.5482	0.2166		
MG063-148	1.896	0.072	0.1781	0.003	0.075	0.0022	0.0489	0.0074	1019	25	25	1057	16	1020	150	1130	82	594	15.6	1.0263	0.2166		
MG063-149	2.64	0.11	0.2231	0.0037	0.098	0.0024	0.0793	0.0038	1111	30	30	1236	30	1200	170	1335	56	286.5	196	0.6075	0.2166		
MG063-150	1.89	0.059	0.1833	0.0028	0.0747	0.0013	0.0487	0.0081	797	14	14	750	17	707	150	1070	28	170	96.8	0.1698	0.2166		
MG063-151	1.34	0.041	0.1498	0.0026	0.0747	0.0011	0.0487	0.0081	797	14	14	750	17	707	150	1070	28	170	96.8	0.1698	0.2166		
MG063-152	0.3	0.0045	0.0056	0.0026	0.0747	0.0011	0.0487	0.0081	797	14	14	750	17	707	150	1070	28	170	96.8	0.1698	0.2166		
MG063-153	1.535	0.037	0.1521	0.003	0.07	0.0018	0.0533	0.0066	915	8.4	8.4	968	16	909	130	986	69	617	13.7	1.5482	0.2166		
MG063-154	1.402	0.042	0.1524	0.0029	0.0753	0.0018	0.0517	0.0062	905	12	12	952	14	902	130	986	69	617	13.7	1.5482	0.2166		
MG063-155	0.333	0.025	0.0453	0.0007	0.0528	0.0039	0.0517	0.0021	290	19	19	285.8	4.4	315	42	200	150	94	29.5	0.3138	0.2166		
MG063-156	1.506	0.088	0.1567	0.0012	0.0707	0.001	0.0503	0.0027	929	36	36	1043	100	1020	150	1200	24.22	20.29	0.8377	0.2166	0.2166		
MG063-157	0.815	0.0062	0.04867	0.00049	0.0511	0.0027	0.0527	0.0019	275	24	24	275.5	3.2	306	32	100	120	43.2	20.31	0.4628	0.2166	0.2166	
MG063-158	0.323	0.012	0.04187	0.00077	0.0559	0.0019	0.0478	0.0018	283.7	9.3	9.3	284.4	4.8	299	37	435	76	471	243.4	0.5168	0.2166	0.2166	
MG063-159	1.54	0.1	0.1512	0.004	0.0735	0.0044	0.052	0.0077	942	42	42	975	15	923	150	1030	120	100	43.2	20.31	0.4628	0.2166	0.2166
MG063-160	1.759	0.034	0.1708	0.004	0.0735	0.0044	0.052	0.0077	942	42	42	975	15	923	150	1030	120	100	43.2	20.31	0.4628	0.2166	0.2166
MG063-161	0.674	0.0089	0.04335	0.00084	0.0526	0.0026	0.0518	0.0016	271.1	6.9	6.9	282.4	3.8	284.4	48	206	37	435	76	471	243.4	0.5168	0.2166
MG063-162	1.538	0.044	0.1521	0.0036	0.0722	0.0015	0.0499	0.0064	905	16	16	913	20	866	120	902	75	163.2	21.74	0.3398	0.2166	0.2166	
MG063-163	1.406	0.042	0.1454	0.0025	0.0702	0.0021	0.0469	0.0062	890	18	18	935	24	892	120	919	69	617	13.7	1.5482	0.2166	0.2166	
MG063-164	1.439	0.034	0.1489	0.003	0.0743	0.0023	0.0469	0.0062	890	18	18	935	24	892	120	919	69	617	13.7	1.5482	0.2166	0.2166	
MG063-165	0.304	0.039	0.0451	0.0002	0.0499	0.0011	0.0459	0.002	248	15	15	274.5	5.7	293	41	210	130	124.4	57.4	0.4814	0.2166	0.2166	
MG063-166	1.531	0.069	0.1516	0.0029	0.0716	0.0011	0.0459	0.002	248	15	15	274.5	5.7	293	41	210	130	124.4	57.4	0.4814	0.2166	0.2166	
MG063-167	1.515	0.015	0.1534	0.0029	0.0717	0.0011	0.0459	0.002	248	15	15	274.5	5.7	293	41	210	130	124.4	57.4	0.4814	0.2166	0.2166	
MG063-168	1.599	0.045	0.1602	0.0023	0.0715	0.0013	0.0522	0.0027	929	36	36	1043	100	1020	150	1200	24.22	20.29	0.8377	0.2166	0.2166		
MG063-169	1.94	0.044	0.1615	0.0024	0.0704	0.0013	0.0522	0.0027	929	36	36	1043	100	1020	150	1200	24.22	20.29	0.8377	0.2166	0.2166		
MG063-170	1.492	0.037	0.1526	0.0039	0.0709	0.0018	0.0487	0.0081	797	14	14	750	17	707	150	1070	28	170	96.8	0.1698	0.2166	0.2166	
MG063-171	1.473	0.047	0.1524	0.0039	0.0709	0.0018	0.0487	0.0081	797	14	14	750	17	707	150	1070	28	170	96.8	0.1698	0.2166	0.2166	
MG063-172	0.313	0.0097	0.0448	0.0006	0.0509	0.0014	0.0436	0.0018	276.2	7.5	7.5	280.5	4.3	270	37	239	59	446	226	0.9067	0.2166	0.2166	
MG063-173	0.323	0.012	0.04507	0.00066	0.0521	0.0021	0.0436	0.0022	284	8.9	8.9	284.2	4.1	292	47	239	59	446	226	0.9067	0.2166	0.2166	

LRW-21

LRW_21_1	0.579	0.0078	0.0394	0.00073	0.0507	0.0011	0.0123	0.0087	23.9	6.3	6.3	23.8	4.5	22.8	17	23	48	1325	305	0.2320
LRW_21_2	0.821	0.0077	0.0394	0.00073	0.0507	0.0011	0.0123	0.0087	23.9	6.3	6.3	23.8	4.5	22.8	17	23	48	1325	305	0.2320
LRW_21_3	0.561	0.01	0.03603	0.00078	0.0517	0.0018	0.0129	0.0081	21.1	6.4	6.4	22.8	4.6	22.7	22	28	75	448	103.8	0.2317
LRW_21_4	0.262	0.01	0.027	0.00073	0.0514	0.0017	0.0117	0.0081	21.62	8.2	8.2	22.42	4.6	23.5	22	249	60	479	108.6	0.2287
LRW_21_5	0.277	0.0087	0.0286	0.00078	0.0514	0.0017	0.0117	0.0081	21.62	8.2	8.2	22.42	4.6	23.5	22	249	60	479	108.6	0.2287
LRW_21_6	0.262	0.01	0.027	0.00073	0.0514	0.0017	0.0117	0.0081	21.62	8.2	8.2	22.42	4.6	23.5	22	249	60	479	108.6	0.2287
LRW_21_7	0.262	0.01	0.027	0.00078	0.0514	0.0017	0.0117	0.0081	21.62	8.2	8.2	22.42	4.6	23.5	22	249	60	479	108.6	0.2287
LRW_21_8	0.262	0.01	0.027	0.00078	0.0514	0.0017	0.0117	0.0081	21.62	8.2	8.2	22.42	4.6	23.5	22	249	60	479	108.6	0.2287
LRW_21_9	0.262	0.01	0.027	0.00078	0.0514	0.0017	0.0117	0.0081	21.62	8.2	8.2	22.42	4.6	23.5	22	249	60	479	108.6	0.2287
LRW_21_10	0.262	0.01	0.027	0.00078	0.0514	0.0017	0.0117	0.0081	21.62	8.2	8.2	22.42	4.6	23.5	22	249	60	479	108.6	0.2287
LRW_21_11	0.262	0.01	0.027	0.00078	0.0514	0.0017	0.0117	0.0081	21.62	8.2	8.2	22.42	4.6	23.5	22	249	60	479	108.6	0.2287
LRW_21_12	0.262	0.01	0.027	0.00078	0.0514	0.0017	0.0117	0.0081	21.62	8.2	8.2	22.42	4.6	23.5	22	249	60	479	108.6	0.2287
LRW_21_13	0.262	0.01	0.027	0.00078	0.0514	0.0017	0.0117	0.0081	21.62	8.2	8.2	22.42	4.6	23.5	22	249	60	479	108.6	0.2287
LRW_21_14	0.262	0.01	0.027	0.00078	0.0514	0.0017	0.0117	0.0081	21.62	8.2	8.2	22.42	4.6	23.5	22	249	60	479	108.6	0.2287
LRW_21_15	0.262	0.01	0.027	0.00078	0.0514	0.0017	0.0117	0.0081	21.62	8.2	8.2	22.42	4.6	23.5	22	249	60	479	108.6	0.2287
LRW_21_16	0.262	0.01	0.027	0.00078	0.0514	0.0017	0.0117	0.0081	21.62	8.2	8.2	22.42	4.6	23.5	22	249	60	479</		

108- 83	1.666	0.064	0.1884	0.0038	0.0756	0.0022	0.0203	0.0045	1994	24	1009	21	991	87	962	60	684	37.4	0.6468
108- 84	1.039	0.044	0.1864	0.0044	0.0752	0.0022	0.0203	0.0045	1994	24	1009	21	991	87	962	60	684	37.4	0.6468
108- 85	0.2612	0.0002	0.03882	0.00069	0.0153	0.0012	0.01213	0.0015	2315	74	2331	4.3	244	23	262	59	444	67.2	0.1514
108- 86	0.1701	0.0002	0.03882	0.00069	0.0153	0.0012	0.01213	0.0015	2315	74	2331	4.3	244	23	262	59	444	67.2	0.1514
108- 87	0.2706	0.011	0.03798	0.00083	0.01511	0.0015	0.01212	0.0015	242.9	9	240.3	5.1	245	29	255	72	252	34.7	0.1937
108- 88	0.264	0.01	0.03738	0.00088	0.01511	0.0016	0.01216	0.0016	237	12	236.6	5.4	233	31	230	110	189.2	21.39	0.1731
108- 89	0.1534	0.001	0.03812	0.001	0.01506	0.0012	0.01216	0.0012	231.1	8.1	230	5.2	238	28	242	49	207	40.1	0.1920
108- 90	0.2652	0.0089	0.03738	0.00077	0.01508	0.001	0.01185	0.00099	238.7	7.1	238.8	4.4	238	30	227	46	117	26.2	0.2346
108- 91	0.1991	0.001	0.03703	0.00077	0.01507	0.001	0.01187	0.001	233.7	7.1	233.7	4.4	238	31	227	46	117	26.2	0.2346
108- 92	0.27	0.018	0.03727	0.0012	0.01517	0.0011	0.0117	0.002	239	14	235.9	4.7	235	40	250	120	186.7	22	0.1178
108- 94	0.247	0.0047	0.03707	0.00105	0.01514	0.0011	0.0117	0.002	246	14	243.3	7.0	246	40	250	120	186.7	22	0.1178
108- 94	0.242	0.014	0.0363	0.00062	0.01483	0.0007	0.01223	0.0007	219	11	219	5.1	246	110	120	110	186.7	22	0.1178
108- 95	0.236	0.015	0.03647	0.00077	0.01508	0.001	0.01216	0.001	238.7	7.1	238.8	4.4	238	30	227	46	117	26.2	0.2346
108- 96	0.1484	0.004	0.0363	0.00074	0.01514	0.0011	0.0117	0.002	239	14	235.9	4.7	235	40	250	120	186.7	22	0.1178
108- 97	0.2605	0.01	0.03652	0.00077	0.01514	0.0011	0.0117	0.002	239	14	235.9	4.7	235	40	250	120	186.7	22	0.1178
108- 98	0.1549	0.0004	0.03673	0.00074	0.01514	0.0011	0.0117	0.002	239	14	235.9	4.7	235	40	250	120	186.7	22	0.1178
108- 99	0.261	0.013	0.03761	0.00093	0.01504	0.0023	0.011	0.0013	215.3	10	228	5.8	220	26	202	30	474	37.7	0.1389
108- 100	0.2599	0.01	0.03676	0.00082	0.01507	0.0015	0.01189	0.0015	214.3	9	227	5.1	219	25	220	26	441	76.8	0.1741
108- 101	0.263	0.012	0.03752	0.00084	0.01507	0.0016	0.01189	0.0016	217.1	9	227.4	5.1	219	25	220	26	441	76.8	0.1741
108- 102	0.283	0.012	0.03886	0.00074	0.01516	0.002	0.01156	0.002	236.6	9.8	234	4.6	232	24	255	83	307	46.7	0.1415
108- 103	0.263	0.015	0.03759	0.001	0.01495	0.0025	0.01196	0.0025	236	10	240.4	5.3	256	34	190	164	124	42.8	0.2010
108- 104	0.895	0.01	0.03685	0.001	0.01485	0.0021	0.01475	0.0021	246	11	246.3	6.7	247	89	282	67	155.3	37	0.2382
108- 105	0.2386	0.0007	0.03705	0.00077	0.0151	0.0016	0.01198	0.0016	233.4	7.9	234.4	4.8	241	25	228	68	457	71.7	0.1969
108- 106	0.152	0.01	0.03657	0.001	0.01512	0.0015	0.01187	0.0015	239	19	238	5.8	247	71	185	42	276	91.9	0.3330
108- 107	0.2655	0.008	0.03687	0.00081	0.01522	0.0012	0.01183	0.00099	238.9	7.9	234.5	5	237.6	59	285	53	840	194.4	0.2279
108- 108	0.2651	0.0004	0.03684	0.001	0.01519	0.0014	0.01188	0.001	238.2	7.9	242.2	5.3	229	27	295	61	138	62.0	0.2259
108- 111	0.2648	0.0003	0.03787	0.0012	0.01506	0.0012	0.01205	0.00097	238.4	7.5	239.6	4.7	242.1	19	217	21	119	23.7	0.2143
108- 114	0.28	0.012	0.03886	0.001	0.01511	0.0011	0.01154	0.0011	250.1	9.7	246.4	6.3	232	21	273	44	505	101.8	0.2004
108- 114	0.28	0.012	0.0384	0.0014	0.01489	0.0017	0.01214	0.0015	231	15	242.7	4.5	245	45	140	150	108	10.18	0.2004
108- 115	0.25	0.024	0.0381	0.0013	0.01488	0.0017	0.01208	0.0023	229	19	243.3	7.8	238	46	180	163	103	103	0.142
108- 116	0.262	0.012	0.03877	0.0009	0.01502	0.0022	0.01165	0.0022	235.8	9.6	238.6	5.6	251	31	205	52	122	37.5	0.1920
108- 117	0.263	0.012	0.03877	0.0009	0.01502	0.0022	0.01165	0.0022	235.8	9.6	238.6	5.6	251	31	205	52	122	37.5	0.1920
108- 118	0.271	0.011	0.03752	0.001	0.01523	0.0017	0.01242	0.0013	241.2	8.8	237.8	5.6	249	26	287	70	392	64.7	0.1651
108- 120	0.254	0.009	0.03614	0.00085	0.01514	0.0015	0.01151	0.0015	210.7	7.9	228.5	5.3	231	36	248	62	46	34.8	0.2546
108- 121	0.262	0.01	0.0384	0.00081	0.01513	0.0016	0.01151	0.0016	212.2	8.8	232.7	5.6	237	45	245	62	46	34.8	0.2546
108- 122	1.8	0.11	0.1871	0.0035	0.01769	0.0036	0.0133	0.0038	1104	36	1105	27	1048	150	1090	90	21.82	21.15	0.9033
108- 123	0.876	0.01	0.03726	0.0008	0.01512	0.0012	0.01165	0.0012	235.5	9.5	235.5	4.5	238	59	278	59	494	75.4	0.1607
108- 124	0.263	0.014	0.03749	0.00094	0.01508	0.0018	0.01108	0.0013	236	12	237.2	5.8	217	26	220	100	174.4	38.2	0.2190

ABH-27

ABH-27- 1	0.474	0.045	0.0625	0.0058	0.0549	0.0012	0.0165	0.0034	392	32	390	35	391	67	405	49	4420	8.10E+03	1.8326
ABH-27- 2	0.676	0.012	0.0733	0.0023	0.0571	0.001	0.0202	0.0032	461.4	76	466	7	505	64	498	36	650	645	0.2923
ABH-27- 3	1.959	0.038	0.0614	0.0022	0.0563	0.0014	0.0177	0.0022	404	102	407	110	412	108	407	65	167	17.6	0.1778
ABH-27- 4	0.586	0.014	0.07307	0.00075	0.0566	0.0014	0.01549	0.0033	467.6	9.2	466.6	4.5	509	64	465	52	329	149	0.4529
ABH-27- 5	0.445	0.013	0.057	0.0045	0.0517	0.0045	0.0132	0.0045	455.7	9.8	390.5	4.5	57	379	439	39	2208	200	0.1118
ABH-27- 6	0.488	0.01	0.07049	0.00109	0.0558	0.0028	0.01088	0.0025	405.5	11	394.4	4.5	613	46	505	278	0.78	0.2943	
ABH-27- 7	0.569	0.017	0.0721	0.0012	0.0574	0.0017	0.0206	0.0043	457	11	448.9	7.2	529	64	407	65	193.6	18.13	0.0847
ABH-27- 8	0.784	0.016	0.0761	0.0023	0.0561	0.0025	0.0138	0.0025	408	11	408	6.7	494	71	407	65	193.6	18.13	0.0847
ABH-27- 9	0.544	0.017	0.0703	0.00099	0.0564	0.0019	0.0248	0.0036	441	11	448	6	494	71	407	65	193.6	18.13	0.0847
ABH-27- 10	2.074	0.017	0.074	0.0019	0.0567	0.0019	0.0248	0.0036	441	11	448	6	494	71	407	65	193.6	18.13	0.0847
ABH-27- 11	0.471	0.016	0.0556	0.0014	0.0556	0.0019	0.0232	0.003	391	11	385.4	8.7	464	60	421	78	177.8	115.2	0.6479
ABH-27- 12	0.549	0.016	0.0688	0.0012	0.0572	0.0016	0.0245	0.0031	444	10	443.9	7.5	460	60	421	78	177.8	115.2	0.6479
ABH-27- 13	0.4013	0.01	0.05817	0.0006	0.05067	0.0007	0.0199	0.0007	449	10	449	7.5	460	60	421	78	177.8	115.2	0.6479
ABH-27- 14	0.267	0.005	0.03942	0.00039	0.05099	0.00082	0.0126	0.0022	240	14	243.3	7.2	253	44	219	37	184	22.91	0.1024
ABH-27- 15	0.431	0.017	0.0518	0.0017	0.0542	0.0015	0.0173	0.0028	413	11	413	7.2	430	43	430	43	130	16.6	0.1378
ABH-27- 16	0.574	0.017	0.07452	0.00088	0.0563	0.0017	0.0236	0.0031	462	12	463.4	5.3	475	61	405	65	189.5	82.6	0.4359
ABH-27- 17	0.513	0.01	0.07076	0.001	0.0561	0.0017	0.0236	0.0031	462	12	463.4	5.3	475	61	405	65	189.5	82.6	0.4359
ABH-27- 18	2.258	0.052	0.0326	0.0035	0.0181	0.0013	0.0257	0.0038	1198	16	1189	19	1227	160	1219	32	134	41.69	0.3111
ABH-27- 19	0.577	0.012	0.0572	0.00097	0.0559	0.0012	0.0243	0.0034	462	8.1	462.9	5.8	485	68	479	68	292	45.1	0.1546
ABH-27- 20	0.532	0.011	0.0577	0.0012	0.0559	0.0012	0.0243	0.0034	462	8.1	462.9	5.8	485	68	479	68	292	45.1	0.1546
ABH-27- 21	0.7479	0.0096	0.0741	0.0011	0.05645	0.00085	0.0225	0.0041	461	6.2	460.6	6.8	489	61	466	49	33	81.3	0.0990
ABH-27- 22	1.523	0.016	0.0585	0.0012	0.0565	0.0012	0.0246	0.0031	462	8.1	462.9	5.8	485	68	479	68	292	45.1	0.1546
ABH-27- 23	0.784	0.014	0.0585	0.0017	0.0564	0.0017	0.0246	0.0031	462	8.1	462.9	5.8	485	68	479	68	292	45.1	0.1546
ABH-27- 24	0.542	0.015	0.0562	0.00095	0.0562	0.001	0.0246	0.0031	462	8.1	462.9	5.8	485	68	479	68	292	45.1	0.1546
ABH-27- 25	0.246	0.01	0.0565	0.001	0.														

Ak-01-8	0.211	0.025	0.02764	0.00908	0.055	0.0061	0.00904	0.0011	192	16	197.7	6.1	182	22	370	230	334	34.6	1.0399
Ak-01-10	0.165	0.053	0.1555	0.032	0.0728	0.0013	0.05	0.0042	994	20	184.1	5.4	184	20	1004	36	139.2	30.5	0.2191
Ak-01-12	0.192	0.018	0.02089	0.0089	0.0483	0.0045	0.0094	0.0014	177	15	180.7	4.8	188	27	90	170	56.4	30.8	0.5481
Ak-01-13	0.203	0.036	0.02226	0.0026	0.0524	0.0009	0.00876	0.0011	187	19	193.6	4.8	176	22	260	150	65	50.9	0.7831
Ak-01-14	0.201	0.035	0.02056	0.0055	0.0495	0.0032	0.0091	0.0011	184	18	184.1	4.7	187	18	110	112	64.2	36.4	0.7952
Ak-01-15	0.233	0.033	0.02096	0.00092	0.0563	0.0075	0.0104	0.0014	209	27	197.0	5.7	209	29	340	250	349	24.6	0.7049
Ak-01-16	0.209	0.022	0.02044	0.00081	0.0522	0.0033	0.0095	0.0011	191	15	187.1	4.9	185	20	160	160	44.4	33.1	0.7500
Ak-01-17	0.2	0.015	0.03011	0.0088	0.0485	0.0037	0.0015	0.0094	184	13	191.2	5.5	184	13	120	150	66.2	97.2	1.4883
Ak-01-18	0.21	0.01	0.03044	0.00077	0.0511	0.0031	0.0088	0.0006	191	24	193.3	6.1	185	25	180	200	60.2	91.7	0.7166
Ak-01-19	0.199	0.019	0.1441	0.0051	0.0784	0.0016	0.0589	0.0046	951	26	187	29	155	40	115	40	116.4	10.8	0.7808
Ak-01-20	0.221	0.074	0.0216	0.0048	0.0801	0.0051	0.0851	0.001	1186	29	1183	26	1274	95	1194	42	105.6	39.9	0.3778
Ak-01-21	0.225	0.034	0.01304	0.00074	0.0542	0.0042	0.0047	0.001	198	32	186.5	5.3	193	32	98	145	20.3	0.584	0.7808
Ak-01-22	0.209	0.02	0.02943	0.00092	0.0519	0.0049	0.0081	0.00099	191	17	184.9	5.7	177	20	230	180	48	55	1.1317
Ak-01-23	0.206	0.027	0.02032	0.00076	0.0532	0.0042	0.0095	0.001	192	20	185.3	4.9	186	20	160	160	55.8	42.6	0.7638
Ak-01-24	0.199	0.016	0.02005	0.0007	0.0507	0.0042	0.00975	0.001	183	13	185.2	4.7	186	21	200	160	17	68.2	1.1905
Ak-01-25	0.208	0.034	0.02938	0.00072	0.0511	0.0056	0.009	0.0012	190	20	186.2	4.5	181	24	130	200	42	37.6	0.8922
Ak-01-26	0.207	0.021	0.02021	0.00071	0.0521	0.0087	0.009	0.001	189	18	184.8	4.5	184	18	130	200	38.2	38.2	0.842
Ak-01-27	0.21	0.023	0.02022	0.0011	0.0527	0.0056	0.00921	0.0011	192	19	185.6	6.8	185	21	240	200	37.5	34.9	0.9397
Ak-01-28	0.209	0.027	0.02028	0.0013	0.0525	0.0052	0.0095	0.0012	192	26	185.3	6.2	193	28	230	230	45.4	40.1	0.9744
Ak-01-29	0.2	0.02	0.0204	0.0062	0.0544	0.0015	0.0561	0.0053	1305	26	1313	32	1294	100	1299	34	394	94.6	0.2401
Ak-01-30	0.198	0.016	0.02032	0.00076	0.0532	0.0049	0.0097	0.001	189	18	185.2	4.5	186	20	160	160	55.2	46.5	0.8222
Ak-01-31	0.198	0.016	0.02034	0.00075	0.053	0.0049	0.00975	0.001	181	29	184	7	188	30	50	300	36.4	31.4	0.9624
Ak-01-32	0.195	0.015	0.02053	0.00081	0.0484	0.0037	0.00952	0.00092	180	12	187.6	5	182	18	130	180	55.5	66.9	1.2054
Ak-01-33	0.188	0.01	0.02027	0.00071	0.0512	0.0046	0.009	0.001	190	20	186.2	4.5	186	20	160	160	55.2	46.5	0.8222
Ak-01-34	0.197	0.023	0.02022	0.00089	0.0506	0.0062	0.01063	0.0011	278	17	185.6	5.6	214	22	1100	160	61.2	67.1	1.0964
Ak-01-35	0.183	0.019	0.02027	0.00076	0.0512	0.0049	0.009	0.001	190	20	177.8	7.2	188	22	400	300	50.9	67.9	1.1268
Ak-01-36	0.183	0.019	0.02027	0.00076	0.0512	0.0049	0.009	0.001	190	20	177.8	7.2	188	22	400	300	50.9	67.9	1.1268
Ak-01-37	0.183	0.019	0.02027	0.00076	0.0512	0.0049	0.009	0.001	190	20	177.8	7.2	188	22	400	300	50.9	67.9	1.1268
Ak-01-38	0.183	0.019	0.02027	0.00076	0.0512	0.0049	0.009	0.001	190	20	177.8	7.2	188	22	400	300	50.9	67.9	1.1268
Ak-01-39	0.183	0.019	0.02027	0.00076	0.0512	0.0049	0.009	0.001	190	20	177.8	7.2	188	22	400	300	50.9	67.9	1.1268
Ak-01-40	0.183	0.019	0.02027	0.00076	0.0512	0.0049	0.009	0.001	190	20	177.8	7.2	188	22	400	300	50.9	67.9	1.1268
Ak-01-41	0.222	0.027	0.02021	0.0012	0.056	0.007	0.0094	0.0014	202	22	185	7.3	188	27	350	240	34.4	27.2	0.7473
Ak-01-42	0.191	0.027	0.02027	0.00079	0.0512	0.0049	0.009	0.001	190	20	177.8	7.2	188	22	400	300	50.9	67.9	1.1268
Ak-01-43	0.1616	0.046	0.1525	0.0033	0.07223	0.00885	0.0479	0.0034	976	17	190	18	946	66	90	24	164	270.8	0.4978
Ak-01-44	0.179	0.027	0.02937	0.0011	0.0544	0.0005	0.0089	0.0012	200	29	186.6	6.8	179	24	200	210	47.3	42.2	0.8922
Ak-01-45	0.199	0.018	0.1812	0.0048	0.0512	0.0027	0.0512	0.0041	199	20	191.2	5.7	197	19	110	110	137.2	66.4	0.7862
Ak-01-46	0.165	0.037	0.1703	0.0036	0.0738	0.0016	0.0482	0.0037	1018	20	1018	20	1028	43	1028	43	1028	43	0.4542
Ak-01-47	0.245	0.045	0.2088	0.0045	0.0644	0.0015	0.045	0.0045	1200	20	1200	20	1200	20	1200	20	1200	20	0.4542
Ak-01-48	0.221	0.024	0.02933	0.00091	0.0552	0.0056	0.01026	0.0011	201	20	188.1	5.7	200	21	340	200	47.6	52.1	1.0946
Ak-01-49	0.183	0.026	0.02938	0.001	0.0543	0.0056	0.01026	0.0011	201	20	188.1	5.7	200	21	340	200	47.6	52.1	1.0946
Ak-01-50	0.213	0.012	0.02938	0.001	0.0543	0.0056	0.01026	0.0011	201	20	188.1	5.7	200	21	340	200	47.6	52.1	1.0946
Ak-01-51	0.165	0.065	0.1681	0.0035	0.0716	0.0027	0.0487	0.004	969	28	969	28	969	28	969	28	969	28	0.5283
Ak-01-52	0.189	0.066	0.1909	0.0044	0.062	0.0042	0.0487	0.004	1170	28	1170	28	1170	28	1170	28	1170	28	0.5283
Ak-01-53	0.189	0.066	0.1909	0.0044	0.062	0.0042	0.0487	0.004	1170	28	1170	28	1170	28	1170	28	1170	28	0.5283
Ak-01-54	0.189	0.066	0.1909	0.0044	0.062	0.0042	0.0487	0.004	1170	28	1170	28	1170	28	1170	28	1170	28	0.5283
Ak-01-55	0.1742	0.076	0.1708	0.0038	0.0741	0.0025	0.0516	0.0048	1001	28	1001	28	1001	28	1001	28	1001	28	0.5283
Ak-01-56	0.205	0.021	0.02978	0.0009	0.0497	0.0004	0.0067	0.0007	188	18	189.6	5.3	173	20	130	180	54.9	42.2	0.7802
Ak-01-57	0.205	0.021	0.02978	0.0009	0.0497	0.0004	0.0067	0.0007	188	18	189.6	5.3	173	20	130	180	54.9	42.2	0.7802
Ak-01-58	0.2	0.024	0.02928	0.00086	0.0544	0.0054	0.0086	0.00099	200	20	188	5.4	188	20	300	200	42.9	37.1	0.8648
Ak-01-59	0.149	0.051	0.1465	0.0075	0.0515	0.005	0.0489	0.008	985	20	985	20	985	20	985	20	985	20	0.2483
Ak-01-60	0.169	0.05	0.142	0.0034	0.0732	0.0017	0.0475	0.0035	981	19	986	19	987	68	1011	49	124.5	156.6	0.8974
Ak-01-61	0.178	0.054	0.1498	0.0075	0.0515	0.005	0.0489	0.008	985	20	985	20	985	20	985	20	985	20	0.2483
Ak-01-62	0.178	0.054	0.1498	0.0075	0.0515	0.005	0.0489	0.008	985	20	985	20	985	20	985	20	985	20	0.2483
Ak-01-63	0.209	0.035	0.02939	0.0007	0.0534	0.0079	0.0088	0.0013	191	28	192.2	6.1	177	20	270	270	55.9	9.1	1.0594
Ak-01-64	0.199	0.036	0.02939	0.0007	0.0534	0.0079	0.0088	0.0013	191	28	192.2	6.1	177	20	270	270	55.9	9.1	1.0594
Ak-01-65	0.219	0.031	0.02988	0.0011	0.0561	0.009	0.0099	0.0015	199	26	193.3	6.7	188	29	300	300	32.3	35.3	0.9005
Ak-01-66	0.164	0.033	0.02933	0.001	0.0552	0.0073	0.0045	0.0039	189	18	189.6	5.3	186	18	105	105	65.6	65.6	0.3598
Ak-01-67	0.233	0.027	0.02969	0.0049	0.0826	0.0046	0.0602	0.0045	1227	23	1232	26	118	27	1256	38	245.8	68.2	0.2775
Ak-01-68	0.225	0.029	0.02969	0.0049	0.0826	0.0046	0.0602	0.0045	1227	23	1232	26	118	27	1256	38	245.8	68.2	0.2775
Ak-01-69	0.233	0.027	0.02969	0.0049	0.0826	0.0046	0.0602	0.0045	1227	23	1232	26	118	27	1256	38	245.8	68.2	0.2775
Ak-01-70	0.233	0.027	0.02969	0.0049	0.0826	0.0046	0.0602	0.0045	1227	23	1232	26	118	27	1256	38	245.8	68.2	0.2775
Ak-01-71	0.233	0.027	0.02969	0.0049	0.0826	0.0046	0.0602	0.0045	1227	23	1232	26	118	27	1256	38	245.8	68.2	0.2775
Ak-01-72	0.217	0.023	0.03914	0.00086	0.0527	0.0058	0.0091	0.0013	197	20	191.4	5.4	183	26	250	210	41.1	33.2	0.8078
Ak-01-73	0.169	0.014	0.1714	0															

CV-1344-54	0.0714	0.0049	0.01072	0.00026	0.0484	0.0037	0.00291	0.00057	69.9	4.6	68.8	1.6	59	11	110	140	227	36.8	0.1621	
CV-1344-55	0.0739	0.0049	0.01072	0.00026	0.0484	0.0037	0.00291	0.00057	72.3	4.8	68.8	1.6	61.3	12	100	140	496	150.3	0.2345	
CV-1344-56	0.0696	0.0027	0.01562	0.00014	0.0476	0.0039	0.00047	0.00047	68.3	2.6	68.13	0.9	78.7	9.4	84	74	948	287	0.3027	
CV-1344-57	0.0735	0.0057	0.01511	0.00027	0.0485	0.0041	0.00433	0.00046	71.0	5.4	71.2	1.7	57	17	110	160	118	21.5	0.1446	
CV-1344-58	0.0598	0.0036	0.02011	0.00023	0.047	0.0038	0.00413	0.00069	68.4	3.4	70.5	1.4	83	14	110	302	302	57.9	0.1917	
CV-1367																				
CV-1367-1	0.0731	0.0043	0.01104	0.00017	0.0481	0.003	0.00114	0.00063	71.6	4.9	70.8	1.1	63.4	11	100	130	310	60.3	0.1945	
CV-1367-2	0.0662	0.0046	0.01569	0.00017	0.0454	0.0032	0.00139	0.00039	65	4.4	69.9	1.2	68.5	13	10	100	257.7	77	0.2988	
CV-1367-3	0.0518	0.0064	0.0088	0.00024	0.044	0.0052	0.00211	0.00051	52.6	3.04	61.3	1.49	44.5	40	40	200	368	43.1	0.4790	
CV-1367-4	0.0551	0.0069	0.00843	0.00023	0.0478	0.0065	0.00214	0.00045	54.1	6.5	54.1	1.5	43.2	40	40	240	134.9	97.1	0.7198	
CV-1367-5	0.0055	0.0088	0.00088	0.00023	0.0508	0.0049	0.00247	0.00064	61.1	5.2	57.7	1.5	49.9	11	100	180	215	39.3	0.1672	
CV-1367-6	0.0421	0.0081	0.00861	0.00051	0.0519	0.0061	0.00433	0.00063	61.8	9.2	58.3	5.09	50.9	100	160	242	322	90.3	0.2054	
CV-1367-7	0.058	0.0027	0.00852	0.00011	0.0492	0.0021	0.00228	0.00019	57.2	2.6	54.7	0.73	45.9	7.9	152	88	1196	484	0.4047	
CV-1367-8	0.0603	0.0048	0.00808	0.00022	0.0481	0.0017	0.00223	0.0004	59.4	3.9	58.3	1.1	44.9	108	158	158	133	696	0.1444	
CV-1367-9	0.094	0.016	0.01446	0.00083	0.0465	0.0045	0.0015	0.0015	91	15	92.5	9.1	91	41	40	260	288	141	0.4896	
CV-1367-10	0.0512	0.0045	0.00803	0.00019	0.0464	0.0042	0.00188	0.00036	50.6	4.3	51.6	1.2	38	7.2	30	160	160	162	0.1225	
CV-1367-11	0.9	0.078	0.1039	0.003	0.0612	0.0064	0.0052	0.004	646	437	637	2.7	527	602	210	243	2219	243	0.2824	
CV-1367-12	0.0575	0.0063	0.00863	0.00017	0.0483	0.0026	0.00215	0.00063	56.7	3.1	55.4	1.1	43.4	11	110	110	421	91	0.2162	
CV-1367-13	0.0628	0.0098	0.00868	0.00019	0.0462	0.0043	0.0025	0.00083	61.8	4.5	55.7	1.5	47	17	260	160	242	61.5	0.1798	
CV-1367-14	0.433	0.028	0.0532	0.0027	0.0614	0.0016	0.018	0.0034	377	19	334	16	359	67	643	55	438	46.2	0.1055	
CV-1367-15	0.0565	0.0062	0.00803	0.00043	0.055	0.0046	0.00264	0.00064	67.1	7.7	58	2.8	52.4	13	390	180	1013	378	0.3731	
CV-1367-16	0.139	0.011	0.01943	0.0017	0.0524	0.0017	0.0168	0.001	131.6	9.4	134	9.4	21	294	71	402.2	58	58	0.1442	
CV-1367-17	0.359	0.022	0.0485	0.001	0.0528	0.0025	0.01545	0.0026	310	16	305.1	6.4	310	52	330	110	146.7	75.3	0.1513	
AP-038																				
AP-038-1	0.033	0.036	0.00899	0.00043	0.0519	0.005	0.02	0.035	26	35	37.7	2.8	380	690	690	410	2750	71.3	0.0259	
AP-038-2	0.037	0.028	0.00894	0.00052	0.029	0.027	0.044	0.022	40	31	61.8	3.3	840	420	840	280	7500	630	0.0812	
AP-038-3	0.027	0.035	0.00962	0.00037	0.0323	0.027	0.056	0.041	29	61	61.7	2.3	1030	170	1700	400	9000	75	0.0277	
AP-038-4	0.0384	0.0091	0.00762	0.00013	0.0339	0.0085	0.0005	0.00098	35.8	9.1	48.96	0.86	90	200	800	220	5000	164.6	0.0329	
AP-038-5	0.026	0.0077	0.00829	0.00021	0.046	0.0065	0.0011	0.0028	51.9	7.4	53.2	1.5	50	670	200	700	200	700	0.0268	
AP-038-6	0.0503	0.0043	0.00796	0.00012	0.0458	0.0038	0.0005	0.00035	49.7	4.2	51.12	0.75	9	71	400	110	660	380	0.0571	
AP-038-7	0.0507	0.0028	0.0078	0.00014	0.0472	0.0024	0.0027	0.0047	50.2	2.7	50.06	0.88	55	250	300	100	11470	243	0.0212	
AP-038-8	0.0482	0.0035	0.007613	0.00013	0.0457	0.0033	0.0014	0.00034	47.7	1.4	49.02	0.48	27	61	410	110	930	246	0.0203	
AP-038-9	0.0496	0.0039	0.00799	0.00012	0.0449	0.0033	0	0.0008	49.1	3.7	51.3	1.3	0	76	300	100	9470	312	0.0329	
AP-038-10	0.031	0.048	0.0101	0.0011	0.031	0.026	0.04	0.04	23	41	49.1	7.1	1300	1300	410	410	4920	120.3	0.0245	
AP-038-11	0.0435	0.0038	0.00796	0.00019	0.0398	0.0034	0.00042	0.00042	41.2	3.8	51.1	1.2	85	130	84	880	208	0.0296	0.0296	
AP-038-12	0.042	0.0049	0.00736	0.00013	0.0434	0.0044	0.0006	0.00042	41.8	4.8	47.27	1.2	85	30	100	100	11140	309	0.0277	
AP-038-13	0.052	0.011	0.00822	0.00031	0.0452	0.0091	0.0027	0.006	51	11	52.7	2	53	210	660	270	7540	353	0.0468	
AP-038-14	0.0498	0.0068	0.00789	0.00016	0.0457	0.0058	0.0015	0.011	49.2	6.4	50.7	1	30	310	180	180	5480	114	0.0228	
AP-038-15	0.0426	0.0048	0.00755	0.00025	0.0447	0.0045	0.0002	0.0048	47.3	4.7	48.4	1.6	4	98	410	190	1351+04	970	0.0524	
AP-038-16	0.0474	0.0057	0.00748	0.00016	0.0459	0.0052	0.00016	0.0088	47	5.6	48	1	30	180	280	110	6540	95.6	0.0146	
AP-038-17	0.0471	0.0028	0.00705	0.00019	0.0451	0.004	0.0004	0.0049	46.7	0.49	48.48	0.64	98	700	700	100	7000	171	0.0222	
AP-038-18	0.0498	0.0033	0.00748	0.00013	0.0457	0.0033	0.00013	0.0054	49.3	3.1	50.74	0.8	21	180	890	187	8930	187	0.0209	
AP-038-19	0.0499	0.0039	0.00745	0.00017	0.0427	0.0034	0.0001	0.0005	43.6	3.8	47.8	1.1	3	360	150	690	242.9	0.0251		
AP-038-20	0.0511	0.0023	0.00765	0.00012	0.0483	0.002	0.00038	0.0048	50.5	2.9	49.12	0.77	77	96	238	40	8830	134.7	0.0153	
AP-038-21	0.0487	0.0051	0.00713	0.00026	0.044	0.0044	0.0005	0.0052	45.2	4.9	48.3	1.6	11	100	480	360	7400	369	0.0486	
AP-038-22	0.0464	0.0059	0.00803	0.00022	0.0416	0.0049	0.0037	0.016	45.9	5.7	45.5	1.4	80	60	490	490	4920	94.7	0.0192	
AP-038-23	1.715	0.077	0.1709	0.0022	0.0724	0.0027	0.021	0.021	1010	28	1017	12	281	1024	300	976	81	281	82.6	0.2940
AP-038-24	1.681	0.024	0.1679	0.0031	0.0728	0.0011	0.04	0.04	1007.7	9.2	1007.7	13	790	1000	400	1000	400	1355	64.1	0.0473
AP-038-25	0.049	0.0077	0.00783	0.00038	0.0463	0.0042	0.0148	0.017	48.3	7.4	50.3	2.5	290	150	900	180	1021+04	327	0.0321	
AP-038-26	0.295	0.048	0.18458	0.00377	0.0488	0.0076	0.01361	0.0034	258	40	275	4.8	69	610	610	110	5281	630	1.1907	
AP-038-27	0.296	0.067	0.0475	0.0013	0.0437	0.0014	0.0134	0.0044	247	56	298.8	7.8	269	87	750	200	295	374	1.2078	
AP-038-28	0.29	0.023	0.0423	0.00063	0.0495	0.0017	0.01264	0.0028	257	19	267.3	3.9	253.8	56	48	98	833	1281	1.5018	
AP-038-29	0.0313	0.0045	0.00802	0.00012	0.0479	0.0031	0.00165	0.00066	52.6	4.3	51.5	1.3	33	370	120	1300	1404+04	180	0.0695	
AP-038-30	0.0511	0.0031	0.008	0.00012	0.0463	0.0028	0.00019	0.00076	50.5	3.2	51.37	0.78	36	150	318	67	7370	121	0.0162	
AP-038-31	0.0484	0.0025	0.007802	0.0001	0.0448	0.0022	0.0003	0.0003	47.9	2.4	50.1	0.65	5	254	65	85	9100	201	0.0219	
AP-038-32	0.0488	0.0042	0.0076	0.00011	0.0466	0.0038	0.0018	0.0006	48.3	4	48.84	0.36	130	411	85	4710	162.1	0.0420		
AP-038-33	0.0487	0.0068	0.00763	0.00011	0.0461	0.0061	0.0013	0.019	48.1	6.6	48.99	0.71	20	690	690	100	1130	84.9	0.0169	
AP-038-34	0.045	0.0066	0.0078	0.00023	0.0416	0.0054	0.0011	0.004	44.6	6.4	49.1	1.5	20	81	190	150	12380	366	0.0268	
AP-038-35	0.0484	0.0027	0.007633	0.000091	0.0461	0.0025	0.0006	0.005	48	2.6	49.02	0.58	11	100	308	88	1155+04	266	0.0231	
AP-038-36	0.0511	0.0031	0.00784	0.00012	0.0473	0.0028	0.0005	0.0045	50.5	3	50.31	0.74	40	91	382	60	950	111.2	0.0169	
AP-038-37	0.0526	0.002	0.00788	0.000098	0.0484	0.0018	0.003	0.0023	52	2	50.6	0.63	61	45	249	57	10210	270.4	0.0265	
AP-038-38	0.0498	0.0097	0.00787	0.00014	0.0461	0.0067	0.0002	0.0083	49	9.5	50.53	0.9	0	170	820	140	8160	332	0.0407	
AP-038-39	0.0559	0.0045	0.00835	0.00																

TABLE A2. Trace elements in Zircon, and Ti-in Zr Thermometry

I.D.	target	²³⁸ U/ ²⁰⁶ Pb	Al	Ti	Y	Zr	Ba	La	Ce	Pr	Nd	Sm	Eu	Gd	Tb	Dy	Ho	Er	Tm	Yb	Lu	Hf	Th	U	Log Ti	Ti/Pb
AP-00																										
AP-09_145	core	51.1	-1.4	1.98	1057	4.80E+05	0.03	0	7.4	0.005	0.29	1.06	0.81	10.1	4.5	68.1	31.6	174.9	49.3	563	114.1	10510	17.6	161.9	0.29466519	616.00
AP-09_146	core	51.7	-1.4	1.98	1057	4.80E+05	0.03	0	7.4	0.005	0.29	1.06	0.81	10.1	4.5	68.1	31.6	174.9	49.3	563	114.1	10510	17.6	161.9	0.29466519	616.00
AP-09_147	core	51.7	4.5	3.06	1220	4.94E+05	0.099	-0.00363	11.8	0.037	2.86	14.1	16.9	64.7	88.6	39.3	202.5	52	592	118.1	8440	30.1	257.7	0.63948499	672.76	
AP-09_148	core	51.2	29	40.05	2620	5.19E+05	0.267	0.018	35.96	0.071	1.64	4.96	3.08	38.8	14.84	20.2	85.4	419.9	99.9	1036	208.5	9630	28.1	105.1	1.00216006	741.26
AP-09_149	core	52.21	-1.1	7.85	1345	4.64E+05	0.063	0.004	25.02	0.06	1.17	3.19	1.09	21.9	8.86	116.4	45.8	242.6	63.5	658	124.2	7500	63.4	194.4	0.84886866	715.98
AP-09_150	core	52.21	3.3	3.17	1115	4.05E+05	0.15	0	11.01	0.016	0.26	0.7	0.88	13.9	5.97	34.6	188	52.7	656	122.7	10390	16.4	188.4	0.8454559	647.83	
AP-09_151	core	54.3	6.1	21.6	831	5.31E+05	0.099	0	4.98	0.006	0.22	0.66	0.38	3.14	5.0	26.75	150.2	41.6	502	103.4	12820	15.8	194.4	0.33445375	621.92	
AP-09_152	core	54.3	6.1	21.6	831	5.31E+05	0.099	0	4.98	0.006	0.22	0.66	0.38	3.14	5.0	26.75	150.2	41.6	502	103.4	12820	15.8	194.4	0.33445375	621.92	
AP-09_153	core	54.3	6.1	21.6	831	5.31E+05	0.099	0	4.98	0.006	0.22	0.66	0.38	3.14	5.0	26.75	150.2	41.6	502	103.4	12820	15.8	194.4	0.33445375	621.92	
AP-09_154	core	54.3	6.1	21.6	831	5.31E+05	0.099	0	4.98	0.006	0.22	0.66	0.38	3.14	5.0	26.75	150.2	41.6	502	103.4	12820	15.8	194.4	0.33445375	621.92	
AP-09_155	core	54.3	6.1	21.6	831	5.31E+05	0.099	0	4.98	0.006	0.22	0.66	0.38	3.14	5.0	26.75	150.2	41.6	502	103.4	12820	15.8	194.4	0.33445375	621.92	
AP-09_156	core	54.3	6.1	21.6	831	5.31E+05	0.099	0	4.98	0.006	0.22	0.66	0.38	3.14	5.0	26.75	150.2	41.6	502	103.4	12820	15.8	194.4	0.33445375	621.92	
AP-09_157	core	54.3	6.1	21.6	831	5.31E+05	0.099	0	4.98	0.006	0.22	0.66	0.38	3.14	5.0	26.75	150.2	41.6	502	103.4	12820	15.8	194.4	0.33445375	621.92	
AP-09_158	core	54.3	6.1	21.6	831	5.31E+05	0.099	0	4.98	0.006	0.22	0.66	0.38	3.14	5.0	26.75	150.2	41.6	502	103.4	12820	15.8	194.4	0.33445375	621.92	
AP-09_159	core	54.3	6.1	21.6	831	5.31E+05	0.099	0	4.98	0.006	0.22	0.66	0.38	3.14	5.0	26.75	150.2	41.6	502	103.4	12820	15.8	194.4	0.33445375	621.92	
AP-09_160	core	54.3	6.1	21.6	831	5.31E+05	0.099	0	4.98	0.006	0.22	0.66	0.38	3.14	5.0	26.75	150.2	41.6	502	103.4	12820	15.8	194.4	0.33445375	621.92	
AP-09_161	core	54.3	6.1	21.6	831	5.31E+05	0.099	0	4.98	0.006	0.22	0.66	0.38	3.14	5.0	26.75	150.2	41.6	502	103.4	12820	15.8	194.4	0.33445375	621.92	
AP-09_162	core	54.3	6.1	21.6	831	5.31E+05	0.099	0	4.98	0.006	0.22	0.66	0.38	3.14	5.0	26.75	150.2	41.6	502	103.4	12820	15.8	194.4	0.33445375	621.92	
AP-09_163	core	54.3	6.1	21.6	831	5.31E+05	0.099	0	4.98	0.006	0.22	0.66	0.38	3.14	5.0	26.75	150.2	41.6	502	103.4	12820	15.8	194.4	0.33445375	621.92	
AP-09_164	core	54.3	6.1	21.6	831	5.31E+05	0.099	0	4.98	0.006	0.22	0.66	0.38	3.14	5.0	26.75	150.2	41.6	502	103.4	12820	15.8	194.4	0.33445375	621.92	
AP-09_165	core	54.3	6.1	21.6	831	5.31E+05	0.099	0	4.98	0.006	0.22	0.66	0.38	3.14	5.0	26.75	150.2	41.6	502	103.4	12820	15.8	194.4	0.33445375	621.92	
AP-09_166	core	54.3	6.1	21.6	831	5.31E+05	0.099	0	4.98	0.006	0.22	0.66	0.38	3.14	5.0	26.75	150.2	41.6	502	103.4	12820	15.8	194.4	0.33445375	621.92	
AP-09_167	core	54.3	6.1	21.6	831	5.31E+05	0.099	0	4.98	0.006	0.22	0.66	0.38	3.14	5.0	26.75	150.2	41.6	502	103.4	12820	15.8	194.4	0.33445375	621.92	
AP-09_168	core	54.3	6.1	21.6	831	5.31E+05	0.099	0	4.98	0.006	0.22	0.66	0.38	3.14	5.0	26.75	150.2	41.6	502	103.4	12820	15.8	194.4	0.33445375	621.92	
AP-09_169	core	54.3	6.1	21.6	831	5.31E+05	0.099	0	4.98	0.006	0.22	0.66	0.38	3.14	5.0	26.75	150.2	41.6	502	103.4	12820	15.8	194.4	0.33445375	621.92	
AP-09_170	core	54.3	6.1	21.6	831	5.31E+05	0.099	0	4.98	0.006	0.22	0.66	0.38	3.14	5.0	26.75	150.2	41.6	502	103.4	12820	15.8	194.4	0.33445375	621.92	
AP-09_171	core	54.3	6.1	21.6	831	5.31E+05	0.099	0	4.98	0.006	0.22	0.66	0.38	3.14	5.0	26.75	150.2	41.6	502	103.4	12820	15.8	194.4	0.33445375	621.92	
AP-09_172	core	54.3	6.1	21.6	831	5.31E+05	0.099	0	4.98	0.006	0.22	0.66	0.38	3.14	5.0	26.75	150.2	41.6	502	103.4	12820	15.8	194.4	0.33445375	621.92	
AP-09_173	core	54.3	6.1	21.6	831	5.31E+05	0.099	0	4.98	0.006	0.22	0.66	0.38	3.14	5.0	26.75	150.2	41.6	502	103.4	12820	15.8	194.4	0.33445375	621.92	
AP-09_174	core	54.3	6.1	21.6	831	5.31E+05	0.099	0	4.98	0.006	0.22	0.66	0.38	3.14	5.0	26.75	150.2	41.6	502	103.4	12820	15.8	194.4	0.33445375	621.92	
AP-09_175	core	54.3	6.1	21.6	831	5.31E+05	0.099	0	4.98	0.006	0.22	0.66	0.38	3.14	5.0	26.75	150.2	41.6	502	103.4	12820	15.8	194.4	0.33445375	621.92	
AP-09_176	core	54.3	6.1	21.6	831	5.31E+05	0.099	0	4.98	0.006	0.22	0.66	0.38	3.14	5.0	26.75	150.2	41.6	502	103.4	12820	15.8	194.4	0.33445375	621.92	
AP-09_177	core	54.3	6.1	21.6	831	5.31E+05	0.099	0	4.98	0.006	0.22	0.66	0.38	3.14	5.0	26.75	150.2	41.6	502	103.4	12820	15.8	194.4	0.33445375	621.92	
AP-09_178	core	54.3	6.1	21.6	831	5.31E+05	0.099	0	4.98	0.006	0.22	0.66	0.38	3.14	5.0	26.75	150.2	41.6	502	103.4	12820	15.8	194.4	0.33445375	621.92	
AP-09_179	core	54.3	6.1	21.6	831	5.31E+05	0.099	0	4.98	0.006	0.22	0.66	0.38	3.14	5.0	26.75	150.2	41.6	502	103.4	12820	15.8	194.4	0.33445375	621.92	
AP-09_180	core	54.3	6.1	21.6	831	5.31E+05	0.099	0	4.98	0.006	0.22	0.66	0.38	3.14	5.0	26.75	150.2	41.6	502	103.4	12820	15.8	194.4	0.33445375	621.92	
AP-09_181	core	54.3	6.1	21.6	831	5.31E+05	0.099	0	4.98	0.006	0.22	0.66	0.38	3.14	5.0	26.75	150.2	41.6	502	103.4	12820	15.8	194.4	0.33445375	621.92	
AP-09_182	core	54.3	6.1	21.6	831	5.31E+05	0.099	0	4.98	0.006	0.22	0.66	0.38	3.14	5.0	26.75	150.2	41.6	502	103.4	12820	15.8	194.4	0.33445375	621.92	
AP-09_183	core	54.3	6.1	21.6	831	5.31E+05	0.099	0	4.98	0.006	0.22	0.66	0.38	3.14	5.0	26.75	150.2	41.6	502	103.4	12820	15.8	194.4	0.33445375	621.92	
AP-09_184	core	54.3	6.1	21.6	831	5.31E+05	0.099	0	4.98	0.006	0.22	0.66	0.38	3.14	5.0	26.75	150.2	41.6	502	103.4	12820	15.8	194.4	0.33445375	621.92	
AP-09_185	core	54.3	6.1	21.6	831	5.31E+05	0.099	0	4.98	0.006	0.22	0.66	0.38	3.14	5.0	26.75	150.2	41.6	502	103.4	12820	15.8	194.4	0.33445375	621.92	
AP-09_186	core	54.3	6.1	21.6	831	5.31E+05	0.099	0	4.98	0.006	0.22	0.66	0.38	3.14	5.0	26.75	150.2	41.6	502	103.4	12820	15.8	194.4	0.33445375	621.92	
AP-09_187	core	54.3	6.1	21.6	831	5.31E+05	0.099	0	4.98	0.0																

AP_010_27	core	828	4.7	8.49	180.7	4.85E+05	0.073	0.014	7.86	0.088	1.31	3.06	0.88	11.74	2.8	21.8	5.92	18.57	3.23	26.34	4.4	37.25	0.9289079	726.64
AP_010_28	rim	945	17.5	9.48	135.7	4.82E+05	0.073	0.009	6.08	1.21	0.59	0.65	1.4	0.46	6.26	11.60	6.7	6.7	1.80	11.90	4.0	26.2	0.9289079	726.64
AP_010_29	rim	919	9.3	10.6	123.0	4.81E+05	0.072	0.011	14.3	0.038	1.3	0.38	1.24	1.72	4.55	6.9	37.4	28.8	65.4	6.9	37.4	28.8	65.4	6.9
AP_010_30	rim	1137	13	11.2	177.7	4.77E+05	0.049	0.021	8.6	0.028	0.8	0.68	1.13	0.429	3.1	1.88	18.92	6.22	4.06	33.7	5.61	11.90	4.31	10.3
AP_010_31	rim	680	42.6	11.5	146.6	5.11E+05	0.075	0.008	4.8	0.026	0.28	0.60	0.28	0.60	1.75	1.44	4.0	4.0	2.9	41.9	5.43	11.90	4.31	21.74
AP_010_32	rim	762	13.3	12.6	150.8	4.66E+05	0.09	0.044	2.35	0.026	0.7	0.37	0.28	3.6	1.47	4.8	14.9	2.8	5.8	2.78	23.0	5.6	11.90	
AP_010_128	rim	549	10.7	12.6	13.6	5.12E+05	0.081	-0.00769	12.55	0.05	0.69	2.32	1.38	8.6	1.8	1.55	4.25	16.8	11.7	43.1	9.7	10.610	11.8	13.17
AP_010_166	rim	1087	1.7	13.6	879	4.98E+05	0.085	0.055	14.36	0.115	1.92	4.03	1.15	23.1	6.94	78.6	28.9	127.7	29.7	305	56.7	10.610	11.8	13.17
AP_010_165	rim	1213	12.7	13.7	55.2	4.98E+05	0.093	0.04	8.41	0.099	0.64	1.3	0.92	0.36	5.84	23.7	18.4	100.6	28.3	35.5	69.7	10.610	11.8	13.17
AP_010_13	rim	1132	1.5	14.1	161.7	4.64E+05	0.034	0	10.24	0.087	2.51	6.4	0.55	3.8	13.8	16.04	58.9	237	48.8	416	7.09	10.930	53.3	13.33
AP_010_90	core	850	22.2	14.2	139.2	4.86E+05	0.05	-0.00955	4.84	0.051	0.68	1.36	0.39	5.29	1.57	16.04	4.46	15.05	28.9	21	3.89	10.930	53.3	13.33
AP_010_61	rim	1095	0.7	14.3	48.3	4.92E+05	0.033	0	10.24	0.087	2.51	6.4	0.55	3.8	13.8	16.04	58.9	237	48.8	416	7.09	10.930	53.3	13.33
AP_010_60	core	829.6	20.2	15.4	202.6	4.82E+05	0.072	0.047	3.71	0.021	0.22	1.02	0.75	8.07	2.54	25.2	6.46	18.2	2.73	18.2	2.95	12.530	15.9	6.37
AP_010_54	core	930	6.6	15.6	166.1	4.79E+05	0.29	-0.00669	5.65	0.065	1.02	2.85	0.97	13.82	3.27	23.9	5.32	14.17	2.05	13.4	1.5	10.880	4.0	24.06
AP_010_26	core	869	26.3	15.7	175.4	4.86E+05	0.15	0.008	15.47	0.057	0.8	1.86	1.6	9.8	2.96	23.3	5.1	17.3	2.05	20.6	3.21	14.40	273.8	3.9
AP_010_68	core	922	5.4	15.7	115.8	4.87E+05	0.068	0.033	3.44	0.009	0.39	0.87	0.53	6.1	1.73	14.98	3.63	9.09	1.41	10.58	1.61	13.110	20.34	5.1
AP_010_53	core	973	5.9	16.2	39.9	4.88E+05	0.017	0.012	3.37	0.01	0.181	0.74	0.468	6.54	2.72	33.1	12.11	49.6	10.63	96.2	18.4	11.990	21.7	75.9
AP_010_100	rim	923	51.4	16.67	157.2	5.02E+05	0.029	0	5.28	0.028	0.27	1.42	0.123	5.76	1.63	17.4	5.04	18.31	3.2	26.8	4.13	10.510	2.1	18.8
AP_010_69	rim	805	29	16.9	181.9	4.59E+05	0.048	0.01	6.05	0.025	1	1.5	0.58	7.2	1.86	20.9	6.06	21.8	4.02	34.8	6.19	10.110	7.41	26.6
AP_010_98	core	761	-0.2	16.9	281	5.03E+05	0.069	0.037	9.09	0.015	0.91	1.87	0.55	8.84	2.89	20.1	9.13	36.3	7.35	62.4	10.9	10.850	4.7	81.9
AP_010_41	core	772	38.7	17	143.2	4.58E+05	0.124	0.023	8.01	0.067	1.46	3.1	0.8	10.3	2.27	17.7	4.49	15.3	3.08	21.0	4.03	10.220	26.3	19.2
AP_010_56	core	1107	6.7	17.4	228.8	4.94E+05	-0.002	0	10.85	0.038	0.54	2.18	0.77	10.4	2.82	27.1	1.86	22.3	3.67	28.5	4.32	11.300	15.8	7.96
AP_010_70	rim	1133	5.1	17.5	23.2	4.83E+05	0.061	0.092	4.4	0.011	0.28	0.75	0.42	5.5	1.95	23.2	7.07	26.4	5.01	40.4	6.35	12.000	4.1	38.8
AP_010_32	core	866	10.6	17.7	36.4	4.89E+05	0.034	0.013	6.15	0.029	0.46	1.46	0.333	7.05	2.32	30.5	11.81	57.2	13.77	144.1	28.45	10.540	10.81	37.9
AP_010_35	core	1003	5.8	17.82	220.6	4.82E+05	0.055	0.059	9.96	0.042	0.78	1.68	0.69	9.53	3.05	26.4	7.51	22.4	3.82	30.1	5.06	11.370	14.8	36.2
AP_010_7	core	1053	15.1	18.1	381	5.03E+05	0.13	0	11.81	0.059	0.96	2.91	1.21	16.5	5.61	49.8	13.41	37.9	6.02	37.3	4.86	12.390	3.3	66.6
AP_010_102	core	909	10.7	18.1	250.3	4.83E+05	0.09	0.024	15.7	0.156	2.7	5.37	0.77	1.8	4.35	33.3	7.92	23.65	4.42	31.9	4.74	11.840	3.2	25.31
AP_010_34	core	911	7.6	18.2	288.3	4.95E+05	0.044	0.005	3.38	0.026	0.3	0.95	0.545	6.87	5.1	28.3	8.95	4.62	31.1	1.47	12.650	13.91	1.99	1.28039
AP_010_119	core	1040	6.2	18.2	398	5.03E+05	0.07	0.118	7.31	0.011	0.41	1.19	0.556	9.4	3.26	36.9	12.3	4.9	9.31	78.9	13.29	12.480	20.9	6.4
AP_010_25	rim	922	27.5	18.5	265	5.07E+05	0.078	0.014	6.07	0.038	0.58	1.49	0.88	8.88	2.67	10.8	6.88	17.6	10.480	2.95	26	1.27935	80.6	24.5
AP_010_51	rim	951	85.8	19.1	233.5	4.80E+05	0.083	0.055	4.7	0.063	0.55	1.22	0.63	5.72	3.96	21.99	4.27	27.2	6.27	43.7	73.3	11.270	3.8	10.68
AP_010_16	core	969	4	19.14	158.8	4.97E+05	0.127	0.026	3.11	0.204	0.32	7.19	0.531	39.3	11.82	15.11	56.6	239.4	50.9	47.8	82.8	13.450	7.97	36.6
AP_010_127	core	1450	2	19.3	424	4.81E+05	0.07	0.071	4.3	0.136	0.64	3.9	1.89	14.51	39.1	41.5	13.69	58.2	12.7	120.6	21.87	60.7	80.4	73.9
AP_010_72	core	914	19	21.9	350	4.89E+05	0.09	0.02	2.59	0.029	0.27	0.94	0.259	0.003	0.97	1.9	1.9	1.9	1.9	1.9	1.9	1.9	1.9	1.9
AP_010_77	core	919	11.3	19.4	194.8	4.94E+05	0.11	0.021	5.38	0.006	0.67	0.87	0.05	4.2	1.86	1.88	6.08	23.4	4.27	42.7	8.3	11.880	55.2	21.8
AP_010_20	core	892	44.7	19.9	17.6	4.76E+05	0.039	0	4.76	0.026	0.71	1.5	0.63	6.1	1.53	12.01	3.25	10.03	1.75	13.64	1.98	9.850	11.65	57.1
AP_010_114	core	1035	2.8	21.3	72.2	4.95E+05	0.057	0.006	4.97	0.035	0.39	0.74	0.18	1.8	1.22	10.78	2.82	10.8	1.85	18.5	1.98	11.880	15.6	42.7
AP_010_130	core	892	15.1	19.9	212.1	4.95E+05	0.14	0.011	3.64	0.027	0.3	0.4	0.5	3.5	1.3	18.1	6.24	29.2	6.54	61.4	10.74	12.900	31.5	42.7
AP_010_23	rim	898	54.2	20.2	107.5	4.74E+05	0.04	0.02	5.77	0.062	1.81	1.61	0.594	5.29	1.56	14.38	3.67	12.08	2.04	16.03	24.1	10.020	24.2	55.2
AP_010_113	core	1014	4	20.2	68.4	4.92E+05	0.04	0	25.1	0.09	0.1	1.36	0.96	1.1	3.6	45.3	3.6	1.9	1.9	1.9	1.9	1.9	1.9	1.9
AP_010_43	core	1075	1	20.6	269.9	4.86E+05	0.09	0.016	11.94	0.037	1.17	2.66	0.86	13.4	4	33	8.88	28.5	4.6	39	6.07	10.740	20.9	7.6
AP_010_87	core	1175	0.1	20.7	207	4.82E+05	0.076	0.017	2.67	0.063	0.143	0.64	0.524	3.61	1.3	16.03	5.94	26.59	5.55	51.7	8.57	11.200	17.36	37.3
AP_010_92	core	1056	21.9	21.3	339	4.83E+05	0.05	0.001	6.41	0.035	0.46	1.25	0.408	6.79	1.55	17.5	4.08	6.79	1.55	17.5	4.08	6.79	1.55	17.5
AP_010_65	rim	999	21.2	21.3	113.9	4.77E+05	0.02	0.0148	6.88	0.05	0.71	1.21	34.1	4.99	1.37	12.1	3.9	14.1	2.88	25.1	4.09	984.0	3	14.9
AP_010_109	core	1055	-1.4	21.3	567.6	4.87E+05	0.032	0	7.21	0.039	0.68	1.93	0.589	10.3	3.76	47	18.61	84.9	19.89	203.2	37.8	11.000	40.4	101.5
AP_010_110	core	1055	-1.4	21.3	567.6	4.87E+05	0.032	0	7.21	0.039	0.68	1.93	0.589	10.3	3.76	47	18.61	84.9	19.89	203.2	37.8	11.000	40.4	101.5
AP_010_28	core	896	80.7	21.4	130.8	4.83E+05	0.053	0.041	5.4	0.068	0.98	1.67	0.841	6.78	1.75	16.3	4.48	15.26	2.74	22.4	3.5	10.240	1.806	13.86
AP_010_115	core	1145	-2	21.9	50.3	4.89E+05	0.051	0	12.38	0.054	0.92	2.28	0.67	14.9	4.85	50.4	16.8	64.3	13.09	113.5	19.31	10.890	47	79.3
AP_010_23	core	896	80.7	21.4	130.8	4.83E+05	0.053	0.041	5.4	0.068	0.98	1.67	0.841	6.78	1.75	16.3	4.48	15.26	2.74	22.4	3.5	10.240	1.806	13.86
AP_010_4	core	891	37.2	22	108.3	4.56E+05	0.088	-0.0041	6.07	0.056	0.78	1.75	0.84	6.9	1.85	11.9	3.59	12.3	2.11	17.5	3.04	10.010	8.26	48.3
AP_010_33	core	918	3.4	22.1	30.4	4.85E+05	-0.009	0.008	4.94	0.06	0.76	1.2	0.78	3.1	10.16	4.03	8.43	73.2	12.28	89.0	6.8	5.9	3.434227	815.67
AP_010_88	core	865	20.2	22.8	148	4.65E+05	0.04	0.018	5.9	0.154	1.28	1.75	1.41	15.4	16.4	3.41	25.4	1.53	19.8	34.2	1.98	14.827	20.9	24.5

AP_048_141	core	1175	0.0	12.41	618.7	4.884e+05	-0.007	0.029	6.29	0.055	0.84	2.22	17.17	13.72	4.79	59.6	22.43	98.2	22.73	210.9	36.75	11890	6.34	75.8	1.0937718	760.16
AP_048_142	core	1180	-0.2	11.11	618.7	4.884e+05	-0.007	0.049	5.16	0.058	0.64	1.54	10.08	8.99	3.12	80.8	14.35	151.1	15.11	142.5	24.1	10580	6.34	75.8	1.0937718	760.16
AP_048_143	core	1055	3	8.98	1113	3.777e+05	0.036	0.013	11.17	0.029	1.09	4.02	2.82	2.82	11.04	38.17	158.4	84.1	110.0	12.13	110.0	12.13	110.0	12.13	0.9327307	751.45
AP_048_144	core	1060	38.4	13.45	93.9	4.78e+05	0.029	0	8.59	0.073	1.12	2.03	3.04	7.59	1.77	13.9	3.16	9.78	1.74	12.32	17.42	13.92	4.59	35.09	1.1287228	767.56
AP_048_145	core	1160	2.2	8.25	699	4.68e+05	0.097	0	10.28	0.039	0.53	1.87	2.09	11.8	4.0	58.7	23.29	112.6	26.8	27.6	48	1140.0	8.0	60.0	0.9169895	742.29
AP_048_146	rim	1127	0.7	10.09	706	4.81e+05	0.025	0.0013	8.43	0.023	0.85	1.77	1.04	11.92	14.1	65.5	24.23	108.5	24.6	23.21	18.79	1180	11.84	64.0	1.0287771	746.72
AP_048_147	rim	963	14.2	10.65	214	4.88e+05	0.047	0	7.54	0.038	0.98	2.49	0.3	10.34	2.7	25.5	7.32	2.7	5.14	4.29	6.59	1078.0	5.55	54.7	1.0273461	746.39
AP_048_148	rim	1042	17.7	11.02	122.5	4.78e+05	0.029	0.0081	8.81	0.077	1.25	2.5	3.82	9.9	2.19	17.61	4.36	12.33	2.26	16.9	2.4	1041.0	5.55	54.7	1.0273461	746.39
AP_048_149	rim	1162	2.2	8.25	699	4.68e+05	0.097	0	10.28	0.039	0.53	1.87	2.09	11.8	4.0	58.7	23.29	112.6	26.8	27.6	48	1140.0	8.0	60.0	0.9169895	742.29
AP_048_150	rim	944	11.3	12.39	184.5	4.73e+05	0.089	0.01	8.91	0.078	1.36	3.05	0.282	13.93	3.36	27.4	6.76	20.3	3.1	26.1	3.68	1156.0	9.65	56.36	1.0937113	760.02
AP_048_151	core	1114	17	17.77	1005	4.61e+05	0.046	0.01	12.02	0.03	0.7	2.39	13.02	5.16	6.26	84.3	149.2	36.9	349	63.5	1085.0	11.61	53.3	0.8515916	712.40	
AP_048_152	rim	1118	1.1	7.16	236.6	4.70e+05	0.085	0.015	8.89	0.053	2.13	5.13	0.227	29.2	10.66	125.4	46	181.7	39.8	366	59.5	980.0	61.8	24.8	1.0287771	746.72
AP_048_153	rim	1280	-0.6	8.4	869	4.81e+05	0.14	0.0074	14.07	0.027	0.69	2.02	0.122	14.71	61.9	79.1	31.03	135.1	31.1	312.5	51.1	1092.0	8.7	404.3	0.9247299	725.73
AP_048_154	core	913	25.4	9.53	204.5	4.74e+05	-0.004	0	7.75	0.05	0.9	2.38	0.178	10.62	2.78	25.7	6.99	23.7	4.29	35	5.04	1114.0	9.89	92.9	0.9790929	736.61
AP_048_155	core	914	48.4	6.62	307.6	4.88e+05	0.167	0.019	9.24	0.043	0.19	0.98	0.27	6.2	6.1	2.16	28	10.4	45.7	11.78	117.8	20.82	1047.0	28.02	0.8205789	705.82
AP_048_156	core	1176	5.4	8.13	1019	4.72e+05	0.023	0	2.59	0.025	0.55	2.19	0.146	16.3	6.2	8.7	34.85	161.8	38.6	376	66.2	1274.0	31.58	32.9	0.9100905	722.95
AP_048_157	core	1269	4.6	7.84	977	4.80e+05	0.058	0	15.39	0.021	0.49	2.02	0.067	14.96	5.9	80.4	32.46	153.9	38.8	384.4	68.5	1245.0	18.85	946	0.8943100	718.87
AP_048_158	core	955	16.5	10.88	163	4.72e+05	0.082	0.008	8.45	0.088	2.7	0.62	8.28	2.39	20.5	5.38	18.8	3.46	30.9	4.36	1047.0	5.56	41.7	1.0287125	746.64	
AP_048_159	core	1185	2.9	6.88	1280	4.78e+05	0.062	0.0032	15.5	0.116	2.84	5.9	0.21	30.7	1.11	132	47.8	190	40.9	36.0	58.3	1063.0	10.7	325	0.8375884	708.98
AP_048_160	core	1087	0.8	6.39	973	4.79e+05	0.003	0.0118	5.53	0.085	1.67	3.58	0.265	25.4	7.79	92.8	34.2	146.2	31.8	296	40.1	945.0	41.5	88.2	0.8055086	702.93
AP_048_161	core	1087	-0.3	8.60	577.1	4.79e+05	0.018	0.0053	7.11	0.043	0.91	1.82	0.384	12.17	4.44	52.5	19.96	87.7	20.43	36.9	34.9	1084.0	9.38	49.9	0.9309178	728.63
AP_048_162	core	921	11.4	7.7	151.8	4.80e+05	0.004	0.035	8.4	0.058	1.26	2.13	0.308	9.33	2.49	2.0	5.53	17.1	2.8	22.5	5	1015.0	6.01	34.23	0.8864073	718.36
AP_048_163	rim	990	11.1	11.1	199.2	4.78e+05	0.04	0	7.43	0.046	1.09	2.17	0.256	9.55	2.82	25.1	6.83	25.3	4.7	40.6	6.3	1120.0	5.65	59.6	1.0432298	730.08
AP_048_164	rim	1223	1.2	6.93	200.4	4.88e+05	0.016	0.0074	6.53	0.0162	0.34	0.96	0.37	6.2	2.03	22.2	6.82	22.9	4.52	4.85	1158.0	25.1	215	0.8070323	705.58	
AP_048_165	core	947	36.2	14.9	89.6	4.84e+05	-0.009	0.0036	8.08	0.086	1.23	2.33	0.435	7.7	1.87	12.85	2.93	8.66	1.41	11.47	15.4	1034.0	43.1	68.8	1.1731827	777.13
AP_048_166	core	969	16.2	14.94	173.4	4.87e+05	0.049	0	8.46	0.107	1.8	3.17	0.303	11.13	2.8	24.8	5.85	17.99	3.2	23.8	3.44	1120.0	7.65	40	1.1743506	777.38
AP_048_167	rim	1262	2.7	8.5	557	4.71e+05	0.024	0.0113	7.76	0.042	0.67	1.84	0.229	11.49	4.2	52.3	19.99	84.8	18.3	177.8	29.75	1070.0	8.65	41.9	0.9772261	736.38
AP_048_168	core	1226	2.1	8.42	980	4.74e+05	0.084	0.0041	12.37	0.0219	0.55	1.83	1.35	15	5.96	80.8	36.6	153.8	37.7	30.9	1289.0	22.1	1040.0	0.9233130	725.53	
AP_048_169	core	1196	3.3	8.11	880	4.86e+05	0.041	0.0078	11.65	0.022	0.83	2.77	0.355	17.1	6.08	80.9	31.5	138.6	32.2	306	5.14	1014.0	68	247	0.9092085	722.74
AP_048_170	rim	1114	-1.4	6.69	473.5	4.79e+05	0.041	0.0016	7.64	0.038	0.63	1.93	0.114	11.3	3.7	48.6	16.16	161.6	26.65	110.0	15.7	64.6	0.8489915	706.88		
AP_048_171	rim	1134	8.5	8.43	494.3	4.81e+05	0.005	0.015	7.12	0.021	0.68	3.23	0.092	1.86	1.1	48.4	14.54	67.4	15.92	27.2	1190.0	27.8	131.2	0.9247828	712.63	
AP_048_172	core	1126	1.8	12.95	490.5	4.91e+05	-0.014	0	5.74	0.028	0.51	1.23	0.091	9.39	3.51	45.3	17.57	7.8	17.76	17.36	30.4	1145.0	21.52	34.51	1.1122697	747.04
AP_048_173	core	1027	51.4	13.09	923	4.80e+05	0.001	0.0063	9.34	0.062	1.33	2.17	0.361	9.04	1.91	13.69	3.17	9.63	1.598	11.01	16.1	1008.0	5.78	37.43	1.1169995	765.06
AP_048_174	core	1122	2.2	9.1	723	4.85e+05	0.004	0.012	11.49	0.055	0.59	1.96	0.12	14.83	2.9	45.0	17.8	146.2	31.8	146.2	31.8	1140.0	7.65	40	1.1743506	777.38
AP_048_175	rim	983	3.8	10.3	391	4.82e+05	-0.005	0.0063	8.09	0.093	0.97	3.17	0.309	15.7	4.71	48.1	13.64	46.8	1.598	63.4	10.21	1119.0	7.2	63.5	1.0128372	743.43
AP_048_176	rim	1255	3.2	7.4	782	4.76e+05	-0.006	0.0017	11.63	0.044	0.72	2.1	0.063	14.8	5.7	68.4	26.8	123.1	27.4	264	44.7	1030.0	6.16	240.8	0.8692372	715.03
AP_048_177	rim	1027	4.2	7.17	116.0	4.83e+05	0.062	0.012	15.26	0.036	0.61	1.87	0.084	21.6	2.2	11.8	11.8	11.8	11.8	11.8	11.8	1110.0	13.92	13.92	1.131	712.63
AP_048_178	core	1081	20.3	8.83	343	4.71e+05	0.66	0	6.73	0.047	0.76	1.99	0.204	8.22	2.89	34.3	12.1	5.2	11.8	11.8	11.8	1110.0	6.6	81.6	0.945607	730.00
AP_048_179	core	1237	2.6	10.34	176.2	4.95e+05	0.059	0.014	16.42	0.228	0.74	7.92	0.369	51.3	16.41	187.2	6.6	257.3	55.7	47.2	72.4	992.0	100.5	108.7	1.0142504	743.77
AP_048_180	core	1126	2.2	9.1	723	4.85e+05	0.004	0.012	11.49	0.055	0.59	1.96	0.12	14.83	2.9	45.0	17.8	146.2	31.8	146.2	31.8	1140.0	7.65	40	1.1743506	777.38
AP_048_181	core	1020	20.5	13.9	192	4.87e+05	-0.011	-0.00596	7.73	0.074	1.45	2.1	0.192	9.81	26.28	3.9	6.75	26.3	5.4	49.6	7.9	1187.0	9.04	107.5	1.1480148	770.62
AP_048_182	core	1290	0.5	9.07	250	4.73e+05	0.065	0	10.78	0.03	0.8	2.57	11.16	14.61	5.36	69.2	25.8	111	24.41	226.6	36.2	1087.0	39.9	151.6	0.9576029	732.31
AP_048_183	core	1123	6	8.0	662	4.80e+05	0.004	0.012	11.49	0.055	0.59	1.96	0.12	14.83	2.9	45.0	17.8	146.2	31.8	146.2	31.8	1140.0	7.65	40	1.1743506	777.38
AP_048_184	core	1123	6	8.0	662	4.80e+05	0.004	0.012	11.49	0.055	0.59	1.96	0.12	14.83	2.9	45.0	17.8	146.2	31.8	146.2	31.8	1140.0	7.65	40	1.1743506	777.38
AP_048_185	core	1304	7.6	14.05	240.0	4.81e+05	0.064	0.032	22.81	0.406	7.49	15.8	0.385	71.6	22.92	256	88.6	358	75.2	641	98.2	1080.0	218	489	1.1747632	771.62
AP_048_186	core	1123	6	8.0	662	4.80e+05	0.004	0.012	11.49	0.055	0.59	1.96	0.12	14.83	2.9	45.0	17.8	146.2	31.8	146.2	31.8	1140.0	7.65	40	1.1743506	777.38
AP_048_187	core	1255	0.4	8.41	964	4.77e+05	0.032	0	14.35	0.055	1.24	4.06	0.203	21.1	7.67	95.1	38.88	145.7	31.7	290.2	46.3	986.0	75.8	161.7	0.9247828	7

DAI_39_78	core	872	21.7	24.3	1387	4.95E+05	0.181	0	41.3	0.161	4.32	7.9	2.24	40	12.55	136.4	48.5	205.8	46.4	45.9	1.37	10900	331	725	1.38560627	825.37	
DAI_39_79	core	916	6.0	39.1	29.9	0.056	0.013	0.008	8.1	0.079	1.49	0.28	0.06	6.68	1.14	11.19	2.79	8.27	1.36	1.37	0.27	9230	1.63	8.4	1.3856072	878.48	
DAI_39_80	core	984	0.7	27.7	1048	0.156	0.036	0.018	48.1	0.186	1.77	1.98	5.28	18.35	14.12	18.2	18.1	147.8	18.2	18.2	0.27	9380	80.9	70.5	1.4427977	893.89	
DAI_39_81	core	933	0.3	40.4	71.3	4.72E+04	0.044	0	14.64	0.065	1.22	2.59	1.23	6.85	1.19	9.83	2.47	7.09	1.102	8.09	1.101	9320	11.33	68.9	1.6068317	880.45	
DAI_39_82	core	903	28.7	32.7	13.2	0.002	0.153	0.037	0.13	0.037	0.33	0.57	0.29	1.17	0.47	1.37	1.47	1.37	1.47	1.37	1.47	1.37	2.28	1.5118104	856.24		
DAI_39_83	core	912	40.3	29.9	68.8	4.76E+05	0.006	0.006	17.96	0.232	2.11	3.67	0.37	1.41	0.51	10.2	2.18	0.44	0.949	1.702	1.174	10030	41.9	104.0	1.4756719	847.19	
DAI_39_84	core	906	40.2	37.6	44.84	4.81E+05	0.023	0.0017	11.42	0.043	1.35	1.64	1.27	5.25	0.949	6.75	1.48	3.88	0.604	3.96	0.534	9660	97.2	26.1	1.5751874	870.23	
DAI_39_85	core	906	30.7	25	56.2	4.75E+05	0.139	0.0053	32.01	0.215	1.77	5.59	1.37	22.2	5.99	6.05	19.12	75.9	16.64	138	21.8	8950	101.1	196.1	1.5794901	828.31	
DAI_39_86	core	884	48.8	36.3	21.1	4.73E+05	0.081	0.0068	10.89	0.099	1.43	0.91	0.51	0.69	0.69	1.43	0.91	0.51	0.69	1.43	0.91	9220	38.1	104.0	1.5629236	869.01	
DAI_39_87	core	913	13.7	13.7	47.2	4.83E+05	0.056	0.027	19.91	0.409	6.57	7.36	2.93	2.33	5.51	5.27	16.66	65.8	13.3	125.7	23.07	8040	38.4	18.4	1.5050926	853.50	
DAI_39_88	core	932	414	71.4	923	4.59E+05	2.06	0.047	58.1	0.862	14.38	16.6	4.7	50.3	11.6	107.5	31.8	124.4	25.36	232.5	38.1	8400	96.3	30.9	1.8536921	940.99	
DAI_39_89	core	921	12.3	33.3	72.2	4.75E+05	0.07	0.0052	30.51	0.27	6.59	9.86	5.14	3.92	8.49	6.11	27.08	10.4	22.88	206.1	34.9	9210	80.9	10.0	1.5250481	859.53	
DAI_39_90	core	916	55.7	34.5	76.6	4.82E+05	0.031	0.0016	11.81	0.088	0.96	1.7	0.836	6.13	1.433	10.6	2.75	7.96	1.34	10.15	1.47	8000	2.91	10.5	1.578191	862.76	
DAI_39_91	core	921	48.5	40.6	55.4	4.80E+05	0.024	0.013	82.13	0.098	1.69	1.95	0.77	5.17	1.039	7.59	1.85	5.52	0.859	6.39	0.987	8260	1.46	80.5	1.6088263	881.01	
DAI_39_92	core	986	52	29.5	1718	4.76E+05	0.55	0.184	36.1	0.57	9.8	19.6	8.81	67.1	20.71	195.2	60.2	229.4	44.7	397	67.3	7540	22.1	37.5	1.6468202	845.75	
DAI_39_93	rim	878	46.4	18.66	128	4.69E+05	0.021	0.0068	10.81	0.081	1.44	2.21	0.949	7.59	1.81	15.3	4.23	16.2	3.2	29.5	5.3	9370	15.1	52.7	1.2709164	798.79	
DAI_39_94	core	894	52.7	36.8	47.83	4.74E+05	0.092	0.02	26.76	0.148	2.12	2.38	0.45	5.89	1.088	7.52	1.522	3.96	0.575	4.91	0.704	10300	32.1	12.2	1.5658472	869.93	
DAI_39_95	core	950	60.3	42.2	55.5	4.70E+05	0.029	0.0115	8.66	0.078	1.36	2.15	1.18	5.03	1.056	7.59	1.63	5.42	0.932	7.04	1.13	8240	6.38	20.49	1.6254875	888.12	
DAI_39_96	core	945	28.3	27.8	46.6	4.86E+05	0.058	0.017	17.3	0.325	4.64	6.33	2.44	20.9	5.15	4.99	15.4	64.2	13.22	130	22.1	7940	26.1	34.5	1.4404488	839.43	
DAI_39_97	core	932.8	51.9	38.9	91.8	4.77E+05	0.017	0	10.44	0.052	1.21	2	0.809	6.41	1.5	12.7	3.12	9.78	1.72	1.66	9400	12.65	66.5	1.5899496	876.16		
DAI_39_98	rim	903	46.6	31.3	53.1	4.74E+05	0.055	0.005	25.58	0.042	7.53	8.82	3.44	25.7	6.39	57.3	17.43	68.8	13.8	122.5	22.37	7880	41.55	45.8	1.4955444	852.12	
DAI_39_99	rim	927	40	37.1	80	4.68E+05	0.026	0.0077	9.86	0.063	1.61	2.26	0.957	7.47	1.57	10.8	2.63	8.2	1.63	13.5	2.41	9460	8.8	35.6	1.5693791	870.83	
DAI_39_100	core	903.2	298.9	38.9	554.1	4.62E+05	0.306	0.021	27.57	0.23	4.34	7.94	2.84	27.6	6.53	62.7	19.3	75.6	16.26	149.7	25.56	8560	13.66	61.4	1.5899496	876.16	
DAI_39_101	core	904	83.8	24.5	418.2	4.77E+05	0.14	0.033	18	0.221	4.37	5.79	2.58	20.2	4.86	47.3	14.87	57.9	12.04	113.3	19.33	8410	32	37.5	1.3891608	825.22	
DAI_39_102	core	967	51.1	39.9	144.5	4.70E+05	0.017	-0.0011	5.96	0.09	1.34	1.77	1.15	7.05	1.69	17.18	4.93	17.94	5.32	29.1	4.37	7940	2.95	6.31	1.6009729	879.03	
DAI_39_103	core	915	66.5	37.8	59.8	4.76E+05	0.037	0.0016	12.21	0.132	1.66	2.27	0.99	6.72	1.408	8.96	2.07	5.53	0.999	7.04	1.12	8650	84.6	22.42	1.5747918	872.93	
DAI_39_104	core	931	73.3	33.9	48.7	4.72E+05	-0.001	0	6.29	0.061	0.92	1.51	0.74	6.12	1.06	7.22	1.66	4.77	0.342	5.01	0.679	679	15.44	1.5301397	863.88		
DAI_39_105	core	930.5	46.4	43	108.8	4.88E+05	0.062	-0.003012	9.63	0.087	1.62	1.98	1.02	1.14	9.08	20.5	16.18	25	11.2	1.254	8.86	1.08	9090	32.44	71.6	1.6345466	887.59
DAI_39_106	core	1022	0.9	25.96	72.4	4.73E+05	0.086	0.027	24.4	0.468	9.53	13.2	4.82	41.4	9.32	8.22	24.95	94.9	19.9	180.9	31	8080	50	1.0	1.4143046	832.23	
DAI_39_107	core	1019	-0.6	32.5	45.8	4.78E+05	0.046	0.0116	15.95	0.164	3.17	5.13	2.66	22.8	5.77	58.2	10.02	72.6	15.39	144.5	24.82	8067	67.4	26.54	1.5118826	856.24	
DAI_39_108	core	1019	44.8	33.3	40.4	4.73E+05	0.047	-0.0067	11.95	0.077	1.15	1.76	0.97	3.84	0.87	2.84	1.43	1.43	0.82	6.07	0.643	9020	18.7	75.0	1.5629236	869.01	
DAI_39_109	rim	914	22.8	33.6	150	4.68E+05	0.012	0.0074	13.08	0.135	2.22	2.98	1.86	8.9	1.57	18.3	4.95	19.9	4.3	39.6	6.7	9490	16.3	64	1.5263238	859.85	
DAI_39_110	rim	877	33	26.4	1493	4.72E+05	0.283	0.026	6.5	0.763	13.72	20.1	5.94	69.9	19.99	159.6	49.53	191.8	39.95	35.7	57.6	8740	196.9	196.5	1.5418039	831.99	
DAI_39_111	rim	907.7	50.7	29.7	87.6	4.76E+05	0.059	0.009	12.89	0.08	1.26	1.66	0.927	8.86	1.13	8.02	6.16	7.15	0.82	8.68	5.67	8880	17.67	10.0	1.5629236	869.01	
DAI_39_112	core	1044	0.5	38.9	495.5	4.66E+05	0.114	-0.0034	24.01	0.175	0.35	5.18	1.76	18.39	5.35	14.1	16.55	68.1	10.51	145.6	25.04	8440	45.5	66.9	1.5899496	876.16	
DAI_39_113	core	902	28.5	32	309.2	4.68E+05	0.088	0.0099	25.69	0.093	1.64	2.33	0.73	10.1	2.83	31.1	10.8	46.7	10.63	110.1	19.9	10000	28.47	56	1.5051498	854.52	
DAI_39_114	core	907.7	42.8	42.8	42.8	4.74E+05	0.047	0.004	19.02	0.047	1.94	4.68	1.102	14.1	1.41	10.2	10.2	10.2	10.2	10.2	10.2	9880	64.7	11.4	1.5327438	861.48	
DAI_39_115	core	925.8	40	30	686	4.64E+05	0.79	0.023	81.2	0.386	6.69	9.4	2.09	29.7	7.26	72.8	23.66	94.7	20.96	202.4	36	8700	114.7	265.2	1.4771225	847.55	
DAI_39_116	core	912	74.2	34.2	259	4.72E+05	0.026	0.0011	9.86	0.017	1.26	1.66	0.927	8.86	1.13	8.02	6.16	7.15	0.82	8.68	5.67	8880	17.67	10.0	1.5629236	869.01	
DAI_39_117	rim	910	70.2	37.8	54.1	4.54E+05	0.032	0.0108	10.29	0.132	1.52	2.15	1.05	6.13	1.14	7.16	1.69	4.83	0.916	7.35	1.124	8470	7.3	10.3	1.5747918	872.93	
DAI_39_118	core	891	19	23.7	50.6	4.74E+05	0.092	0.067	20.2	0.262	5.04	6.89	2.64	20.6	5.41	52.9	16.4	66.4	15.1	137.6	24.39	9030	46.6	9.5	1.3748383	822.80	
DAI_39_119	core	913	74.2	34.2	259	4.72E+05	0.026	0.0011	9.86	0.017	1.26	1.66	0.927	8.86	1.13	8.02	6.16	7.15	0.82	8.68	5.67	8880	17.67	10.0	1.5629236	869.01	
DAI_39_120	core	937	43.3	40	59.8	4.67E+05	0.055	0.0031	9.37	0.038	1.33	2.25	1.22	6.26	1.27	8.59	2.04	5.82	1.07	8.3	13.3	8410	5.59	15.31	1.6020599	879.32	
DAI_39_121	core	913	46.8	25.7	83.4	4.65E+05	0.041	0.0129	10.09	0.072	1.67	1.8	1.4	6.42	1.42	11.33	2.63	8.88	1.59	13.8	2.31	8500	67.9	20.9	1.4099312	831.18	
DAI_39_122	core	913	46.8	25.7	83.4	4.65E+05	0.041	0.0129	10.09	0.072	1.67	1.8	1.4	6.42	1.42	11.33	2.63	8.88	1.59	13.8	2.31	8500	67.9	20.9	1.4099312	831.18	
DAI_39_123	core	937	43.3	40	59.8	4.67E+05	0.055	0.0031	9.37	0.038	1.33	2.25	1.22	6.26	1.27	8.59	2.04	5.82	1.07	8.3	13.3	8410	5.59	15.31	1.6020599	879.32	
DAI_39_124	core	913	46.8	25.7	83.4	4.65E+05	0.041	0.0129	10.09	0.072	1.67	1.8	1.4	6.42	1.42	11.33	2.63	8.88	1.59	13.8	2.31	8500	67.9	20.9	1.4099312	831.18	

MPP_33A_07	core	266.1	15	4.15	1555	4.66E+05	0.25	0.013	23.8	0.117	2.51	5.09	1.24	30.8	10.94	131.9	61.5	249	572	111.0	9070	255.4	30.5	0.6180481	668.99	
MPP_33A_5	core	271.6	9	6.34	1950	0.09E+06	0.046	0.21	0.13	1.9	5.26	1.27	28.7	11.58	153.6	61.2	307.2	72.8	77.8	143.9	989	399	610	0.6180481	673.45	
MPP_33A_7	core	280	1.4	4.4	1950	1.13E+05	0.15	0.051	23.25	0.41	7.8	10.2	16.4	17.2	18.6	65.2	309	72.2	36.8	137.8	8620	139	810	0.6180481	673.45	
MPP_33A_71	rim	290.1	-1.6	4.43	528	4.82E+05	0.073	0	10.1	0.037	0.57	1.41	0.39	9.82	3.25	42.9	17.46	86.9	21.3	21.3	41.3	9090	84.4	0.6180481	673.98	
MPP_33A_7	rim	302.3	4.7	4.65	737	4.80E+05	0.048	0.009	21.02	0.031	8.81	2.32	0.88	13.7	4.96	59.1	24.42	47.7	28.6	15.6	41.3	9090	84.4	0.6180481	673.98	
MPP_33A_26	core	278.4	6.4	4.52	1381	4.84E+05	0.15	0.012	38.65	0.188	1.55	8.88	8.17	15.8	8.81	124.2	47.5	21.1	90.9	94.5	200.0	44.4	208	0.6180481	675.42	
MPP_33A_40	core	268.6	-1.4	4.53	934	4.84E+05	0.082	0.0088	25.1	0.099	1.95	4.57	1.17	33.2	12.66	162.1	61.56	311.9	71.4	71.3	133.4	9070	570	0.6180481	676.03	
MPP_33A_34	core	267.6	3.5	4.55	1056	4.82E+05	0.096	0.016	10.63	0.1	1.72	3.23	0.94	18.5	6.55	85	35.5	17.42	44.5	47.5	93.6	10550	140.4	0.6180481	676.03	
MPP_33A_105	core	279.4	6.4	4.45	1498	4.81E+05	0.31	0.1	2.9	1.28	9.3	3.18	3.81	11.9	21.3	54.4	24.1	54.8	53.8	99.9	139.9	87.0	184.4	0.6180481	676.87	
MPP_33A_107	core	256.9	4.5	4.64	1821	4.73E+05	0.28	0.104	23.8	0.098	1.75	4.04	0.92	31.1	15.91	154.2	63.3	28.4	66.6	63.1	123.9	9560	58.4	0.6180481	677.54	
MPP_33A_30	rim	290.3	6.4	4.66	834	4.87E+05	0.141	0.006	18.18	0.047	0.66	1.92	0.88	12.7	5.03	16.1	26.8	13.2	32	33.1	63.2	9940	151.7	0.6180481	677.87	
MPP_33A_35	core	276.8	17	4.68	4210	4.79E+05	0.075	0.026	12.55	0.149	2.98	7.11	0.87	56.5	23.84	39.7	139.7	665	155.3	153.6	278.2	10270	65.1	0.6180481	678.20	
MPP_33A_99	core	274.7	0.7	4.91	1646	4.81E+05	0.104	0.069	29.7	0.134	1.93	3.82	0.94	26.3	9.87	129.4	53.1	26.6	64.5	66.2	129.1	9130	279	0.6180481	678.93	
MPP_33A_93	core	265.8	45	5.2	1280	4.85E+05	0.173	0.051	10.01	0.277	5.23	7.51	2.3	35.5	10.5	121.4	43.2	19.7	43.3	41.8	77.6	8050	1139	0.6180481	684.43	
MPP_33A_22	rim	245	11.1	5.3	744	4.66E+05	0.14	0.08	14.46	0.055	0.52	1.73	0.62	22.2	4.2	58.3	24.7	120.4	28.5	29.1	98.2	950	139.2	0.6180481	687.93	
MPP_33A_41	core	252.1	12.5	5.3	2670	4.65E+05	0.218	0.019	59.9	0.177	2.53	7.07	4.09	48.9	1.81	237	93.1	430	99.8	92.3	169.7	12670	1590	0.6180481	687.93	
MPP_33A_36	rim	278.6	5.3	5.3	1041	4.82E+05	0.221	1.14	34.6	2.03	14.9	11.3	1.03	29.3	8.79	98.6	34.9	160.3	37	359	64.9	9150	204.3	0.6180481	687.93	
MPP_33A_74	core	280.4	6.1	5.37	1258	4.70E+05	0.09	0.032	12.3	0.15	2.42	4.52	1.29	26.6	8.57	107.8	6.2	159	4.8	48.1	93.4	8640	172	0.6180481	688.97	
MPP_33A_6	rim	286.5	15.3	5.5	592	4.85E+05	0.186	0.014	10.6	0.065	0.88	2.2	0.567	10.17	3.74	47.6	20.33	95.7	22.68	21.6	46.6	8740	109.7	0.6180481	690.86	
MPP_33A_101	core	543	37	5.5	848	4.60E+05	0.26	0.02	10.37	0.046	0.84	1.82	0.209	12.5	5	68.6	27.8	136.4	34.4	34.6	66	14490	279	0.6180481	690.86	
MPP_33A_17	core	275.7	10.3	5.56	2163	4.70E+05	0.117	0.179	16.24	0.093	8.28	13.1	0.87	63.9	15.9	214.9	77.4	64.2	72.2	64.1	120.1	8270	23.2	0.6180481	691.73	
MPP_33A_37	core	272.8	17.8	5.81	4300	4.76E+05	0.115	0.094	39.2	0.89	16.4	28.3	3.79	131.1	40.3	443	44.3	131.6	121.2	212.6	7880	665	63	0.6180481	695.24	
MPP_33A_75	rim	267	1.3	5.89	1132	4.80E+05	0.135	0.002	26.84	0.097	1.96	3.68	1.4	23.5	7.93	98.4	37.3	176.3	41.5	408.8	40	40.8	40.8	0.6180481	696.34	
MPP_33A_42	core	263.9	5.3	5.92	1197	4.76E+05	0.026	0	12.24	0.028	0.65	1.78	0.531	13.25	6.09	88.4	42	21.3	55.7	61.6	123.1	10300	64.6	0.6180481	697.69	
MPP_33A_78	rim	302.6	4.3	6.02	1196	4.87E+05	0.086	0.0027	15.52	0.083	1.25	2.96	0.83	18.5	7.0	21.2	96.2	90	389	88.7	841	156	8560	53.1	0.6180481	697.69
MPP_33A_100	core	271.1	31	6.14	2760	4.77E+05	0.07	0.049	19.5	0.365	6.74	12.6	2.29	70.9	23.3	273	98.4	430	91.6	367	154.1	8300	335	0.6180481	698.69	
MPP_33A_16	core	265.8	17	6.3	1326	4.83E+05	0.29	0.018	29.7	0.088	2.98	4.37	1.05	25.7	8.65	114.7	44.2	212.8	94.6	48.4	95.5	9460	88.8	0.6180481	701.77	
MPP_33A_82	core	281.1	4	6.3	2607	5.01E+05	0.16	0.054	32.8	0.872	8.6	15.1	4.86	70	22.33	246	90.5	410	92.3	895	167.2	8690	473	0.6180481	701.77	
MPP_33A_32	core	290.4	3.1	6.3	1605	4.78E+05	0.088	0.012	14.5	0.096	0.52	1.57	4.06	14.4	24.7	81.2	106.6	43.2	201	48.1	99.6	816	154	0.6180481	704.45	
MPP_33A_81	core	299	0.61	6.41	1303	4.81E+05	0.072	0	16.9	0.083	1.77	3.23	1.1	12.5	9.1	61.2	32.5	91.1	52.2	103	82.0	160	139	0.6180481	705.69	
MPP_33A_84	core	276.8	2.1	6.75	1712	4.76E+05	0.076	0.006	15.94	0.102	3.23	4.87	1.43	31.1	11.19	143.3	57.6	27.4	62.2	60.4	112	7970	202.9	0.6180481	707.41	
MPP_33A_9	core	276.1	1.4	6.86	1990	4.77E+05	0.086	0.0046	28.5	0.151	2.72	5.9	1.57	35.3	15.2	108	67	32.2	76	770	154	7630	451	0.6180481	708.74	
MPP_33A_17	core	297.7	0.61	7.68	248	4.80E+05	0.06	0.013	19.3	0.031	0.9	1.82	0.29	24.0	5.58	118.4	24.0	55.4	65.2	80.6	144.2	860	362	0.6180481	708.86	
MPP_33A_98	rim	274.6	60.8	6.94	1339	4.79E+05	0.284	0.032	22.27	0.064	1.84	3.37	1.11	22.3	8.3	109.3	43.72	209.5	50	497	96.4	8510	339.7	0.6180481	708.74	
MPP_33A_31	core	289.9	30	7.1	2130	4.97E+05	0.5	0.025	37.6	0.217	4.72	8	2.43	44.4	14.3	181	72.4	338	84.9	898	167	9220	446	0.6180481	711.59	
MPP_33A_10	core	289.3	2.9	7.1	288	4.87E+05	0.06	0.019	19.3	0.031	0.9	1.82	0.29	24.0	5.58	118.4	24.0	55.4	65.2	80.6	144.2	860	362	0.6180481	715.48	
MPP_33A_19	rim	270.1	27.6	7.11	2283	4.95E+05	0.257	0.204	24.57	0.697	9.05	13.9	3.42	54.9	18.4	110	216.7	77.5	355	82.6	83.1	150.7	820	641	0.6180481	711.70
MPP_33A_103	core	270.4	12.6	7.28	3460	4.76E+05	0.241	0.076	61.1	0.499	8.29	16	3.68	77.6	24.5	295	113.7	531	124.6	1248	246.1	7910	1126	0.6180481	713.67	
MPP_33A_27	core	294.4	20.2	7.34	4230	4.86E+05	0.73	0.39	30.5	0.92	1.24	2.99	1.04	24.1	8.23	82.5	42.1	82.1	82.5	92.4	91.5	1160	443	0.6180481	715.48	
MPP_33A_65	rim	270.1	15.2	7.5	2184	4.94E+05	0.19	0.031	27.4	0.21	4.99	6.28	1.68	38.2	14.93	181.5	73.2	347.1	81.2	799	151	8590	1122	0.6180481	716.15	
MPP_33A_44	rim	289.4	29	7.6	1751	4.94E+05	0.149	0.013	21.2	0.074	1.5	3.88	1.11	25.7	9.95	133.6	57.2	290.5	69.5	718	143.5	9490	226	0.6180481	717.26	
MPP_33A_12	rim	289.3	1.1	7.68	248	4.80E+05	0.06	0.013	19.3	0.031	0.9	1.82	0.29	24.0	5.58	118.4	24.0	55.4	65.2	80.6	144.2	860	362	0.6180481	725.42	
MPP_33A_13	core	283	4	8.27	1055	4.86E+05	0.121	0.056	22.28	0.122	1.15	4.11	1.23	20.9	7.28	91.3	35.74	166.4	38.2	387	73.4	8640	336.9	0.6180481	724.40	
MPP_33A_28	core	279.9	2.5	8.29	1239	4.89E+05	0.075	0.016	15.37	0.083	1.45	2.35	1.19	20	7.28	99.3	41.5	198	47.9	489	98.2	7650	120.7	0.6180481	724.40	
MPP_33A_18	rim	271.1	0.1	8.4	1723	4.86E+05	0.096	0.0129	22.82	0.082	1.87	4.52	1.46	26.2	10.6	139.2	57.4	275.9	66.7	675	132.2	8440	263.7	0.6180481	725.73	
MPP_33A_86	core	981	8.4	8.7	607	4.84E+05	0.084	0.027	8.79	0.058	0.92	2.43	0.144	12.1	4.62	55.5	21.4	93.2	21.1	205	35.9	11690	103	0.6180481	728.73	
MPP_33A_17	core	297.7	0.61	7.68	248	4.80E+05	0.06	0.013	19.3	0.031	0.9	1.82	0.29	24.0	5.58	118.4	24.0	55.4	65.2	80.6	144.2	860	362	0.6180481	728.73	
MPP_33A_18	rim	287.1	4.7	9.6	2313	5.03E+05	0.118	0.012	38.6	0.113	2.54	5.26	1.81	39.8	14.6	190.1	77.6	364	85.4	925	162.4	8950	626	0.6180481	737.25	
MPP_33A_108	core	280.8	0.7	9.8	1758	4.81E+05	0.1	0.02	24.48	0.115	21.6	4.16	2.05	25.5	10.61	144.6	57.5	278.5	65.8	669	133	8020	231.6	0.		

MG_Q03_106	rim	911	47.6	7.6	417	4.666+05	0.274	7.76	0.023	1.75	0.145	8.7	2.42	38.1	15.7	66.4	14.81	147.3	25.1	137.9	35.08	94.6	0.88081359	717.26	
MG_Q03_105	core	912	1.1	5.1	37	3863	4.795+05	0.051	1.17	0.98	2.24	1.77	82.7	22.8	280.3	72.4	271.5	53.2	44.4	72.2	99.9	74.4	112.7	0.71382988	685.79
MG_Q03_104	rim	913	2.02	14.9	736	4.727+05	0.196	5.4	0.196	5.4	0.196	5.4	0.196	5.4	0.196	5.4	0.196	5.4	0.196	5.4	0.196	5.4	0.196	5.4	0.196
MG_Q03_103	rim	914	2.1	2.38	1519	4.85+05	0.107	21.29	0.186	8.93	3.6	44.6	14.29	164.1	56.9	233.6	50.9	45.5	7.5	12.89	24.3	54.3	0.37657696	628.61	
MG_Q03_163	core	915	303	5.84	2607	4.831+05	1.18	16.7	0.77	20.9	1.36	99.3	30.1	100.7	60.9	80.2	60.7	110.4	107.7	131.1	30.3	40.9	0.71382988	685.65	
MG_Q03_155	rim	914	13.9	207	3381	4.807+05	0.188	20.9	0.967	28.9	1.8	12.5	38.4	39.1	12.3	44.2	92.2	76.2	114.3	20.8	151.1	107.9	0.93979728	809.68	
MG_Q03_91	rim	915	1	19.7	530	4.83+05	0.069	4.85	0.058	3.56	0.5	17.2	52.3	18.8	81.1	16.42	149.8	25.92	98.30	19.37	37.3	1.29446623	804.14		
MG_Q03_171	core	915.6	1.3	27.1	3848	4.90+05	0.031	13.01	0.87	22.3	1.49	80.4	22.83	229.7	7.9	272.9	54.6	447.8	7.4	10.98	88.7	137.7	0.42129629	617.73	
MG_Q03_80	rim	916	2.3	7.14	890	4.87+05	0.041	8.51	0.075	19.1	0.9	20.7	55.5	19.1	40.9	17.7	15.6	26.36	107.9	19.9	30.9	0.71382988	685.65		
MG_Q03_88	rim	916	0.6	9.88	504.7	4.83+05	0.067	7.7	0.057	2.96	0.224	14.8	44.7	52.3	17.78	73.8	15.67	144.8	24.94	110.90	27.69	72.2	0.99475696	739.76	
MG_Q03_51	core	917	483	4.6	612	4.68+05	0.003	15.5	0.055	3.07	0.181	17.7	5.61	60.7	21.1	84.3	17.62	146.8	23.97	1217.0	72.8	164.7	1.66475783	895.41	
MG_Q03_78	core	917	0.5	9.8	452.2	4.86+05	0.002	7	0.052	2.41	0.169	13.3	4.42	48.2	16.45	65.9	14.15	134	22.58	114.3	24.63	80.9	0.926951	729.41	
MG_Q03_139	rim	918	11.8	11.87	441.6	4.90+05	0.054	6.88	0.111	2.58	0.27	13.7	43.1	46	16.35	67.4	13.92	129.1	20.96	106.60	23.56	57.2	1.07445072	756.12	
MG_Q03_25	rim	918	11.3	13.8	680	4.98+05	0.087	6.86	0.125	3.9	0.448	20.6	6.69	72.9	25.38	104.2	10.98	191	32.78	102.50	27.48	56.3	1.13897890	769.95	
MG_Q03_66	rim	919	20.5	6.17	930	4.80+05	0.015	2.28	0.053	1.34	0.153	14.5	6.9	85.9	19.31	102.8	19.5	154.2	23.82	140.20	53.3	95.6	0.79208116	701.08	
MG_Q03_5	core	919	1.4	2.9	1304	4.70+05	0.044	9.73	0.496	13.9	1.07	56.3	15.62	159.2	50.6	194.1	38.2	323	53	101.60	56	85.9	0.462398	642.56	
MG_Q03_21	rim	920	1.3	7.3	374.7	4.89+05	0.023	7.42	0.288	11.5	0.179	9.81	31.7	36.9	13.74	58.6	12.74	122	21.31	116.60	30	88.8	0.86332286	713.89	
MG_Q03_188	rim	920	3.1	12.5	564	4.88+05	0.009	7.57	0.089	3.62	0.253	17.1	5.55	61.3	20.64	87.5	18.17	158.7	17.33	104.50	27.65	61.4	1.09691001	760.82	
MG_Q03_172	core	921	-1.6	5.11	245	4.92+05	0.021	5.68	0.034	1.35	0.138	6.85	1.96	23.8	8.58	36.1	8.15	7.9	13.8	10.80	34.3	76.8	0.3904209	685.06	
MG_Q03_126	core	921	2.1	4.94	1686	4.94+05	0.082	11.88	0.84	18.2	1.26	69.9	20	205.6	64.2	250.5	48.5	413	64.3	99.90	80.3	125.5	0.6972695	684.21	
MG_Q03_115	rim	922	1.8	5.89	422.7	4.82+05	0.032	6.74	0.054	2.19	0.238	11.9	3.97	46.1	15.16	65.2	13.83	127.9	21.58	108.70	24.09	61.6	0.77011529	696.34	
MG_Q03_40	core	922.7	-3.1	8.47	747	4.87+05	0.089	12.1	0.127	4.48	0.337	23.3	7.39	83.3	28.21	111.5	23.7	205.4	34.65	112.20	47	116.5	0.92788341	725.43	
MG_Q03_42	core	923	4.4	4.99	2333	4.89+05	0.093	14.2	1.06	21.1	1.46	92.2	26.42	278	88.8	340	65.9	552	90.6	102.00	111.1	181.7	0.6810055	683.19	
MG_Q03_101	rim	923	11.3	11.3	831	4.79+05	0.097	20.0	24.8	27.5	1.53	44.3	10.28	99.9	31.16	125.5	24.52	216.5	36.1	108.30	49	102.1	0.96142109	733.07	
MG_Q03_70	rim	923	-5	7	941	4.90+05	0.034	10.27	0.236	8.36	0.68	34.6	10.21	103.2	34.6	134.2	27	240	38.5	103.20	52	91.3	0.84509804	710.41	
MG_Q03_118	core	923	1.9	18.3	434	4.94+05	0.04	4.37	0.059	2.39	0.328	13	4.01	18.3	15.67	65.6	13.41	124.4	21.39	103.00	24.7	91.3	1.24245109	798.88	
MG_Q03_29	core	923	-2.3	12.5	613	4.91+05	0.054	6.28	0.102	3.87	0.51	20.9	6.21	67.4	22.1	93.2	18.95	170.1	28.32	102.80	26.7	102.9	0.9691001	760.82	
MG_Q03_29	core	924	1.8	7.2	940	4.79+05	0.117	5.88	7.1	62.1	17.7	114	35.2	126.2	25.1	25.7	20.7	35.1	105.60	50	84.4	0.8737325	712.75		
MG_Q03_26	core	924	2.6	8.8	702	4.83+05	0.052	10.4	0.18	5.78	0.455	22.9	10.4	13.9	65.3	44.2	71.9	119.9	31.2	107.10	38.6	80.5	0.84448207	754.9	
MG_Q03_37	core	926	2.6	1.2	1506	4.84+05	0.039	12.1	0.95	15.7	1.27	64.9	17.18	180.2	56.4	214	21.3	359.4	37.6	103.30	59.9	91.8	1.01600334	686.43	
MG_Q03_69	rim	926	3.2	8.3	1194	4.80+05	0.064	10.53	0.415	12.1	0.84	48.7	14.04	140	46.48	180.6	36	305.2	50	1017.0	61.8	101.6	0.91597809	724.70	
MG_Q03_24	core	926	3.4	6.27	768	4.86+05	0.028	11.82	0.88	11.82	0.88	11.82	0.88	11.82	0.88	11.82	0.88	11.82	0.88	11.82	0.88	11.82	0.88	11.82	0.88
MG_Q03_99	rim	926	2.4	9.1	1390	4.83+05	0.066	23.95	0.181	6.93	0.444	37.1	11.99	139	49.6	211.1	45.3	42.1	72.3	116.00	159.9	248.9	0.93904139	732.60	
MG_Q03_108	rim	926	-0.7	12.2	551	4.88+05	0.009	7.31	0.067	3.06	0.284	18.2	5.47	60.4	20.36	84	17.24	156.9	25.37	107.00	28.55	63.8	1.08435983	758.61	
MG_Q03_117	rim	926	7.6	7.68	760	4.89+05	0.026	11.82	0.88	11.82	0.88	11.82	0.88	11.82	0.88	11.82	0.88	11.82	0.88	11.82	0.88	11.82	0.88	11.82	0.88
MG_Q03_147	rim	926	-17	12.6	667	5.01+05	0.008	8.19	0.07	4.04	0.453	19.9	6.61	72.7	24.99	99.8	20.63	177.4	31.4	106.10	35.71	120.8	1.10037055	761.55	
MG_Q03_9	core	927	-1.5	5.9	738	4.70+05	0.046	11.55	0.105	5.03	0.843	24.9	7.61	81.3	28.12	114.2	21	209.7	34	110.00	48.7	124.3	0.77852021	696.47	
MG_Q03_112	core	927	-1.2	4.48	1442.9	4.88+05	0.022	6.91	0.24	6.91	0.24	6.91	0.24	6.91	0.24	6.91	0.24	6.91	0.24	6.91	0.24	6.91	0.24	6.91	0.24
MG_Q03_119	core	927	1.5	7.63	623	4.78+05	0.019	11.56	0.091	3.77	0.268	20.3	6.62	71.7	24.52	100.8	20.07	185.8	30.88	115.20	45.5	119.3	0.88254254	717.59	
MG_Q03_157	rim	927	2.8	22.3	1145	4.89+05	0.057	6.01	0.455	13.1	2.83	54.8	13.7	137.8	42.4	157.9	31.1	267.3	43.6	95.10	20.49	26.56	1.34830886	816.58	
MG_Q03_5	core	928	1.5	21	82	4.92+05	0.005	6.91	0.24	6.91	0.24	6.91	0.24	6.91	0.24	6.91	0.24	6.91	0.24	6.91	0.24	6.91	0.24	6.91	0.24
MG_Q03_63	core	929	3.2	8.26	812	4.75+05	0.044	10.13	0.36	7.51	0.52	28.9	8.54	94.1	30.79	124	24.92	216.9	35.9	103.70	48.9	90.5	0.9168005	724.29	
MG_Q03_75	rim	930	1.7	13.6	678	4.74+05	0.015	7.42	0.149	4.94	0.46	23.9	8.84	74.5	24.6	95.9	20.4	179	30.5	99.70	31.8	5.88	1.13358391	768.59	
MG_Q03_79	rim	930	2.1	7.2	148	4.90+05	0.024	10.27	0.236	8.36	0.68	34.6	10.21	103.2	34.6	134.2	27	240	38.5	103.20	52	91.3	0.84509804	710.41	
MG_Q03_79	rim	931	3	11.7	781	4.79+05	0.025	7.46	0.144	5.63	0.7	28.1	8.44	90.5	28.87	111.5	22.63	201.5	33.9	104.30	37.93	62.8	1.06818586	754.81	
MG_Q03_120	core	931	0.9	12.9	545	4.90+05	0.023	7	0.057	3.25	0.239	16.8	5.34	56.3	20.3	82	16.97	152.3	25.59	105.10	26.95	59.9	1.11058971	763.71	
MG_Q03_120	core	931	1.3	7.2	939	4.88+05	0.005	6.79	0.057	4.58	0.167	13.3	81.1	81.1	81.1	81.1	81.1	81.1	81.1	81.1	81.1	81.1	81.1	81.1	81.1
MG_Q03_129	core	931	-2	12.08	480	4.82+05	0.034	6.73	0.043	2.66	0.284	14.45	5.48	50.7	17.67	72.6	10.59	136.7	23.56	102.20	24.92	64.2	0.53	1.08206693	757.71
MG_Q03_64	core	932	-1.7	5.8	1164	4.83+05	0.023	9.59	0.48	11.3	0.81	44.7	13	137	43.5	174.8	34.4	290.5	48.1	98.40	54.7	84.3	0.76142399	695.10	
MG_Q03_112	core	932	1.1	15.1	63	4.88+05	0.036	4.58	0.046	4.58	0.046	4.58	0.046	4.58	0.046	4.58	0.046	4.58	0.046	4.58	0.046	4.58	0.046	4.58	0.046
MG_Q03_4	rim	933	1.3	13.2	329	4.90+05	0.033	4.66	0.046	4.66	0.046	4.66	0.046	4.66											

AKH_27_6	rim	394.4	36.8	3.66	825	4.78E+05	0.192	1431	0.021	1.37	1.061	11.72	4.96	66.67	27.97	131.8	32.36	313.9	56.2	11790	283.1	774	0.56348109	659.56		
AKH_27_7	rim	464.1	86.1	3.11	3.32	685	5.02E+05	0.152	1627	1.307	1.265	38.2	17.6	22.6	90.7	44.6	102.8	106.3	106.3	106.3	106.3	106.3	106.3	106.3		
AKH_27_17	rim	426.6	86.1	3.11	3.32	685	5.02E+05	0.152	1627	1.307	1.265	38.2	17.6	22.6	90.7	44.6	102.8	106.3	106.3	106.3	106.3	106.3	106.3	106.3		
AKH_27_28	rim	427.1	783	87.3	1364	4.88E+05	3.03	1.58	0.067	4.1	3.008	26.8	11.42	13.85	46.9	186	37.56	334.7	61.5	1080	29.63	361	1.9410424	975.32		
AKH_27_24	rim	484	81	8.1	1571	4.81E+05	0.345	0.796	0.00589	1.41	1.058	12.44	7.17	11.58	5.6	281.2	74.5	770	144.3	1240	15.32	212.1	0.9180701	124.70		
AKH_27_12	rim	436.0	190	17.4	1014	4.81E+05	0.249	1.1	0.088	1.8	1.36	21.1	8.17	16.1	15.1	15.1	15.1	15.1	15.1	15.1	15.1	15.1	15.1	15.1		
AKH_27_9	rim	418	14.1	4.78	1776	4.88E+05	0.022	1.49	0.011	1.37	1.18	14.8	8.25	13.01	62.6	306.4	7.59	751	133	12880	19.36	1017	0.6794279	679.84		
AKH_27_41	rim	448.1	129	30.5	3760	4.98E+05	1.2	31.9	0.19	9	0.97	5.3	23.02	32.3	130	612	139.7	1311	225.6	2000	6.36	450	1.4482994	843.3		
AKH_27_25	rim	448.2	133	15.62	4045	4.81E+05	0.249	1.1	0.088	1.8	1.36	21.1	8.17	16.1	15.1	15.1	15.1	15.1	15.1	15.1	15.1	15.1	15.1	15.1		
AKH_27_28	rim	446.2	7.9	4.23	804	4.94E+05	0.051	0.1625	0.096	3.65	3.14	19	6.49	75.6	29.2	129.9	28.2	267	48.7	10100	48.4	74.5	0.6234037	670.45		
AKH_27_28	rim	407.7	1.9	6.05	982	4.85E+05	0.056	0.076	0.029	2.51	2.066	16	6.52	84.9	33.77	155.5	35.87	399.6	58.7	11380	226.8	459.1	0.1735737	608.50		
AKH_27_27	rim	448.9	1480	12.1	1538	4.69E+05	8.9	1.28	0.0076	1.28	1.164	12.54	7.21	114.4	53.3	268	7.1	743	133.6	1400	17.9	14	0.6890511	1022.74		
AKH_27_29	rim	448.2	9.2	4.6	699	4.82E+05	0.072	0.155	0.077	3.07	2.174	12.6	6.41	60.5	23.68	113.7	26.9	255.6	47.6	11000	138.8	48.7	0.6277283	678.87		
AKH_27_53	rim	448.3	13.3	5.02	3420	5.02E+05	0.025	2.45	0.055	3.21	0.389	32.9	16.7	263	114.3	579	138.2	1329	239	12540	48.5	327	0.7070737	683.66		
AKH_27_26	rim	448.6	12.9	4.26	2666	4.89E+05	0.008	1.82	0.032	2.5	0.228	21.7	12.71	202.8	93.5	402	153.3	1380	212.7	13170	32.35	242	0.629406	670.88		
AKH_27_20	rim	450.1	7.9	6.22	2854	4.84E+05	0.106	22.5	0.183	8.39	0.66	53.5	21.02	263.4	102.3	460	103	949	165.5	11610	446	554	0.7937908	700.74		
AKH_27_48	rim	450.2	0	6.44	1403	4.96E+05	0.13	34.7	0.059	4.05	1.24	25.2	98.7	125	49.5	224.3	49.7	480	85.5	11500	218	250	0.8088887	703.56		
AKH_27_20	rim	451.2	12.2	6.32	2143	4.89E+05	0.134	48	0.173	6.49	0.205	40.3	15.3	188.2	72	326	71	644	107.4	11740	1528	1554	0.8002708	702.03		
AKH_27_20	rim	452.6	47	11.5	1956	4.90E+05	0.6	36.2	1.99	9.2	1.99	44.4	15.17	183	71.8	308	68.4	633	108.4	10550	1360	176.2	1.0606974	753.26		
AKH_27_30	rim	452.7	17.4	3.69	2624	4.82E+05	0.196	1.7	0.0113	2.22	0.225	21.8	12.85	196.8	90.2	457	120.3	1231	222.1	13600	273.9	401.9	0.5670237	660.16		
AKH_27_22	rim	452.8	2360	28	1438	4.69E+05	1.11	1.26	0.031	1.08	0.218	11.45	7.2	102.8	41.9	235	57.8	573	104.7	12580	146.5	174.1	1.4475203	801.26		
AKH_27_22	rim	456	179	7.55	1462	4.71E+05	1.47	7.2	5.3	11.1	0.55	36.3	12.1	142	52.3	225.7	49.1	446	73.9	10480	677	715	0.8779495	716.71		
AKH_27_19	rim	458.7	400	49	3458	4.89E+05	1.2	28	0.069	4.18	0.363	33.1	17.39	264.7	115.4	569	140.5	1323	234	12350	45.5	326.6	1.6910908	902.83		
AKH_27_25	rim	458.8	17.3	3.71	3794	4.83E+05	0.024	2.58	0.061	2.57	0.335	33.8	18.91	295.1	125.6	642	170	1743	310	13860	81.1	912	0.5697391	607.58		
AKH_27_52	rim	461.9	17.8	7.82	2145	4.97E+05	0.204	11.2	0.101	3.49	0.541	25.9	12.12	169	75.4	372.4	87.6	865	152.2	11750	109.8	274	0.8932075	719.66		
AKH_27_16	rim	463.3	3.7	6.44	526.3	4.86E+05	0.036	11.08	0.032	0.84	0.073	7.69	3.19	43.9	18.16	85.5	17.2	197.7	132.1	11320	86.5	218.7	0.8088887	703.56		
AKH_27_49	core	465.8	27.6	6.97	2332	4.97E+05	0.038	2.63	0.016	2.7	0.222	23.4	11.99	175.1	74.7	379	90.2	854	154	12000	26.99	156.5	0.8423232	731.06		
AKH_27_49	core	466.6	2.7	5.28	625	4.78E+05	0.007	11.48	0.031	1.2	1.027	8.91	3.77	48.2	100.2	23.52	225.4	384.6	11030	148.9	384	108	0.7263392	687.63		
AKH_27_39	core	471	0.7	3.75	1290	4.88E+05	0.094	22.5	0.37	3.87	1.178	18.5	7.55	106.8	43	206.9	48.6	465	81	11580	889	1203	0.5740317	661.37		
AKH_27_34	core	493	12.2	5.71	1484	4.97E+05	0.044	8.8	0.058	5.0	1.7	24.1	8.89	117.1	44.1	50.1	50.1	50.1	50.1	11670	83.8	108	0.7566611	694.30		
AKH_27_47	rim	512	83	48	1001	4.88E+05	3.06	13.56	0.099	2.35	0.37	13.3	18.2	34.2	172.8	40.8	412	79.9	13420	192	70	1.6112421	690.48			
AKH_27_45	rim	816	84	8.3	2140	4.97E+05	0.3	10.2	0.29	13.4	2.17	64.1	23.7	76.4	314	65.6	594	95.9	8710	182.1	191	0.9190780	724.70			
AKH_27_34	core	829	2.7	6.97	475	4.78E+05	0.071	6.21	0.1	3.38	0.405	16.9	6.17	7.1	26.8	114.5	25.5	230.5	41.4	9860	19.69	505	0.8442278	731.06		
AKH_27_22	core	848	2.9	8.44	516	4.81E+05	0.078	6.21	0.1	3.38	0.405	16.9	6.17	7.1	26.8	114.5	25.5	230.5	41.4	9860	19.69	505	0.8442278	731.06		
AKH_27_5	core	1046	-0.7	10.9	416	4.82E+05	0.056	4.81	0.055	3.52	0.28	13.4	41.9	14.7	57.6	12.88	11.6	18.7	9080	29.3	10	0.8506642	711.47			
AKH_27_3	core	1087	7.3	7.33	885	4.85E+05	0.098	10.78	0.059	2.43	0.152	14.6	6.39	79.6	31.7	148.1	34.3	338.7	59.6	10110	71.7	187.2	0.8651039	714.24		
AKH_27_18	rim	1189	-4.1	5.44	1252	4.87E+05	0.045	11.48	0.087	2.43	0.158	19.8	6.42	11.7	48.2	17.2	133.7	14.8	1137	14.8	14.8	14.8	14.8	14.8		
AKH_27_21	rim	837	252	36.7	1052	4.81E+05	6.63	13.37	0.504	2.54	0.589	14.6	6.33	85.3	34.09	160	42.72	449	87.1	12780	139.9	1720	1.5646606	806.62		
AKH_27_25	rim	1118	61	13.7	844	4.84E+05	0.188	0.071	0.072	1.41	0.17	37.2	17.1	16.1	29.9	17.7	36.7	34.9	100.3	37	100.3	37	100.3	37		
AKH_27_37	rim	857	10020	840	1070	4.30E+05	25.7	8.8	0.194	3.65	0.312	22.6	7.86	36.1	169.1	37.7	360	67.1	9520	126.5	336	2.9242729	1373.14			
AKH_27_42	rim	633	220	23.7	1051	4.99E+05	0.67	7.6	0.076	3.1	0.149	16.9	6.56	87.1	34.4	166.9	39.1	382	72.5	10970	79.4	41	1.3747485	822.80		
AKH_27_46	rim	1017	5.1	8.4	4445	4.81E+05	1.1	1.6	0.046	4.44	0.355	26.2	11.6	151.6	59.2	277.2	59.3	564	96.5	10370	74.2	208.1	1.1038037	762.28		
AKH_27_60	rim	2020	7.77E+04	960	0.61	1.3	15.37	0.048	0.048	0.048	0.048	0.048	0.048	0.048	0.048	0.048	0.048	0.048	0.048	0.048	0.048	0.048	0.048	0.048		
AK_01	core	1172	740	30.2	2850	4.55E+05	1.71	20	65.4	5.7	32.3	30.1	4.09	75.8	24.4	275	97.2	408	83.1	710	119.4	9360	390	964	1.4500004	848.26
AK_01_3	rim	189.2	4.9	11.68	571	4.87E+05	0.081	0	12.29	0.34	0.45	13.2	0.247	9.49	3.4	45.8	18.5	87.3	21.06	204.9	38.4	10680	30.44	44	1.0674424	754.66
AK_01_3	rim	190.0	39.5	16	3966	4.80E+05	0.076	0.0081	0.642	0.147	32.1	7.74	0.347	9.49	3.4	45.8	18.5	87.3	21.06	204.9	38.4	10680	30.44	44	1.0674424	754.66
AK_01_4	rim	670.1	6.47	6.67	666	4.89E+05	0.045	0.086	0.82	0.50	0.66	98.2	14.2	14.9	60.3	24.52	63.0	14.2	13.9	66.0	26.1	62	0.707466	707.60		
AK_01_5	rim	1020	-3	8.43	708	4.78E+05	0.05	0.0023	14.42	0.05	0.99	2.1	0.122	13.8	5.09	6.3	25.11	108.5	24.27	227.7	40.1	10970	74.7	269.2	0.9258257	726.03
AK_01_6	rim	191.6	7.6	14.2	1444	4.80E+05	0.208	5.9	32.1	2.24	12.3	7.75	1.8	35.2	11.2	133.3	47.8	214.2	47.3	493.3	78.7	91.7	0.1522834	727.61		
AK_01_7	rim	191.6	7.6	14.2	1444	4.80E+05	0.208	5.9	32.1	2.24	12.3	7.75	1.8	35.2	11.2	133.3	47.8	214.2	47.3	493.3	78.7	91.7	0.1522834	727.61		
AK_01_8	rim	175.7	3.7	12.16	2805	4.84E+05	0.214	0.036	3.13	0.563	9.31	15.2	8.7	72.3	22.71	258.8	95.4	424.8	92.8	883	165.8	7200				

AK_01_133	core	5937	2140	25	1200	4.41e+05	8.19	0.543	19.5	0.293	2.86	4.75	0.407	25	8.23	105	41.7	199.1	48.6	494	94.9	10280	438	1167	1.39794001	828.31		
AK_01_141	core	1817	470	16.7	678	4.71e+05	2.5	0.31	10.5	0.095	0.69	2.96	0.77	15.9	5.6	10.2	10.8	22.1	208.3	37.1	9640	28.4	8163	1.22277647	797.883			
AK_01_155	core	1889	8	8.92	499	0.000000	0.182	0.066	13.38	0.0000	0.47	1.89	2.19	0.49	0.49	40.3	16.78	81.1	19.52	194.9	37.4	12050	31.72	57	0.000000	37.4		
AK_01_166	core	1001	81	6.41	2820	4.88e+05	0.07	0.108	14.71	0.147	7.58	16.1	3.75	19.4	29.4	310	101.8	40.4	78.4	663	110.7	9290	435	403	0.86880803	703.18		
AK_01_117	core	1857	10	15.44	1107	4.65e+05	0.056	0.073	16.20	0.097	2.73	4.62	0.67	23.6	7.96	10.3	36.7	17.2	26.2	68.7	65.7	9760	64.2	61.7	1.18792029	862.36		
AK_01_118	core	1001	19	10.71	757	0.81e+05	0.107	0.005	24.2	0.043	0.66	2.08	0.56	11.76	27.3	270	48.3	11.70	27.0	77.0	270	48.3	11.70	27.0	77.0	0.81789747	768.89	
AK_01_119	rim	1047	19.5	12.2	1266	5.19e+05	2.8	0.068	16.2	0.22	1.92	3.9	0.19	2.0	8.02	106.5	41.7	219.3	53.8	576	95.3	14190	132.9	416	1.08639883	758.61		
AK_01_120	core	491	37.8	16	798	4.50e+05	1.48	0.62	7.4	0.75	6.3	4.24	1.05	16.4	4.74	60.9	26.4	143.2	38.4	449	94.5	15410	84.7	2045	1.20413998	783.89		
AK_01_121	core	1857	0.4	13.82	827	4.84e+05	0.073	0.046	17.43	0.045	1.06	3.05	0.36	15.2	5.89	70.7	27.49	129.5	29.25	279.6	51.4	10410	50.1	63.1	1.12329808	768.88		
AK_01_122	core	1000	22.8	11.9	1283	4.77e+05	0.153	0.027	15.64	0.072	14.3	3.41	0.186	2.4	9.13	117.1	45.4	201.8	44.3	409.3	71.9	11240	139.1	448	1.07546696	756.35		
AK_01_123	core	1292	160	24.3	622	4.76e+05	0.152	0.0036	8.64	0.039	0.63	1.8	0.197	11.8	4.21	56.1	21.54	99.8	22.37	218.2	38.95	10790	52.5	178.1	1.38560627	825.37		
AK_01_124	rim	1295	12	19.9	1185	4.83e+05	0.093	0.3	36.7	0.37	1.44	0.3	0.241	25.5	7.4	20.1	25.4	39.3	28.1	114.7	20.2	11470	20.2	11470	20.2	11470	0.81789747	768.89
AK_01_125	rim	1864	5.4	4.66	1124	4.90e+05	0.238	0.0054	35.35	0.037	1.11	2.97	0.52	17.4	6.98	95.9	18.26	42.8	41.6	71.1	11520	64.9	71.1	11520	64.9	71.1	0.66888592	677.87
AK_01_126	core	1794	0.3	9.4	950	4.80e+05	0.105	0.1	20.94	0.036	0.92	2.39	0.5	16.2	6.03	81.2	32.25	152.5	35.5	345.1	64.1	10960	65.2	63.4	0.97497199	675.78		
AK_01_127	core	1120	21	11.06	920	4.79e+05	0.071	0.0088	14.05	0.223	1.75	3.4	0.81	12.2	7.02	86.4	33.21	150.9	33.4	321.1	59.3	9430	62.1	48.7	1.18792029	780.25		
AK_01_128	rim	193.3	51.4	14.5	982	4.88e+05	0.32	0.1	12.99	0.085	1.86	3.21	0.64	19.3	6.81	85	33.4	151.5	35.2	339.4	63	10250	35.5	45.5	1.161368	774.57		
AK_01_129	core	1867	0.4	9.04	1225	4.88e+05	0.038	0.1	16.77	0.045	11.7	3.43	0.64	23.7	8.59	106.7	42.1	187.2	40.9	391.5	69.3	10520	65.1	61.8	0.95616843	732.03		
AK_01_130	core	1001	24.49	24.12	4.82e+05	0.057	0.015	16.91	0.218	4.25	7.66	1.34	29.5	11.33	123.6	48.2	218.8	46.1	42.1	75.7	9790	66.9	61.9	1.16108239	774.50			
AK_01_131	core	1872	13.5	14.2	885	4.76e+05	0.045	0.1	11.92	0.027	10.5	3.03	0.5	16.8	6.03	75.8	29.7	136.4	30.7	294.5	54.7	1070	32.9	40.6	1.15228384	772.61		
AK_01_132	core	1006	3.2	11.31	1598	4.74e+05	0.192	0.176	28.81	0.144	1.76	4.15	0.429	27.2	10.77	139.4	54.7	246.9	56.5	525	93.2	12150	202.3	59.1	1.03436226	751.76		
AK_01_133	core	184.5	-0.5	16.5	835	4.83e+05	0.05	0.1	15.69	0.088	1.32	3.56	0.54	17.3	5.89	74.6	28.26	127.2	29.4	278.8	51.7	9850	46.76	61.6	1.21748394	786.84		
AK_01_134	rim	183	-2.7	1.96	2690	4.87e+05	0.146	3.78	3.7	1.52	9.1	5.05	0.54	31.2	13.17	195	88.2	446	112.4	1169	22.3	12560	123.4	382.8	0.29252607	615.31		
AK_01_135	core	1804	1.6	10.8	1997	4.84e+05	0.065	0.31	13.8	0.23	1.63	3.45	0.59	25.1	10.28	151.2	63.1	321	78.8	784	152.7	11520	86.7	257	0.03134276	576.83		
AK_01_136	core	185.5	-1.7	15.4	978	4.79e+05	0.071	0.0088	14.05	0.223	1.75	3.4	0.81	12.2	7.02	86.4	33.21	150.9	33.4	321.1	59.3	9430	62.1	48.7	1.18792029	780.25		
AK_01_137	core	1869	8	11.17	750	4.89e+05	0.363	0.1	14.75	0.04	0.6	1.83	0.388	12.6	4.4	60.1	24.24	115.8	27.25	270.8	51	11330	37.3	48.5	1.05604885	752.47		
AK_01_138	rim	909	4.4	13.7	821	4.82e+05	0.135	0.312	22.53	0.161	1.19	2.71	0.61	16.6	5.54	72.2	28.2	131.4	30.2	292.6	55.8	12530	73.4	198	1.05757606	752.23		
AK_01_139	core	1026	209	887	1214	4.92e+05	0.73	0.105	13.3	0.086	1.63	4.4	0.75	25.7	8.84	111.9	42	188.2	40.4	383	69.2	10700	43.4	99.6	0.94792362	730.39		
AK_01_140	core	1027	490	29	884	4.91e+05	0.404	0.043	18.99	0.105	1.54	3.48	0.83	19.8	8.82	83	31.1	133.3	29.3	277	49.5	10060	35.2	64.7	1.4822898	843.92		
AK_01_141	rim	1857	2	10.02	2470	4.79e+05	0.236	0.92	37.9	1.14	12.6	14.9	5.45	66.2	20.9	42	88.5	383	81.3	738	135.4	10310	112.4	86.9	1.0008772	741.00		
AK_01_142	core	155	1980	100	149	4.89e+05	15.2	5.7	31.3	2.2	10.2	10.2	10.2	10.2	10.2	10.2	10.2	10.2	10.2	10.2	10.2	10.2	10.2	10.2	10.2	1.0008772	741.00	
AK_01_143	core	1318	6	12.8	2740	4.80e+05	0.1	0.0015	8.45	0.205	5.13	12.3	0.62	27.7	24.8	188	95.3	396	78.6	697	112.9	11400	161	349	1.10720939	800.41		
AK_01_144	rim	190	107	1.55	1400	5.12e+05	0.17	3.7	21.5	1.99	13.3	4.9	0.68	19.7	6.83	98.3	45.3	240.6	60.5	680	140.5	11400	61.3	61.8	0.190317	599.75		
AK_01_145	rim	187.1	-1.3	13.1	827	4.96e+05	0.074	0.0087	16.8	0.042	0.9	2.48	0.379	14.16	5.16	70.3	27.38	129	29.19	282.2	54	11010	45.6	107.6	1.1172713	765.13		
AK_01_146	core	185.5	-0.5	16.5	835	4.83e+05	0.05	0.1	15.69	0.088	1.32	3.56	0.54	17.3	5.89	74.6	28.26	127.2	29.4	278.8	51.7	9850	46.76	61.6	1.21748394	786.84		
AK_01_147	core	1007	-1.7	15.52	683	4.89e+05	0.034	0.0045	9.85	0.043	0.94	2.14	0.307	14.8	5.11	61.2	23.69	103.8	22.96	214.8	37.7	10640	38.1	105.7	1.1980871	780.60		
AK_01_148	rim	1008	-5.9	9.7	1362	4.93e+05	0.084	0.0086	14.71	0.033	0.87	2.29	0.455	17	6.89	93.8	38.61	186	44.1	439.8	83.2	11560	209.9	64.9	0.99869516	749.56		
AK_01_149	core	184.5	-0.5	16.5	835	4.83e+05	0.05	0.1	15.69	0.088	1.32	3.56	0.54	17.3	5.89	74.6	28.26	127.2	29.4	278.8	51.7	9850	46.76	61.6	1.21748394	786.84		
AK_01_150	core	973	-2.5	12.3	720	5.09e+05	0.057	0.0086	14.64	0.085	1.14	3.62	1.15	19.5	6.01	72	27.7	117.8	25.4	236.8	43	10340	23.7	48.8	1.0899011	759.35		
CVI-1344																												
AK_1344_48	core	64.1	10	9.3	1022	4.86e+05	0.172	0.1	1.41	0.042	1.74	0.5	14.2	5.47	76.8	12.84	160.1	38.88	423	84.6	8230	81.9	429	0.0684289	734.48			
AK_1344_5	core	64.3	7.3	9.3	990	4.87e+05	0.021	1.17	0.042	1.07	0.47	35.1	54.9	75.1	32.5	160.7	37.7	397	81.5	760	76.9	859	340	488	0.0684289	734.48		
AK_1344_4	core	67.2	46	4.82	1422	4.84e+05	0.061	2.67	0.069	2.5	0.57	20.8	8.06	110.3	47.2	237.3	56.9	595	122.4	9300	145.8	61.5	68.304704	680.49				
AK_1344_36	core	67.3	6.8	11.57	1094	4.92e+05	0.01	1.31	0.031	1.66	0.51	14.6	5.97	80.2	35	173.3	41.82	440.9	8.96	7720	89.4	402.1	1.06333336	753.80				
AK_1344_8	core	75	3	10.8	1163	4.91e+05	0.01	1.64	0.043	1.93	0.88	1.9	0.89	5.5	10.8	10.8	10.8	10.8	10.8	10.8	10.8	10.8	10.8	10.8	10.8	1.4822898	843.92	
AK_1344_20	rim	67.7	3.4	9.24	961	4.81e+05	-0.009	0.0083	1.248	0.028	1.35	0.394	11.7	4.76	70.5	30.92	156.3	37.92	401.1	82.1	8180	76.5	397.7	0.96567199	732.92			
AK_1344_35	rim	68	210	7.4	1268	4.76e+05	0.77	2.76	0.058	2.08	0.5	17	6.61	96	10.5	40.27	196.7	47.4	484	1000	10520	138.7	67.2	0.96567199	732.92			
AK_1344_38	core	69	23	6.0	1000	4.91e+05	0.166	1.91	0.066	1.91	0.066	1.91	0.066	1.91	0.066	1.91	0.066	1.91	0.066	1.91	0.066	1.91	0.066	1.91	0.066	1.91	0.066	1.91
AK_1344_50	core	68.1	-0.1	6	1285	4.88e+05	0.044	2.18	0.031	3.2	0.52	20.6	8.17	106.3														

AP_038_1	rim	57.7	270	132	1330	4.33E+05	5.07	4.44	0.69	5.77	0.9	26.2	11.7	125	34.4	137	32.7	335	53.6	2.68E+04	74.2	3270	2.12057393	1032.96
AP_038_2	core	61.7	2.31E+04	29.8	2970	3.95E+05	8.6	3.45	0.369	4.64	0.75	40.6	20.7	260	85	374	103	1120	204	22260	93	4730	1.47421626	846.83
AP_038_2	core	63.8	2590	80	8830	4.83E+05	12.4	36	4.36	43	7.6	179	71.1	820	268	1216	305	3260	595	21330	670	9680	1.95424251	979.39
AP_038_10	core	69.1	2920	28.2	2440	4.32E+05	15.9	8.9	1.44	15.8	1.98	71.7	27.2	254	60	206.8	47.2	482	66.8	34250	127.8	5740	1.45024911	840.95
AP_038_28	core	267.1	132	11.3	2123	4.72E+05	0.614	104.4	0.59	11.32	2.68	58	17.01	205.5	75.3	321.9	69.1	639	114.6	9130	1358	903	1.05307844	751.68
AP_038_26	core	275	64	15.3	1709	4.70E+05	0.301	61.1	0.292	817	2.56	41	13.28	156.2	59.8	255	57.8	554	97.8	7950	629	530	1.18466143	779.63
AP_038_27	rim	298.8	-0.1	10.75	936	4.73E+05	0.138	35.7	0.12	4.16	1.99	22.1	7.44	85.7	32.54	144.9	33.3	335	61.7	7280	392	310	1.03140846	747.22
AP_038_24	core	1000	17.4	29.4	1577	4.74E+05	0.045	4.8	0.51	2.06	0.233	18.2	8.76	122.8	51.3	247.9	63.2	666	118.3	13860	64.9	1408	1.46834733	845.39
AP_038_43	rim	1004	16.6	38.4	1479	4.74E+05	0.001	3.07	0.064	3.3	0.161	23.4	9.47	125.2	49.7	213.7	47.3	413	70.2	12890	85.4	594	1.58433132	874.70
AP_038_23	rim	1017	12.6	30.8	1347	4.82E+05	0.029	2.38	0.065	3.06	0.074	24.8	9.88	121.7	44.5	186.4	38.9	352	58.9	12890	85.8	297.2	1.48855072	850.38
AP_038_42	core	1170	3.2	36.3	1810	4.72E+05	0.037	18.66	0.133	5.14	1.04	34	12.48	158.4	62.9	289.5	66	641	116.9	9630	188	357	1.55990663	868.40

TABLE A3. Representative electron microprobe analysis of garnet, biotite, and plagioclase in the La Seta creta garnet-bearing mylonite samples.

CVI-1388 Amphibolite																															
Hornblende																				Plagioclase											
No Spot.	13	14	15	16	17	18	19	21	22	23	24	25	51	52	53	Average	No Spot.	54	55	56	58	59	60	61	Average						
SiO2	40.83	40.89	40.88	40.30	40.46	40.24	41.07	40.77	40.65	40.18	40.50	41.29	40.71	41.85	42.09	40.89	SiO2	59.22	59.07	60.13	60.23	58.17	58.78	58.46	59.15						
TiO2	0.34	0.31	0.36	0.45	0.40	0.51	0.11	0.26	0.51	0.52	0.30	0.36	1.47	1.21	0.99	0.86	TiO2	0.02	0.00	0.02	0.02	0.00	0.00	0.00	0.01						
Al2O3	17.39	17.59	17.22	18.40	18.06	18.57	17.75	18.80	18.05	18.32	18.80	18.89	18.05	18.25	14.79	17.85	Al2O3	25.21	25.77	24.87	25.09	26.17	26.17	25.90	25.58						
FeO	18.25	18.20	18.27	18.61	18.76	18.48	18.30	18.60	18.30	18.60	18.70	18.28	18.28	18.28	18.28	18.28	FeO	0.05	0.00	0.01	0.01	0.01	0.01	0.01	0.01						
MnO	0.56	0.46	0.44	0.07	0.09	0.10	0.10	0.07	0.07	0.15	0.07	0.12	0.10	0.08	0.09	0.09	MnO	0.03	0.02	0.01	0.01	0.01	0.01	0.01	0.02						
MgO	6.46	6.15	6.52	7.17	6.08	5.95	6.06	5.91	6.05	5.98	6.06	6.46	6.04	7.40	7.39	6.28	MgO	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00						
CaO	11.57	11.54	11.88	11.50	11.51	11.51	11.51	11.61	11.61	11.61	11.56	11.58	11.43	11.48	11.54	11.54	CaO	7.01	7.78	6.20	6.64	7.98	7.65	7.12	7.25						
Na2O	1.37	1.34	1.38	1.32	1.36	1.38	1.46	1.34	1.32	1.32	1.35	1.41	1.31	1.19	1.29	1.34	Na2O	7.80	7.67	8.08	8.15	7.12	7.41	7.36	7.66						
K2O	0.03	0.75	0.00	0.02	0.78	0.00	0.11	0.82	0.76	0.82	0.75	0.65	0.81	0.80	0.78	0.80	K2O	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00						
Cr2O3	0.00	0.01	0.00	0.01	0.02	0.04	0.00	0.02	0.00	0.01	0.01	0.01	0.01	0.02	0.01	0.01	Cr2O3	0.00	0.01	0.00	0.00	0.00	0.00	0.00	0.00						
Total	97.00	96.85	97.26	97.17	96.95	97.74	97.90	98.22	97.31	98.21	97.27	96.52	99.96	96.96	97.07	97.36	Total	99.41	99.89	99.70	100.24	99.68	100.16	99.81	99.84						
Y	13.7558	3.4036	2.0987	1.8472	5.4007	5.0935	2.062	6.4249	5.939	6.2449	5.939	5.7933	5.5653	5.8703	5.7339	5.7339	Y	14.5321	14.5321	14.5321	14.5321	14.5321	14.5321	14.5321	14.5321						
Z	10.7271	10.7271	10.7271	10.7271	10.7271	10.7271	10.7271	10.7271	10.7271	10.7271	10.7271	10.7271	10.7271	10.7271	10.7271	10.7271	Z	10.7265	10.7265	10.7265	10.7265	10.7265	10.7265	10.7265	10.7265						

Garnet																																								
Profile 1																				Profile 2																				
No Spot.	104	105	106	107	108	109	110	111	112	113	114	115	116	117	118	119	120	121	122	123	No Spot.	124	125	126	127	128	129	130	131	132	133	134	135	136	137	138	139	140		
SiO2	36.90	36.78	36.96	36.88	36.92	36.88	36.79	37.04	36.80	36.76	36.76	36.90	36.87	36.76	36.74	36.84	36.68	36.54	36.36	36.14	SiO2	36.71	36.55	36.79	36.85	36.92	36.62	36.77	36.91	36.92	36.83	36.92	36.85	36.90	36.76	37.00	36.86	36.47		
TiO2	0.11	0.12	0.00	0.00	0.10	0.09	0.11	0.11	0.09	0.09	0.10	0.11	0.11	0.09	0.08	0.11	0.09	0.06	0.00	0.02	TiO2	0.01	0.02	0.01	0.00	0.01	0.01	0.01	0.01	0.01	0.01	0.01	0.01	0.01	0.01	0.01	0.01	0.01	0.01	
Al2O3	20.25	20.25	20.32	20.29	20.31	20.31	20.17	20.31	20.31	20.31	20.30	20.34	20.29	20.19	20.19	20.30	20.22	20.28	20.54	20.37	Al2O3	20.27	20.56	20.39	20.21	20.29	20.21	20.09	20.37	20.41	20.38	20.29	20.38	20.37	20.21	20.17	20.31	20.16		
FeO	26.36	26.28	26.19	26.23	26.03	26.06	26.44	26.45	26.45	26.45	26.45	26.45	26.45	26.45	26.45	26.45	26.45	26.45	26.45	26.45	FeO	26.38	26.28	26.41	26.21	26.19	26.19	26.19	26.19	26.19	26.19	26.19	26.19	26.19	26.19	26.19	26.19	26.19	26.19	
MnO	0.61	0.61	0.61	0.62	0.61	0.62	0.61	0.63	0.63	0.63	0.61	0.62	0.61	0.62	0.61	0.62	0.61	0.62	0.61	0.58	MnO	1.53	1.15	1.05	1.07	0.63	0.66	0.66	0.59	0.59	0.53	0.56	0.77	0.18	0.70	0.82	1.24	1.84		
MgO	3.04	3.07	3.10	3.09	3.02	3.15	3.13	3.14	3.18	3.16	3.12	3.13	3.11	3.09	3.06	2.98	2.83	1.88	1.85	1.52	MgO	1.74	1.66	2.65	3.17	3.22	3.17	3.17	3.20	3.27	3.27	3.20	2.80	2.84	2.31	3.04	2.70	1.80		
CaO	11.28	11.02	11.28	11.20	11.24	11.24	11.20	11.24	11.20	11.24	11.24	11.24	11.24	11.24	11.24	11.24	11.24	11.24	11.24	11.24	CaO	9.28	9.28	10.58	10.24	10.53	10.56	10.44	10.56	10.44	10.56	10.44	10.56	10.44	10.56	10.44	10.56	10.44		
Na2O	0.02	0.00	0.01	0.01	0.01	0.02	0.00	0.01	0.02	0.00	0.02	0.01	0.01	0.01	0.00	0.01	0.01	0.01	0.01	0.00	Na2O	0.01	0.01	0.00	0.00	0.00	0.00	0.00	0.01	0.01	0.01	0.01	0.01	0.01	0.01	0.01	0.01	0.01	0.00	
K2O	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	K2O	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	
Cr2O3	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	Cr2O3	0.00	0.02	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	
Total	98.44	98.18	98.49	98.42	98.13	98.21	98.21	98.38	98.47	98.32	98.01	98.55	98.33	98.54	98.30	98.24	98.11	98.40	98.35	98.30	Total	99.32	99.36	99.28	99.33	99.32	99.12	99.29	99.23	99.23	99.23	99.23	99.23	99.23	99.23	99.23	99.23	99.23	99.23	
Y	13.7051	3.4036	2.0987	1.8472	5.4007	5.0935	2.062	6.4249	5.939	6.2449	5.939	5.7933	5.5653	5.8703	5.7339	5.7339	5.7339	5.7339	5.7339	Y	13.7051	3.4036	2.0987	1.8472	5.4007	5.0935	2.062	6.4249	5.939	6.2449	5.939	5.7933	5.5653	5.8703	5.7339	5.7339	5.7339	5.7339		
Z	10.7271	10.7271	10.7271	10.7271	10.7271	10.7271	10.7271	10.7271	10.7271	10.7271	10.7271	10.7271	10.7271	10.7271	10.7271	10.7271	10.7271	10.7271	10.7271	10.7271	Z	10.7271	10.7271	10.7271	10.7271	10.7271	10.7271	10.7271	10.7271	10.7271	10.7271	10.7271	10.7271	10.7271	10.7271	10.7271	10.7271	10.7271	10.7271	10.7271

Garnet																																																		
Profile 3																				Profile 4																														
No Spot.	144	145	146	147	148	149	150	151	152	153	154	155	156	157	158	159	160	161	162	163	No Spot.	164	165	166	167	168	169	170	171	172	173	174	175	176	177	178	179	180	181	182	183	184	185	186	187	188	189	190	191	192
SiO2	36.71	37.34	37.44	37.44	37.40	37.37	37.50	37.39	35.57	37.48	37.34	37.38	37.29	37.38	37.51	37.59	37.46	37.57	37.50	37.45	SiO2	37.23	37.30	37.51	37.67	37.62	37.51	37.55	37.56	37.50	37.59	37.59	37.60	37.61	37.45	37.65	37.73	37.71	37.56	37.33										
TiO2	0.09	0.09	0.09	0.09	0.08	0.08	0.09	0.09	0.09	0.09	0.09	0.09	0.09	0.09	0.09	0.09	0.09	0.09	0.09	0.09	TiO2	0.09	0.09	0.09	0.09	0.09	0.09	0.09	0.09	0.09	0.09	0.09	0.09	0.09	0.09	0.09	0.09	0.09	0.09	0.09										
Al2O3	20.27	20.56	20.39	20.21	20.29	20.21	20.09	20.37	20.41	20.38	20.29	20.38	20.37	20.21	20.17	20.31	20.16	20.31	20.16	20.31	Al2O3	20.28	20.56	20.38	20.22	20.21	20.21	20.09	20.37	20.41	20.38	20.29	20.38	20.37	20.21	20.17	20.31	20.16	20.31	20.16										
FeO	26.38	26.28	26.41	26.21	26.19	26.19	26.19	26.19	26.19	26.19	26.19	26.19	26.19	26.19	26.19	26.19	26.19	26.19	26.19	26.19	FeO	26.38	26.28	26.41	26.21	26.19	26.19	26.19	26.19	26.19	26.19	26.19	26.19	26.19	26.19	26.19	26.19	26.19	26.19	26.19										
MnO	1.53	1.15	1.05	1.07	0.63	0.66	0.66	0.59	0.59	0.53	0.56	0.77	0.18	0.70	0.82	1.24	1.84	1.84	1.84	1.84	MnO	1.53	1.15	1.05	1.07	0.63	0.66	0.66	0.59	0.59	0.53	0.56	0.77	0.18	0.70	0.82	1.24	1.84	1.84	1.84										
MgO	1.74	1.66	2.65	3.17	3.22	3.17	3.17	3.20	3.27	3.27	3.20	2.80	2.84	2.31	3.04	2.70	1.80	1.80	1.80	1.80	MgO	1.74	1.66	2.65	3.17	3.22	3.17	3.17	3.20	3.27	3.27	3.20	2.80	2.84	2.31	3.04	2.70	1.80	1.80	1.80										
CaO	9.28	9.28	10.58	10.24	10.53	10.56	10.44	10.56	10.44	10.56	10.44	10.56	10.44	10.56	10.44	10.56	10.44	10.56	10.44	10.56	CaO	9.28	9.28	10.58	10.24	10.53	10.56	10.44	10.56	10.44	10.56	10.44	10.56																	

		X-ray Map 2 Profile 3																				
No Spot	59	60	61	62	63	64	65	66	67	68	69	70	71	72	73	74	75	76	77	78	79	80
SiO2	37.86	37.83	37.88	37.29	37.33	37.86	37.32	37.41	37.60	37.89	38.03	38.87	38.83	38.76	38.92	38.79	38.99	37.96	37.34	37.13	37.31	37.88
TiO2	0.11	0.07	0.05	0.02	0.00	0.00	0.00	0.01	0.01	0.05	0.06	0.05	0.10	0.08	0.07	0.11	0.09	0.06	0.02	0.04	0.02	0.06
Al2O3	21.26	21.60	21.63	20.88	21.02	20.90	20.99	21.09	20.96	20.99	21.17	21.59	21.50	21.67	21.50	21.55	21.52	21.31	21.06	21.02	21.16	20.90
FeO	28.03	30.65	30.47	30.85	30.43	30.10	30.05	30.22	29.44	28.88	27.52	25.25	22.50	22.32	22.33	22.52	22.61	26.41	30.07	30.87	29.61	29.39
MnO	2.41	3.25	3.52	3.79	4.26	4.22	4.22	4.21	3.14	2.29	2.33	1.04	0.64	0.65	0.73	0.75	0.68	1.71	3.65	3.21	2.47	2.11
MgO	2.81	3.06	3.19	3.38	3.30	3.37	3.36	3.48	4.06	4.64	4.96	6.68	8.02	8.11	8.03	8.07	8.01	5.16	3.94	3.23	2.95	3.12
CaO	7.79	4.57	4.50	3.71	3.90	4.17	4.28	3.96	4.99	5.41	6.19	6.85	7.90	8.00	8.17	7.92	7.90	7.11	4.59	4.42	6.29	6.37
Na2O	0.00	0.02	0.03	0.02	0.00	0.00	0.00	0.00	0.00	0.00	0.02	0.01	0.00	0.02	0.00	0.02	0.01	0.00	0.00	0.00	0.00	0.00
K2O	0.01	0.00	0.00	0.00	0.00	0.01	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.01	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.01
Cr2O3	0.04	0.05	0.04	0.05	0.07	0.05	0.05	0.06	0.07	0.04	0.04	0.04	0.02	0.02	0.06	0.05	0.05	0.02	0.04	0.07	0.02	0.05
Total	100.44	100.32	100.30	100.10	100.33	100.48	100.25	100.46	100.29	100.22	100.11	99.98	99.52	99.62	99.81	99.78	99.86	99.56	100.10	100.05	99.87	99.38
X	-31.9795	-31.9975	-32.0289	-32.0667	-32.1399	-32.184	-32.2205	-32.2461	-32.2664	-32.2816	-32.2834	-32.2816	-32.2813	-32.2858	-32.2764	-32.2785	-32.8166	-32.8646	-32.9301	-32.9767	-33.0425	-33.0633
Y	5.8007	5.8662	5.9798	5.8823	5.895	5.9973	5.9021	5.9014	5.9166	5.1128	5.1163	5.1177	5.1125	5.117	5.1168	5.1168	5.1185	5.1145	5.114	5.1125	5.1117	5.1122
Z	10.7405	10.7425	10.744	10.7445	10.745	10.744	10.7435	10.745	10.744	10.7445	10.744	10.745	10.7435	10.744	10.743	10.743	10.743	10.743	10.743	10.743	10.743	10.743

		Profile 4																			
No Spot	84	85	86	87	88	89	90	91	92	93	94	95	96	97	98	99	100				
SiO2	37.60	38.96	38.86	38.91	38.92	39.13	39.10	39.11	39.11	39.18	39.21	39.15	39.21	39.15	38.50	37.94	37.71				
TiO2	0.02	0.10	0.10	0.12	0.10	0.03	0.09	0.09	0.09	0.07	0.09	0.08	0.08	0.09	0.08	0.01	0.01				
Al2O3	21.20	21.61	21.68	21.64	21.63	21.79	21.63	21.83	21.74	21.78	21.79	21.75	21.79	21.75	21.13	21.13	21.15				
FeO	30.39	32.80	32.00	32.77	33.08	32.51	32.46	32.35	23.48	23.35	23.35	23.50	23.58	22.19	28.84	30.36					
MnO	1.58	1.58	0.61	0.58	0.59	0.50	0.50	0.47	0.43	0.43	0.44	0.44	0.41	0.86	2.70	3.87					
MgO	3.38	3.38	8.07	8.20	8.10	8.10	8.10	8.14	8.93	8.96	9.08	9.07	9.14	9.11	8.39	4.13	3.21				
CaO	4.66	7.80	7.60	7.19	6.83	6.84	6.46	6.44	6.32	6.22	6.22	6.22	6.22	6.11	5.97	4.57					
Na2O	0.00	0.00	0.00	0.01	0.02	0.01	0.01	0.00	0.02	0.00	0.00	0.00	0.00	0.00	0.00	0.00					
K2O	0.01	0.00	0.00	0.01	0.01	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.01	0.00	0.00					
Cr2O3	0.04	0.07	0.04	0.03	0.06	0.04	0.05	0.03	0.07	0.05	0.08	0.07	0.05	0.05	0.06	0.09					
Total	100.88	99.99	100.29	99.75	99.88	100.10	100.13	100.69	100.38	100.27	100.50	100.49	100.86	100.86	100.83	100.83					
X	-20.7463	-20.644	-20.5514	-20.4054	-20.2383	-20.0317	-19.8722	-19.5968	-19.5109	-19.3432	-19.3105	-19.2455	-19.2172	-19.2139	-19.1656	-19.2011					
Y	2.7366	2.8021	2.8757	2.9804	3.0384	3.2616	3.3666	3.0715	3.758	3.9189	3.9816	4.1248	4.2033	4.2365	4.2973	4.3857					
Z	10.76	10.76	10.761	10.761	10.761	10.761	10.761	10.761	10.761	10.761	10.761	10.761	10.761	10.761	10.761	10.761					

CVI-1385 Bt+Ms+Grt Schist

		Biotite				Plagioclase			
No Spot	25	26	27	28	29	30			
SiO2	35.19	26.81	26.88	36.05	37.72	98.61			
TiO2	1.61	0.08	0.07	1.52	0.80	0.01			
Al2O3	18.62	21.27	20.67	17.36	18.21	0.03			
FeO	17.96	21.88	23.71	19.32	17.36	0.23			
MnO	0.07	0.11	0.19	0.16	0.10	0.01			
MgO	11.35	16.37	16.25	10.31	11.91	0.00			
CaO	0.15	0.18	0.09	0.04	0.08	0.00			
Na2O	0.17	0.03	0.00	0.09	0.15	0.00			
K2O	8.74	0.02	0.06	9.01	8.92	0.02			
Cr2O3	0.03	0.02	0.01	0.04	0.00	0.03			
Total	93.90	86.76	87.92	93.91	96.26	98.94			
X	37.4839	37.3558	29.8817	29.5243	31.2525	30.2122			
Y	18.2219	18.5467	17.6441	18.1685	19.2434	19.1264			
Z	10.7715	10.772	10.7805	10.7805	10.7805	10.7805			

		X-ray Map 1 Profile 1																									
No Spot	238	239	240	241	242	243	244	245	246	247	248	249	250	251	252	253	254	255	256	257	258	259	260	261	262	263	264
SiO2	37.17	37.07	37.21	30.75	36.76	36.63	36.68	36.67	36.62	36.77	33.24	36.40	36.42	36.86	36.71	36.68	36.54	36.78	36.67	36.63	36.79	36.73	37.22				
TiO2	0.03	0.05	0.07	0.05	0.06	0.04	0.01	0.02	0.00	0.00	0.00	0.02	0.01	0.01	0.01	0.00	0.00	0.00	0.00	0.01	0.01	0.01	0.00	0.02	0.00	0.01	
Al2O3	20.98	20.88	18.63	20.52	20.42	20.75	20.76	20.61	20.82	20.77	20.73	20.57	18.74	20.70	20.62	20.76	20.80	20.71	20.72	20.71	20.66	20.71	20.73	20.78	20.87	21.08	
FeO	31.96	31.11	31.42	30.12	30.97	30.71	32.55	32.19	32.03	31.36	30.67	30.57	30.69	30.09	30.74	30.51	30.47	30.67	30.50	30.24	30.14	30.60	30.78	30.94	31.03	31.55	31.71
MnO	1.81	1.76	1.97	2.26	2.47	2.99	4.57	5.13	5.38	5.89	6.50	6.29	6.45	6.15	6.59	6.41	6.44	6.31	6.41	6.45	6.46	6.48	6.54	6.24	6.23	5.99	
MgO	3.20	3.02	2.96	1.73	2.67	2.81	2.96	2.98	2.92	2.84	2.76	2.74	2.76	2.57	2.71	2.78	2.83	2.85	2.88	2.77	2.79	2.77	2.79	2.81	2.84	2.87	
CaO	4.47	5.40	5.07	4.89	5.48	5.22	2.15	1.80	1.64	2.11	2.49	2.39	2.22	1.97	2.04	1.98	2.27	2.31	2.31	2.52	2.49	2.37	2.11	1.96	1.77	1.65	
Na2O	0.00	0.04	0.01	0.01	0.00	0.03	0.01	0.00	0.01	0.00	0.00	0.00	0.02	0.03	0.00	0.02	0.01	0.00	0.01	0.01	0.00	0.04	0.00	0.02	0.00	0.00	
K2O	0.00	0.01	0.00	0.01	0.01	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	
Cr2O3	0.01	0.01	0.01	0.02	0.03	0.03	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	
Total	99.63	99.31	99.58	86.07	98.70	98.60	99.78	99.65	99.74	99.50	99.35	99.49	92.77	98.99	98.78	99.65	99.70	99.54	99.27	99.41	99.68	99.68	99.31	99.54	99.77	99.87	
X	29.8443	29.8225	29.846	29.8355	29.8179	29.8274	29.8124	29.8006	29.7841	29.7702	29.7569	29.744	29.7281	29.7178	29.7011	29.6827	29.6718	29.6527	29.6397	29.6288	29.6186	29.6084	29.5968	29.5811	29.5695	29.5504	29.5361
Y	17.6636	17.6737	17.6655	17.705	17.712	17.7242	17.7457	17.7718	17.7866	17.8089	17.8265	17.827	17.868	17.8864	17.9155	17.9394	17.9735	18.0008	18.018	18.0403	18.0273	18.0778	18.0962	18.1244	18.133	18.1453	18.1543
Z	10.7805	10.7805	10.7805	10.7805	10.7805	10.7805	10.7805	10.7805	10.7805	10.7805	10.7805	10.7805	10.7805	10.7805	10.7805	10.7805	10.7805	10.7805	10.7805	10.7805	10.7805	10.7805	10.7805	10.7805	10.7805	10.7805	10.7805

		Profile 2																										
No Spot	265	266	267	268	269	270	271	272	273	274	275	276	277	278	279	280	281	282	283	284	285	286	287	288	289	290	291	292
SiO2	37.44	37.20	37.09	37.13	37.02	36.80	36.94	36.18	37.03	37.03	36.99	36.89	36.97	37.00	36.88	36.71	36.80	36.78	36.68	36.74	36.69	36.73	36.78	36.76	36.76	36.76	37.17	95.41
TiO2	0.01	0.00	0.02	0.02	0.02	0.02	0.01	0.11	0.02	0.05	0.04	0.03	0.02	0.00	0.00	0.												

TABLE A4. Whole rock geochemistry used in this study

SAMPLE	AP-038	AP-010	AP-009	CJ-91	CVI-1367	MPR-33	CVI-1388	CVI-13108	LRW-21
Ba	833	472	1005	365	937	288	90.2	852	519
Ce	18.4	38.6	29.3	6.4	28.3	98.2	180	40.4	94.1
Cr	20	30	0	0	150	30	40	160	210
Cs	1.17	0.56	0.29	0.13	2.81	0.97	0.16	1.43	1.45
Dy	2.47	6.69	1.32	0.3	2.57	5.97	19.9	5.06	9.48
Er	1.94	3.94	0.94	0.19	1.3	2.86	12.75	2.39	4.79
Eu	0.42	2.01	0.74	0.59	0.85	2.71	5.93	1.29	2.18
Ga	17.2	21.4	18.7	22.6	17.7	27	21.1	26.7	31.4
Gd	2.21	7.04	1.79	0.25	2.5	7.69	19.1	6.08	11.85
Hf	1.6	3.7	3.1	0.4	2.8	6.9	30.7	1.5	2.3
Ho	0.59	1.35	0.3	0.08	0.53	1.08	4.35	0.94	1.91
La	9.2	17.1	14.5	3.6	14.1	47.5	81.4	16.8	38
Lu	0.33	0.53	0.16	0.02	0.2	0.36	1.79	0.27	0.62
Nb	4.5	7	7.1	1.3	3	9.9	20.1	13.4	23.2
Nd	8.8	25.4	13.8	1.9	13.2	42.4	99.1	24.2	52.5
Pr	2.2	5.66	3.53	0.62	3.46	11.3	23.9	5.53	12.35
Rb	109	16.3	38.4	4.6	47.1	27.3	4.8	114	118.5
Sm	2.11	6.85	2.36	0.27	2.56	8.63	18.9	5.7	12.75
Sn	1	2	1	1	1	3	1	1	2
Sr	207	452	500	1550	303	931	404	196	319
Ta	0.5	0.3	0.4	0.1	0.2	0.7	0.9	0.6	1
Tb	0.41	1.06	0.23	0.05	0.36	1.12	3.13	0.99	1.79
Th	4.34	0.06	1.95	0.35	4.26	11.65	5.73	1.16	3.01
Tm	0.28	0.59	0.15	0.02	0.21	0.36	1.85	0.3	0.68
U	2.94	0.09	0.86	0.12	1.87	3.95	1.32	1.26	1.7
V	12	133	41	21	153	214	164	238	238
W	1	1	1	3	1	1	1	2	1
Y	16.9	35.9	8.7	1.7	13.2	32	119	25.9	49.7
Yb	2.35	3.68	1.08	0.15	1.35	2.51	11.65	2.13	4.29
Zr	38	124	110	10	98	276	1400	57	76
SiO2	74.7	59.7	70.3	59.8	54.9	52.9	40.2	55.1	45.5
Al2O3	14.2	15.3	16.05	25.1	15	15.85	19.1	17.1	19.15
Fe2O3t	1.21	9.27	2.61	0.91	5.55	9.47	22.1	9.19	16.1
CaO	1.03	5.76	3.03	7.87	6.43	8.8	11.4	5.7	6.22
MgO	0.07	2.45	0.66	0.19	4.24	4.45	3.18	5.44	4.1
Na2O	3.92	4.26	4.21	6.31	3.99	4.02	1.44	0.96	2.43
K2O	4.56	1	2.05	0.35	1.48	1.28	0.39	3.11	2.31
Cr2O3	0	0	0	0	0.02	0.01	0.01	0.02	0.03
TiO2	0.03	1.25	0.23	0.04	0.61	0.87	1.71	1.42	2.58
MnO	0.1	0.17	0.1	0.02	0.11	0.2	0.49	0.13	0.32
P2O5	0.02	0.46	0.09	0.01	0.21	0.39	1.14	0.65	0.9
SrO	0.03	0.06	0.07	0.19	0.04	0.11	0.06	0.03	0.04
BaO	0.1	0.06	0.11	0.04	0.11	0.03	0.01	0.1	0.06
LOI	0.41	1.15	1.72	0.54	7.27	1.57	-0.88	1.34	0.3
Total	100.38	100.89	101.23	101.37	99.96	99.95	100.35	100.29	100.04
Ag	0	0	0	0	0	0	0	0	0
As	0	0	6	0	0	0	5	0	0
Cd	0	0	0	0	0	0	0.9	0	0
Co	1	17	3	1	17	22	23	27	41
Cu	1	33	3	2	40	586	22	5	43
Li	0	10	10	0	30	10	0	10	30
Mo	1	1	0	1	0	1	3	0	0
Ni	1	9	1	1	27	23	15	57	68
Pb	29	4	6	8	8	16	5	11	5
Sc	2	22	3	1	18	13	40	23	34
Tl	0	0	0	0	0	0	0	0	0
Zn	16	92	63	18	78	110	174	168	248
FeO	0.9254443	7.0899741	1.9962063	0.6959953	4.2448065	7.2429401	16.902743	7.0287877	12.313763
Fe2O3	0.1815	1.3905	0.3915	0.1365	0.8325	1.4205	3.315	1.3785	2.415
mg	0.05947605	0.22414398	0.2165563	0.18581993	0.45506714	0.3393465	0.13591054	0.39285672	0.21775196

TABLE A5. ANACOMP Database Results, geochemical analysis from major Oxides, the First 6 samples are control from the database

This table includes NORMAL wt. % and LOG wt. % comparisons

*** Output of program and database A N A C O M P from C. Roehr, version Mar. 1991 ***

Date: 5.11.2016. Analyses read from file EXAMPLE.DAT (recalculated to 100%). The database holds 3591 analyses for comparison.

Abs. diff. of wt.% for most similiar: 3.36. Comparison using NORMAL wt.% oxides. No oxides ommited from comparison.

No.	dif.	name of rock	location							source	SiO2	TiO2
			Al2O3	FeOt	MnO	MgO	CaO	Na2O	K2O	P2O5		
		Metagabbro, KTB VB									48.8	1.80
			15.7	12.0	.17	8.1	9.5	2.5	.9	.45		

1	100	basalt. scoria lava									Schmincke+Sunkel_87_	48.4	1.52
			15.9	11.3	.19	8.0	9.7	3.7	1.0	.22			
2	132	basalt									Wilson,M._(1989)_____	48.5	1.98
			15.6	12.3	.21	6.6	11.3	2.3	.8	.43			
3	134	Amphibolite									Villagomez_2011_____	48.8	1.94
			14.7	12.3	.20	8.0	10.5	3.3	.2	.17			
4	139	Basalt									Wilson,M._(1989)_____	49.8	1.62
			16.0	11.3	.18	7.3	10.2	2.8	.5	.25			
5	143	Latit-Basalt									Wimmenauer_(1985)____	49.0	1.93
			15.0	10.4	.15	9.2	9.9	2.6	1.5	.35			
6	146	altered pillow lava									Amstutz_(ed)_(1974)_	48.9	3.06
			16.4	11.4	.15	7.9	8.1	2.8	.8	.53			
7	151	Arquia									Cochrane_2014a_____	49.1	1.63
			14.9	11.8	.21	8.6	11.2	2.2	.2	.11			
8	154	basalt									Best_82/LeMaitre_76_	50.1	1.88
			16.0	10.8	.20	6.9	9.7	3.0	1.1	.36			
9	160	Amph									Cochrane_2014_____	50.7	1.22
			15.1	12.2	.15	8.4	9.0	2.8	.3	.08			
10	161	basalt. scoria lava									Schmincke+Sunkel_87_	48.2	2.03
			16.4	11.3	.17	7.2	9.2	3.7	1.5	.33			

Date: 5.11.2016. Analyses read from file EXAMPLE.DAT (recalculated to 100%). The database holds 3591 analyses for comparison.

Abs. diff. of wt.% for most similiar: 2.12. Comparison using NORMAL wt.% oxides. No oxides ommited from comparison.

No.	dif.	name of rock	location							source	SiO2	TiO2	
			Al2O3	FeOt	MnO	MgO	CaO	Na2O	K2O	P2O5			
		Paragneis 33a, Nordrach, Schwarzwald										70.4	.74
			14.0	4.9	.12	2.5	2.4	2.7	2.2	.07			

1	100	graywacke, Proterozoic									Chab_and_Pelc_1973__	70.1	.59

holds 3591 analyses for comparison.

Abs. diff. of wt.% for most similiar: 3.14. Comparison using NORMAL wt.% oxides. No oxides ommited from comparison.

No.	dif.	name of rock	location								source	SiO2	TiO2	
			Al2O3	FeOt	MnO	MgO	CaO	Na2O	K2O	P2O5				
		Silikatmarmor, Val Strona										7.0	.02	
			1.5	1.7	.54	1.1	88.0	.1	.1	.02				
1	100	limestone									Rhein.Schiefergeb.Germany	Schulz-Dobrick(1975)	6.9	.07
	1.9		.6	.34	.9	88.7	.3	.3	.02					
2	178	limestone									Rhein.Schiefergeb.Germany	Schulz-Dobrick(1975)	6.1	.14
	3.2		1.1	.07	.9	87.4	.0	1.0	.02					
3	187	limestone									Rhein.Schiefergeb.Germany	Schulz-Dobrick(1975)	8.1	.02
			.3	.5	.10	1.0	89.8	.0	.1	.07				
4	257	lake chalk (Seekreide)									Gross Drewitz,Germany	Niggli,_P._(1952)___	7.7	.
	1.5		4.7	.36	.5	85.1	.	.	.					
5	267	limestone									Rhein.Schiefergeb.Germany	Schulz-Dobrick(1975)	8.4	.03
			.5	2.5	1.78	1.5	84.9	.2	.0	.27				
6	288	limestone									Rhein.Schiefergeb.Germany	Schulz-Dobrick(1975)	11.3	.05
			.8	.8	.13	1.0	85.5	.1	.2	.03				
7	296	limestone									Rhein.Schiefergeb.Germany	Schulz-Dobrick(1975)	11.0	.05
	1.2		.6	.22	1.4	85.1	.1	.2	.15					
8	351	limestone									Rhein.Schiefergeb.Germany	Schulz-Dobrick(1975)	3.7	.03
			.7	1.0	.23	.7	93.4	.1	.1	.05				
9	385	limestone									Rhein.Schiefergeb.Germany	Schulz-Dobrick(1975)	11.3	.13
	2.5		1.1	.46	.8	82.9	.5	.3	.07					
10	419	limestone									Rhein.Schiefergeb.Germany	Schulz-Dobrick(1975)	3.0	.05
	1.2		.5	.07	.5	94.2	.0	.4	.					

Date: 5.11.2016. Analyses read from file EXAMPLE.DAT (recalculated to 100%). The database holds 3591 analyses for comparison.

Abs. diff. of wt.% for most similiar: 3.04. Comparison using NORMAL wt.% oxides. No oxides ommited from comparison.

No.	dif.	name of rock	location								source	SiO2	TiO2	
			Al2O3	FeOt	MnO	MgO	CaO	Na2O	K2O	P2O5				
		Serpentinit, Foehrenbuehl, EGZ										46.8	.33	
			3.5	9.1	.11	37.6	2.1	.2	.2	.06				
1	100	Lherzolith									Ivrea-Zone,Italien	Wimmenauer_(1985)___	46.5	.09
	3.2		8.7	.16	37.9	3.2	.2	.0	.					
2	175	Spinell-Lherzolith									S. de Ronda,Spanien	Wimmenauer_(1985)___	48.8	.06
	2.7		7.9	.13	37.5	2.8	.1	.	.					
3	176	Raspas Complex									peridotite	John_2010_____	47.2	.05
	2.1		8.7	.12	40.0	1.9	.0	.	.					
4	203	Rio Panupili Mafic									Harzburgite	Bosch_2002_____	44.5	.09
	3.2		9.0	.12	39.8	2.6	.5	.0	.08					
5	250	El Toro Unit UM									Serpentinite	Bosch_2002_____	44.1	.10
	3.1		9.1	.11	40.7	2.8	.1	.0	.01					

6	270	serpentinite	Peräpohja, Finland	Amstutz_(ed)_(1974)_	45.8	.39
2.3	12.6	.18	38.1	.5	.1	.02
7	292	Granat-Peridotit	Böhmen,CSFR	Wimmenauer_(1985)___	44.0	.14
2.5	9.2	.16	42.0	2.0	.	.
8	295	Komatiit	Munro Township,Kanada	Wimmenauer_(1985)___	44.2	.19
6.6	7.4	.11	38.1	2.9	.2	.1
9	299	Al-spinel lherzolite		Wilson,M._(1989)___	44.6	.18
1.8	8.9	.14	41.9	2.3	.1	.0
10	307	Arquia	Serpentinite	Cochrane_2014a_____	46.4	.05
1.7	8.3	.10	42.4	1.0	.	.01

Date: 5.11.2016. Analyses read from file EXAMPLE.DAT (recalculated to 100%). The database holds 3591 analyses for comparison.

Abs. diff. of wt.% for most similiar: 32.30. Comparison using NORMAL wt.% oxides. No oxides ommited from comparison.

No.	dif.	name of rock	location	source	SiO2	TiO2	
Al2O3	FeOt	MnO	MgO	CaO	Na2O	K2O	P2O5
		Chloritschiefer Stbr.	Mariestollen, EGZ		32.1	2.13	
20.1	12.1	.39	32.7	.4	.1	.0	

1	100	harzburgite (serpentin.)	Harzburg massif, Germany	Vinx,_R._(1982)_____	43.1	.15
6.9	11.3	.19	35.1	2.8	.2	.2
2	104	mela-olivinegabbronorite	Harzburg massif, Germany	Vinx,_R._(1982)_____	42.2	.17
7.2	10.3	.19	34.6	4.5	.7	.0
3	104	feldspar peridotite	Lahn-Dill, Germany	Amstutz_(ed)_(1974)_	43.2	.66
5.5	15.2	.19	32.2	2.7	.2	.1
4	105	Kimberlit	Iron Mountain,Wyoming	Wimmenauer_(1985)___	39.2	3.50
3.3	12.5	.21	33.2	6.0	.1	1.6
5	108	mela-olivinenorite	Harzburg massif, Germany	Vinx,_R._(1982)_____	44.3	.36
7.2	10.3	.17	31.8	5.1	.4	.2
6	112	picrite	Rhein.Schiefergeb.Germany	Schulz-Dobrick(1975)	41.8	.79
6.0	15.7	.23	30.2	4.1	.8	.3
7	114	Komatiit	Munro Township,Kanada	Wimmenauer_(1985)___	44.2	.23
6.0	10.0	.14	34.5	4.5	.3	.1
8	116	peridotite	(average)	Best_82/LeMaitre_76_	44.7	.67
4.5	10.4	.43	33.0	5.3	.5	.4
9	118	Peridotit	Skaergaard,Grönland	Wimmenauer_(1985)___	41.6	.07
8.8	13.1	.16	27.3	6.6	.7	.1
10	120	kimberlite		Wilson,M._(1989)_____	38.1	2.37
3.1	9.8	.24	33.8	9.7	.2	1.3

Date: 5.11.2016. Analyses read from file EXAMPLE.DAT (recalculated to 100%). The database holds 3591 analyses for comparison.

Abs. diff. of wt.% for most similiar: 3.64. Comparison using NORMAL wt.% oxides. No oxides ommited from comparison.

No.	dif.	name of rock	location	source	SiO2	TiO2	
Al2O3	FeOt	MnO	MgO	CaO	Na2O	K2O	P2O5
		AP-010			59.9	1.25	
15.4	9.3	.17	2.5	5.8	4.3	1.0	

1	100	Andesit						San Salvador	Wimmenauer_(1985)___	58.7	1.19
	15.5	9.4	.20	2.5	6.2	4.0	2.1	.19			
2	128	Andesite, calc-alk.series						South Sandwich arc	Wilson,M._(1989)___	60.4	1.12
	14.2	10.6	.22	2.1	6.1	3.8	1.2	.26			
3	143	spilite						Schirmeck, Vosges, France	Amstutz_(ed)_(1974)_	61.4	1.30
	14.4	10.2	.16	2.3	5.6	3.4	1.3	.			
4	146	acid andesite,med-K,TH						New Britain	Gill,_J.B._(1981)___	58.9	.89
	15.5	8.7	.18	3.2	7.0	3.8	1.5	.25			
5	152	Andesit						Thingmuli,Island	Wimmenauer_(1985)___	61.9	1.28
	15.4	8.0	.19	1.8	5.1	4.4	1.6	.44			
6	153	Clemesi						andesite	Mamani_et_al.,_2008_	59.3	1.03
	16.3	8.6	.15	2.7	6.3	3.4	2.0	.23			
7	176	Pyroxen-Bio.-Qrz.-Andesit						Saar-Nahe	Wimmenauer_(1985)___	59.7	1.79
	16.2	7.5	.09	3.2	6.2	3.3	1.7	.35			
8	178	Andesite,tholeiit. series						South Sandwich arc	Wilson,M._(1989)___	62.8	1.01
	14.2	8.8	.18	2.1	6.2	4.1	.4	.17			
9	179	spilit.tuffite (Schalst.)						Schirmeck, Vosges, France	Amstutz_(ed)_(1974)_	60.2	1.12
	14.5	10.7	.16	3.9	3.9	4.5	.9	.			
10	185	Misti						andesite	Mamani_et_al.,_2008_	58.9	.90
	16.9	7.5	.12	3.2	6.0	4.3	1.8	.29			

Date: 5.11.2016. Analyses read from file EXAMPLE.DAT (recalculated to 100%). The database holds 3591 analyses for comparison.

Abs. diff. of wt.% for most similiar: 2.85. Comparison using NORMAL wt.% oxides. No oxides ommited from comparison.

No.	dif.	name of rock	location							source	SiO2	TiO2
			Al2O3	FeOt	MnO	MgO	CaO	Na2O	K2O	P2O5		
		CVI-1367									59.3	.66
			16.2	6.0	.12	4.6	6.9	4.3	1.6	.23		

1	100	Lascar						andesite	Mamani_et_al.,_2008_	59.0	.76
	16.4	7.0	.12	4.4	6.9	3.5	1.8	.19			
2	114	Lascar						andesite	Mamani_et_al.,_2009_	58.8	.75
	16.1	7.3	.12	4.7	7.0	3.3	1.7	.19			
3	115	Lascar						andesite	Mamani_et_al.,_2009_	58.7	.75
	16.4	7.1	.12	4.7	7.1	3.4	1.7	.19			
4	117	Precordillera						andesite	Haschke_2002_____	59.5	.81
	16.6	5.9	.11	3.6	6.9	3.8	2.5	.31			
5	119	Andesite						South-west Pacific	Wilson,M._(1989)___	59.3	.73
	16.9	6.7	.13	3.8	7.1	3.4	1.7	.23			
6	128	Negrillar						andesite	Mamani_et_al.,_2009_	58.5	1.08
	16.5	6.6	.09	4.1	6.6	4.1	2.0	.39			
7	136	Lascar						andesite	Mamani_et_al.,_2009_	58.4	.76
	16.4	7.2	.12	4.8	7.2	3.4	1.6	.19			
8	136	Las Cuevas						andesite	Worner_1992_____	59.3	.95
	16.8	6.5	.09	3.8	5.9	4.4	2.0	.26			
9	138	Base Misti						andesite	Mamani_et_al.,_2008_	58.7	.84
	16.6	7.1	.11	4.3	6.5	3.6	1.9	.22			
10	143	acid andesite,med-K, CA						Fiji	Gill,_J.B._(1981)___	60.1	.69
	17.2	6.0	.14	3.3	7.2	3.9	1.3	.20			

Date: 5.11.2016. Analyses read from file EXAMPLE.DAT (recalculated to 100%). The database holds 3591 analyses for comparison.

Abs. diff. of wt.% for most similiar: 3.96. Comparison using NORMAL wt.% oxides. No oxides ommited from comparison.

No.	dif.	name of rock	location							source	SiO2	TiO2
Al2O3	FeOt	MnO	MgO	CaO	Na2O	K2O	P2O5					

		MPR-33								53.8	.89	
		16.1	9.6	.20	4.5	9.0	4.1	1.3	.40			

1	100	Kalk-Alkali-Basalt						San Salvador	Wimmenauer_(1985)___	53.3	1.23	
17.3	9.9	.20	4.0	9.1	3.3	1.4	.20					
2	103	Leon Muerto						basalt	Trumbull_1999_____	54.1	1.26	
17.2	10.0	.17	3.8	8.2	3.7	1.4	.32					
3	111	Kalk-Alkali-Basalt						San Salvador	Wimmenauer_(1985)___	53.0	.97	
17.1	10.1	.19	4.5	9.7	3.0	1.3	.20					
4	111	Kalk-Alkali-Basalt						San Salvador	Wimmenauer_(1985)___	53.0	.97	
17.1	10.1	.19	4.5	9.7	3.0	1.3	.20					
5	111	Antofagasta						gabbro	Mamani_et_al.,_2008_	53.7	1.37	
15.8	11.2	.16	4.6	8.4	3.2	1.4	.22					
6	117	Nicholson						basaltic andesite	Delacour_2007_____	52.7	1.47	
16.6	9.5	.12	5.4	8.4	3.7	1.7	.36					
7	119	Las Cuevas						basalt	Worner_1992_____	52.9	1.42	
17.1	9.2	.11	5.2	8.7	3.6	1.5	.27					
8	119	Rhyolite						Riolita de Golero	Colmenares_2009_____	54.2	1.27	
16.1	9.9	.17	5.8	8.5	2.6	1.1	.34					
9	119	Santa Marta						Qz-Andesite	Quandt_2013_____	54.2	1.27	
16.1	9.9	.17	5.8	8.5	2.6	1.1	.34					
10	120	Nicholson						basalt	Mamani_et_al.,_2008_	52.5	1.47	
16.5	9.6	.14	5.5	8.5	3.7	1.8	.35					

Date: 5.11.2016. Analyses read from file EXAMPLE.DAT (recalculated to 100%). The database holds 3591 analyses for comparison.

Abs. diff. of wt.% for most similiar: 14.32. Comparison using NORMAL wt.% oxides. No oxides ommited from comparison.

No.	dif.	name of rock	location							source	SiO2	TiO2
Al2O3	FeOt	MnO	MgO	CaO	Na2O	K2O	P2O5					

		CVI-1388								39.7	1.69	
		18.9	21.9	.48	3.1	11.3	1.4	.4	1.13			

1	100	volcaniclastic rock						E Iceland, drill hole	Schmincke_et_al.1982	41.1	4.22	
17.9	20.6	.18	3.5	7.3	1.5	3.2	.44					
2	105	olivinegabbro						Harzburg massif, Germany	Vinx,_R._(1982)_____	45.7	1.84	
16.8	21.9	.51	1.2	8.1	2.5	.6	.73					
3	122	volcaniclastic rock						E Iceland, drill hole	Schmincke_et_al.1982	45.2	3.19	
19.5	16.1	.18	4.3	11.0	.2	.1	.27					
4	132	olivinegabbronorite						Harzburg massif, Germany	Vinx,_R._(1982)_____	42.9	3.70	


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LRW-21                                     45.7 2.59
19.2 16.2 .32 4.1 6.2 2.4 2.3 .90
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1 100 pillow lava                Lahn-Dill, Germany      Schmincke+Sunkel_87_ 46.7 3.72
18.6 14.5 .15 4.6 4.9 3.4 2.6 .92
2 106 volcaniclastic rock       E Iceland, drill hole  Schmincke_et_al.1982 47.3 3.39
20.0 17.0 .27 3.5 3.9 2.1 2.3 .19
3 125 pillow lava                Lahn-Dill, Germany      Schmincke+Sunkel_87_ 46.7 4.06
19.2 13.4 .12 4.3 5.0 2.1 4.4 .73
4 134 volcaniclastic rock       E Iceland, drill hole  Schmincke_et_al.1982 45.2 4.77
17.8 18.4 .41 3.0 5.8 1.3 2.9 .33
5 145 volcaniclastic rock       E Iceland, drill hole  Schmincke_et_al.1982 47.2 2.86
16.4 14.6 .31 6.5 6.5 2.0 3.4 .26
6 145 spilite                    Rhein.Schiefergeb.Germany Schulz-Dobrick(1975) 45.2 3.46
15.5 17.5 .24 5.7 7.1 3.2 1.6 .45
7 146 Diabas                     Lahn-Dill, Germany      Hentschel_(1970)_____ 45.4 2.84
18.5 14.0 .   5.4 10.4 2.3 .8 .32
8 147 volcaniclastic rock       E Iceland, drill hole  Schmincke_et_al.1982 48.1 3.54
16.0 16.5 .32 5.6 6.6 1.8 1.2 .33
9 151 basalt. volcaniclastic    Lahn-Dill, Germany      Schmincke+Sunkel_87_ 47.9 3.51
19.9 14.1 .10 3.3 3.8 2.7 4.0 .69
10 152 albitized basalt         Deccan, Bombay, India   Amstutz_(ed)_(1974)_ 47.9 2.32
15.1 16.1 .   5.9 7.5 3.2 2.0 .
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Date: 5.11.2016. Analyses read from file EXAMPLE.DAT (recalculated to 100%). The database holds 3591 analyses for comparison.

Abs. diff. of wt.% for most similiar: 1.09. Comparison using NORMAL wt.% oxides. No oxides ommited from comparison.

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No. dif. name of rock            location                source                  SiO2 TiO2
Al2O3 FeOt MnO MgO  CaO  Na2O  K2O P2O5
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AP-038                                     74.8
.03 14.2 1.2 .10 .1 1.0 3.9 4.6 .02
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1 100 Yura                        rhyolite                Paquereau_2008_____ 74.9 .19
13.8 1.2 .08 .3 1.1 4.0 4.5 .04
2 111 Chili                       rhyolite                Paquereau_2008_____ 74.7 .20
13.9 1.3 .08 .3 1.0 4.0 4.5 .05
3 143 Vitor                       rhyolite                Paquereau_2008_____ 74.8 .20
13.8 1.2 .08 .2 1.0 4.4 4.3 .04
4 150 Yura                        rhyolite                Paquereau_2008_____ 75.1 .20
13.6 1.3 .08 .2 1.0 3.8 4.6 .04
5 152 Chili                       rhyolite                Paquereau_2008_____ 75.1 .19
13.7 1.2 .07 .3 1.1 4.1 4.3 .04
6 153 Guallatiri                 rhyolite                Worner_1992_____ 74.3 .19
14.0 1.1 .06 .3 1.1 4.3 4.6 .02
7 164 Yura                        rhyolite                Paquereau_2008_____ 74.4 .23
13.9 1.5 .08 .3 1.1 4.1 4.4 .04
8 164 Vitor                       rhyolite                Paquereau_2008_____ 75.0 .19
13.7 1.2 .07 .2 1.0 4.3 4.3 .04
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9	167	Chachani								rhyolite	Mamani_et_al.,_2008_	74.5	.19
13.7	1.3	.07	.3	1.0	4.0	5.0	.05						
10	169	Chili								rhyolite	Paquereau_2008_____	74.5	.14
14.2	1.1	.09	.3	1.2	3.4	5.0	.05						

Date: 5.11.2016. Analyses read from file EXAMPLE.DAT (recalculated to 100%). The database holds 3591 analyses for comparison.

Abs. diff. of wt.% for most similiar: 3.19. Comparison using NORMAL wt.% oxides. No oxides ommited from comparison.

No.	dif.	name of rock	location					source				SiO2	TiO2
Al2O3	FeOt	MnO	MgO	CaO	Na2O	K2O	P2O5						

AP-009												70.8		
.23	16.2	2.6	.10	.7	3.0	4.2	2.1	.09						

1	100	Santa Marta								Granodiorite	Quandt_2013_____	70.9	.37	
15.4	2.6	.06	1.3	2.3	4.3	2.7	.13							
2	104	Mylonites								Santa Marta P-Triassic	Cardona_2010c_____	70.8	.19	
16.4	1.9	.09	.3	2.5	4.9	2.8	.06							
3	107	Santa Marta								Granodiorite	Quandt_2013_____	69.9	.22	
17.0	2.0	.04	.6	3.6	4.6	1.9	.09							
4	113	dacite								Saint-Bel-Sourcieux,Rhone	Davoine,_P._(1968)___	70.4	.25	
16.2	2.2	.08	1.3	4.0	3.2	2.2	.13							
5	124	Granodiorit								Sierra Nevada,USA	Wimmenauer_(1985)___	70.3	.38	
15.5	2.1	.06	.7	2.7	4.3	3.7	.14							
6	124	Mylonites								Santa Marta P-Triassic	Cardona_2010c_____	71.2	.23	
15.7	3.1	.03	.4	2.9	5.4	1.0	.07							
7	134	Sara Sara								rhyolite	Mamani_et_al.,_2008_	70.3	.37	
15.5	2.5	.05	.6	2.3	4.5	3.7	.15							
8	145	Pichu Pichu								rhyolite	Mamani_et_al.,_2009_	70.8	.30	
15.1	2.4	.05	.8	2.3	4.1	4.2	.10							
9	145	Huanynaputina								dacite	Mamani_et_al.,_2008_	69.1	.40	
15.7	3.0	.06	1.0	2.8	4.5	3.2	.14							
10	145	Santa Marta								Granodiorite	Quandt_2013_____	69.2	.46	
15.6	3.1	.06	1.2	3.0	4.1	3.1	.15							

Date: 5.11.2016. Analyses read from file EXAMPLE.DAT (recalculated to 100%). The database holds 3591 analyses for comparison.

Abs. diff. of wt.% for most similiar: 11.01. Comparison using NORMAL wt.% oxides. No oxides ommited from comparison.

No.	dif.	name of rock	location					source				SiO2	TiO2
Al2O3	FeOt	MnO	MgO	CaO	Na2O	K2O	P2O5						

CJJ-91												59.4		
.04	25.0	.9	.02	.2	7.8	6.3	.4	.01						

1	100	Anorthosit								Adirondack,New York	Wimmenauer_(1985)___	55.4	.87	
24.0	2.4	.01	1.4	9.1	5.8	1.0	.							
2	124	Anorthosit								Labrador	Wimmenauer_(1985)___	57.0	.79	
22.2	4.6	.06	1.4	7.1	5.4	1.4	.12							

No.	dif.	name of rock								location	source	SiO2	TiO2
Al2O3	FeOt	MnO	MgO	CaO	Na2O	K2O	P2O5						
		EMP-001										70.5	.30
	14.7	3.3	.11	1.1	3.0	3.9	3.0	.11					
1	100	graywacke								Franciscan, California	Pettijohn,_F.J._1963	70.6	.26
	14.9	3.1	.	2.0	2.3	4.0	2.8	.15					
2	100	Holocene feldspar sand								Salton Basin, California	v.d.Kamp_et_al_1976_	71.1	.42
	14.3	3.1	.06	1.4	2.9	3.3	3.4	.05					
3	117	arkose								Santa Ynes Mt.,California	v.d.Kamp_et_al_1976_	70.5	.38
	14.5	3.7	.06	1.8	1.8	4.0	3.3	.08					
4	121	Santa Marta								Granodiorite	Quandt_2013_____	69.2	.46
	15.6	3.1	.06	1.2	3.0	4.1	3.1	.15					
5	125	Holocene feldspar sand								Salton Basin, California	v.d.Kamp_et_al_1976_	71.1	.37
	14.9	2.5	.05	1.2	3.1	3.1	3.6	.03					
6	129	Salar de la Isla								rhyolite	Siebel_et_al.,_2001_	70.5	.45
	14.6	2.7	.05	.7	2.9	3.4	4.5	.12					
7	129	Ibague								Granite	Cochrane_2014a_____	71.5	.35
	14.8	2.7	.05	1.1	2.4	3.5	3.5	.09					
8	132	Santa Marta								Dacite	Quandt_2013_____	70.6	.42
	15.3	2.9	.08	.8	2.1	3.7	3.9	.12					
9	133	Santa Marta								Granodiorite	Quandt_2013_____	70.9	.37
	15.4	2.6	.06	1.3	2.3	4.3	2.7	.13					
10	133	Santa Marta								Granite	Quandt_2013_____	69.1	.47
	15.6	3.2	.06	1.2	3.4	4.0	2.8	.19					

Date: 5.11.2016. Analyses read from file EXAMPLE.DAT (recalculated to 100%). The database holds 3591 analyses for comparison.

Abs. diff. of wt.% for most similiar: 1.88. Comparison using NORMAL wt.% oxides. No oxides ommited from comparison.

No.	dif.	name of rock								location	source	SiO2	TiO2
Al2O3	FeOt	MnO	MgO	CaO	Na2O	K2O	P2O5						
		EMP-270										74.2	.07
	14.8	1.4	.03	.1	1.2	4.5	3.8	.01					
1	100	Arequipa								rhyolite	Paquereau_2008_____	74.1	.22
	14.0	1.5	.08	.3	1.2	4.5	4.0	.06					
2	106	Ignimbrite								Ignimbrita de los Clavos	Colmenares_2009_____	73.4	.29
	14.7	1.6	.04	.5	1.0	4.5	3.9	.08					
3	107	Aritinca Salar de Surire								dacite	Mamani_et_al.,_2008_	73.8	.24
	14.2	1.4	.07	.4	1.1	4.5	4.2	.04					
4	115	Chucal								rhyolite	Mamani_et_al.,_2009_	74.7	.21
	14.0	1.2	.05	.3	1.0	4.7	3.8	.06					
5	119	Chucal								rhyolite	Mamani_et_al.,_2009_	74.8	.22
	14.0	1.2	.05	.2	1.0	4.7	3.8	.06					
6	123	Santa Marta								Granite	Quandt_2013_____	74.3	.14
	14.7	1.4	.05	.3	1.9	4.3	2.9	.06					
7	130	Caraveli								rhyolite	Mamani_et_al.,_2009_	74.0	.24

No.	dif.	name of rock	location	source	SiO2	TiO2
13.8	1.7	.09	.3	1.2	4.4	4.2 .07
8	139	Abitagua	Monzogranite	Cochrane_2014a_____	74.8	.17
14.3	1.5	.04	.2	1.5	3.7	3.8 .05
9	143	Guallatiri	rhyolite	Worner_1992_____	74.3	.19
14.0	1.1	.06	.3	1.1	4.3	4.6 .02
10	148	Oberer Oxaya	rhyolite	Mamani_et_al.,_2009_	74.2	.18
14.0	1.4	.07	.6	1.8	3.9	3.8 .04

Date: 5.11.2016. Analyses read from file EXAMPLE.DAT (recalculated to 100%). The database holds 3591 analyses for comparison.
 Abs. diff. of wt.% for most similiar: 4.58. Comparison using NORMAL wt.% oxides. No oxides ommited from comparison.

No.	dif.	name of rock	location	source	SiO2	TiO2
16-JDMA-028-028					50.9	
.86	17.7	11.5	.15	6.0	7.3	4.1 1.3 .25

1	100	spilite	Rhein.Schiefergeb.Germany	Schulz-Dobrick(1975)	51.2	1.49
17.3	12.0	.14	4.3	8.2	4.0	1.2 .23
2	116	La Chilca Unit mafic	Eclogite metabasalt	Bosch_2002_____	51.5	2.50
17.3	9.6	.18	6.1	7.2	4.2	.9 .52
3	117	spilite, Jurassic-Cretac.	Sakhalin, USSR	Amstutz_(ed)_(1974)_	53.1	.78
16.3	11.3	.18	6.1	6.3	4.2	1.6 .10
4	120	Spilit	Kellerwald	Wimmenauer_(1985)___	49.3	1.82
17.3	11.9	.17	6.9	7.4	4.4	.5 .28
5	127	Punta de Bombon	andesite	Mamani_et_al.,_2008_	51.4	.81
17.9	10.8	.09	5.9	8.4	2.2	2.3 .17
6	128	Diabas	Lahn-Dill, Germany	Hentschel_(1970)___	50.4	1.56
16.8	11.1	.16	7.0	7.7	5.0	.2 .21
7	131	Spilite, Av.92		Amstutz_(ed)_(1974)_	52.5	1.66
16.9	10.1	.16	5.4	7.0	4.5	1.4 .28
8	134	Spilite	Nundle, New South Wales	Amstutz_(ed)_(1974)_	51.5	2.04
15.6	12.4	.19	5.9	7.5	3.9	.6 .37
9	140	Basalt,High-K-calc-alk.	Sunda arc	Wilson,M._(1989)___	50.3	1.13
18.1	9.8	.18	5.6	9.7	3.7	1.2 .25
10	142	spilite, Jurassic-Cretac.	Sakhalin, USSR	Amstutz_(ed)_(1974)_	52.8	.90
16.5	9.8	.22	6.5	7.1	4.9	1.3 .10

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COMPARISON-----

Date: 5.11.2016. Analyses read from file EXAMPLE.DAT (recalculated to 100%). The database holds 3591 analyses for comparison.
 Comparison using LOG wt.-% oxides. No oxides ommited from comparison.

No.	dif.	name of rock	location	source	SiO2	TiO2
Al2O3	FeOt	MnO	MgO	CaO	Na2O	K2O P2O5

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AP-010                                     59.9 1.25
15.4  9.3  .17  2.5  5.8  4.3  1.0  .46
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1  100 Andesit                               Thingmuli,Island           Wimmenauer_(1985)___ 61.9 1.28
15.4  8.0  .19  1.8  5.1  4.4  1.6  .44
2  123 Andesite, calc-alk.series South Sandwich arc           Wilson,M._(1989)___ 60.4 1.12
14.2 10.6  .22  2.1  6.1  3.8  1.2  .26
3  126 spilite                               Schirmeck, Vosges, France Amstutz_(ed)_(1974)_ 61.4 1.30
14.4 10.2  .16  2.3  5.6  3.4  1.3  .
4  143 Santa Marta                           Granodiorite             Quandt_2013_____ 56.3 1.01
17.7  8.8  .16  3.5  5.6  4.4  1.8  .65
5  143 Santa Marta                           Andesite                 Quandt_2013_____ 56.3 1.01
17.7  8.8  .16  3.5  5.6  4.4  1.8  .65
6  144 CdL                                   andesite                 Bock_et_al.,_2000___ 56.6 1.20
18.2  8.1  .12  3.9  6.3  4.4  .8  .33
7  145 Leon Muerto                           basalt                   Trumbull__1999_____ 54.1 1.26
17.2 10.0  .17  3.8  8.2  3.7  1.4  .32
8  145 Quarz-Andesit, Mean                   Chile                     Wimmenauer_(1985)___ 57.3  .94
17.5  7.4  .16  4.6  6.8  3.9  1.1  .37
9  147 Basaltic andesites                    Andes,southern volc. zone Wilson,M._(1989)___ 55.5 1.35
16.7 10.1  .19  3.3  7.3  4.5  .9  .21
10 151 Andesit                               San Salvador             Wimmenauer_(1985)___ 58.7 1.19
15.5  9.4  .20  2.5  6.2  4.0  2.1  .19
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Date: 5.11.2016. Analyses read from file EXAMPLE.DAT (recalculated to 100%). The database holds 3591 analyses for comparison.

Comparison using LOG wt.-% oxides. No oxides ommited from comparison.

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No. dif. name of rock           location           source           SiO2 TiO2
Al2O3 FeOt MnO MgO  CaO  Na2O  K2O P2O5
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CVI-1367                                     59.3  .66
16.2  6.0  .12  4.6  6.9  4.3  1.6  .23
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1  100 Lascar                               andesite                 Mamani_et_al.,_2009_ 58.7  .75
16.4  7.1  .12  4.7  7.1  3.4  1.7  .19
2  103 Andesite                             South-west Pacific       Wilson,M._(1989)___ 59.3  .73
16.9  6.7  .13  3.8  7.1  3.4  1.7  .23
3  104 Lascar                               andesite                 Mamani_et_al.,_2008_ 59.0  .76
16.4  7.0  .12  4.4  6.9  3.5  1.8  .19
4  104 Lascar                               andesite                 Mamani_et_al.,_2009_ 58.4  .76
16.4  7.2  .12  4.8  7.2  3.4  1.6  .19
5  107 Lascar                               andesite                 Mamani_et_al.,_2009_ 58.8  .75
16.1  7.3  .12  4.7  7.0  3.3  1.7  .19
6  108 Lascar                               andesite                 Mamani_et_al.,_2009_ 58.0  .86
16.8  7.1  .11  4.8  7.0  3.7  1.5  .24
7  111 Lascar                               andesite                 Mamani_et_al.,_2009_ 58.2  .76
16.4  7.2  .12  4.9  7.3  3.4  1.6  .19
8  118 Base Misti                           andesite                 Mamani_et_al.,_2008_ 58.7  .84
16.6  7.1  .11  4.3  6.5  3.6  1.9  .22
9  126 Lascar                               andesite                 Mamani_et_al.,_2009_ 60.2  .73
17.0  6.5  .11  3.6  6.1  3.6  1.9  .22
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No.	dif.	name of rock	location	source	SiO2	TiO2
10	130	Lascar	andesite	Mamani_et_al.,_2009_	57.2	.95
17.1	7.9	.12	4.3	6.9	3.6	1.5 .24

Date: 5.11.2016. Analyses read from file EXAMPLE.DAT (recalculated to 100%). The database holds 3591 analyses for comparison.

Comparison using LOG wt.-% oxides. No oxides ommited from comparison.

No.	dif.	name of rock	location	source	SiO2	TiO2
Al2O3	FeOt	MnO	MgO	CaO	Na2O	K2O P2O5

		MPR-33			53.8	.89
	16.1	9.6	.20	4.5	9.0	4.1 1.3 .40

1	100	Basalt	Andes	Wilson,M._(1989)_____	51.2	1.14
18.6	8.6	.16	5.6	8.9	4.0	1.4 .38
2	108	Quarz-Andesit, Mean	Chile	Wimmenauer_(1985)_____	57.3	.94
17.5	7.4	.16	4.6	6.8	3.9	1.1 .37
3	109	basic andesite,med-K, TH	Bagana volc.,Bougainville	Gill,_J.B._(1981)_____	55.8	.81
17.8	8.0	.18	3.4	8.0	4.0	1.6 .35
4	112	Leon Muerto	basalt	Trumbull__1999_____	54.1	1.26
17.2	10.0	.17	3.8	8.2	3.7	1.4 .32
5	118	Kalk-Alkali-Basalt	San Salvador	Wimmenauer_(1985)_____	53.0	.97
17.1	10.1	.19	4.5	9.7	3.0	1.3 .20
6	118	Kalk-Alkali-Basalt	San Salvador	Wimmenauer_(1985)_____	53.0	.97
17.1	10.1	.19	4.5	9.7	3.0	1.3 .20
7	121	Basaltic andesite	South-west Pacific	Wilson,M._(1989)_____	54.4	.83
17.1	8.6	.16	5.3	9.1	2.9	1.3 .26
8	129	Basalt,High-K-calc-alk.	Sunda arc	Wilson,M._(1989)_____	50.3	1.13
18.1	9.8	.18	5.6	9.7	3.7	1.2 .25
9	130	Kalk-Alkali-Basalt	San Salvador	Wimmenauer_(1985)_____	53.3	1.23
17.3	9.9	.20	4.0	9.1	3.3	1.4 .20
10	136	basalt, Late Cretaceous	Kurile islands, USSR	Amstutz_(ed)_(1974)_	53.7	1.10
18.9	7.8	.19	3.8	8.6	4.6	1.1 .27

Date: 5.11.2016. Analyses read from file EXAMPLE.DAT (recalculated to 100%). The database holds 3591 analyses for comparison.

Comparison using LOG wt.-% oxides. No oxides ommited from comparison.

No.	dif.	name of rock	location	source	SiO2	TiO2
Al2O3	FeOt	MnO	MgO	CaO	Na2O	K2O P2O5

		CVI-1388			39.7	1.69
	18.9	21.9	.48	3.1	11.3	1.4 .4 1.13

1	100	olivinegabbro	Harzburg massif, Germany	Vinx,_R._(1982)_____	45.7	1.84
16.8	21.9	.51	1.2	8.1	2.5	.6 .73
2	121	olivinegabbbronorite	Harzburg massif, Germany	Vinx,_R._(1982)_____	42.9	3.70
12.3	22.7	.41	5.1	8.4	1.8	1.0 1.71
3	130	Glimmerdiabas	Lahn-Dill, Germany	Hentschel_(1970)_____	45.5	1.62
21.5	12.1	.	4.6	12.1	1.5	.8 .36
4	132	spilite, average		Amstutz_(ed)_(1974)_	49.8	2.39

LRW-21 45.7 2.59
 19.2 16.2 .32 4.1 6.2 2.4 2.3 .90

No.	dif.	name of rock	location	source	SiO2	TiO2
1	100	Basalt	Columbia River Province	Wilson,M._(1989)_____	54.6	2.91
					13.6	12.9
					.24	2.8
					6.2	3.3
					2.6	.88
2	104	pillow lava	Lahn-Dill, Germany	Schmincke+Sunkel_87_	46.7	3.72
					18.6	14.5
					.15	4.6
					4.9	3.4
					2.6	.92
3	114	Tholeiit-Basalt	Washington,USA	Wimmenauer_(1985)_____	54.0	3.03
					13.4	13.0
					.18	3.0
					6.5	3.4
					2.6	.91
4	124	High P2O5-TiO2 Volcanite	Parana	Wilson,M._(1989)_____	54.0	2.93
					13.9	12.7
					.19	3.0
					7.6	3.0
					2.0	.74
5	128	Ignimbrite	Ignimbrita de los Clavos	Colmenares_2009_____	54.4	1.60
					17.1	9.9
					.17	4.0
					5.7	3.7
					2.5	.93
6	130	Diabas	Lahn-Dill, Germany	Hentschel_(1970)_____	48.0	2.01
					19.6	10.3
					.17	6.0
					7.9	2.4
					2.9	.66
7	130	Basalt	Ascension Island Atlantic	Wilson,M._(1989)_____	48.3	3.15
					16.3	11.7
					.19	5.1
					8.4	4.0
					1.9	1.02
8	130	Basalt	Columbia River Province	Wilson,M._(1989)_____	51.2	3.21
					13.5	14.2
					.22	4.5
					8.4	2.9
					1.3	.68
9	133	Basalt	Columbia River Province	Wilson,M._(1989)_____	51.0	3.56
					12.9	15.0
					.22	4.4
					8.5	2.5
					1.4	.57
10	134	spilite	Rhein.Schiefergeb.Germany	Schulz-Dobrick(1975)	45.2	3.46
					15.5	17.5
					.24	5.7
					7.1	3.2
					1.6	.45

Date: 5.11.2016. Analyses read from file EXAMPLE.DAT (recalculated to 100%). The database holds 3591 analyses for comparison.

Comparison using LOG wt.-% oxides. No oxides ommited from comparison.

No.	dif.	name of rock	location	source	SiO2	TiO2
					Al2O3	FeOt
					MnO	MgO
					CaO	Na2O
					K2O	P2O5

AP-038 74.8
 .03 14.2 1.2 .10 .1 1.0 3.9 4.6 .02

No.	dif.	name of rock	location	source	SiO2	TiO2
1	100	Azafran	Granodiorite	Cochrane_2014a_____	76.0	.07
					13.6	1.0
					.09	.0
					.6	4.1
					4.4	.02
2	118	Azafran	Monzogranite	Cochrane_2014a_____	76.7	.12
					13.2	.8
					.06	.1
					.8	3.6
					4.7	.03
3	118	Salar de la Isla	rhyolite	Siebel_et_al.,_2001_	76.1	.10
					13.6	.8
					.07	.1
					1.1	3.5
					4.5	.04
4	118	Yura	rhyolite	Mamani_et_al.,_2008_	75.9	.15
					13.1	1.1
					.07	.2
					1.0	4.1
					4.4	.03
5	119	C° Villacollo	rhyolite	Mamani_et_al.,_2009_	75.1	.12
					13.8	.9
					.07	.2
					1.0	3.7
					5.0	.02
6	119	Yura	rhyolite	Paquereau_2008_____	75.9	.16
					13.3	1.1
					.08	.2
					.8	3.9
					4.6	.03
7	122	Salar de Antofalla	rhyolite	Siebel_et_al.,_2001_	75.7	.11
					13.0	.8
					.08	.1
					1.6	3.0
					5.6	.03
8	123	calcalkaline rhyolite	after Johannsen (1952)	Davoine,_P._(1968)___	74.1	.18
					14.7	1.0
					.	.1
					1.0	3.0
					5.9	.03
9	124	Yura	rhyolite	Paquereau_2008_____	75.7	.16
					13.3	1.0
					.07	.2
					1.1	4.1
					4.4	.03
10	129	Lauca_Pérez	rhyolite	Mamani_et_al.,_2008_	76.4	.12

12.8 .9 .09 .2 .6 3.5 5.4 .03

Date: 5.11.2016. Analyses read from file EXAMPLE.DAT (recalculated to 100%). The database holds 3591 analyses for comparison.

Comparison using LOG wt.-% oxides. No oxides ommited from comparison.

No.	dif.	name of rock	location								source	SiO2	TiO2	
Al2O3	FeOt	MnO	MgO	CaO	Na2O	K2O	P2O5							

		AP-010										59.9		
		1.25	15.4	9.3	.17	2.5	5.8	4.3	1.0	.46				

1	100	Andesit									Thingmuli,Island	Wimmenauer_(1985)____	61.9	1.28
15.4	8.0	.19	1.8	5.1	4.4	1.6	.44							
2	123	Andesite, calc-alk.series									South Sandwich arc	Wilson,M._(1989)____	60.4	1.12
14.2	10.6	.22	2.1	6.1	3.8	1.2	.26							
3	126	spilite									Schirmeck, Vosges, France	Amstutz_(ed)_(1974)_	61.4	1.30
14.4	10.2	.16	2.3	5.6	3.4	1.3	.							
4	143	Santa Marta									Granodiorite	Quandt_2013_____	56.3	1.01
17.7	8.8	.16	3.5	5.6	4.4	1.8	.65							
5	143	Santa Marta									Andesite	Quandt_2013_____	56.3	1.01
17.7	8.8	.16	3.5	5.6	4.4	1.8	.65							
6	144	CdL									andesite	Bock_et_al.,_2000____	56.6	1.20
18.2	8.1	.12	3.9	6.3	4.4	.8	.33							
7	145	Leon Muerto									basalt	Trumbull__1999_____	54.1	1.26
17.2	10.0	.17	3.8	8.2	3.7	1.4	.32							
8	145	Quarz-Andesit, Mean									Chile	Wimmenauer_(1985)____	57.3	.94
17.5	7.4	.16	4.6	6.8	3.9	1.1	.37							
9	147	Basaltic andesites									Andes,southern volc. zone	Wilson,M._(1989)____	55.5	1.35
16.7	10.1	.19	3.3	7.3	4.5	.9	.21							
10	151	Andesit									San Salvador	Wimmenauer_(1985)____	58.7	1.19
15.5	9.4	.20	2.5	6.2	4.0	2.1	.19							

Date: 5.11.2016. Analyses read from file EXAMPLE.DAT (recalculated to 100%). The database holds 3591 analyses for comparison.

Comparison using LOG wt.-% oxides. No oxides ommited from comparison.

No.	dif.	name of rock	location								source	SiO2	TiO2	
Al2O3	FeOt	MnO	MgO	CaO	Na2O	K2O	P2O5							

		AP-009											70.8	
		.23	16.2	2.6	.10	.7	3.0	4.2	2.1	.09				

1	100	Ocoña									granite	Mamani_et_al.,_2008_	72.5	.24
14.2	2.9	.07	.8	3.0	3.6	2.7	.08							
2	127	Pausa									rhyolite	Mamani_et_al.,_2008_	71.6	.35
14.4	2.7	.08	.7	2.4	4.1	3.6	.10							
3	127	Santa Marta									Granodiorite	Quandt_2013_____	69.9	.22
17.0	2.0	.04	.6	3.6	4.6	1.9	.09							
4	155	Abitagua									Granodiorite	Cochrane_2014a_____	63.8	.20
19.8	3.2	.13	.5	2.8	5.5	3.9	.09							

16.1 4.1 .07 1.4 4.6 5.7 .6 .11

No.	dif.	name of rock	location	source	SiO2	TiO2		
1	100	Amphibolic Gneiss	SK-5003-RA	Lopez_2012_____	70.2	.38		
	13.5	4.9	.08	2.0	4.8	3.4	.6	.09
2	144	Ilo	granodiorite	Mamani_et_al.,_2008_	67.0	.45		
	16.1	4.5	.07	2.0	4.4	3.6	1.7	.13
3	154	Punta de Bombon	granodiorite	Mamani_et_al.,_2008_	67.2	.39		
	15.9	4.0	.06	1.7	4.6	3.3	2.8	.12
4	157	dacite	after Johannsen (1952)	Davoine,_P._(1968)__	64.6	.38		
	17.3	4.1	.09	2.3	5.0	4.1	2.0	.10
5	159	Granodiorite	Cordoba Pluton	Villagomez_2011_____	62.3	.67		
	17.5	5.7	.09	1.7	5.8	5.2	.7	.20
6	165	Zamora	Granodiorite	Cochrane_2014a_____	66.1	.35		
	17.5	4.1	.10	1.4	4.4	3.5	2.5	.09
7	172	Ocoña dike	amphibolite	Mamani_et_al.,_2009_	65.3	.41		
	17.9	3.8	.12	1.9	4.9	3.8	1.7	.14
8	174	Amphibolic Gneiss	DRR-6-052A	Lopez_2012_____	69.0	.27		
	14.1	5.5	.07	1.5	4.9	4.3	.3	.04
9	174	El Abra	monzonite	Ruprecht__2007_____	65.5	.49		
	17.1	3.7	.06	1.3	4.2	4.9	2.5	.19
10	176	Anorthosit	Labrador	Wimmenauer_(1985)___	57.0	.79		
	22.2	4.6	.06	1.4	7.1	5.4	1.4	.12

Date: 5.11.2016. Analyses read from file EXAMPLE.DAT (recalculated to 100%). The database holds 3591 analyses for comparison.

Comparison using LOG wt.-% oxides. No oxides omitted from comparison.

No.	dif.	name of rock	location	source	SiO2	TiO2					
		Al2O3	FeO	MnO	MgO	CaO	Na2O	K2O	P2O5		

		EMP-001									70.5	.30
		14.7	3.3	.11	1.1	3.0	3.9	3.0	.11			

1	100	Salar de la Isla	dacite	Siebel_et_al.,_2001_	69.1	.46						
	15.7	3.2	.11	1.3	3.0	3.1	3.8	.14				
2	102	Segovia	Monzogranite	Cochrane_2014a_____	69.1	.51						
	15.2	3.9	.13	1.2	3.4	3.7	2.7	.13				
3	114	Granodiorite	Coastal Batholith of Peru	Wilson,M._(1989)___	69.5	.42						
	15.1	3.0	.07	1.2	2.9	3.5	4.1	.10				
4	117	Huanynaputina	dacite	Mamani_et_al.,_2008_	69.1	.40						
	15.7	3.0	.06	1.0	2.8	4.5	3.2	.14				
5	118	Rhyolite	Saldaña Fm.	Villagomez_2011_____	66.1	.47						
	16.6	3.8	.09	1.3	3.0	4.4	4.1	.12				
6	120	Pausa	rhyolite	Mamani_et_al.,_2008_	71.6	.35						
	14.4	2.7	.08	.7	2.4	4.1	3.6	.10				
7	120	Misti	dacite	Mamani_et_al.,_2009_	68.1	.43						
	15.9	3.5	.07	1.3	3.4	3.9	3.4	.15				
8	121	Santa Marta	Granodiorite	Quandt_2013_____	69.2	.46						
	15.6	3.1	.06	1.2	3.0	4.1	3.1	.15				
9	123	Ocoña	granite	Mamani_et_al.,_2008_	72.5	.24						
	14.2	2.9	.07	.8	3.0	3.6	2.7	.08				
10	126	Zamora	Granodiorite	Cochrane_2014a_____	66.1	.35						
	17.5	4.1	.10	1.4	4.4	3.5	2.5	.09				

Date: 5.11.2016. Analyses read from file EXAMPLE.DAT (recalculated to 100%). The database holds 3591 analyses for comparison.

Comparison using LOG wt.-% oxides. No oxides ommited from comparison.

No.	dif.	name of rock	location							source	SiO2	TiO2
			Al2O3	FeOt	MnO	MgO	CaO	Na2O	K2O	P2O5		
		EMP-270									74.2	.07
			14.8	1.4	.03	.1	1.2	4.5	3.8	.01		
1	100	Saldana fm.										
			13.2	1.5	.03	.2	1.0	4.0	3.3	.04		
											76.6	.19
2	108	Lauca/Pérez-Ignimbrit										
			12.4	.8	.03	.1	.8	4.5	4.9	.02		
											76.2	.13
3	111	Abitagua										
			14.3	1.5	.04	.2	1.5	3.7	3.8	.05		
											74.8	.17
4	116	Salar de la Isla										
			13.6	.8	.07	.1	1.1	3.5	4.5	.04		
											76.1	.10
5	119	Azafran										
			13.2	.8	.06	.1	.8	3.6	4.7	.03		
											76.7	.12
6	123	Yura										
			13.1	1.1	.07	.2	1.0	4.1	4.4	.03		
											75.9	.15
7	124	Arequipa										
			12.9	1.0	.05	.1	.8	4.3	4.8	.02		
											75.7	.16
8	124	Salar de Antofalla										
			12.7	.9	.06	.2	.8	4.0	4.6	.03		
											76.5	.15
9	125	Salar de Antofalla										
			13.2	.7	.05	.1	1.4	3.5	5.0	.02		
											76.0	.10
10	126	Yura										
			13.3	1.0	.07	.2	1.1	4.1	4.4	.03		
											75.7	.16

Date: 5.11.2016. Analyses read from file EXAMPLE.DAT (recalculated to 100%). The database holds 3591 analyses for comparison.

Comparison using LOG wt.-% oxides. No oxides ommited from comparison.

No.	dif.	name of rock	location							source	SiO2	TiO2
			Al2O3	FeOt	MnO	MgO	CaO	Na2O	K2O	P2O5		
		16-JDMA-028-028									50.9	
			.86	17.7	11.5	.15	6.0	7.3	4.1	1.3	.25	
1	100	Basalt,High-K-calc-alk.										
			18.1	9.8	.18	5.6	9.7	3.7	1.2	.25		
											50.3	1.13
2	101	Lascar										
			16.9	8.1	.13	5.4	7.9	3.5	1.2	.23		
											56.0	.84
3	101	Lascar										
			17.2	7.8	.13	4.9	7.8	3.6	1.3	.24		
											56.2	.85
4	109	Basaltic andesite										
			17.1	8.6	.16	5.3	9.1	2.9	1.3	.26		
											54.4	.83
5	118	Spilite, Av.92										
											52.5	1.66

16.9	10.1	.16	5.4	7.0	4.5	1.4	.28				
6	119	spilite						Rhein.Schiefergeb.Germany	Schulz-Dobrick(1975)	51.2	1.49
17.3	12.0	.14	4.3	8.2	4.0	1.2	.23				
7	123	Linga Yarabamba						gabbro	Mamani_et_al.,_2009_	56.3	.86
17.1	8.2	.14	4.4	7.1	4.1	1.6	.19				
8	124	Basaltic andesites						Andes NVZ	Wilson,M._(1989)_____	56.1	.90
17.0	7.9	.10	5.2	7.6	3.9	1.1	.23				
9	127	Andesit, Mean						Chile	Wimmenauer_(1985)____	53.2	1.07
18.4	8.6	.14	5.3	8.2	4.1	.8	.26				
10	128	amygdaloidal spilite						Schirmeck, Vosges, France	Amstutz_(ed)_(1974)_	52.1	1.05
16.7	10.3	.22	7.7	6.2	4.5	1.3	.				

TABLE 6. LOCATION, DESCRIPTION, OF SAMPLES ANALYZED BY U-Pb LA-ICP-MS

Sample Code	Coordinates		Elevation (m)	Lithology	Stratigraphic division	Unit
	Lat (°N)	Long (°W)				
CVI13-108	10.93455475	-74.1374652	545	Grt-Bt Mylonite	Triassic	La Secreta Mylonites
MPR-33A	11.07315844	-74.0739827	1000	Hbl+Plg+Qtz, Orthogneiss	Permian	El Encanto Orthogneiss
GLV-11	11.02193134	-74.1710838	112	Ms+Bt+Grt schist	Permian	Gaira Schists
MG-063	11.24749859	-73.7426375	25	Bt+Qtz schist	Permian	Gaira Schists
LRW-21	10.91291052	-74.1444955	141	Plg+Grt+Qtz gneiss	Triassic	La Secreta Mylonites
CP-055	11.31664991	-74.0792714	6	Psephite	Upper Cretaceous	Cinto Formation
ABH-27	11.16934363	-74.031142	1320	Bt+Ms schist	Jurassic	Lower San Lorenzo Schists

AP-009	10.80251034	-74.0296855	779	Granodiorite	Eocene	Sevilla Stock
CVI1367	11.30005034	-74.1891957	69	Meta-andesite	Paleogene	Concha Formation
CVI1344	11.31848483	-74.0688197	160	Dioritic dike	Paleogene	Concha Formation
AK-01	10.88033347	-74.1140673	487	Metapelite	Jurassic	Lower San Lorenzo Schists
AP-038	11.19673862	-73.8712134	110	Leucosome	Eocene	Guachaca Migmatites
AP-011	10.87057863	-74.0463749	663	Gneiss	Neoproterozoic	Lower San Lorenzo Schists
DAA-039	11.16191682	-73.7866815	168	Gneiss	Neoproterozoic	Lower San Lorenzo Schists
AP-010	10.81871073	-74.0283751	923	Migmatite	Neoproterozoic	Lower San Lorenzo Schists
AP-048	11.26295572	-73.5921871	9	Gneiss	Neoproterozoic	Lower San Lorenzo Schists

APPENDIX CHAPTER 2

TABLE A1. ZFT Analytical data from the Sierra Nevada de Santa Marta

Sample	Lithology	Stratigraphic Age	Elevation(m)	Lat	Long	Grains	RhoS Track/cm2	RhoI Track/cm2	RhoD Track/cm2	U(ppm)	P(x2) Binomfit	Pooled fission Track Age (Ma)	$\pm 2\sigma$	Central fission Track Age (Ma)	$\pm 2\sigma$
CV113-108	Gr-Bio Filonite	Triassic	545	-74.1375	10.9346	20	4.28E+06	2.67E+06	3.094E+05	431	0	34.9	2.1	34.7	1.9
AP-009	Tonalite	Eocene	779	-74.0297	10.8025	20	3.07E+06	1.84E+06	3.097E+05	297	1.2	36.3	2.2	36.1	1.3
MPR-33A	Orthogneiss	Permian	1000	-74.0740	11.0732	22	3.61E+06	2.16E+06	3.094E+05	349	61.8	36.3	2.9	36.3	1
MPR-34A	Schist	Permian	1054	-74.0777	11.0699	21	3.65E+06	2.17E+06	3.094E+05	351	5.7	36.6	3.7	36.6	1.8
CP-055	Psephite	per Cretaceoi	6	-74.0793	11.3166	20	8.67E+05	7.41E+06	3.095E+05	85	63.7	42.3	3.4	42.4	1.3
AP-042	Tonalite	Eocene	89	-73.7238	11.2029	23	1.05E+06	5.07E+05	3.097E+05	82	58.3	44.9	4.5	45	2
AP-010	Migmatite	leoproterozoi	923	-74.0284	10.8187	25	1.02E+06	4.92E+05	3.097E+05	79	61.1	45.3	4.5	45.3	1.9
AP-048	Gneiss	leoproterozoi	9	-73.5922	11.2630	24	2.14E+06	1.01E+06	3.096E+05	163	0	46.1	3.7	45.8	3.8
CV113-120	Mylonite	Jurassic	426	-74.1076	10.9860	24	1.36E+06	6.23E+05	3.095E+05	101	0	47.3	3.8	47.5	2.8
AP-011	Gneiss	Neoproterozo	663	-74.0464	10.8706	24	5.29E+05	2.34E+05	3.097E+05	38	78.9	49.1	5.9	49.1	2.9
DAA-039	Gneiss	leoproterozoi	168	-73.7867	11.1619	25	1.16E+06	4.99E+05	3.095E+05	81	0	50.6	5.1	49.3	3
MG-063	Schist	Permian	25	-73.7426	11.2475	20	2.12E+06	8.94E+05	3.094E+05	145	8	51.6	5.2	53.8	3.1

TABLE A2. AFT Analytical data from the Sierra Nevada de Santa Marta

Sample	Lithology	Stratigraphic Age	Elevation(m)	Lat	Long	Grains	RhoS Track/cm2	RhoI Track/cm2	RhoD Track/cm2	U(ppm)	P(x2) Binomfit	Pooled fission Track Age (Ma)	± 2σ	Central fission Track Age (Ma)	± 2σ
CVI13-108	Gr-Bio Filonite	Triassic	545	-74.1375	10.9346	20	2.32E+05	2.79E+06	1.335E+06	31	93.2	15	2	14.96	0.9
	Gneiss Hbn-														
LRW-21	Gr	Triassic	141	-74.1445	10.9129	25	9.33E+04	1.22E+06	1.330E+06	14	100	13.69	2.45	13.7	1.2
MPR-33A	Orthogneiss	Permian	1000	-74.0740	11.0732	22	2.43E+05	2.53E+06	1.326E+06	29	97.3	17.17	2.33	17.2	1.1
GLV-11	Musc-Gr-Schis	Permian	112	-74.1711	11.0219	20	2.77E+05	3.50E+06	1.322E+06	40	99.8	14.11	2.1	14.11	0.97
CVI13-120	Mylonite	Jurassic	426	-74.1076	10.9860	21	1.32E+05	1.46E+06	1.318E+06	17	54	16.2	3.25	16.2	1.6
CVI13-67	Meta tuff	Paleogene	69	-74.1892	11.3001	22	1.31E+05	5.74E+05	1.313E+06	7	14.7	40.21	7.08	40.5	3.7
CVI13-44	Phyllite	Paleogene	160	-74.0688	11.3185	19	2.30E+05	1.90E+06	1.309E+06	22	69.9	21.35	3.25	21.3	1.5
CP-055	Psephite	per Cretaceo	6	-74.0793	11.3166	20	8.67E+05	7.41E+06	1.301E+06	85	63.7	20.6	1.77	20.6	0.66
AP-048	Gneiss	leoproterozoi	9	-73.5922	11.2630	25	1.96E+05	9.19E+05	1.297E+06	11	75.7	37.25	5.15	37.2	2.3
AP-042	Tonalite	Eocene	89	-73.7238	11.2029	22	3.11E+05	1.31E+06	1.292E+06	15	80.3	41.3	5.12	41.3	2.3
AP-009	Tonalite	Eocene	779	-74.0297	10.8025	24	1.31E+05	6.10E+05	1.284E+06	7	48.8	37.17	6.15	37.4	3.1

TABLE A3. (U-Th-Sm)/He analytical data from the Sierra Nevada de Santa Marta

Sample	Coordinates		Grain ID	He fmol	238U fmol	U ppm	232Th fmol	Th ppm	147Sm fmol	Sm ppm	Mass microng	eU ppm	RAW AGE Ma	Ft EU	Ft EL	Ft CY	SHAPE	CORR.AGE Ma	age error Ma (1σ)	AVG Age Ma	error Ma (1σ)	replicate #	U blk %	Th blk %	Rs EU micromm	Rs EL micromm	Rs CY micromm
	Lat	Long																									
AP-010		923	ap010a1	0.17	11.18	1.46	6.17	1.10	136.09	113.71	1.35	1.72	12.57	0.86	0.87	0.52	CY	13.61	0.96	13.61	0.96	1	29.79	28.89	42.24	48.97	42.24
			ap010a2	0.16	12.28	0.43	1.54	0.05	472.73	67.12	6.90	0.44	8.16	0.91	0.91	0.96	CY	8.70	0.43	19.84	119.65	66.66	67.06	119.67	67.06	119.67	
			ap010a3	0.03	61.17	9.04	6.29	0.90	118.15	71.38	1.62	9.25	0.38	0.81	0.82	0.89	CY	0.42	0.18	3.98	29.29	41.00	44.26	73.75	41.00	44.26	73.75
			ap010a4	0.01	0.93	0.19	24.62	4.99	60.20	51.51	1.15	1.17	1.62	0.78	0.80	0.88	EU	2.08	1.10	263.85	7.46	36.03	39.54	65.54	36.03	39.54	65.54
			ap010a5	0.07	5.78	0.65	0.24	0.03	199.20	91.89	2.12	0.66	7.50	0.89	0.90	0.94	EU	8.43	0.99	42.21	3090.88	42.44	48.58	78.96	42.44	48.58	78.96
AP-042	89		ap042a1	1.36	174.77	9.49	314.84	16.54	244.65	54.29	4.42	13.37	4.25	0.79	0.81	0.89	EU	5.35	0.11	6.84	0.89	5	1.40	1.95	57.09	61.78	102.94
			ap042a2	4.82	294.11	11.73	658.06	25.41	514.57	83.93	6.01	17.71	8.34	0.81	0.82	0.90	EU	10.28	0.07	0.83	0.93	61.84	64.78	113.79	61.84	64.78	113.79
			ap042a3	1.50	153.82	9.23	310.73	18.05	229.32	56.26	3.99	13.47	5.12	0.79	0.80	0.88	EU	6.52	0.12	1.60	1.98	55.18	60.04	99.36	55.18	60.04	99.36
			ap042a4	1.44	149.79	8.01	311.16	16.11	227.10	49.67	4.48	11.80	5.02	0.79	0.79	0.88	EU	6.39	0.03	1.03	1.73	56.01	58.36	103.18	56.01	58.36	103.18
			ap042a5	0.38	47.00	2.97	82.04	5.02	62.59	16.28	2.79	4.15	4.40	0.78	0.78	0.88	EU	5.67	0.07	3.19	6.58	52.25	54.34	97.81	52.25	54.34	97.81
AP-048	9		ap048a1	12.65	320.21	12.07	1111.04	40.55	4418.92	681.16	6.36	21.60	16.37	0.89	0.90	0.94	EL	18.22	0.07	18.30	1.69	5	0.77	0.55	63.08	72.33	114.85
			ap048a2	13.56	411.29	15.48	325.54	11.86	3463.27	532.83	6.37	18.27	20.81	0.90	0.91	0.95	CY	22.00	0.09	0.60	1.89	63.92	67.88	116.52	63.92	67.88	116.52
			ap048a3	6.25	296.35	16.14	198.54	10.46	2784.09	603.85	4.40	18.60	13.67	0.89	0.90	0.94	CY	14.52	0.10	0.83	1.10	56.77	60.93	102.61	56.77	60.93	102.61
			ap048a4	8.18	194.17	10.59	388.35	20.50	3736.36	833.14	4.40	15.41	20.91	0.89	0.90	0.94	CY	22.20	0.11	1.26	1.58	56.60	57.60	102.70	56.60	57.60	102.70
			ap048a5	4.53	200.27	10.73	181.98	9.44	2526.00	553.23	4.47	12.95	13.76	0.90	0.90	0.94	CY	14.58	0.14	1.23	3.38	58.23	61.03	103.55	58.23	61.03	103.55
CV113-108	545		cv113108a1	0.99	102.83	13.81	-1.15	-0.13	49.91	23.43	2.09	11.78	7.42	0.75	0.76	0.87	EL	9.56	0.37	5.29	552.51	42.88	48.51	79.01	42.88	48.51	79.01
			cv113108a2	1.39	147.70	15.28	-1.26	-0.13	92.68	39.18	2.32	15.25	7.27	0.78	0.79	0.88	CY	8.28	0.25	1.66	490.75	45.79	49.18	82.73	45.79	49.18	82.73
			cv113108a3	0.32	48.69	3.36	-3.73	-0.25	51.21	14.46	3.47	3.30	5.18	0.83	0.84	0.91	CY	5.71	0.50	5.04	164.94	52.24	55.35	95.17	52.24	55.35	95.17
			cv113108a4	0.30	40.48	6.55	0.00	0.00	36.23	23.98	1.48	6.55	5.67	0.76	0.77	0.87	EL	7.39	0.10	3.82	195.13	39.20	40.73	71.56	39.20	40.73	71.56
			cv113108a5	0.21	18.47	2.89	0.00	0.00	39.10	25.04	1.53	2.89	8.75	0.82	0.83	0.90	EL	10.57	0.37	8.17	176.60	40.40	43.73	72.16	40.40	43.73	72.16
CV113-120	426		cv113120a1	12.29	480.31	8.81	951.71	36.89	352.28	26.40	11.08	12.77	13.57	0.83	0.84	0.91	EU	16.27	0.06	11.73	3.08	3	0.51	0.65	80.86	87.24	147.06
			cv113120a2	7.39	418.07	11.53	886.58	22.58	772.50	83.30	9.11	16.83	8.85	0.84	0.84	0.91	CY	9.71	0.06	0.56	0.69	71.88	73.75	130.90	71.88	73.75	130.90
			cv113120a3	1.78	150.44	2.00	37.96	0.49	134.07	7.30	17.99	2.12	8.64	0.89	0.89	0.94	CY	9.21	0.22	1.63	16.20	91.10	93.15	164.70	91.10	93.15	164.70
			DAA03a1	0.16	24.42	1.19	2.30	0.11	115.06	22.87	4.93	1.21	4.84	0.89	0.90	0.94	EL	5.36	0.25	19.83	163.34	57.99	64.94	105.60	57.99	64.94	105.60
			DAA03a2	0.04	18.51	0.59	0.09	0.00	134.60	22.24	5.94	0.58	1.98	0.92	0.92	0.95	EL	2.04	0.36	33.37	5273.12	63.46	68.52	115.57	63.46	68.52	115.57
LRW-21	141		DAA03a3	0.13	11.62	0.65	16.72	0.91	38.84	8.48	4.26	0.87	6.28	0.83	0.83	0.91	CY	6.90	0.48	41.66	22.36	54.39	56.02	101.34	54.39	56.02	101.34
			DAA03a4	0.17	45.98	3.72	54.65	4.27	51.09	16.88	2.97	4.72	2.44	0.77	0.77	0.87	CY	2.57	0.09	3.36	9.87	49.44	51.75	93.80	49.44	51.75	93.80
			lrw21a1	7.75	664.58	10.72	393.50	6.14	2360.00	142.42	14.86	11.16	7.83	0.91	0.91	0.95	CY	8.25	0.03	0.17	0.47	84.77	85.43	154.51	84.77	85.43	154.51
			lrw21a2	1.74	194.47	10.17	31.14	1.58	476.27	101.82	4.58	10.54	6.59	0.87	0.87	0.93	EL	7.54	0.05	1.25	5.90	58.19	60.72	104.40	58.19	60.72	104.40
			lrw21a3	3.45	301.23	4.79	63.03	0.97	571.30	37.16	15.07	5.02	8.88	0.80	0.91	0.95	EL	9.21	0.04	0.81	2.51	83.35	83.46	153.30	83.35	83.46	153.30
MPR-33A	1000		lrw21a4	9.50	770.27	7.77	343.64	3.56	3280.71	195.32	23.76	8.56	8.49	0.93	0.93	0.96	EL	9.11	0.03	0.20	1.57	100.89	105.61	180.49	105.61	180.49	
			lrw21a5	2.66	269.14	6.73	108.80	2.63	1113.72	113.79	9.59	7.35	6.87	0.90	0.91	0.95	CY	7.26	0.03	0.57	4.96	73.75	76.88	133.54	73.75	76.88	133.54
			mpr33a1	0.84	65.28	4.83	69.72	4.99	48.02	14.51	3.24	6.00	7.99	0.76	0.78	0.87	EU	10.51	0.12	0.83	0.43	50.91	56.40	92.07	50.91	56.40	92.07
			mpr33a2	1.67	270.14	25.59	320.03	29.34	137.33	53.17	2.53	32.48	8.24	0.72	0.74	0.85	EU	11.40	0.04	0.90	0.57	47.07	52.31	85.16	47.07	52.31	85.16
			mpr33a3	2.75	233.72	28.24	260.60	30.48	112.59	55.61	1.98	35.40	7.24	0.69	0.70	0.83	EU	10.45	0.06	1.04	0.70	42.70	44.88	78.82	42.70	44.88	78.82
mpr33a4	7.58	561.40	34.10	422.29	24.82	1111.34	27.65	3.95	39.93	8.90	0.75	0.75	0.86	CY	10.14	0.04	0.43	0.43	55.57	57.08	99.21	55.57	57.08	99.21			
mpr33a5	5.19	451.19	20.69	383.08	17.63	174.28	33.87	5.04	24.83	7.67	0.78	0.78	0.88	EU	9.85	0.04	0.56	0.48	59.51	61.26	107.79	59.51	61.26	107.79			

APPENDIX CHAPTER 3

EMP-70	2.894	0.002	0.0216	0.0052	0.088	0.0019	0.0715	0.0056	1356	22	1342	27	1356	110	1377	41	415	116.7	0.2812	1342	27
EMP-71	2.876	0.002	0.0203	0.0051	0.0869	0.0019	0.0699	0.0055	1318	17	1318	16	1318	110	1339	32	329	116.4	0.279	1318	16
EMP-72	1.158	0.006	0.247	0.0071	0.0207	0.0075	0.07	0.025	1446	23	1423	39	1366	84	1480	46	500	60.0	0.6075	1423	39
EMP-73	1.262	0.001	0.1732	0.0023	0.072	0.0022	0.082	0.0025	1000	17	1000	17	1000	110	1021	33	28	43.2	0.4023	1000	17
EMP-74	1.85	0.001	0.1017	0.0011	0.052	0.0011	0.062	0.0012	1000	17	1000	17	1000	110	1021	33	28	43.2	0.4023	1000	17
EMP-75	1.85	0.001	0.1017	0.0011	0.052	0.0011	0.062	0.0012	1000	17	1000	17	1000	110	1021	33	28	43.2	0.4023	1000	17
EMP-76	1.85	0.001	0.1017	0.0011	0.052	0.0011	0.062	0.0012	1000	17	1000	17	1000	110	1021	33	28	43.2	0.4023	1000	17
EMP-77	1.707	0.004	0.1703	0.0036	0.0535	0.0036	0.0622	0.0046	1006	24	1014	27	1027	110	1043	37	42.4	43.6	0.3062	1014	27
EMP-78	1.85	0.001	0.1017	0.0011	0.052	0.0011	0.062	0.0012	1000	17	1000	17	1000	110	1021	33	28	43.2	0.4023	1000	17
EMP-79	2.427	0.078	0.2528	0.005	0.0626	0.0018	0.0685	0.0057	1249	24	1242	27	1237	110	1256	34	170	21.8	0.281	1242	27
EMP-80	2.91	0.002	0.113	0.0063	0.087	0.0021	0.0663	0.0057	1326	19	1326	19	1326	110	1347	40	449	18.1	0.282	1326	19
EMP-81	1.859	0.034	0.166	0.0032	0.074	0.0034	0.0485	0.0043	1003	19	989	18	977	83	1015	35	207	67.6	0.324	989	18
EMP-82	1.854	0.032	0.1664	0.0032	0.074	0.0032	0.0485	0.0043	1006	19	992	18	980	84	1022	34	208	68.4	0.324	992	18
EMP-83	1.854	0.032	0.1664	0.0032	0.074	0.0032	0.0485	0.0043	1006	19	992	18	980	84	1022	34	208	68.4	0.324	992	18
EMP-84	1.854	0.032	0.1664	0.0032	0.074	0.0032	0.0485	0.0043	1006	19	992	18	980	84	1022	34	208	68.4	0.324	992	18
EMP-85	1.854	0.032	0.1664	0.0032	0.074	0.0032	0.0485	0.0043	1006	19	992	18	980	84	1022	34	208	68.4	0.324	992	18
EMP-86	1.854	0.032	0.1664	0.0032	0.074	0.0032	0.0485	0.0043	1006	19	992	18	980	84	1022	34	208	68.4	0.324	992	18
EMP-87	1.854	0.032	0.1664	0.0032	0.074	0.0032	0.0485	0.0043	1006	19	992	18	980	84	1022	34	208	68.4	0.324	992	18
EMP-88	1.854	0.032	0.1664	0.0032	0.074	0.0032	0.0485	0.0043	1006	19	992	18	980	84	1022	34	208	68.4	0.324	992	18
EMP-89	1.854	0.032	0.1664	0.0032	0.074	0.0032	0.0485	0.0043	1006	19	992	18	980	84	1022	34	208	68.4	0.324	992	18
EMP-90	1.854	0.032	0.1664	0.0032	0.074	0.0032	0.0485	0.0043	1006	19	992	18	980	84	1022	34	208	68.4	0.324	992	18
EMP-91	1.854	0.032	0.1664	0.0032	0.074	0.0032	0.0485	0.0043	1006	19	992	18	980	84	1022	34	208	68.4	0.324	992	18
EMP-92	2.12	0.14	0.193	0.004	0.085	0.0038	0.0734	0.0272	913	19	727	14	1410	140	1387	59	11.8	33.2	0.1956	727	14
EMP-93	1.854	0.032	0.1664	0.0032	0.074	0.0032	0.0485	0.0043	1006	19	992	18	980	84	1022	34	208	68.4	0.324	992	18
EMP-94	1.854	0.032	0.1664	0.0032	0.074	0.0032	0.0485	0.0043	1006	19	992	18	980	84	1022	34	208	68.4	0.324	992	18
EMP-95	1.854	0.032	0.1664	0.0032	0.074	0.0032	0.0485	0.0043	1006	19	992	18	980	84	1022	34	208	68.4	0.324	992	18
EMP-96	2.552	0.038	0.1844	0.0029	0.0768	0.0012	0.0572	0.004	1068	13	1101	16	1124	77	1114	33	349	217.1	0.484	1068	13
EMP-97	1.784	0.051	0.1921	0.0052	0.0921	0.0057	0.0644	0.0057	1075	16	1068	14	1108	54	1108	54	187	18.7	0.281	1075	16
EMP-98	1.858	0.038	0.1844	0.0029	0.0768	0.0012	0.0572	0.004	1068	13	1101	16	1124	77	1114	33	349	217.1	0.484	1068	13
EMP-99	1.858	0.038	0.1844	0.0029	0.0768	0.0012	0.0572	0.004	1068	13	1101	16	1124	77	1114	33	349	217.1	0.484	1068	13
EMP-100	1.858	0.038	0.1844	0.0029	0.0768	0.0012	0.0572	0.004	1068	13	1101	16	1124	77	1114	33	349	217.1	0.484	1068	13
EMP-101	1.858	0.038	0.1844	0.0029	0.0768	0.0012	0.0572	0.004	1068	13	1101	16	1124	77	1114	33	349	217.1	0.484	1068	13
EMP-102	1.858	0.038	0.1844	0.0029	0.0768	0.0012	0.0572	0.004	1068	13	1101	16	1124	77	1114	33	349	217.1	0.484	1068	13
EMP-103	1.858	0.038	0.1844	0.0029	0.0768	0.0012	0.0572	0.004	1068	13	1101	16	1124	77	1114	33	349	217.1	0.484	1068	13
EMP-104	1.858	0.038	0.1844	0.0029	0.0768	0.0012	0.0572	0.004	1068	13	1101	16	1124	77	1114	33	349	217.1	0.484	1068	13
EMP-105	1.858	0.038	0.1844	0.0029	0.0768	0.0012	0.0572	0.004	1068	13	1101	16	1124	77	1114	33	349	217.1	0.484	1068	13
EMP-106	1.858	0.038	0.1844	0.0029	0.0768	0.0012	0.0572	0.004	1068	13	1101	16	1124	77	1114	33	349	217.1	0.484	1068	13
EMP-107	1.758	0.032	0.1709	0.0028	0.0746	0.0032	0.0444	0.0039	1030	12	1017	17	1015	75	1066	32	229.9	68.09	0.2960	1017	17
EMP-108	1.120	0.006	0.1042	0.0043	0.0783	0.0014	0.0444	0.0033	74	28	38	25	801	44	1111	36	471	110.7	0.2360	801	25
EMP-109	1.120	0.006	0.1042	0.0043	0.0783	0.0014	0.0444	0.0033	74	28	38	25	801	44	1111	36	471	110.7	0.2360	801	25
EMP-110	1.120	0.006	0.1042	0.0043	0.0783	0.0014	0.0444	0.0033	74	28	38	25	801	44	1111	36	471	110.7	0.2360	801	25
EMP-111	2.18	0.16	0.174	0.011	0.0604	0.0015	0.0520	0.0049	1162	48	1030	59	1041	94	1030	59	1431	33	0.4178	1162	48
EMP-112	1.737	0.032	0.1671	0.0031	0.0726	0.0032	0.0415	0.0026	1015	19	1011	19	1006	86	1006	86	208	68.4	0.324	1015	19
EMP-113	1.737	0.032	0.1671	0.0031	0.0726	0.0032	0.0415	0.0026	1015	19	1011	19	1006	86	1006	86	208	68.4	0.324	1015	19
EMP-114	1.746	0.036	0.1686	0.0034	0.0736	0.0034	0.0420	0.0029	1019	27	1018	27	1018	27	1018	27	214	14.2	0.4178	1019	27
EMP-115	1.58	0.055	0.16	0.0029	0.0719	0.0022	0.05	0.0043	967	21	956	16	986	83	984	58	52.8	27.3	0.1950	956	16
EMP-116	1.67	0.031	0.1667	0.0031	0.0727	0.0031	0.0467	0.0028	1018	12	1007	12	1007	12	1007	12	214	14.2	0.4178	1018	12
EMP-117	1.925	0.001	0.02659	0.00029	0.0489	0.0021	0.0692	0.0079	1785	6.9	1817	2.2	179	16	140	89	52.8	27.3	0.1950	1785	6.9
EMP-118	1.737	0.032	0.1671	0.0031	0.0726	0.0032	0.0415	0.0026	1015	19	1011	19	1006	86	1006	86	208	68.4	0.324	1015	19
EMP-119	1.737	0.032	0.1671	0.0031	0.0726	0.0032	0.0415	0.0026	1015	19	1011	19	1006	86	1006	86	208	68.4	0.324	1015	19
EMP-120	1.852	0.031	0.1635	0.003	0.0733	0.0032	0.0415	0.0026	1015	19	1011	19	1006	86	1006	86	208	68.4	0.324	1015	19
EMP-121	1.852	0.031	0.1635	0.003	0.0733	0.0032	0.0415	0.0026	1015	19	1011	19	1006	86	1006	86	208	68.4	0.324	1015	19
EMP-122	0.216	0.03	0.02604	0.00098	0.0558	0.0026	0.0096	0.0015	136	25	173.5	6.1	191	30	209	26	20.1	19.3	0.7633	136	25
EMP-123	1.756	0.044	0.1682	0.0034	0.0736	0.0034	0.0420	0.0029	1019	27	1018	27	1018	27	1018	27	214	14.2	0.4178	1019	27
EMP-124	2.783	0.055	0.232	0.0037	0.087	0.0014	0.0704	0.0051	1350	15	1345	20	1376	96	1358	31	203	119.3	0.4712	1345	20
EMP-125	1.859	0.046	0.1882	0.0037	0.0768	0.0012	0.0572	0.004	1068	13	1101	16	1124	77	1114	33	349	217.1	0.484	1068	13
EMP-126	1.859	0.046	0.1882	0.0037	0.0768	0.0012	0.0572	0.004	1068	13	1101	16	1124	77	1114	33	349	217.1	0.484	1068	13
EMP-127	1.859	0.046	0.1882	0.0037	0.0768	0.0012	0.0572	0.004	1068	13	1101	16	1124	77	1114	33	349	217.1	0.484	1068	13
EMP-128	1.859	0.046	0.1882	0.0037	0.0768	0.0012	0.0572	0.004	1068	13	1101	16	1124	77	1114	33	349	217.1	0.484	1068	13
EMP-129	1.859	0.046	0.1882	0.0037	0.0768	0.0012	0.0572	0.004	1068	13	1101	16	1124	77	1114	33	349	217.1	0.484</		

MI_CP88-2	3.381	0.059	0.1338	0.0465	0.2072	0.0077	0.0389	0.0073	878	22	207	808	25	710	140	1014	407	22	3	0.1384	808	25	25
MI_CP88-1	2.242	0.071	0.2042	0.0832	0.0302	0.0063	0.0866	0.0063	1300	13	1003	1035	8.8	802	10	1003	8.8	802	81	0.0707	1035	8.8	81
MI_CP88-4	2.522	0.067	0.2188	0.0822	0.0285	0.0068	0.0862	0.0079	1380	19	1275	1289	16	1307	87	1060	1286	16	137	0.0550	1275	16	137
MI_CP88-5	2.582	0.068	0.2265	0.0835	0.0285	0.0068	0.0861	0.0086	1356	19	1275	1289	16	1307	87	1060	1286	16	137	0.0550	1275	16	137
MI_CP88-7	2.722	0.067	0.2325	0.0845	0.0285	0.0068	0.0862	0.0082	1334	18	1275	1289	16	1307	87	1060	1286	16	137	0.0550	1275	16	137
MI_CP88-8	2.722	0.067	0.2325	0.0845	0.0285	0.0068	0.0862	0.0082	1334	18	1275	1289	16	1307	87	1060	1286	16	137	0.0550	1275	16	137
MI_CP88-9	2.622	0.065	0.2192	0.0831	0.0285	0.0068	0.0861	0.0086	1311	18	1275	1289	16	1307	87	1060	1286	16	137	0.0550	1275	16	137
MI_CP88-10	2.622	0.065	0.2192	0.0831	0.0285	0.0068	0.0861	0.0086	1311	18	1275	1289	16	1307	87	1060	1286	16	137	0.0550	1275	16	137
MI_CP88-11	2.670	0.063	0.2225	0.0836	0.0285	0.0068	0.0862	0.0085	1320	18	1275	1289	16	1307	87	1060	1286	16	137	0.0550	1275	16	137
MI_CP88-12	2.977	0.067	0.2577	0.0877	0.0285	0.0068	0.0862	0.0087	1344	19	1275	1289	16	1307	87	1060	1286	16	137	0.0550	1275	16	137
MI_CP88-13	2.54	0.063	0.2209	0.0830	0.0285	0.0068	0.0861	0.0073	1227	19	1228	126	20	1320	140	1216	24	7	114	0.0500	1228	26	7
MI_CP88-14	2.076	0.065	0.2045	0.0807	0.0285	0.0068	0.0861	0.0067	1143	18	1275	1289	16	1307	87	1060	1286	16	137	0.0550	1275	16	137
MI_CP88-15	2.48	0.073	0.2294	0.0852	0.0285	0.0068	0.0862	0.0078	1217	20	121	28	120	1320	140	1216	24	7	114	0.0500	1228	26	7
MI_CP88-16	2.788	0.074	0.2447	0.0861	0.0285	0.0068	0.0862	0.0081	1268	20	121	28	120	1320	140	1216	24	7	114	0.0500	1228	26	7
MI_CP88-17	1.7	0.044	0.1692	0.0703	0.0276	0.0066	0.0846	0.0059	1015	17	107	107	22	900	120	1019	21	134	77	0.1746	1007	22	77
MI_CP88-18	1.856	0.056	0.1874	0.0784	0.0279	0.0067	0.0851	0.0061	1079	18	107	107	22	900	120	1019	21	134	77	0.1746	1007	22	77
MI_CP88-19	1.151	0.029	0.1081	0.0486	0.0246	0.0046	0.0707	0.0044	140	16	139	5.1	1.41	29	40	160	18	22	1.2222	1034	5.1	1.41	
MI_CP88-20	1.151	0.029	0.1081	0.0486	0.0246	0.0046	0.0707	0.0044	140	16	139	5.1	1.41	29	40	160	18	22	1.2222	1034	5.1	1.41	
MI_CP88-21	1.151	0.029	0.1081	0.0486	0.0246	0.0046	0.0707	0.0044	140	16	139	5.1	1.41	29	40	160	18	22	1.2222	1034	5.1	1.41	
MI_CP88-22	1.42	0.044	0.1702	0.0403	0.0275	0.0058	0.0505	0.0042	1023	17	1013	22	100	1020	130	1027	16	239	133	0.4448	1013	22	133
MI_CP88-23	1.42	0.044	0.1702	0.0403	0.0275	0.0058	0.0505	0.0042	1023	17	1013	22	100	1020	130	1027	16	239	133	0.4448	1013	22	133
MI_CP88-24	2.38	0.089	0.2035	0.0848	0.0284	0.0064	0.0852	0.0064	1233	21	114	24	114	24	24	144	24	24	24	1.1111	1013	22	24
MI_CP88-25	2.38	0.089	0.2035	0.0848	0.0284	0.0064	0.0852	0.0064	1233	21	114	24	114	24	24	144	24	24	24	1.1111	1013	22	24
MI_CP88-26	2.026	0.084	0.2014	0.0801	0.0289	0.0063	0.0848	0.0065	1135	12	197	8.1	1.68	30	125	129	30	15	0.5000	1016	8.1	1.68	
MI_CP88-27	1.700	0.054	0.1622	0.0702	0.0251	0.0049	0.0697	0.0045	1062	19	126	26	172	20	150	138	20	200	102	0.5100	1226	26	200
MI_CP88-28	2.454	0.082	0.2026	0.0828	0.0289	0.0063	0.0848	0.0078	1253	19	126	26	172	20	150	138	20	200	102	0.5100	1226	26	200
MI_CP88-29	1.700	0.054	0.1622	0.0702	0.0251	0.0049	0.0697	0.0045	1062	19	126	26	172	20	150	138	20	200	102	0.5100	1226	26	200
MI_CP88-30	1.700	0.054	0.1622	0.0702	0.0251	0.0049	0.0697	0.0045	1062	19	126	26	172	20	150	138	20	200	102	0.5100	1226	26	200
MI_CP88-31	1.700	0.054	0.1622	0.0702	0.0251	0.0049	0.0697	0.0045	1062	19	126	26	172	20	150	138	20	200	102	0.5100	1226	26	200
MI_CP88-32	2.026	0.084	0.2014	0.0801	0.0289	0.0063	0.0848	0.0078	1253	19	126	26	172	20	150	138	20	200	102	0.5100	1226	26	200
MI_CP88-33	2.026	0.084	0.2014	0.0801	0.0289	0.0063	0.0848	0.0078	1253	19	126	26	172	20	150	138	20	200	102	0.5100	1226	26	200
MI_CP88-34	2.026	0.084	0.2014	0.0801	0.0289	0.0063	0.0848	0.0078	1253	19	126	26	172	20	150	138	20	200	102	0.5100	1226	26	200
MI_CP88-35	2.026	0.084	0.2014	0.0801	0.0289	0.0063	0.0848	0.0078	1253	19	126	26	172	20	150	138	20	200	102	0.5100	1226	26	200
MI_CP88-36	1.975	0.054	0.1556	0.0688	0.0253	0.0048	0.0675	0.0045	1104	19	111	26	110	120	130	102	110	18	22	1.2222	1215	26	110
MI_CP88-37	1.975	0.054	0.1556	0.0688	0.0253	0.0048	0.0675	0.0045	1104	19	111	26	110	120	130	102	110	18	22	1.2222	1215	26	110
MI_CP88-38	1.975	0.054	0.1556	0.0688	0.0253	0.0048	0.0675	0.0045	1104	19	111	26	110	120	130	102	110	18	22	1.2222	1215	26	110
MI_CP88-39	1.975	0.054	0.1556	0.0688	0.0253	0.0048	0.0675	0.0045	1104	19	111	26	110	120	130	102	110	18	22	1.2222	1215	26	110
MI_CP88-40	1.975	0.054	0.1556	0.0688	0.0253	0.0048	0.0675	0.0045	1104	19	111	26	110	120	130	102	110	18	22	1.2222	1215	26	110
MI_CP88-41	1.975	0.054	0.1556	0.0688	0.0253	0.0048	0.0675	0.0045	1104	19	111	26	110	120	130	102	110	18	22	1.2222	1215	26	110
MI_CP88-42	1.975	0.054	0.1556	0.0688	0.0253	0.0048	0.0675	0.0045	1104	19	111	26	110	120	130	102	110	18	22	1.2222	1215	26	110
MI_CP88-43	1.975	0.054	0.1556	0.0688	0.0253	0.0048	0.0675	0.0045	1104	19	111	26	110	120	130	102	110	18	22	1.2222	1215	26	110
MI_CP88-44	1.975	0.054	0.1556	0.0688	0.0253	0.0048	0.0675	0.0045	1104	19	111	26	110	120	130	102	110	18	22	1.2222	1215	26	110
MI_CP88-45	1.975	0.054	0.1556	0.0688	0.0253	0.0048	0.0675	0.0045	1104	19	111	26	110	120	130	102	110	18	22	1.2222	1215	26	110
MI_CP88-46	1.975	0.054	0.1556	0.0688	0.0253	0.0048	0.0675	0.0045	1104	19	111	26	110	120	130	102	110	18	22	1.2222	1215	26	110
MI_CP88-47	1.975	0.054	0.1556	0.0688	0.0253	0.0048	0.0675	0.0045	1104	19	111	26	110	120	130	102	110	18	22	1.2222	1215	26	110
MI_CP88-48	1.975	0.054	0.1556	0.0688	0.0253	0.0048	0.0675	0.0045	1104	19	111	26	110	120	130	102	110	18	22	1.2222	1215	26	110
MI_CP88-49	1.975	0.054	0.1556	0.0688	0.0253	0.0048	0.0675	0.0045	1104	19	111	26	110	120	130	102	110	18	22	1.2222	1215	26	110
MI_CP88-50	1.975	0.054	0.1556	0.0688	0.0253	0.0048	0.0675	0.0045	1104	19	111	26	110	120	130	102	110	18	22	1.2222	1215	26	110
MI_CP88-51	1.975	0.054	0.1556	0.0688	0.0253	0.0048	0.0675	0.0045	1104	19	111	26	110	120	130	102	110	18	22	1.2222	1215	26	110
MI_CP88-52	1.975	0.054	0.1556	0.0688	0.0253	0.0048	0.0675	0.0045	1104	19	111	26	110	120	130	102	110	18	22	1.2222	1215	26	110
MI_CP88-53	1.975	0.054	0.1556	0.0688	0.0253	0.0048	0.0675	0.0045	1104	19	111	26	110	120	130	102	110	18	22	1.2222	1215	26	110
MI_CP88-54	1.975	0.054	0.1556	0.0688	0.0253	0.0048	0.0675	0.0045	1104	19	111	26	110	120	130	102	110	18	22	1.2222	1215	26	110
MI_CP88-55	1.975	0.054	0.1556	0.0688	0.0253	0.0048	0.0675	0.0045	1104	19	111	26	110	120	130	102	110	18	22	1.2222	1215	26	110
MI_CP88-56	1.975	0.054	0.1556	0.0688	0.0253	0.0048	0.0675	0.0045	1104	19	111	26	110	120	130	102	110	18	22	1.2222	1215	26	110
MI_CP88-57	1.975	0.054	0.1556	0.0688	0.0253	0.0048	0.0675	0.0045	1104	19	111	26	110	120	130	102	110	18	22	1.2222	1215	26	110
MI_CP88-58	1.975	0.054	0.1556	0.0688	0.0253	0.0048																	

TABLE A2. ZIRCON AND APATITE FISSION TRACK DATA, FROM THE SEDIMENTARY ROCKS OF THE ARACATACA AND PALOMINO BASINS

Zircon Fission Track data																					
Sample	Lithology	Stratigraphic age	Elevation (m)	Latitude	Longitude	U (ppm)	2σ	Grains n	Ns	Ni	Counted squares	ζ estimate (Ma.cm ²)	se(ζ) std error on ζ estimate (Ma.cm ²)	P1	2σ se(t) (Frac)	P2	2σ se(t) (Frac)	P3	2σ se(t) (Frac)	P4	2σ se(t) (Frac)
CP-088	Coarse Sandstone	Miocene	180	10°31'33"	74°03'13"	216	8	96	20034	5105	4606	141.03	2.32	32.9 ± 15 (2%)	64.9 ± 5 (37.4%)	103.6 ± 6.9 (60.7%)					
CVI-1302	Conglomerate	Miocene	150	10°35'18"	74°05'06"	226	10	68	11865	2566	2249	141.03	2.32	29.2 ± 3 (6.4%)	52 ± 8 (18.9%)	103.8 ± 6.6 (74.7%)					
EMP-16	Conglomerate	Lower Miocene	176	10°32'21"	74°03'50"	180	8	98	13572	2457	3441	128.02	1.87		56.8 ± 11 (11.8%)	95.9 ± 10 (72.5%)	140.9 ± 29 (15.7%)				
EMP-49A	Conglomerate	Lower Miocene	55	11°12'33"	73°30'17"	180	8	97	17574	3314	3996	128.02	1.87		74.4 ± 7 (37.1%)	107.4 ± 14 (52.8%)	158.4 ± 52 (10.1%)				
AP-045	Conglomerate	Miocene	55	11°13'12"	73°28'39"	89	4	101	17531	2781	6886	128.02	1.87	41 ± 13.1 (1.4%)	89.6 ± 9.4 (38.6%)	131.6 ± 11.6 (60%)					
AP-046	Conglomerate	Miocene	65	11°13'7.99"	73°28'41.4"	97	5	158	13884	2172	4417	126.82	2.05	32.9 ± 6.3 (5.1%)	79.6 ± 15.2 (9%)	129.4 ± 12.1 (74.1%)	250.6 ± 68.7 (11.7%)				
AP-050	Conglomerate	Miocene	14	11°15'29.78"	73°36'47.4"	225	8	93	17166	5347	4696	126.82	2.05	38.9 ± 4.5 (14.6%)	54.8 ± 3.7 (47.3%)	94 ± 14.6 (20.1%)	154.8 ± 27.5 (18%)				
EMP-35b	Conglomerate	Miocene	0	11°15'34.31"	73°29'53.1"	129	6	99	15535	2288	3856	128.02	1.87		55.6 ± 14.5 (1.8%)	104.9 ± 7 (56.4)	170.5 ± 21 (41.7%)				
Detrital Apatite																					
Apatite Fission Track data																					
Sample	Lithology	Stratigraphic age	Elevation (m)	Latitude	Longitude	U (ppm)	2σ	Grains n	Ns	Ni	Counted squares	ζ estimate (Ma.cm ²)	se(ζ) std error on ζ estimate (Ma.cm ²)	P1	2σ se(t) (Frac)	P2	2σ se(t) (Frac)	P3	2σ se(t) (Frac)		
CVI-1302	Conglomerate	Miocene	150	10°35'18"	74°05'06"	41	1	35	1382	7662	3350	270.07	6.38	19.2 ± 4.2 (17.3%)	29.8 ± 4.8 (48.8%)	42 ± 5.7 (33.9%)					
EMP-16	Conglomerate	Lower Miocene	176	10°32'21"	74°03'50"	14	1	38	557	2169	3570	284.52	5.65		21.8 ± 4 (41%)	59.9 ± 9.3 (59%)					
AP-045	Conglomerate	Miocene	55	11°13'12"	73°28'39"	10	0	72	685	2369	5478	284.52	5.65		26.8 ± 6.5 (31.3%)	52.3 ± 6.3 (68.7%)					
EMP-49A	Conglomerate	Miocene	55	11°12'33"	73°30'17"	12	1	8	143	494	780	270.07	6.38		33.4 ± 10.6 (72.6%)	59.2 ± 36 (27.4%)					

Note: n = total number of grains counted; binomial peak-fit ages are given ± 2SE. The percentage of grains in a specific peak is also given.
 All samples were counted at 1250 dry (100 objective, 1.25 tube factor, 10 oculars). Apatite samples by A. Piraquive (CN-1) of 284.52 ± 5.65 (1SE). Samples CVI-1302, EMP-16, AP-045, EMP-49A
 Zircon Samples EMP-16, EMP-49A, AP-45, AP-46, AP-50 and EMP-35b were counted by E. Pinzón (CN-1 zeta 128.02 ± 1.87).
 Zircon Samples CP-088 and CVI-1302 were counted by A. Piraquive (CN-1 zeta 141.03 ± 2.32). Depositional ages after Tschanz (1969)
 Chi-squared test: values greater than 5% are considered to pass the test and represent a single age population.

Merged dataset:

C:\BH2\Edna\6-2015\AP-046\AP_046-2.FTZ

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NEW PARAMETERS - ZETA METHOD

EFFECTIVE TRACK DENSITY FOR FLUENCE MONITOR (tracks/cm²): 3,06E+05
 RELATIVE ERROR (%): 1,16
 EFFECTIVE URANIUM CONTENT OF MONITOR (ppm): 50,00
 ZETA FACTOR AND STANDARD ERROR (yr cm²): 126,82 2,05
 SIZE OF COUNTER SQUARE (cm²): 8,30E-07

GRAIN AGES IN ORIGINAL ORDER

Grain no.	RhoS (cm ⁻²)	(Ns)	RhoI (cm ⁻²)	(Ni)	Squares	U+/-2s	Grain Age (Ma)		
							Age	--95% CI--	
1	1,33E+06	(33)	1,20E+05	(3)	30	20 21	200.3	66.5	1004.1
2	8,66E+06	(115)	2,26E+05	(3)	16	37 40	670.3	243.2	2896.0
3	4,02E+06	(70)	2,98E+06	(52)	21	488 135	26.0	17.9	38.0
4	2,85E+06	(71)	3,61E+05	(9)	30	59 38	148.8	75.7	339.0
5	4,50E+06	(56)	4,82E+05	(6)	15	79 62	174.2	77.6	494.5
6	4,22E+06	(84)	1,05E+06	(21)	24	172 75	76.6	47.4	130.4
7	6,43E+06	(320)	1,06E+06	(53)	60	174 48	114.9	86.0	153.4
8	7,04E+06	(146)	9,64E+05	(20)	25	158 70	139.0	87.7	234.3
9	2,83E+06	(47)	6,02E+05	(10)	20	99 61	89.3	45.2	199.1
10	1,53E+06	(61)	2,01E+05	(8)	48	33 23	143.6	70.0	348.2
11	5,02E+06	(150)	4,35E+05	(13)	36	71 39	217.3	125.6	416.2
12	2,98E+06	(89)	3,68E+05	(11)	36	60 36	152.9	83.0	317.6
13	3,50E+06	(93)	4,14E+05	(11)	32	68 40	159.7	86.8	331.0
14	1,73E+06	(86)	3,82E+05	(19)	60	62 28	86.5	52.7	151.0
15	7,11E+06	(354)	2,11E+06	(105)	60	345 68	64.7	51.9	80.6
16	1,03E+07	(256)	1,65E+06	(41)	30	269 84	119.5	86.1	170.5
17	1,79E+06	(89)	1,61E+05	(8)	60	26 18	208.3	103.9	495.4
18	3,31E+06	(275)	3,25E+05	(27)	100	53 20	193.4	131.3	298.2
19	3,61E+06	(54)	1,00E+06	(15)	18	164 84	68.9	38.7	131.8
20	1,76E+06	(146)	1,69E+05	(14)	100	28 15	196.9	115.6	368.3
21	6,53E+06	(130)	7,53E+05	(15)	24	123 63	164.2	97.4	301.7
22	9,17E+06	(426)	6,45E+05	(30)	56	106 38	268.1	187.0	400.5
23	1,48E+06	(60)	1,97E+05	(8)	49	32 22	141.3	68.8	342.8
24	1,72E+06	(143)	1,93E+05	(16)	100	32 16	169.4	102.2	304.1
25	3,66E+06	(76)	7,23E+05	(15)	25	118 60	96.6	55.7	181.5
26	6,27E+06	(104)	1,02E+06	(17)	20	168 80	116.5	70.2	208.0
27	9,98E+05	(58)	1,20E+05	(7)	70	20 14	155.5	72.7	404.2
28	7,68E+06	(153)	1,36E+06	(27)	24	222 85	108.4	72.1	169.9
29	5,83E+06	(121)	1,01E+06	(21)	25	166 72	110.0	69.4	184.4
30	3,64E+06	(151)	4,82E+05	(20)	50	79 35	143.7	90.8	241.9
31	3,66E+06	(146)	6,02E+05	(24)	48	99 40	116.2	75.7	187.3
32	1,90E+06	(158)	1,93E+05	(16)	100	32 16	186.9	113.2	334.4
33	5,78E+06	(288)	8,43E+05	(42)	60	138 42	131.1	95.1	185.8
34	1,12E+07	(84)	2,81E+06	(21)	9	460 199	76.6	47.4	130.4
35	4,88E+06	(385)	5,83E+05	(46)	95	95 28	159.7	118.1	221.6
36	5,71E+06	(166)	4,82E+05	(14)	35	79 41	223.4	131.7	416.0
37	3,04E+06	(53)	4,02E+05	(7)	21	66 48	142.2	66.1	371.8
38	1,08E+06	(44)	1,23E+05	(5)	49	20 17	163.7	67.4	528.9
39	2,11E+06	(70)	4,82E+05	(16)	40	79 39	83.6	48.5	154.6
40	7,48E+06	(621)	1,19E+06	(99)	100	195 39	119.8	96.8	148.3
41	8,49E+06	(296)	1,20E+06	(42)	42	197 61	134.7	97.8	190.8
42	4,63E+06	(73)	5,07E+05	(8)	19	83 57	171.4	84.6	411.7
43	2,70E+06	(179)	3,92E+05	(26)	80	64 25	131.4	87.5	206.6
44	3,23E+06	(268)	6,63E+05	(55)	100	108 29	92.9	69.5	124.1
45	2,98E+06	(99)	4,52E+05	(15)	40	74 38	125.5	73.5	233.0
46	2,25E+06	(140)	1,77E+05	(11)	75	29 17	238.8	132.3	486.8
47	6,25E+06	(363)	3,96E+05	(23)	70	65 27	296.9	197.5	471.3
48	3,83E+06	(127)	3,61E+05	(12)	40	59 33	199.5	112.4	395.2
49	1,89E+06	(47)	5,22E+05	(13)	30	85 47	69.1	37.2	139.7
50	1,03E+06	(77)	1,61E+05	(12)	90	26 15	121.8	66.9	246.4
51	4,78E+06	(119)	1,61E+05	(4)	30	26 25	532.4	215.6	1856.3
52	3,49E+06	(290)	4,22E+05	(35)	100	69 23	158.0	111.8	231.0

Merged dataset:

C:\BH2\Edna\6-2015\AP-046\AP_046-2.FTZ

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Grain no.	RhoS (cm ⁻²)	(Ns)	RhoI (cm ⁻²)	(Ni)	Squares	U+/-2s		Grain Age (Ma)		
								Age	--95% CI--	
53	3,33E+06	(276)	4,94E+05	(41)	100	81	25	128.7	93.0	183.3
54	9,24E+05	(46)	1,41E+05	(7)	60	23	17	123.7	56.8	326.0
55	3,99E+06	(53)	6,78E+05	(9)	16	111	72	111.5	55.5	258.0
56	2,61E+06	(26)	2,01E+05	(2)	12	33	42	231.0	62.7	1885.8
57	5,18E+06	(86)	1,14E+06	(19)	20	187	85	86.5	52.7	151.0
58	2,51E+06	(75)	2,01E+05	(6)	36	33	26	232.1	105.3	647.6

Merged dataset:

C:\BH2\Edna\6-2015\AP-046\AP_046-2.FTZ

C:\BH2\Edna\6-2015\AP-046\AP_046-1.ftz

NEW PARAMETERS - ZETA METHOD

EFFECTIVE TRACK DENSITY FOR FLUENCE MONITOR (tracks/cm²): 3,06E+05
 RELATIVE ERROR (%): 1,17
 EFFECTIVE URANIUM CONTENT OF MONITOR (ppm): 50,00
 ZETA FACTOR AND STANDARD ERROR (yr cm²): 126,82 2,05
 SIZE OF COUNTER SQUARE (cm²): 8,30E-07

Grain no.	RhoS (cm ⁻²)	(Ns)	RhoI (cm ⁻²)	(Ni)	Squares	U+/-2s	Grain Age (Ma)		
							Age	--95% CI--	
59	6,02E+06	(250)	8,19E+05	(34)	50	134 46	140.5	98.6	207.4
60	4,06E+06	(101)	3,61E+05	(9)	30	59 38	210.7	109.3	472.4
61	2,53E+06	(147)	3,96E+05	(23)	70	65 27	122.1	79.0	198.8
62	6,69E+06	(50)	9,37E+05	(7)	9	153 112	134.5	62.2	352.6
63	1,97E+06	(80)	1,72E+05	(7)	49	28 21	213.6	101.9	545.7
64	6,02E+06	(100)	1,45E+06	(24)	20	236 96	79.9	51.1	130.8
65	6,61E+06	(351)	1,45E+06	(77)	64	237 54	87.3	68.2	111.8
66	9,29E+05	(27)	2,41E+05	(7)	35	39 29	73.2	31.7	200.5
67	4,76E+06	(79)	7,83E+05	(13)	20	128 70	115.6	64.8	227.2
68	3,27E+06	(114)	5,45E+05	(19)	42	89 40	114.6	70.8	197.5
69	3,10E+06	(108)	5,45E+05	(19)	42	89 40	108.6	66.9	187.6
70	3,07E+06	(252)	4,75E+05	(39)	99	78 25	123.7	88.5	178.2
71	4,70E+06	(195)	7,23E+05	(30)	50	118 43	124.3	84.9	189.3
72	1,14E+06	(57)	2,41E+05	(12)	60	39 22	90.5	48.7	186.1
73	1,08E+06	(27)	2,81E+05	(7)	30	46 34	73.2	31.7	200.5
74	6,27E+06	(156)	1,04E+06	(26)	30	171 66	114.8	76.0	181.3
75	7,92E+06	(322)	1,20E+06	(49)	49	197 56	125.9	93.4	173.7
76	5,16E+06	(60)	1,12E+06	(13)	14	183 100	88.1	48.4	175.5
77	1,45E+06	(42)	7,23E+05	(21)	35	118 51	38.5	22.4	68.7
78	6,20E+06	(72)	1,46E+06	(17)	14	239 114	81.1	47.7	147.1
79	1,42E+06	(59)	1,69E+05	(7)	50	28 20	158.3	74.1	411.2
80	4,31E+06	(143)	3,92E+05	(13)	40	64 35	207.6	119.7	398.3
81	3,15E+06	(94)	3,68E+05	(11)	36	60 36	161.6	87.9	334.8
82	1,05E+06	(21)	4,02E+05	(8)	24	66 45	50.1	21.6	131.7
83	5,16E+06	(210)	8,85E+05	(36)	49	145 48	111.8	78.6	164.0
84	2,28E+06	(34)	2,68E+05	(4)	18	44 41	157.4	58.5	610.2
85	1,41E+06	(35)	3,61E+05	(9)	30	59 38	74.0	35.4	176.1
86	3,51E+06	(70)	7,03E+05	(14)	24	115 60	95.4	53.9	184.0
87	3,21E+06	(24)	1,34E+05	(1)	9	22 36	396.5	75.5	8889.8
88	7,03E+06	(70)	1,10E+06	(11)	12	180 107	120.8	64.6	253.7
89	3,66E+06	(73)	6,02E+05	(12)	24	98 56	115.7	63.3	234.6
90	3,60E+06	(209)	4,30E+05	(25)	70	70 28	159.3	105.9	251.6
91	7,23E+06	(108)	1,27E+06	(19)	18	208 94	108.6	66.9	187.6
92	8,78E+05	(51)	4,99E+05	(29)	70	82 30	34.0	21.2	55.6
93	6,57E+06	(229)	8,89E+05	(31)	42	145 52	141.1	97.4	212.5
94	2,74E+06	(91)	1,57E+06	(52)	40	256 71	33.8	23.8	48.6
95	1,78E+06	(59)	9,34E+05	(31)	40	153 54	36.7	23.5	58.8
96	4,95E+06	(411)	6,75E+05	(56)	100	110 29	139.6	105.7	184.3
97	7,88E+06	(157)	6,53E+05	(13)	24	107 58	227.6	131.7	435.1
98	4,47E+06	(89)	9,54E+05	(19)	24	156 71	89.6	54.7	156.2
99	1,87E+06	(31)	2,41E+05	(4)	20	39 37	143.7	52.9	561.3
100	7,11E+06	(354)	1,20E+06	(60)	60	197 51	112.6	85.6	147.8
POOLED	3,79E+06	(13884)	5,92E+05	(2172)	4417	97 5	122.7	115.6	130.2

CHI² PROBABILITY (%): 0.0

POOLED AGE W/ 68% CONF. INTERVAL(Ma): 122.7, 119.0 -- 126.4 (-3.6 +3.8)
 95% CONF. INTERVAL(Ma): 115.6 -- 130.2 (-7.0 +7.5)

CENTRAL AGE W/ 68% CONF. INTERVAL(Ma): 117.6, 111.0 -- 124.5 (-6.6 +6.9)
 95% CONF. INTERVAL(Ma): 105.1 -- 131.6 (-12.5 +14.0)
 AGE DISPERSION (%): 47.2

Merged dataset:

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FIT OPTION: Best-fit peaks using the binomial model of Galbraith and Green

INITIAL GUESS FOR MODEL PARAMETERS (number of peaks to fit = 4)

Peak #.	Peak Age	Theta	Fraction(%)	Count
1.	64.80	0.771	9.6	9.65
2.	86.80	0.819	20.8	20.82
3.	122.60	0.865	41.5	41.48
4.	202.50	0.914	16.2	16.21

Total range for grain ages: 26,0 to 609,7 Ma
 Number of active grains (Num. used for fit): 100
 Number of removed grains: 0
 Degrees of freedom for fit: 93
 Average of the SE(Z)'s for the grains: 0,31
 Estimated width of peaks in PD plot in Z units: 0,36

PARAMETERS FOR BEST-FIT PEAKS

- * Standard error for peak age includes group error
- * Peak width is for PD plot assuming a kernel factor = 0.60

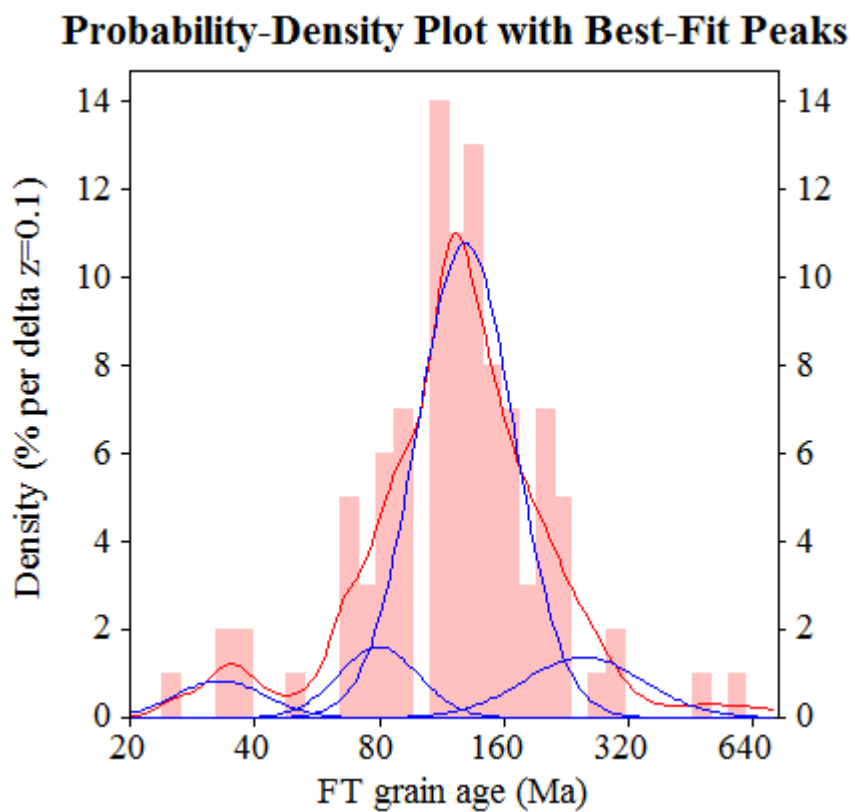
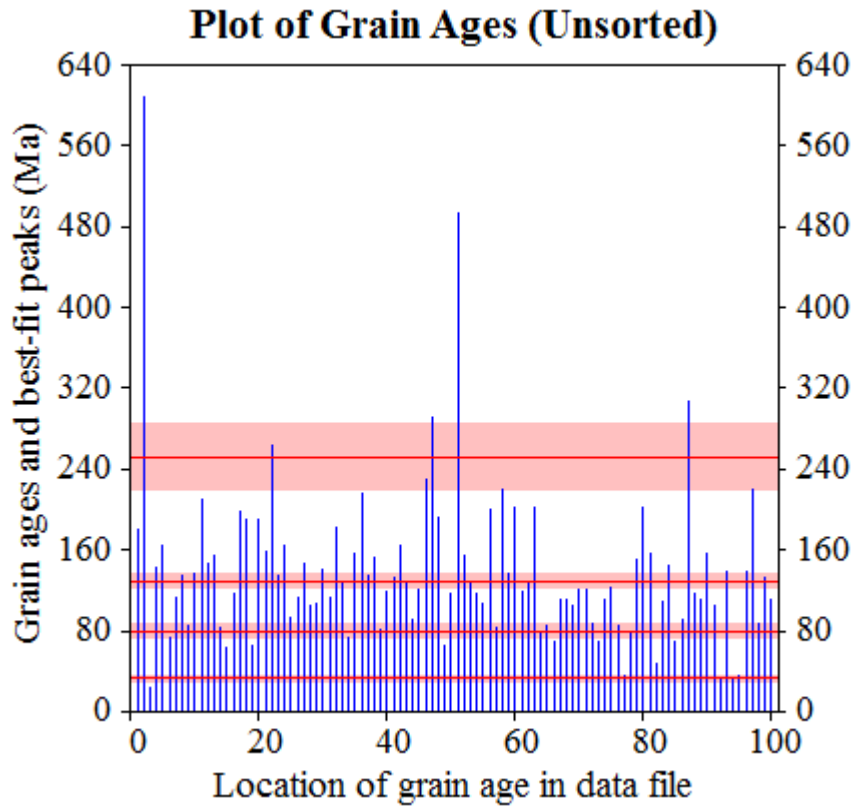
#.	Peak Age(Ma)	68%CI	95%CI	W(Z)	Frac(%)	SE,%	Count
1.	32.9	-3,0 ...+3,3	-5,7 ...+6,9	0.25	5.1	2.3	5.1
2.	79.6	-7,3 ...+8,1	-13,7 ...+16,5	0.22	9.0	6.1	9.0
3.	129.4	-6,0 ...+6,3	-11,6 ...+12,7	0.27	74.1	8.0	74.1
4.	250.6	-30,8 ...+35,1	-56,9 ...+73,3	0.34	11.7	6.2	11.7

Log-likelihood for best fit: -324,328
 Chi-squared value for best fit: 93,580
 Reduced chi-squared value: 1,006
 Probability for F test: 4%
 Condition number for COVAR matrix: 51,79
 Number of iterations: 17

Merged dataset:

C:\BH2\Edna\6-2015\AP-046\AP_046-2.FTZ

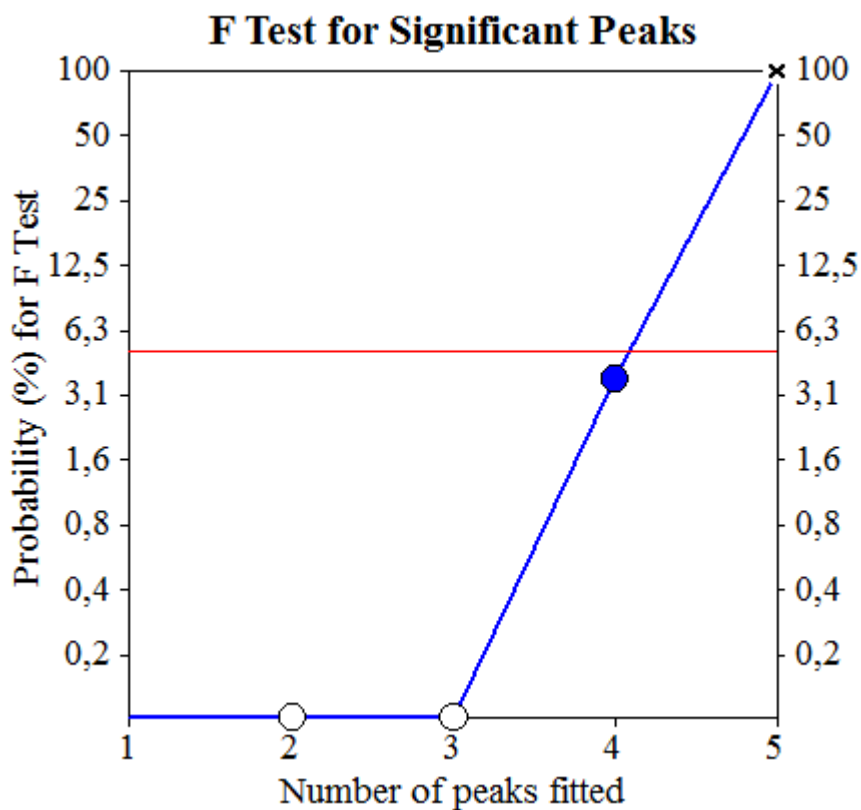
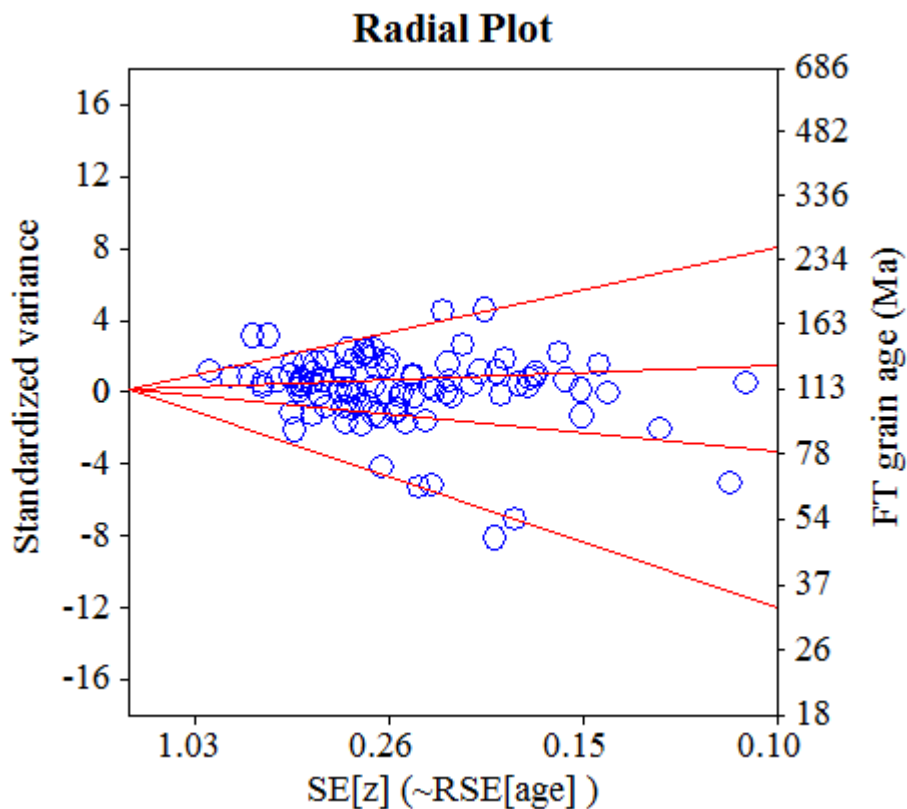
C:\BH2\Edna\6-2015\AP-046\AP_046-1.ftz



Merged dataset:

C:\BH2\Edna\6-2015\AP-046\AP_046-2.FTZ

C:\BH2\Edna\6-2015\AP-046\AP_046-1.ftz



Merged dataset:

C:\BH2\Edna\6-2015\AP-050\AP_050-2A1FTZ

C:\BH2\Edna\6-2015\AP-050\AP_050-1.ftz

NEW PARAMETERS - ZETA METHOD

EFFECTIVE TRACK DENSITY FOR FLUENCE MONITOR (tracks/cm²): 3,04E+05
 RELATIVE ERROR (%): 1,14
 EFFECTIVE URANIUM CONTENT OF MONITOR (ppm): 50,00
 ZETA FACTOR AND STANDARD ERROR (yr cm²): 126,82 2,05
 SIZE OF COUNTER SQUARE (cm²): 8,30E-07

GRAIN AGES IN ORIGINAL ORDER

Grain no.	RhoS (cm ⁻²)	(Ns)	RhoI (cm ⁻²)	(Ni)	Squares	U+/-2s	Grain Age (Ma)		
							Age	--95% CI--	
1	4,13E+06	(144)	2,01E+05	(7)	42	33 24	377.2	185.1	934.5
2	3,84E+06	(255)	1,93E+06	(128)	80	317 56	38.2	30.8	47.3
3	9,94E+06	(165)	2,23E+06	(37)	20	366 120	85.2	59.6	125.4
4	4,53E+06	(263)	1,45E+06	(84)	70	237 52	59.8	46.7	76.5
5	6,66E+06	(221)	1,08E+06	(36)	40	178 59	117.0	82.4	171.4
6	6,49E+06	(264)	6,64E+05	(27)	49	109 42	185.0	125.5	285.6
7	2,14E+06	(71)	2,71E+05	(9)	40	45 29	148.2	75.4	337.7
8	5,12E+06	(153)	1,41E+06	(42)	36	231 71	69.7	49.5	100.7
9	5,66E+06	(94)	3,01E+05	(5)	20	49 42	342.8	148.3	1053.0
10	4,82E+06	(120)	1,85E+06	(46)	30	303 89	50.1	35.4	72.0
11	1,70E+06	(113)	2,11E+05	(14)	80	35 18	152.4	88.5	287.7
12	5,71E+06	(166)	7,23E+05	(21)	35	119 51	149.8	95.8	248.3
13	4,72E+06	(47)	1,00E+06	(10)	12	165 102	89.0	45.1	198.4
14	4,80E+06	(319)	5,87E+05	(39)	80	96 31	155.4	111.9	222.5
15	3,17E+06	(158)	1,51E+06	(75)	60	247 57	40.3	30.6	53.1
16	6,55E+06	(136)	1,30E+06	(27)	25	214 82	96.0	63.6	151.2
17	2,86E+06	(190)	5,57E+05	(37)	80	92 30	98.0	69.0	143.6
18	3,41E+06	(198)	1,70E+06	(99)	70	280 56	38.3	30.0	48.8
19	5,69E+06	(85)	1,54E+06	(23)	18	253 104	70.6	44.4	117.5
20	4,39E+06	(91)	1,54E+06	(32)	25	253 89	54.5	36.2	84.4
21	4,42E+06	(33)	6,69E+05	(5)	9	110 94	122.8	49.2	405.1
22	3,98E+06	(132)	1,69E+06	(56)	40	277 74	45.3	32.9	63.1
23	4,60E+06	(382)	1,57E+06	(130)	100	257 45	56.3	46.0	68.8
24	8,43E+06	(84)	1,71E+06	(17)	12	280 134	94.0	55.9	169.2
25	2,33E+06	(31)	7,53E+05	(10)	16	124 76	58.9	28.5	135.4
26	5,01E+06	(208)	1,37E+06	(57)	50	226 60	69.9	52.1	95.5
27	4,46E+06	(148)	1,96E+06	(65)	40	322 80	43.5	32.5	58.2
28	7,99E+06	(232)	1,20E+06	(35)	35	198 67	126.2	88.7	185.6
29	4,81E+06	(399)	9,76E+05	(81)	100	160 36	93.8	73.8	119.2
30	6,44E+06	(171)	7,91E+05	(21)	32	130 56	154.2	98.8	255.5
31	5,20E+06	(108)	1,73E+06	(36)	25	285 95	57.5	39.2	86.4
32	8,19E+06	(408)	2,53E+06	(126)	60	416 75	62.0	50.6	75.8
33	3,41E+06	(198)	6,88E+05	(40)	70	113 36	94.5	67.3	136.5
34	4,99E+06	(145)	5,51E+05	(16)	35	90 45	171.1	103.3	306.9
35	5,16E+06	(214)	1,13E+06	(47)	50	186 54	87.1	63.5	122.2
36	3,04E+06	(252)	1,00E+06	(83)	100	164 36	58.0	45.2	74.4
37	4,55E+06	(378)	1,49E+06	(124)	100	245 44	58.3	47.5	71.6
38	5,62E+06	(112)	1,46E+06	(29)	24	239 88	73.8	49.0	115.4
39	1,87E+06	(124)	3,31E+05	(22)	80	54 23	107.2	68.3	177.5
40	2,54E+06	(76)	6,36E+05	(19)	36	104 47	76.3	46.0	133.9
41	8,73E+06	(87)	1,51E+06	(15)	12	247 126	110.0	63.9	205.4
42	5,09E+06	(296)	2,08E+06	(121)	70	342 63	46.9	37.8	58.0
43	5,31E+06	(141)	2,90E+06	(77)	32	476 109	35.0	26.5	46.3
44	4,61E+06	(268)	1,03E+06	(60)	70	170 44	85.0	64.2	112.4
45	4,50E+06	(56)	1,61E+06	(20)	15	264 117	53.5	31.8	94.4
46	5,50E+06	(219)	8,53E+05	(34)	48	140 48	122.7	85.7	181.7
47	5,83E+06	(121)	9,64E+05	(20)	25	158 70	115.0	71.9	195.1
48	6,02E+06	(80)	2,33E+06	(31)	16	383 137	49.5	32.4	77.6
49	4,75E+06	(197)	2,55E+06	(106)	50	420 82	35.6	28.1	45.2
50	6,78E+06	(563)	2,22E+06	(184)	100	364 54	58.6	49.5	69.5
51	3,54E+06	(294)	1,30E+06	(108)	100	214 41	52.1	41.7	65.1
52	5,83E+06	(237)	2,26E+06	(92)	49	372 78	49.3	38.7	62.8

Merged dataset:

C:\BH2\Edna\6-2015\AP-050\AP_050-2A.FTZ

C:\BH2\Edna\6-2015\AP-050\AP_050-1.ftz

Grain no.	RhoS (cm ⁻²)	(Ns)	RhoI (cm ⁻²)	(Ni)	Squares	U+/-2s		Grain Age (Ma)		
								Age	--95% CI--	
53	4,75E+06	(197)	1,54E+06	(64)	50	253	63	58.7	44.2	77.8
54	4,67E+06	(388)	1,42E+06	(118)	100	234	43	62.9	51.1	77.5
55	4,55E+06	(151)	1,75E+06	(58)	40	287	75	50.0	36.8	68.9
56	5,42E+06	(135)	1,73E+06	(43)	30	284	86	60.2	42.5	87.0
57	4,92E+06	(204)	1,88E+06	(78)	50	309	70	50.0	38.5	64.9
58	4,71E+06	(313)	1,70E+06	(113)	80	280	53	53.0	42.7	65.9
59	3,52E+06	(292)	1,61E+06	(134)	100	265	46	41.8	34.0	51.4
60	3,10E+06	(90)	1,10E+06	(32)	35	181	64	53.9	35.8	83.5
61	1,19E+07	(158)	1,43E+06	(19)	16	235	107	157.4	98.6	268.2
62	1,05E+07	(174)	4,28E+06	(71)	20	703	167	46.8	35.5	61.7
63	5,07E+06	(202)	2,11E+06	(84)	48	346	76	46.0	35.6	59.4
64	5,64E+06	(117)	7,23E+05	(15)	25	119	60	147.4	87.0	271.9
65	6,02E+06	(300)	2,41E+06	(120)	60	396	73	47.9	38.7	59.3
66	4,54E+06	(226)	1,65E+06	(82)	60	270	60	52.7	40.9	67.9
67	4,42E+06	(55)	8,03E+05	(10)	15	132	81	103.9	53.4	229.6
68	4,72E+06	(274)	1,22E+06	(71)	70	201	48	73.6	56.6	95.5
69	2,65E+06	(77)	1,96E+06	(57)	35	322	85	26.0	18.2	37.3
70	4,08E+06	(339)	2,14E+06	(178)	100	352	53	36.6	30.4	44.0
71	3,95E+06	(164)	2,19E+06	(91)	50	360	76	34.5	26.7	44.7
72	4,82E+06	(240)	1,79E+06	(89)	60	294	62	51.6	40.4	65.9
73	5,06E+06	(126)	7,23E+05	(18)	30	119	55	132.7	81.5	231.3
74	2,41E+06	(40)	2,41E+05	(4)	20	40	37	183.7	69.4	703.2
75	4,73E+06	(110)	4,73E+05	(11)	28	78	46	187.7	102.9	386.3
76	4,04E+06	(67)	2,17E+06	(36)	20	356	118	35.8	23.6	55.3
77	4,30E+06	(125)	7,23E+05	(21)	35	119	51	113.2	71.6	189.4

Merged dataset:

C:\BH2\Edna\6-2015\AP-050\AP_050-2A.FTZ

C:\BH2\Edna\6-2015\AP-050\AP_050-1.ftz

NEW PARAMETERS - ZETA METHOD

EFFECTIVE TRACK DENSITY FOR FLUENCE MONITOR (tracks/cm²): 3,05E+05
 RELATIVE ERROR (%): 1,14
 EFFECTIVE URANIUM CONTENT OF MONITOR (ppm): 50,00
 ZETA FACTOR AND STANDARD ERROR (yr cm²): 126,82 2,05
 SIZE OF COUNTER SQUARE (cm²): 8,30E-07

Grain no.	RhoS (cm ⁻²)	(Ns)	RhoI (cm ⁻²)	(Ni)	Squares	U+/-2s	Grain Age (Ma)		
							Age	--95% CI--	
78	5,27E+06	(328)	1,70E+06	(106)	75	279 55	59.3	47.5	73.9
79	3,57E+06	(296)	1,19E+06	(99)	100	196 39	57.3	45.5	72.0
80	3,19E+06	(238)	6,56E+05	(49)	90	108 31	93.0	68.4	129.2
81	3,23E+06	(268)	1,36E+06	(113)	100	223 42	45.5	36.4	56.8
82	3,60E+06	(239)	1,17E+06	(78)	80	193 44	58.6	45.3	75.7
83	4,56E+06	(159)	2,18E+06	(76)	42	358 82	40.1	30.4	52.7
84	3,98E+06	(330)	1,51E+06	(125)	100	247 44	50.6	41.1	62.3
85	3,96E+06	(263)	1,45E+06	(96)	80	237 49	52.5	41.5	66.4
86	1,39E+06	(115)	3,37E+05	(28)	100	55 21	78.6	51.9	123.6
87	5,46E+06	(136)	2,17E+06	(54)	30	356 97	48.4	35.1	67.7
88	2,74E+06	(91)	6,93E+05	(23)	40	114 47	75.6	47.8	125.5
89	6,58E+06	(131)	2,61E+06	(52)	24	428 119	48.4	34.9	68.2
90	2,60E+06	(54)	1,20E+06	(25)	25	198 78	41.5	25.5	69.7
91	1,36E+06	(54)	4,52E+05	(18)	48	74 35	57.4	33.4	104.3
92	3,92E+06	(130)	1,54E+06	(51)	40	252 70	49.0	35.3	69.2
93	4,63E+06	(123)	2,00E+06	(53)	32	327 90	44.6	32.1	62.9
POOLED	4,40E+06	(17166)	1,37E+06	(5347)	4696	225 8	61.7	58.7	64.8

CHI^2 PROBABILITY (%): 0.0

POOLED AGE W/ 68% CONF. INTERVAL(Ma): 61.7, 60.1 -- 63.2 (-1.5 +1.6)
 95% CONF. INTERVAL(Ma): 58.7 -- 64.8 (-3.0 +3.1)

CENTRAL AGE W/ 68% CONF. INTERVAL(Ma): 65.6, 62.4 -- 69.0 (-3.2 +3.4)
 95% CONF. INTERVAL(Ma): 59.5 -- 72.4 (-6.2 +6.8)
 AGE DISPERSION (%): 41.2

Merged dataset:

C:\BH2\Edna\6-2015\AP-050\AP_050-2A.FTZ

C:\BH2\Edna\6-2015\AP-050\AP_050-1.ftz

FIT OPTION: Best-fit peaks using the binomial model of Galbraith and Green

INITIAL GUESS FOR MODEL PARAMETERS (number of peaks to fit = 4)

Peak #.	Peak Age	Theta	Fraction(%)	Count
1.	36.60	0.655	17.4	16.17
2.	48.90	0.718	41.5	38.57
3.	61.70	0.762	32.4	30.10
4.	361.60	0.951	1.3	1.21

Total range for grain ages: 26,0 to 361,6 Ma
 Number of active grains (Num. used for fit): 93
 Number of removed grains: 0
 Degrees of freedom for fit: 86
 Average of the SE(Z)'s for the grains: 0,21
 Estimated width of peaks in PD plot in Z units: 0,24

PARAMETERS FOR BEST-FIT PEAKS

- * Standard error for peak age includes group error
- * Peak width is for PD plot assuming a kernel factor = 0.60

#.	Peak Age(Ma)	68%CI	95%CI	W(Z)	Frac(%)	SE,%	Count
1.	38.9	-2,2 ...+2,3	-4,2 ...+4,7	0.15	14.6	5.5	13.6
2.	54.8	-1,8 ...+1,9	-3,5 ...+3,8	0.16	47.3	6.9	44.0
3.	94.0	-7,2 ...+7,7	-13,5 ...+15,8	0.22	20.1	6.4	18.7
4.	154.8	-13,5 ...+14,7	-25,3 ...+30,1	0.29	18.0	6.2	16.7

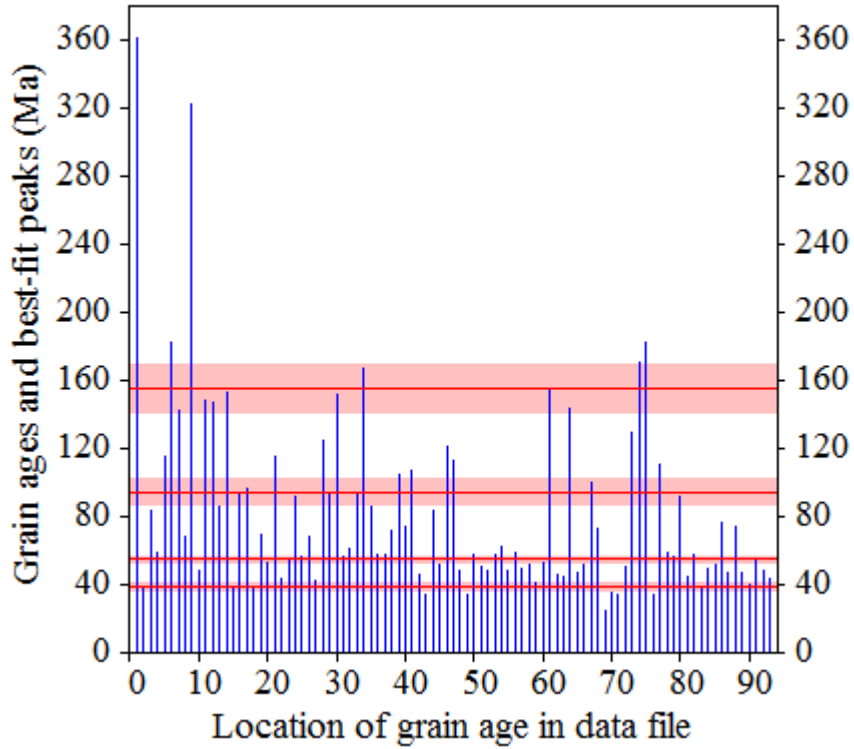
Log-likelihood for best fit: -380,676
 Chi-squared value for best fit: 93,978
 Reduced chi-squared value: 1,093
 Probability for F test: 0%
 Condition number for COVAR matrix: 20,75
 Number of iterations: 14

Merged dataset:

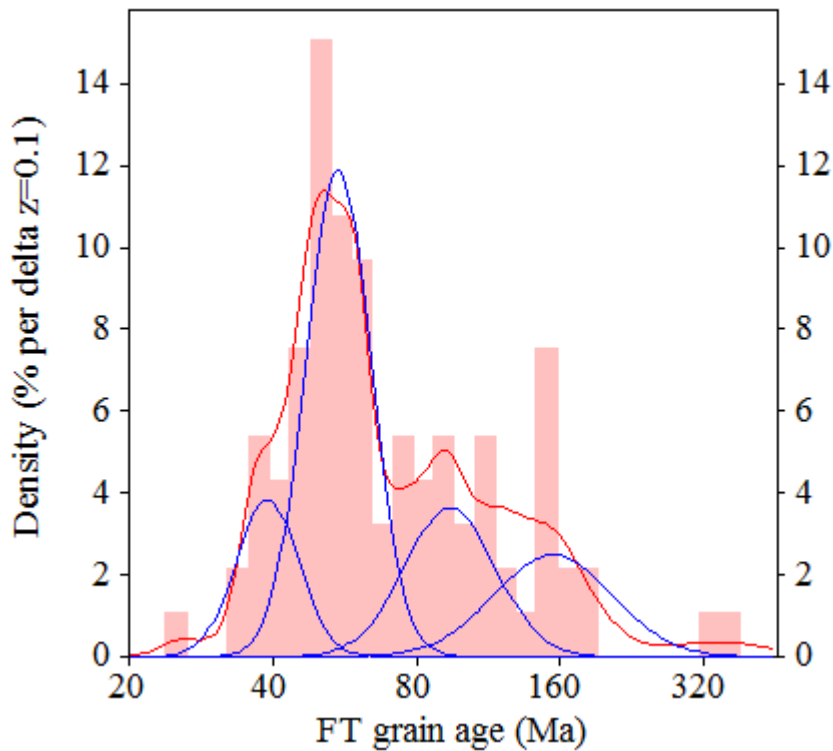
C:\BH2\Edna\6-2015\AP-050\AP_050-2A.FTZ

C:\BH2\Edna\6-2015\AP-050\AP_050-1.ftz

Plot of Grain Ages (Unsorted)



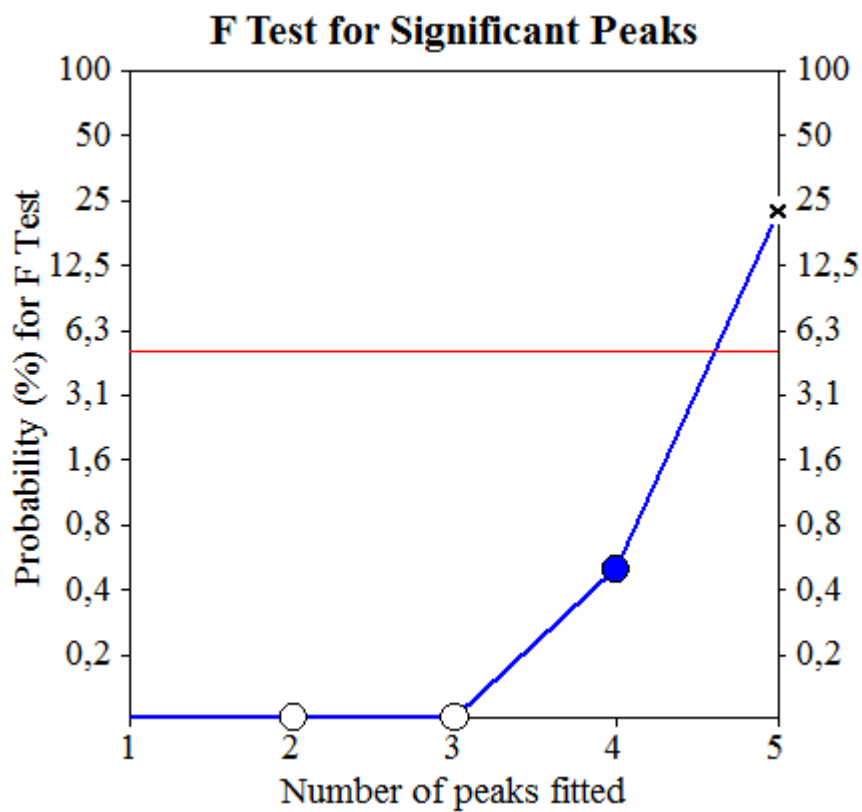
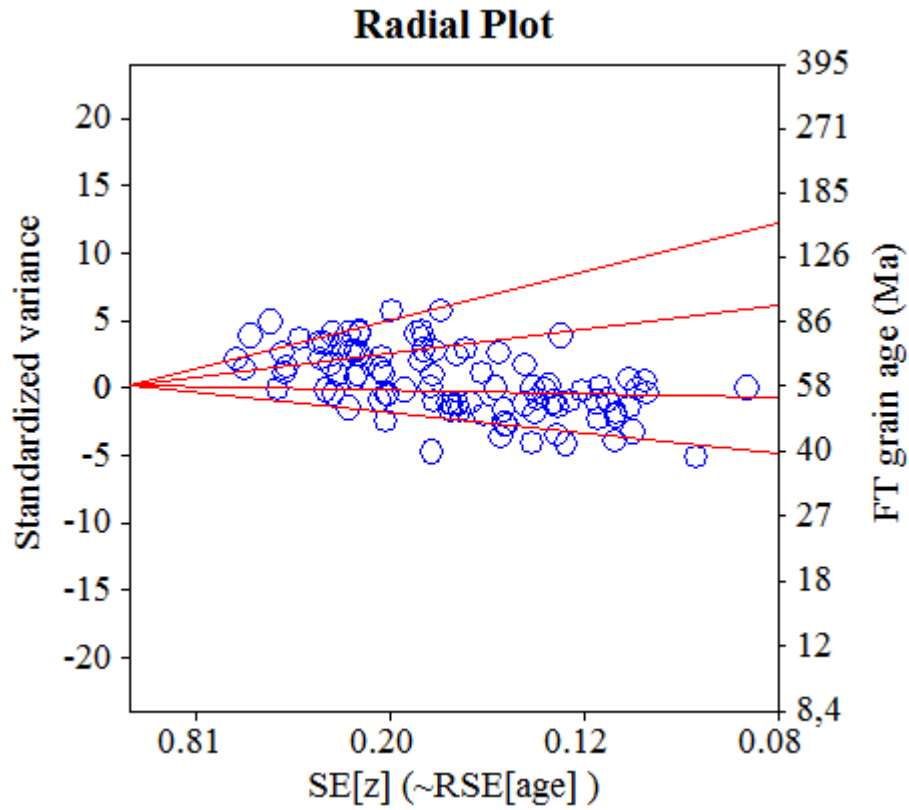
Probability-Density Plot with Best-Fit Peaks



Merged dataset:

C:\BH2\Edna\6-2015\AP-050\AP_050-2A.FTZ

C:\BH2\Edna\6-2015\AP-050\AP_050-1.ftz



NEW PARAMETERS - ZETA METHOD

EFFECTIVE TRACK DENSITY FOR FLUENCE MONITOR (tracks/cm²): 2,74E+05
 RELATIVE ERROR (%): 1,20
 EFFECTIVE URANIUM CONTENT OF MONITOR (ppm): 50,00
 ZETA FACTOR AND STANDARD ERROR (yr cm²): 128,02 1,87
 SIZE OF COUNTER SQUARE (cm²): 8,30E-07

GRAIN AGES IN ORIGINAL ORDER

Grain no.	RhoS (cm ⁻²)	(Ns)	RhoI (cm ⁻²)	(Ni)	Squares	U+/-2s	Grain Age (Ma)		
							Age	--95% CI--	
1	7,95E+05	(66)	1,08E+05	(9)	100	20 13	125.2	63.3	286.7
2	3,71E+06	(148)	5,77E+05	(23)	48	106 44	111.0	71.8	180.8
3	3,49E+06	(87)	3,21E+05	(8)	30	59 40	184.3	91.8	439.7
4	2,09E+06	(97)	3,44E+05	(16)	56	63 31	104.4	61.8	190.3
5	2,23E+06	(185)	3,01E+05	(25)	100	55 22	127.5	84.4	202.3
6	4,43E+06	(92)	4,34E+05	(9)	25	79 51	173.7	89.6	391.7
7	3,22E+06	(267)	3,86E+05	(32)	100	70 25	143.8	100.1	214.5
8	1,36E+06	(113)	2,29E+05	(19)	100	42 19	102.5	63.3	177.0
9	9,29E+05	(27)	2,07E+05	(6)	35	38 30	76.7	31.8	229.3
10	1,31E+06	(109)	1,45E+05	(12)	100	26 15	155.2	86.8	309.8
11	4,19E+06	(87)	9,64E+05	(20)	25	176 78	75.2	46.2	129.5
12	4,60E+06	(191)	5,30E+05	(22)	50	97 41	149.2	96.7	243.8
13	5,07E+06	(202)	1,15E+06	(46)	48	211 62	76.2	55.3	107.5
14	2,75E+06	(137)	5,02E+05	(25)	60	92 36	94.7	62.0	151.7
15	2,80E+06	(232)	4,22E+05	(35)	100	77 26	114.6	80.5	168.6
16	1,05E+06	(87)	8,43E+04	(7)	100	15 11	209.6	100.4	534.6
17	1,13E+06	(94)	1,45E+05	(12)	100	26 15	134.1	74.4	269.3
18	2,24E+06	(67)	2,68E+05	(8)	36	49 34	142.4	69.9	344.2
19	3,81E+06	(316)	5,06E+05	(42)	100	93 28	129.9	94.5	183.9
20	2,97E+06	(74)	2,81E+05	(7)	30	51 38	178.7	84.9	459.6
21	5,22E+06	(52)	9,04E+05	(9)	12	165 107	98.9	49.2	229.4
22	2,77E+06	(92)	3,61E+05	(12)	40	66 37	131.3	72.8	263.8
23	4,18E+06	(104)	7,23E+05	(18)	30	132 62	99.6	60.6	175.0
24	2,01E+06	(100)	1,81E+05	(9)	60	33 21	188.6	97.7	423.7
25	2,74E+06	(205)	3,21E+05	(24)	90	59 24	146.9	96.9	234.6
26	3,06E+06	(203)	7,08E+05	(47)	80	129 38	75.0	54.5	105.4
27	6,36E+06	(132)	1,06E+06	(22)	25	194 82	103.5	66.2	171.1
28	3,94E+06	(327)	6,14E+05	(51)	100	112 31	111.0	82.7	152.2
29	4,56E+06	(227)	5,02E+05	(25)	60	92 36	156.1	104.0	246.2
30	2,18E+06	(152)	5,88E+05	(41)	84	108 33	64.4	45.5	93.4
31	9,11E+06	(121)	1,66E+06	(22)	16	303 128	95.0	60.5	157.5
32	3,92E+06	(130)	2,41E+05	(8)	40	44 30	273.2	138.6	640.2
33	3,78E+06	(314)	4,94E+05	(41)	100	90 28	132.2	95.8	187.9
34	1,83E+06	(152)	1,93E+05	(16)	100	35 17	162.7	98.3	291.8
35	9,64E+05	(80)	9,64E+04	(8)	100	18 12	169.6	84.1	406.5
36	2,37E+06	(197)	2,05E+05	(17)	100	37 18	198.0	122.3	346.0
37	4,47E+06	(371)	5,54E+05	(46)	100	101 30	139.2	102.8	193.5
38	2,09E+06	(111)	4,89E+05	(26)	64	89 35	74.0	48.2	118.4
39	1,98E+06	(164)	2,41E+05	(20)	100	44 19	141.0	89.3	236.9
40	2,68E+06	(80)	4,69E+05	(14)	36	86 45	98.3	56.0	188.5
41	7,63E+06	(95)	1,12E+06	(14)	15	206 108	116.6	67.0	221.8
42	3,86E+06	(96)	7,63E+05	(19)	30	139 63	87.2	53.4	151.6
43	3,11E+06	(258)	6,75E+05	(56)	100	123 33	79.5	59.5	106.0
44	6,21E+05	(33)	5,65E+04	(3)	64	10 11	181.2	60.1	913.6
45	1,55E+06	(129)	2,89E+05	(24)	100	53 21	92.9	60.2	150.5
46	1,99E+06	(165)	2,89E+05	(24)	100	53 21	118.6	77.6	190.5
47	2,60E+06	(216)	5,42E+05	(45)	100	99 29	83.2	60.4	117.6
48	3,29E+06	(134)	2,21E+05	(9)	49	40 26	251.3	131.8	557.8
49	1,80E+06	(149)	3,98E+05	(33)	100	73 25	78.3	53.6	118.0
50	3,52E+06	(146)	7,95E+05	(33)	50	145 50	76.7	52.5	115.8
51	7,57E+06	(157)	1,35E+06	(28)	25	247 93	97.0	65.0	150.8
52	4,24E+06	(88)	1,06E+06	(22)	25	194 82	69.3	43.3	116.4
53	5,28E+06	(438)	1,01E+06	(84)	100	185 40	90.1	71.3	113.9

Grain no.	RhoS (cm ⁻²)	(Ns)	RhoI (cm ⁻²)	(Ni)	Squares	U+/-2s	Grain Age (Ma)		
							Age	--95% CI--	
54	6,84E+06	(142)	1,20E+06	(25)	25	220 87	98.1	64.3	157.0
55	2,43E+06	(202)	3,98E+05	(33)	100	73 25	105.8	73.4	158.1
56	1,95E+06	(162)	3,01E+05	(25)	100	55 22	111.8	73.7	178.1
57	2,25E+06	(56)	2,41E+05	(6)	30	44 35	157.6	70.1	448.3
58	4,26E+06	(99)	6,02E+05	(14)	28	110 58	121.4	70.0	230.6
59	4,59E+06	(137)	1,94E+06	(58)	36	355 93	41.2	30.1	57.0
60	3,94E+06	(229)	9,98E+05	(58)	70	183 48	68.2	51.1	90.9
61	3,42E+06	(284)	4,46E+05	(37)	100	81 27	132.5	94.4	192.0
62	4,85E+06	(322)	1,02E+06	(68)	80	187 45	81.8	62.9	106.2
63	1,10E+06	(91)	1,33E+05	(11)	100	24 14	141.4	76.7	293.7
64	2,46E+06	(204)	3,86E+05	(32)	100	70 25	110.2	76.1	165.5
65	2,91E+06	(145)	5,02E+05	(25)	60	92 36	100.2	65.7	160.2
66	1,93E+06	(160)	3,49E+05	(29)	100	64 24	95.4	64.4	147.3
67	4,26E+06	(212)	6,02E+05	(30)	60	110 40	122.0	83.5	185.3
68	3,33E+06	(276)	6,75E+05	(56)	100	123 33	85.0	63.8	113.1
69	4,54E+06	(377)	4,58E+05	(38)	100	84 27	170.7	122.9	244.8
70	7,33E+06	(219)	9,71E+05	(29)	36	177 65	130.2	88.8	199.1
71	4,58E+06	(380)	5,30E+05	(44)	100	97 29	149.0	109.4	208.4
72	2,95E+06	(88)	4,02E+05	(12)	36	73 42	125.6	69.5	253.0
73	1,77E+06	(88)	3,01E+05	(15)	60	55 28	101.0	58.7	188.6
74	5,45E+06	(226)	8,43E+05	(35)	50	154 52	111.6	78.4	164.4
75	3,73E+06	(310)	7,35E+05	(61)	100	134 34	87.7	66.6	115.3
76	3,69E+06	(230)	4,66E+05	(29)	75	85 31	136.7	93.4	208.7
77	6,52E+06	(379)	1,22E+06	(71)	70	223 53	92.1	71.5	118.7
78	2,43E+06	(97)	4,02E+05	(16)	48	73 36	104.4	61.8	190.3
79	4,92E+06	(204)	7,47E+05	(31)	50	137 49	113.7	78.1	171.8
80	3,86E+06	(192)	4,42E+05	(22)	60	81 34	150.0	97.2	245.0
81	1,27E+06	(21)	6,63E+05	(11)	20	121 72	33.1	15.4	76.3
82	4,10E+06	(340)	5,06E+05	(42)	100	93 28	139.7	101.7	197.4
83	1,54E+06	(51)	3,31E+05	(11)	40	61 36	79.7	41.6	170.5
84	4,64E+06	(385)	8,07E+05	(67)	100	148 36	99.1	76.5	128.4
85	8,31E+06	(69)	9,64E+05	(8)	10	176 121	146.6	72.1	353.8
86	2,57E+06	(213)	5,66E+05	(47)	100	104 30	78.6	57.3	110.4
87	3,71E+06	(154)	5,54E+05	(23)	50	101 42	115.5	74.8	187.8
88	2,67E+06	(222)	3,49E+05	(29)	100	64 24	132.0	90.0	201.7
89	4,47E+06	(297)	7,38E+05	(49)	80	135 38	105.0	77.7	145.1
90	2,92E+06	(242)	3,61E+05	(30)	100	66 24	139.0	95.6	210.5
91	1,66E+06	(138)	2,65E+05	(22)	100	48 20	108.2	69.3	178.5
92	3,69E+06	(306)	5,42E+05	(45)	100	99 29	117.6	86.2	164.7
93	2,79E+06	(139)	7,43E+05	(37)	60	136 44	65.2	45.3	96.6
94	1,74E+06	(52)	4,02E+05	(12)	36	73 42	74.6	39.8	154.3
95	5,85E+06	(238)	8,61E+05	(35)	49	157 53	117.5	82.6	172.8
96	2,98E+06	(247)	3,37E+05	(28)	100	62 23	151.8	103.3	233.1
97	4,62E+06	(230)	8,23E+05	(41)	60	151 47	97.1	69.8	139.1
98	4,28E+06	(213)	4,62E+05	(23)	60	84 35	159.1	104.3	256.3
99	2,00E+06	(166)	3,49E+05	(29)	100	64 24	99.0	66.8	152.6
100	1,88E+06	(156)	3,13E+05	(26)	100	57 22	103.6	68.6	163.8
101	6,02E+06	(45)	4,02E+05	(3)	9	73 79	245.5	83.7	1201.9

POOLED 3,07E+06(17531) 4,87E+05(2781) 6886 89 4 109.4 103.6 115.5

CHI^2 PROBABILITY (%): 0.0

POOLED AGE W/ 68% CONF. INTERVAL(Ma): 109.4, 106.4 -- 112.5 (-3.0 +3.1)
 95% CONF. INTERVAL(Ma): 103.6 -- 115.5 (-5.8 +6.1)

CENTRAL AGE W/ 68% CONF. INTERVAL(Ma): 109.7, 105.7 -- 113.8 (-4.0 +4.1)
 95% CONF. INTERVAL(Ma): 102.1 -- 117.9 (-7.6 +8.2)
 AGE DISPERSION (%): 23.4

FIT OPTION: Best-fit peaks using the binomial model of Galbraith and Green

INITIAL GUESS FOR MODEL PARAMETERS (number of peaks to fit = 3)

Peak #.	Peak Age	Theta	Fraction(%)	Count
1.	40.90	0.701	2.0	2.06
2.	109.40	0.863	43.1	43.56
3.	263.30	0.939	2.9	2.96

Total range for grain ages: 32,6 to 263,3 Ma
 Number of active grains (Num. used for fit): 101
 Number of removed grains: 0
 Degrees of freedom for fit: 96
 Average of the SE(Z)'s for the grains: 0,25
 Estimated width of peaks in PD plot in Z units: 0,29

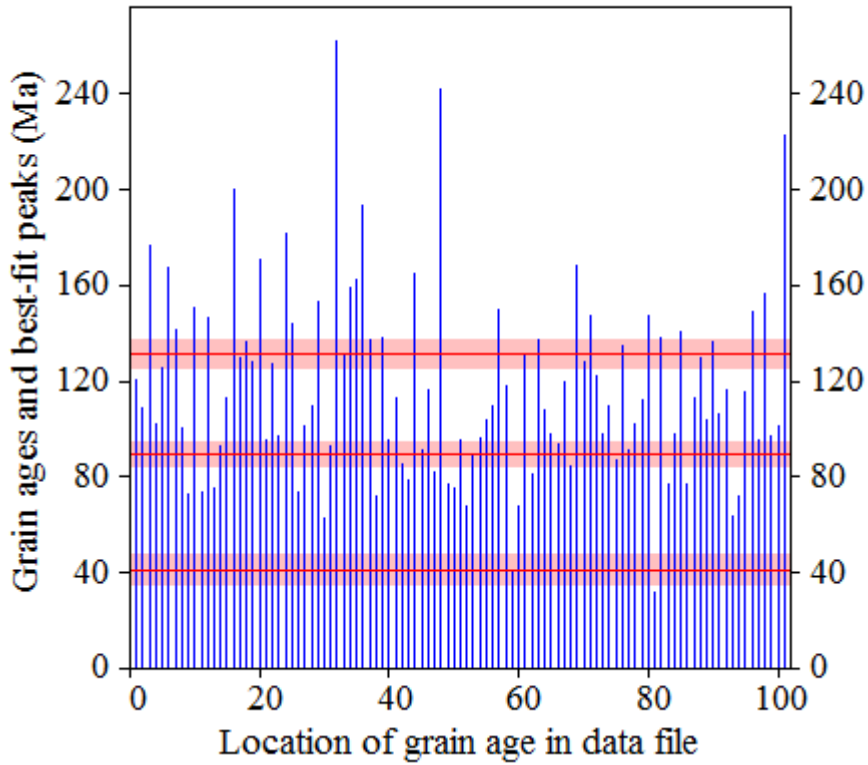
PARAMETERS FOR BEST-FIT PEAKS

- * Standard error for peak age includes group error
- * Peak width is for PD plot assuming a kernel factor = 0.60

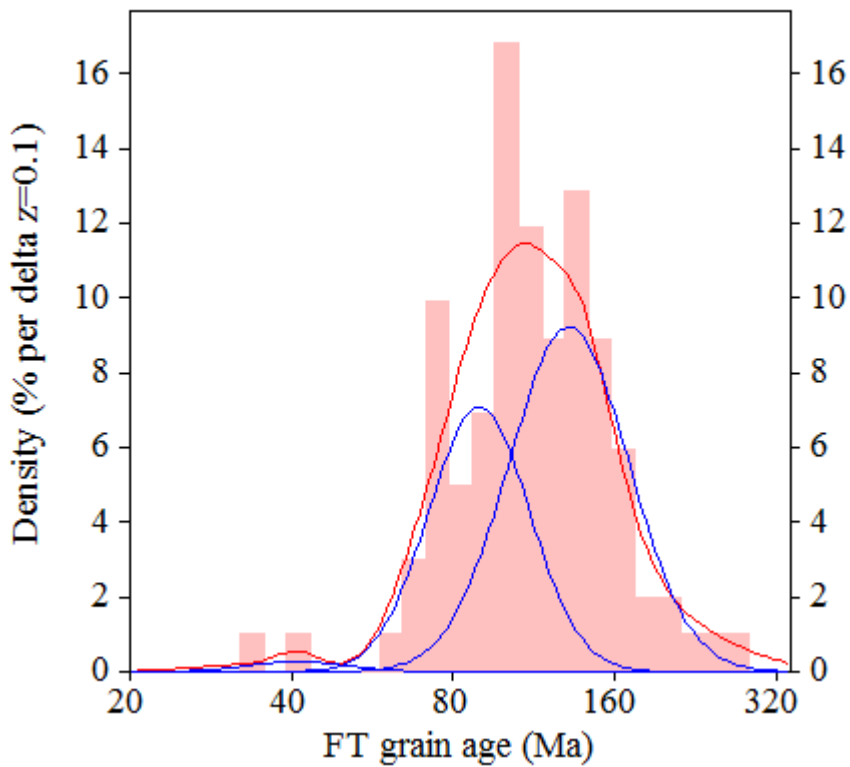
#.	Peak Age(Ma)	68%CI	95%CI	W(Z)	Frac(%)	SE,%	Count
1.	41.0	-6,1 ...+7,1	-11,0 ...+15,1	0.21	1.4	1.3	1.4
2.	89.6	-4,7 ...+4,9	-8,9 ...+9,9	0.22	38.6	10.2	39.0
3.	131.6	-5,8 ...+6,0	-11,1 ...+12,1	0.26	60.0	10.2	60.6

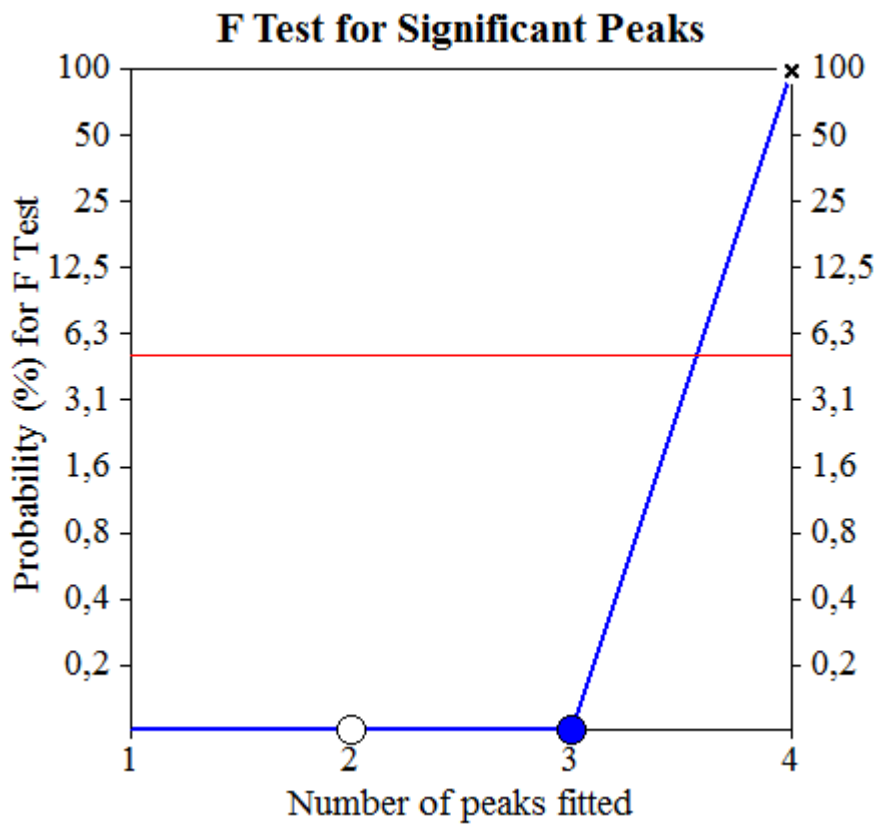
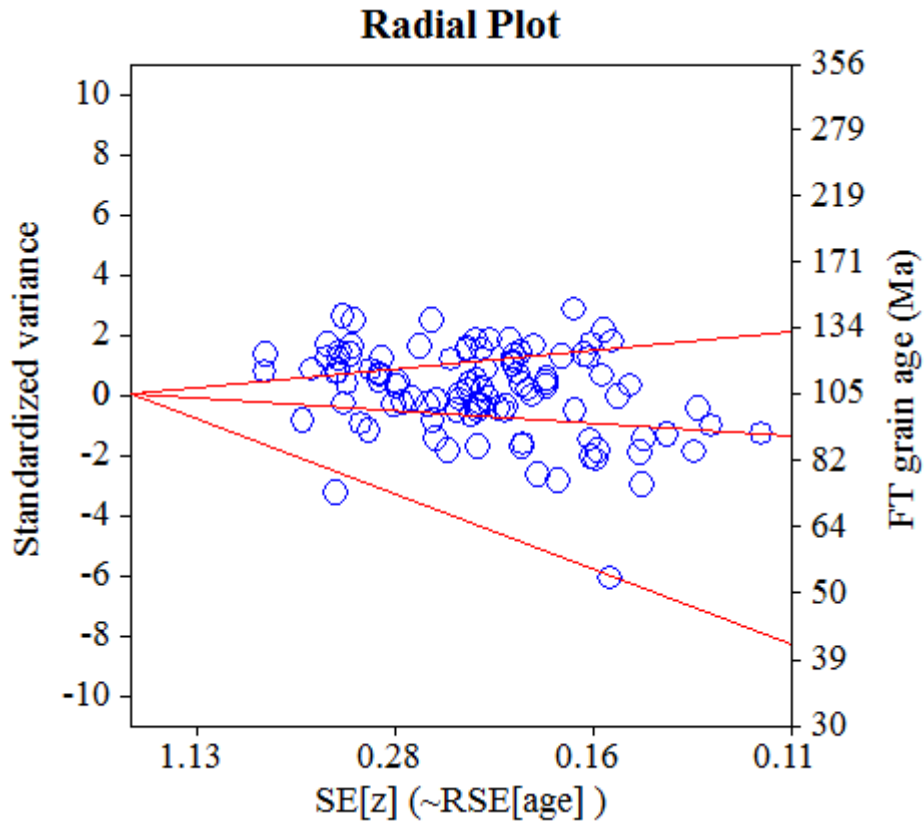
Log-likelihood for best fit: -326,691
 Chi-squared value for best fit: 98,578
 Reduced chi-squared value: 1,027
 Probability for F test: 0%
 Condition number for COVAR matrix: 152,47
 Number of iterations: 34

Plot of Grain Ages (Unsorted)



Probability-Density Plot with Best-Fit Peaks





NEW PARAMETERS - ZETA METHOD

EFFECTIVE TRACK DENSITY FOR FLUENCE MONITOR (tracks/cm²): 3,10E+05
 RELATIVE ERROR (%): 1,12
 EFFECTIVE URANIUM CONTENT OF MONITOR (ppm): 50,00
 ZETA FACTOR AND STANDARD ERROR (yr cm²): 141,03 2,32
 SIZE OF COUNTER SQUARE (cm²): 8,30E-07

GRAIN AGES IN ORIGINAL ORDER

Grain no.	RhoS (cm ⁻²)	(Ns)	RhoI (cm ⁻²)	(Ni)	Squares	U+/-2s	Grain Age (Ma)		
							Age	--95% CI--	
1	7,72E+06	(205)	2,37E+06	(63)	32	383 97	70.1	52.8	93.0
2	6,31E+06	(110)	2,24E+06	(39)	21	361 115	61.1	42.2	90.6
3	1,02E+07	(676)	2,00E+06	(133)	80	324 56	109.6	90.8	132.2
4	5,75E+06	(229)	2,54E+06	(101)	48	410 82	49.1	38.7	62.1
5	6,10E+06	(253)	1,78E+06	(74)	50	288 67	73.7	56.8	95.5
6	7,73E+06	(77)	2,21E+06	(22)	12	357 151	75.5	46.8	127.7
7	5,61E+06	(466)	1,18E+06	(98)	100	191 39	102.4	82.3	127.5
8	5,72E+06	(57)	1,10E+06	(11)	12	178 105	110.8	58.4	234.9
9	4,21E+06	(339)	1,08E+06	(87)	97	175 38	84.0	66.3	106.4
10	6,53E+06	(65)	1,41E+06	(14)	12	227 119	99.6	56.0	192.7
11	4,63E+06	(384)	1,37E+06	(114)	100	222 42	72.8	58.9	89.9
12	6,43E+06	(48)	4,42E+06	(33)	9	714 247	31.6	19.9	50.9
13	1,69E+07	(126)	3,75E+06	(28)	9	606 227	97.0	64.4	152.0
14	7,07E+06	(176)	1,77E+06	(44)	30	285 86	86.5	62.1	123.3
15	8,97E+06	(521)	4,06E+06	(236)	70	656 87	47.9	40.9	56.1
16	9,04E+06	(225)	2,37E+06	(59)	30	383 100	82.0	61.6	109.2
17	8,67E+06	(180)	1,59E+06	(33)	25	257 89	117.5	81.2	175.9
18	7,19E+06	(179)	1,24E+06	(31)	30	201 72	124.3	85.1	188.3
19	8,35E+06	(104)	1,12E+06	(14)	15	182 96	158.5	91.7	299.9
20	1,32E+07	(219)	4,40E+06	(73)	20	710 166	64.7	49.6	84.4
21	5,69E+06	(236)	1,18E+06	(49)	50	191 54	104.0	76.5	144.6
22	7,49E+06	(261)	1,86E+06	(65)	42	301 75	86.4	65.8	113.4
23	5,34E+06	(124)	1,12E+06	(26)	28	181 70	102.7	67.4	163.6
24	4,73E+06	(393)	1,06E+06	(88)	100	171 37	96.2	76.3	121.3
25	9,37E+06	(389)	3,23E+06	(134)	50	522 91	62.8	51.5	76.6
26	8,19E+06	(340)	2,63E+06	(109)	50	424 82	67.4	54.2	83.8
27	1,65E+06	(41)	2,01E+05	(5)	30	32 28	171.7	70.4	555.8
28	4,53E+06	(376)	9,28E+05	(77)	100	150 34	105.0	82.1	134.2
29	6,37E+06	(148)	1,42E+06	(33)	28	229 79	96.8	66.4	145.9
30	5,22E+06	(303)	1,48E+06	(86)	70	239 52	76.0	59.7	96.7
31	6,39E+06	(371)	1,41E+06	(82)	70	228 50	97.4	76.6	123.8
32	6,99E+06	(58)	3,49E+06	(29)	10	564 208	43.4	27.4	70.4
33	5,35E+06	(311)	1,67E+06	(97)	70	270 55	69.3	55.0	87.1
34	6,39E+06	(191)	1,10E+06	(33)	36	178 62	124.6	86.3	186.2
35	4,54E+06	(264)	1,02E+06	(59)	70	164 43	96.1	72.5	127.4
36	7,59E+06	(126)	2,35E+06	(39)	20	380 121	69.9	48.7	103.0
37	9,64E+06	(200)	2,46E+06	(51)	25	397 111	84.8	62.3	117.8
38	7,90E+06	(236)	1,37E+06	(41)	36	222 69	124.0	89.2	177.3
39	6,31E+06	(220)	1,66E+06	(58)	42	269 71	81.6	61.1	108.9
40	4,40E+06	(219)	1,02E+06	(51)	60	165 46	92.8	68.4	128.6
41	4,37E+06	(58)	1,13E+06	(15)	16	182 93	83.2	47.0	158.4
42	6,12E+06	(61)	2,81E+06	(28)	12	454 170	47.2	29.8	76.8
43	6,59E+06	(164)	1,49E+06	(37)	30	240 79	95.7	67.0	140.8
44	6,48E+06	(113)	2,07E+06	(36)	21	334 111	67.9	46.5	101.9
45	3,01E+05	(25)	2,17E+05	(18)	100	35 16	30.2	15.9	58.7
46	2,92E+06	(242)	6,75E+05	(56)	100	109 29	92.8	69.4	124.0
47	4,10E+06	(170)	1,23E+06	(51)	50	199 56	72.2	52.7	100.8
48	4,87E+06	(101)	1,64E+06	(34)	25	265 90	64.3	43.4	98.0
49	5,27E+06	(70)	1,43E+06	(19)	16	231 105	79.4	47.7	140.0
50	5,16E+06	(107)	6,75E+05	(14)	25	109 57	163.0	94.4	308.1
51	5,72E+06	(285)	1,91E+06	(95)	60	308 63	64.8	51.3	81.9
52	5,66E+06	(169)	1,81E+06	(54)	36	292 79	67.8	49.8	94.0
53	8,39E+06	(174)	2,46E+06	(51)	25	397 111	73.9	53.9	103.1

Datafile: C:\BH2\Alejandro\6-2015_zircon\CP_088b\CP_088b.ftz

Title: Sample No. CP-088b Irr 6-2015-30

Grain no.	RhoS (cm ⁻²)	(Ns)	RhoI (cm ⁻²)	(Ni)	Squares	U+/-2s		Grain Age (Ma)		
								Age	--95% CI--	
54	3,75E+06	(28)	2,54E+06	(19)	9	411	186	32.0	17.3	60.7
55	4,22E+06	(70)	1,02E+06	(17)	20	165	79	88.6	52.1	161.0
56	1,08E+06	(90)	2,41E+05	(20)	100	39	17	96.9	59.7	166.3
57	3,45E+06	(258)	9,91E+05	(74)	90	160	37	75.2	58.0	97.4
58	5,27E+06	(437)	1,17E+06	(97)	100	189	38	97.1	77.8	121.1
59	7,83E+06	(156)	1,61E+06	(32)	24	260	91	105.1	71.9	159.0
60	4,01E+06	(266)	1,36E+06	(90)	80	219	46	63.8	50.2	81.2
61	4,98E+06	(413)	1,57E+06	(130)	100	253	45	68.7	56.3	83.9
62	1,65E+06	(137)	4,58E+05	(38)	100	74	24	78.0	54.3	114.9
63	4,82E+06	(40)	1,81E+06	(15)	10	292	148	57.5	31.4	112.5
64	4,74E+06	(118)	9,64E+05	(24)	30	156	63	105.8	68.3	171.9
65	5,98E+06	(124)	9,64E+05	(20)	25	156	69	133.0	83.4	225.3
66	7,11E+06	(590)	1,11E+06	(92)	100	179	37	137.7	110.5	171.6
67	9,80E+06	(813)	1,64E+06	(136)	100	265	46	128.7	107.1	154.6
68	5,82E+06	(116)	1,36E+06	(27)	24	219	84	92.7	60.9	146.7
69	7,52E+06	(131)	2,01E+06	(35)	21	324	109	80.9	55.6	121.2
70	3,73E+06	(155)	1,11E+06	(46)	50	179	53	72.9	52.4	103.8
71	5,82E+06	(464)	1,23E+06	(98)	96	199	40	102.0	81.9	127.0
72	1,33E+06	(55)	2,65E+05	(11)	50	43	25	106.9	56.2	227.3
73	7,41E+06	(369)	1,87E+06	(93)	60	302	63	85.6	68.1	107.5
74	5,37E+06	(446)	1,13E+06	(94)	100	183	38	102.2	81.7	127.8
75	4,52E+06	(375)	9,76E+05	(81)	100	158	35	99.7	78.3	126.8
76	1,30E+06	(54)	2,41E+05	(10)	50	39	24	115.3	59.1	254.6
77	2,94E+06	(122)	5,30E+05	(22)	50	86	36	119.2	75.9	197.3
78	4,96E+06	(103)	1,73E+06	(36)	25	280	93	62.0	42.2	93.4
79	6,25E+06	(83)	2,33E+06	(31)	16	377	135	58.0	38.1	90.8
80	4,56E+06	(121)	1,69E+06	(45)	32	274	81	58.3	41.2	84.1
81	6,02E+06	(45)	1,20E+06	(9)	9	195	126	106.7	52.4	249.1
82	2,17E+06	(63)	3,10E+05	(9)	35	50	32	148.7	75.1	340.4
83	6,67E+06	(277)	1,04E+06	(43)	50	167	51	138.6	100.8	195.8
84	4,39E+06	(182)	5,54E+05	(23)	50	90	37	169.4	110.6	273.6
85	4,51E+06	(262)	1,26E+06	(73)	70	203	48	77.3	59.6	100.3
86	4,37E+06	(254)	9,64E+05	(56)	70	156	42	97.4	73.0	129.9
87	3,45E+06	(172)	8,63E+05	(43)	60	139	42	86.5	61.8	123.9
88	4,49E+06	(261)	2,05E+06	(119)	70	331	61	47.5	38.2	59.1
89	5,49E+06	(164)	2,04E+06	(61)	36	330	84	58.3	43.3	79.7
90	4,48E+06	(186)	6,02E+05	(25)	50	97	39	159.5	105.7	252.5
91	2,11E+06	(105)	5,02E+05	(25)	60	81	32	90.6	58.5	146.4
92	3,82E+06	(57)	7,36E+05	(11)	18	119	70	110.8	58.4	234.9
93	6,29E+06	(47)	1,61E+06	(12)	9	260	147	84.1	44.5	174.7
94	8,19E+06	(68)	1,33E+06	(11)	10	214	126	131.9	70.4	277.0
95	5,49E+06	(456)	1,36E+06	(113)	100	220	42	87.1	70.7	107.2
96	3,33E+06	(116)	3,44E+05	(12)	42	56	31	205.1	115.2	407.1
POOLED	5,24E+06	(20034)	1,34E+06	(5105)	4606	216	8	85.1	81.0	89.4

CHI^2 PROBABILITY (%): 0.0

POOLED AGE W/ 68% CONF. INTERVAL(Ma): 85.1, 83.0 -- 87.2 (-2.1 +2.2)
 95% CONF. INTERVAL(Ma): 81.0 -- 89.4 (-4.1 +4.3)

CENTRAL AGE W/ 68% CONF. INTERVAL(Ma): 84.3, 81.1 -- 87.6 (-3.2 +3.3)
 95% CONF. INTERVAL(Ma): 78.2 -- 90.9 (-6.1 +6.6)
 AGE DISPERSION (%): 27.0

FIT OPTION: Best-fit peaks using the binomial model of Galbraith and Green

INITIAL GUESS FOR MODEL PARAMETERS (number of peaks to fit = 3)

Peak #.	Peak Age	Theta	Fraction(%)	Count
1.	30.80	0.586	2.9	2.80
2.	47.90	0.688	11.6	11.16
3.	85.10	0.797	42.1	40.45

Total range for grain ages: 30,0 to 200,3 Ma
 Number of active grains (Num. used for fit): 96
 Number of removed grains: 0
 Degrees of freedom for fit: 91
 Average of the SE(Z)'s for the grains: 0,2
 Estimated width of peaks in PD plot in Z units: 0,24

PARAMETERS FOR BEST-FIT PEAKS

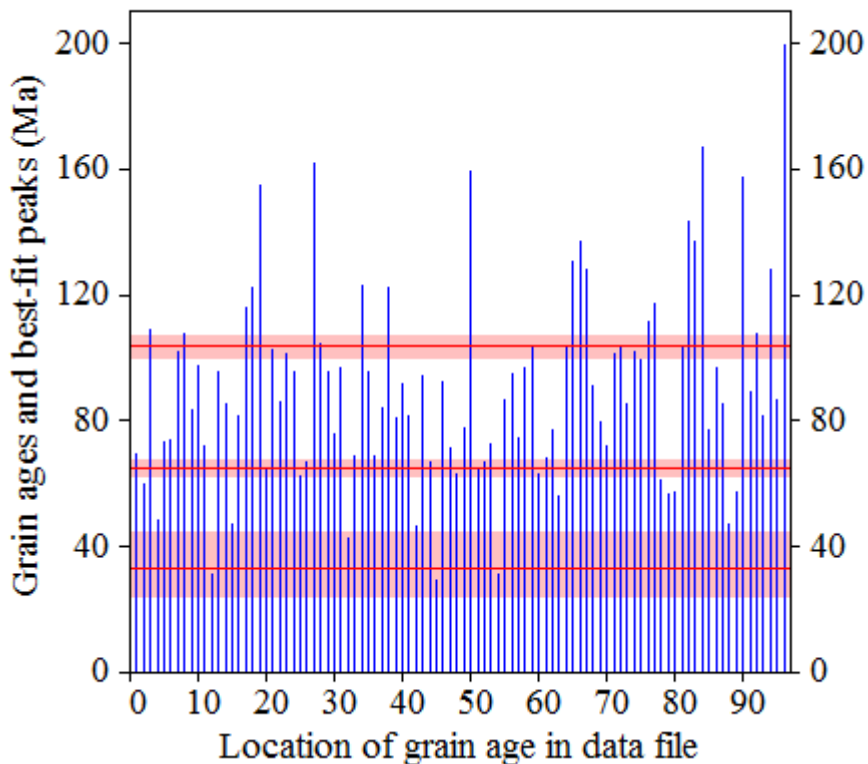
* Standard error for peak age includes group error

* Peak width is for PD plot assuming a kernel factor = 0.60

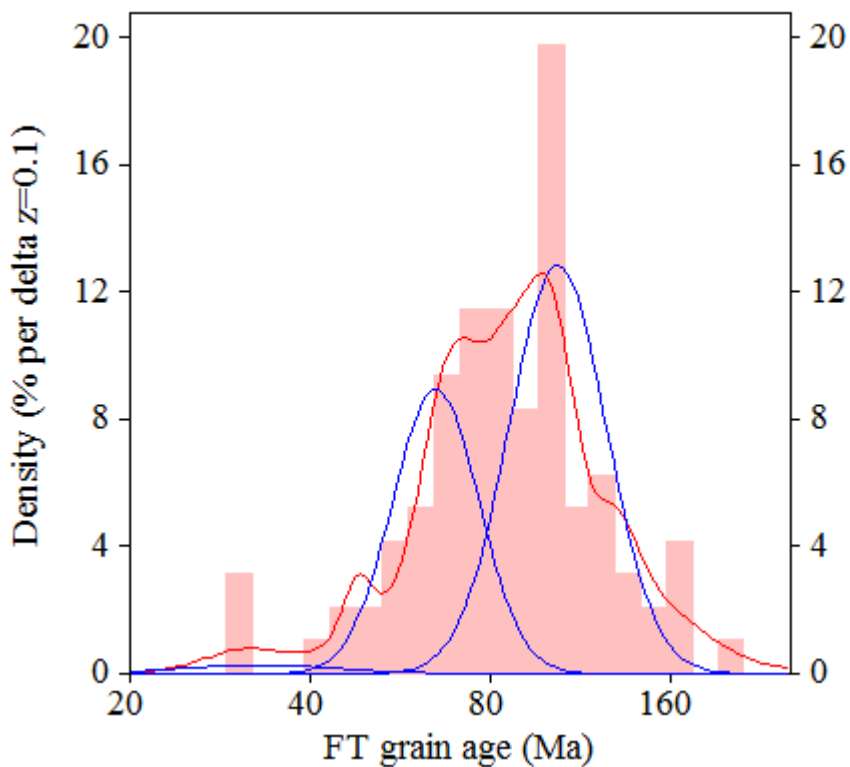
#.	Peak Age(Ma)	68%CI	95%CI	W(Z)	Frac(%)	SE,%	Count
1.	32.9	-8,5 ...+11,4	-14,5 ...+26,0	0.29	2.0	1.9	1.9
2.	64.9	-2,5 ...+2,6	-4,7 ...+5,1	0.17	37.4	6.7	35.9
3.	103.6	-3,4 ...+3,5	-6,5 ...+6,9	0.19	60.7	6.7	58.3

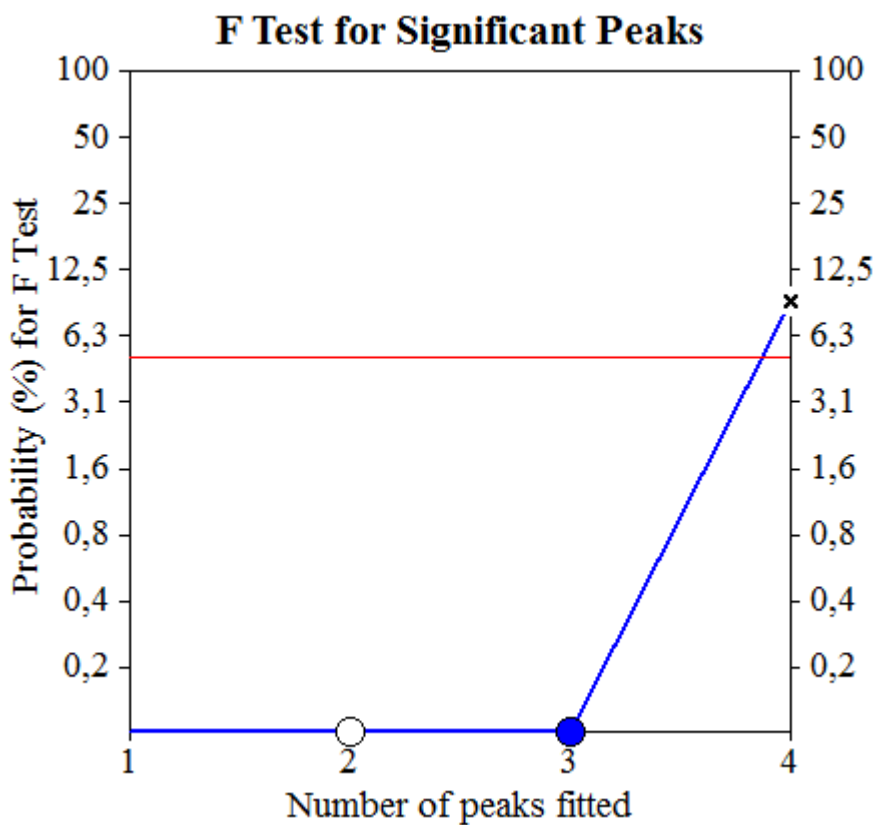
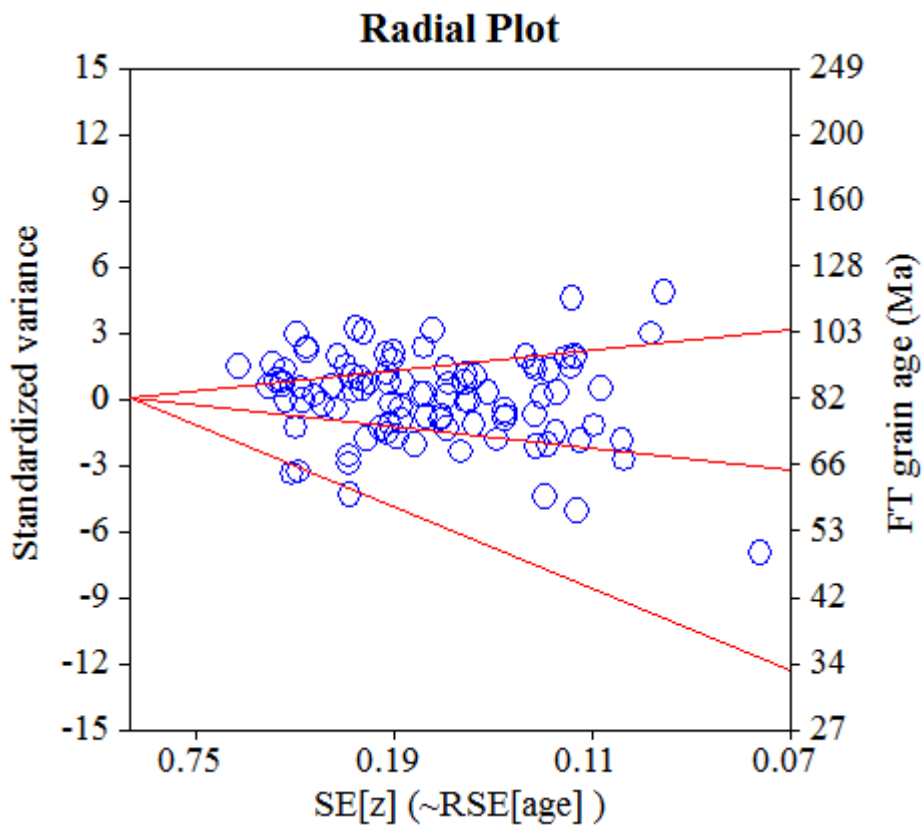
Log-likelihood for best fit: -377,072
 Chi-squared value for best fit: 98,995
 Reduced chi-squared value: 1,088
 Probability for F test: 0%
 Condition number for COVAR matrix: 282,89
 Number of iterations: 14

Plot of Grain Ages (Unsorted)



Probability-Density Plot with Best-Fit Peaks





NEW PARAMETERS - ZETA METHOD

EFFECTIVE TRACK DENSITY FOR FLUENCE MONITOR (tracks/cm²): 3,04E+05
 RELATIVE ERROR (%): 1,13
 EFFECTIVE URANIUM CONTENT OF MONITOR (ppm): 50,00
 ZETA FACTOR AND STANDARD ERROR (yr cm²): 126,82 2,05
 SIZE OF COUNTER SQUARE (cm²): 8,30E-07

GRAIN AGES IN ORIGINAL ORDER

Grain no.	RhoS (cm ⁻²)	(Ns)	RhoI (cm ⁻²)	(Ni)	Squares	U+/-2s	Grain Age (Ma)		
							Age	--95% CI--	
1	6,75E+06	(112)	1,14E+06	(19)	20	188 85	111.8	69.1	192.9
2	4,82E+06	(64)	9,79E+05	(13)	16	161 88	93.3	51.5	185.2
3	2,41E+06	(100)	4,34E+05	(18)	50	71 33	105.4	64.1	185.4
4	5,30E+06	(176)	9,64E+05	(32)	40	159 56	104.7	72.0	157.9
5	7,83E+06	(78)	1,61E+06	(16)	12	264 130	92.5	54.1	170.3
6	6,10E+06	(253)	1,49E+06	(62)	50	246 62	77.6	58.7	102.4
7	5,76E+06	(287)	1,02E+06	(51)	60	168 47	107.3	79.7	147.6
8	6,20E+06	(144)	2,32E+06	(54)	28	382 104	51.1	37.2	71.3
9	8,13E+06	(81)	2,11E+06	(21)	12	347 150	73.5	45.4	125.4
10	4,55E+06	(34)	2,41E+06	(18)	9	396 185	36.2	20.0	68.1
11	8,48E+06	(352)	1,61E+06	(67)	50	266 65	99.8	76.8	129.5
12	8,67E+06	(144)	2,83E+06	(47)	20	466 136	58.7	42.0	83.5
13	4,69E+06	(35)	1,47E+06	(11)	9	242 143	60.4	30.4	132.6
14	7,93E+06	(79)	1,61E+06	(16)	12	264 130	93.7	54.9	172.3
15	6,48E+06	(86)	1,66E+06	(22)	16	272 115	74.5	46.5	125.3
16	8,96E+06	(119)	1,66E+06	(22)	16	272 115	102.8	65.4	170.5
17	8,98E+06	(149)	2,23E+06	(37)	20	367 120	76.9	53.6	113.6
18	6,69E+06	(333)	1,20E+06	(60)	60	198 51	105.3	80.0	138.4
19	6,25E+06	(83)	9,79E+05	(13)	16	161 88	120.6	67.8	236.5
20	6,82E+06	(283)	1,28E+06	(53)	50	210 58	101.9	76.0	139.4
21	8,04E+06	(667)	1,55E+06	(129)	100	256 45	98.5	81.4	119.2
22	5,54E+06	(69)	1,93E+06	(24)	15	317 128	54.9	34.3	91.6
23	5,60E+06	(372)	1,11E+06	(74)	80	183 43	95.6	74.4	122.6
24	2,96E+06	(59)	1,61E+06	(32)	24	264 93	35.4	22.7	56.3
25	5,25E+06	(122)	1,33E+06	(31)	28	219 78	75.1	50.6	115.5
26	4,59E+06	(61)	1,20E+06	(16)	16	198 98	72.5	41.7	135.2
27	9,13E+06	(379)	1,30E+06	(54)	50	214 58	132.7	99.9	176.1
28	5,30E+06	(132)	7,63E+05	(19)	30	126 57	131.6	81.9	225.6
29	7,63E+06	(57)	3,48E+06	(26)	9	572 223	42.0	26.1	69.7
30	3,25E+06	(135)	2,17E+06	(90)	50	357 75	28.7	21.9	37.5
31	6,33E+06	(105)	1,27E+06	(21)	20	208 90	95.1	59.6	160.3
32	5,28E+06	(92)	1,95E+06	(34)	21	321 109	51.8	34.7	79.3
33	6,46E+06	(134)	1,40E+06	(29)	25	230 85	88.1	59.0	136.7
34	5,29E+06	(123)	1,68E+06	(39)	28	276 88	60.3	41.9	89.0
35	5,42E+06	(72)	1,13E+06	(15)	16	186 94	91.1	52.3	171.6
36	4,96E+06	(103)	8,67E+05	(18)	25	143 66	108.6	66.1	190.7
37	5,54E+06	(216)	1,05E+06	(41)	47	173 54	100.4	72.0	144.0
38	2,61E+06	(26)	1,00E+05	(1)	12	17 27	425.6	81.8	9232.3
39	6,08E+06	(106)	2,29E+06	(40)	21	377 119	50.8	35.1	75.1
40	3,68E+06	(55)	1,54E+06	(23)	18	253 105	45.7	27.8	78.2
41	4,92E+06	(49)	3,21E+06	(32)	12	528 186	29.4	18.5	47.5
42	6,28E+06	(198)	1,33E+06	(42)	38	219 67	90.0	64.5	128.8
43	9,23E+06	(268)	1,62E+06	(47)	35	266 77	108.7	79.8	151.6
44	8,28E+06	(165)	8,53E+05	(17)	24	140 67	182.9	112.3	321.0
45	4,39E+06	(153)	1,15E+06	(40)	42	189 59	73.1	51.5	106.4
46	4,42E+06	(33)	2,41E+06	(18)	9	396 185	35.1	19.3	66.3
47	8,96E+06	(223)	2,01E+06	(50)	30	330 93	85.2	62.7	118.3
48	8,82E+06	(183)	3,33E+06	(69)	25	547 132	50.6	38.3	66.7
49	8,89E+06	(118)	1,73E+06	(23)	16	285 118	97.6	62.5	160.1
50	4,88E+06	(81)	2,05E+06	(34)	20	337 115	45.6	30.3	70.4
51	7,86E+06	(652)	1,33E+06	(110)	100	218 42	112.8	92.0	138.2
52	5,77E+06	(115)	9,04E+05	(18)	24	149 69	121.1	74.1	211.7
53	6,02E+06	(250)	9,88E+05	(41)	50	162 51	116.1	83.6	165.8

Grain no.	RhoS (cm ⁻²)	(Ns)	RhoI (cm ⁻²)	(Ni)	Squares	U+/-2s		Grain Age (Ma)		
								Age	--95% CI--	
54	7,20E+06	(251)	8,61E+05	(30)	42	142	51	158.5	109.2	239.7
55	9,69E+06	(201)	1,30E+06	(27)	25	214	82	141.2	95.0	219.5
56	8,28E+06	(110)	3,01E+05	(4)	16	50	47	491.1	198.1	1727.0
57	3,88E+06	(116)	8,70E+05	(26)	36	143	56	85.0	55.6	135.8
58	3,40E+06	(79)	2,62E+06	(61)	28	432	111	24.9	17.6	35.4
59	7,23E+06	(150)	1,16E+06	(24)	25	190	77	118.7	77.4	191.2
60	7,57E+06	(628)	1,22E+06	(101)	100	200	40	118.2	95.7	146.0
61	5,87E+06	(78)	9,79E+05	(13)	16	161	88	113.4	63.5	223.0
62	5,48E+06	(364)	1,31E+06	(87)	80	216	46	79.7	63.0	100.8
63	5,98E+06	(124)	9,64E+05	(20)	25	159	70	117.6	73.7	199.4
64	5,98E+06	(248)	1,23E+06	(51)	50	202	57	92.8	68.7	128.2
65	7,13E+06	(355)	1,20E+06	(60)	60	198	51	112.1	85.3	147.3
66	5,36E+06	(267)	9,64E+05	(48)	60	159	46	106.0	78.1	147.4
67	5,62E+06	(56)	1,20E+06	(12)	12	198	112	88.4	47.5	181.9
68	6,85E+06	(273)	1,33E+06	(53)	48	219	60	98.3	73.3	134.6
POOLED	6,36E+06	(11865)	1,37E+06	(2566)	2249	226	10	88.5	83.6	93.7

CHI^2 PROBABILITY (%): 0.0

POOLED AGE W/	68% CONF. INTERVAL(Ma):	88.5,	86.0 --	91.1 (-2.5 +2.6)
	95% CONF. INTERVAL(Ma):		83.6 --	93.7 (-4.9 +5.2)
CENTRAL AGE W/	68% CONF. INTERVAL(Ma):	80.9,	76.1 --	85.9 (-4.8 +5.1)
	95% CONF. INTERVAL(Ma):		71.8 --	91.1 (-9.1 +10.3)
	AGE DISPERSION (%):	43.2		

FIT OPTION: Best-fit peaks using the binomial model of Galbraith and Green

INITIAL GUESS FOR MODEL PARAMETERS (number of peaks to fit = 3)

Peak #.	Peak Age	Theta	Fraction(%)	Count
1.	24.90	0.564	3.9	2.62
2.	88.50	0.822	37.4	25.45
3.	99.10	0.838	51.6	35.10

Total range for grain ages: 24,9 to 456,8 Ma
 Number of active grains (Num. used for fit): 68
 Number of removed grains: 0
 Degrees of freedom for fit: 63
 Average of the SE(Z)'s for the grains: 0,24
 Estimated width of peaks in PD plot in Z units: 0,28

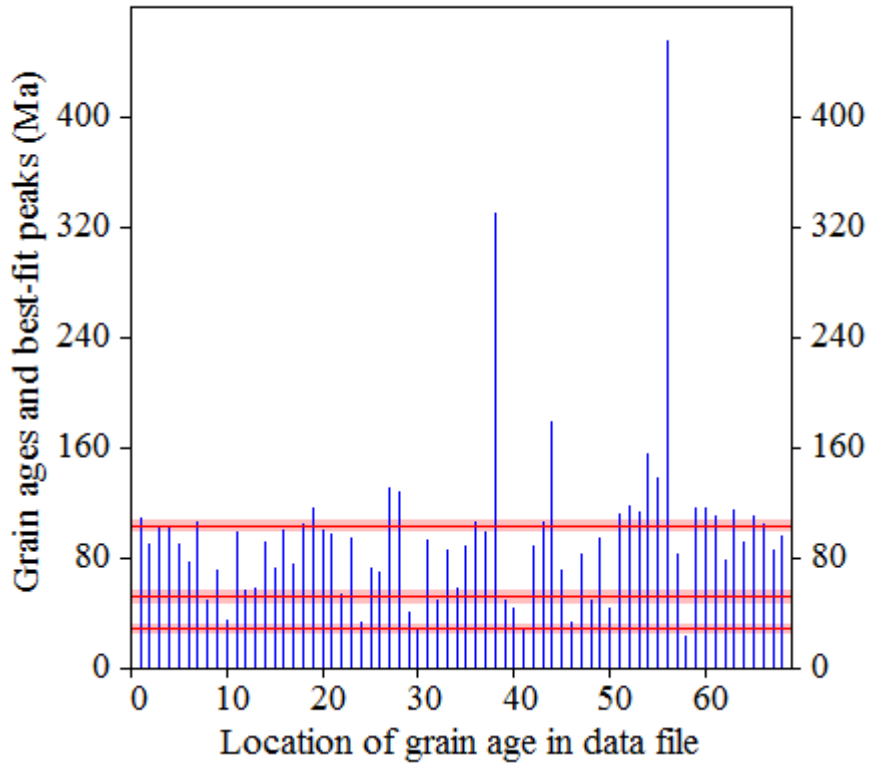
PARAMETERS FOR BEST-FIT PEAKS

- * Standard error for peak age includes group error
- * Peak width is for PD plot assuming a kernel factor = 0.60

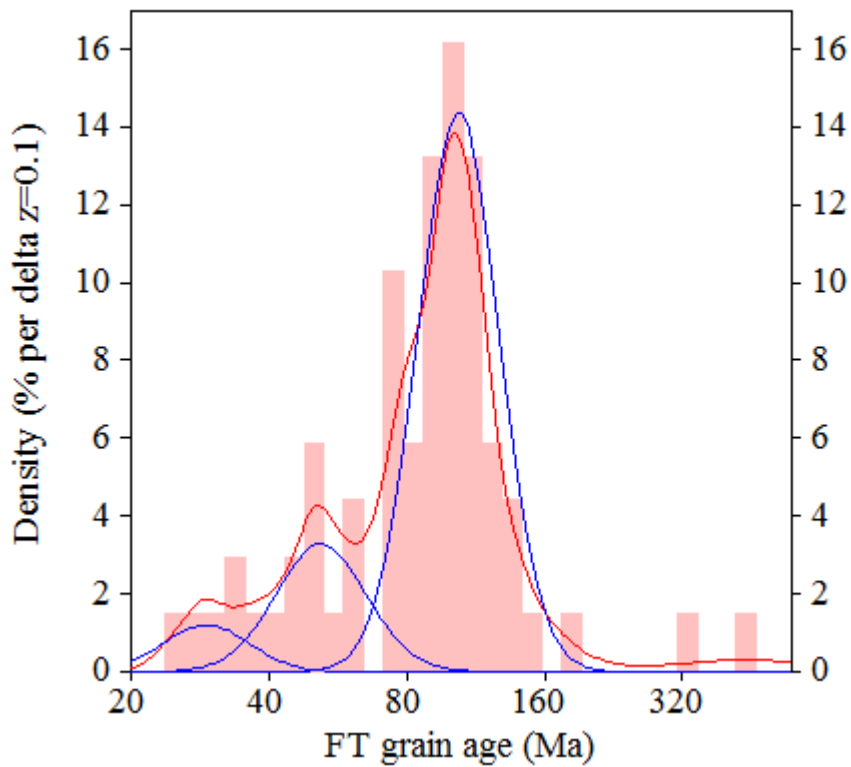
#.	Peak Age(Ma)	68%CI	95%CI	W(Z)	Frac(%)	SE,%	Count
1.	29.2	-2,9 ...+3,2	-5,4 ...+6,6	0.21	6.4	3.7	4.3
2.	52.0	-3,9 ...+4,2	-7,3 ...+8,5	0.23	18.9	5.6	12.9
3.	103.8	-3,3 ...+3,4	-6,4 ...+6,8	0.21	74.7	5.8	50.8

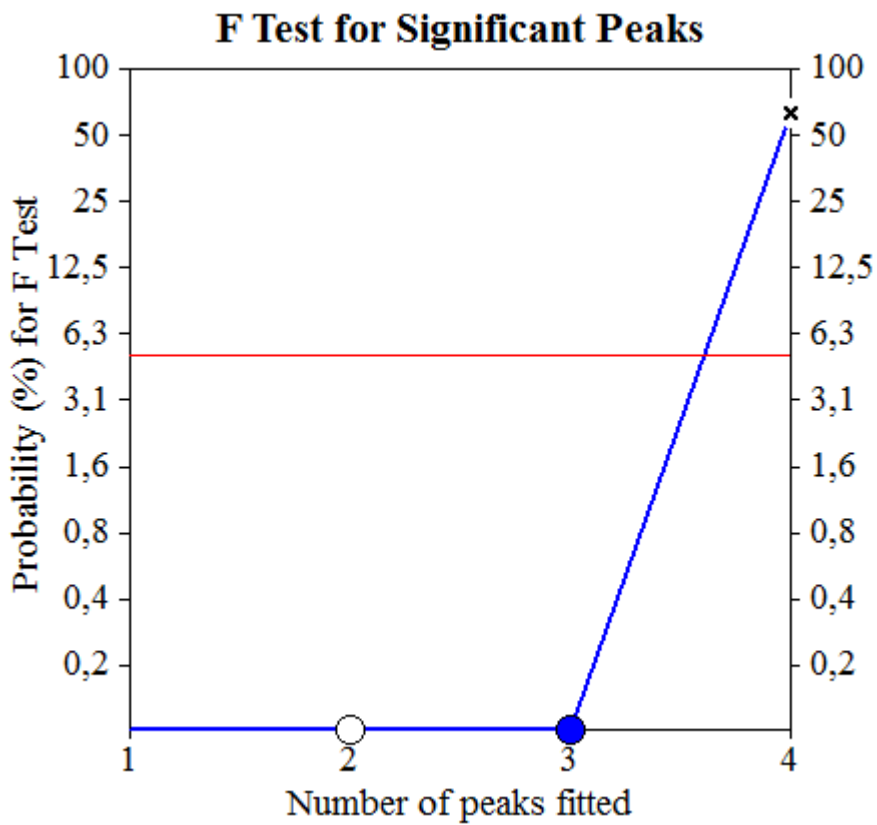
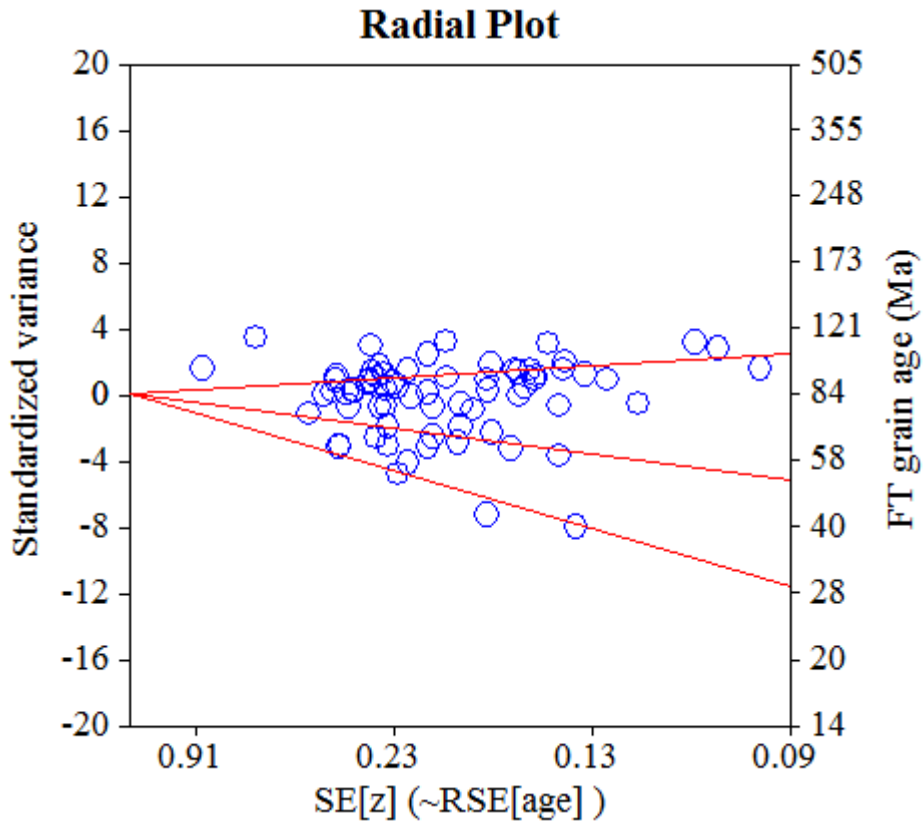
Log-likelihood for best fit: -252,314
 Chi-squared value for best fit: 67,569
 Reduced chi-squared value: 1,073
 Probability for F test: 0%
 Condition number for COVAR matrix: 17,08
 Number of iterations: 18

Plot of Grain Ages (Unsorted)



Probability-Density Plot with Best-Fit Peaks





NEW PARAMETERS - ZETA METHOD

EFFECTIVE TRACK DENSITY FOR FLUENCE MONITOR (tracks/cm²): 2,77E+05
 RELATIVE ERROR (%): 1,23
 EFFECTIVE URANIUM CONTENT OF MONITOR (ppm): 50,00
 ZETA FACTOR AND STANDARD ERROR (yr cm²): 128,02 1,87
 SIZE OF COUNTER SQUARE (cm²): 8,30E-07

GRAIN AGES IN ORIGINAL ORDER

Grain no.	RhoS (cm ⁻²)	(Ns)	RhoI (cm ⁻²)	(Ni)	Squares	U+/-2s	Grain Age (Ma)		
							Age	--95% CI--	
1	7,28E+06	(302)	1,20E+06	(50)	50	217 61	106.1	78.7	146.2
2	6,60E+06	(274)	1,37E+06	(57)	50	248 66	84.1	63.2	111.8
3	5,76E+06	(478)	1,48E+06	(123)	100	267 49	68.4	56.0	83.5
4	3,80E+06	(63)	6,02E+05	(10)	20	109 67	109.4	56.7	240.1
5	6,07E+06	(247)	9,34E+05	(38)	49	168 54	114.0	81.2	165.0
6	4,73E+06	(165)	4,02E+05	(14)	42	72 38	203.8	120.0	380.0
7	4,03E+06	(164)	4,92E+05	(20)	49	89 39	143.0	90.6	240.3
8	9,04E+06	(225)	1,37E+06	(34)	30	246 84	116.0	81.1	171.7
9	6,55E+06	(87)	9,04E+05	(12)	16	163 92	126.0	69.7	253.8
10	5,54E+06	(230)	6,02E+05	(25)	50	109 43	160.4	106.9	252.9
11	5,24E+06	(87)	7,23E+05	(12)	20	130 74	126.0	69.7	253.8
12	8,07E+06	(134)	1,81E+06	(30)	20	326 118	78.5	52.8	121.1
13	7,67E+06	(191)	1,53E+06	(38)	30	275 89	88.3	62.4	128.8
14	5,75E+06	(191)	7,53E+05	(25)	40	136 54	133.5	88.5	211.6
15	7,23E+06	(210)	1,65E+06	(48)	35	298 86	77.0	56.2	107.8
16	8,29E+06	(344)	1,40E+06	(58)	50	252 66	103.6	78.5	136.7
17	2,37E+06	(59)	2,41E+05	(6)	30	43 34	168.3	75.1	477.0
18	6,60E+06	(219)	1,42E+06	(47)	40	255 74	82.0	59.8	115.0
19	2,74E+06	(91)	4,22E+05	(14)	40	76 40	113.3	65.0	215.9
20	8,06E+06	(107)	1,43E+06	(19)	16	258 117	98.5	60.7	170.4
21	7,95E+06	(66)	8,43E+05	(7)	10	152 111	162.0	76.4	418.9
22	5,73E+06	(476)	1,08E+06	(90)	100	195 41	92.7	73.9	116.2
23	4,56E+06	(265)	1,41E+06	(82)	70	254 56	56.8	44.3	72.8
24	3,93E+06	(261)	8,58E+05	(57)	80	155 41	80.1	60.2	106.6
25	7,11E+06	(118)	1,39E+06	(23)	20	250 103	89.9	57.6	147.7
26	3,32E+06	(248)	4,28E+05	(32)	90	77 27	135.6	94.2	202.6
27	4,82E+06	(168)	1,15E+06	(40)	42	207 65	73.9	52.3	107.3
28	8,76E+06	(109)	2,01E+06	(25)	15	362 144	76.6	49.5	123.7
29	7,20E+06	(299)	1,57E+06	(65)	50	282 70	80.6	61.6	105.3
30	7,44E+06	(247)	1,99E+06	(66)	40	358 88	65.6	50.0	86.1
31	4,22E+06	(210)	4,42E+05	(22)	60	80 34	166.2	108.0	270.7
32	6,40E+06	(85)	1,20E+06	(16)	16	217 107	92.9	54.6	170.3
33	6,69E+06	(50)	1,47E+06	(11)	9	265 157	79.3	41.3	169.8
34	8,55E+06	(213)	1,57E+06	(39)	30	282 90	95.9	68.3	138.8
35	6,90E+06	(573)	1,06E+06	(88)	100	191 41	114.0	91.0	142.7
36	7,04E+06	(146)	1,73E+06	(36)	25	313 104	71.4	49.5	106.0
37	4,50E+06	(112)	2,81E+05	(7)	30	51 37	272.2	132.0	685.4
38	5,93E+06	(246)	9,64E+05	(40)	50	174 55	107.9	77.4	154.9
39	6,51E+06	(216)	1,51E+06	(50)	40	271 77	76.1	55.9	105.8
40	5,92E+06	(59)	1,51E+06	(15)	12	271 138	68.9	39.0	131.3
41	3,06E+06	(142)	5,38E+05	(25)	56	97 38	99.5	65.2	159.2
42	9,12E+06	(227)	1,37E+06	(34)	30	246 84	117.0	81.8	173.2
43	4,05E+06	(84)	6,75E+05	(14)	25	122 64	104.7	59.8	200.2
44	5,53E+06	(225)	1,25E+06	(51)	49	226 63	77.7	57.3	107.6
45	6,14E+06	(255)	1,57E+06	(65)	50	282 70	68.8	52.4	90.3
46	4,13E+06	(137)	1,48E+06	(49)	40	266 76	49.4	35.5	70.0
47	2,30E+06	(153)	4,82E+05	(32)	80	87 31	84.0	57.4	127.3
48	3,34E+06	(222)	4,22E+05	(28)	80	76 29	138.6	94.0	213.3
49	6,10E+06	(76)	1,12E+06	(14)	15	203 107	94.8	53.8	182.1
50	7,13E+06	(142)	1,26E+06	(25)	24	226 90	99.5	65.2	159.2
51	2,82E+06	(187)	3,92E+05	(26)	80	71 27	125.8	83.8	197.7
52	8,03E+06	(80)	1,20E+06	(12)	12	217 123	116.0	63.8	234.5
53	8,76E+06	(218)	1,20E+06	(30)	30	217 79	127.2	87.2	193.1

Grain no.	RhoS (cm ⁻²)	(Ns)	RhoI (cm ⁻²)	(Ni)	Squares	U+/-2s	Grain Age (Ma)		
							Age	--95% CI--	
54	1,54E+06	(46)	1,67E+05	(5)	36	30 26	156.9	64.8	506.6
55	3,33E+06	(69)	3,37E+05	(7)	25	61 44	169.2	80.0	436.6
56	2,71E+06	(90)	4,52E+05	(15)	40	81 41	104.7	61.0	195.3
57	4,82E+06	(64)	8,28E+05	(11)	16	149 88	101.3	53.8	213.8
58	2,31E+06	(94)	4,43E+05	(18)	49	80 37	91.4	55.3	161.3
59	3,33E+06	(83)	4,82E+05	(12)	30	87 49	120.3	66.3	242.7
60	3,13E+06	(78)	7,63E+05	(19)	30	138 62	72.0	43.5	126.3
61	2,81E+06	(28)	9,04E+05	(9)	12	163 106	54.3	25.3	131.8
62	7,98E+06	(265)	1,48E+06	(49)	40	266 76	95.1	70.1	131.8
63	3,10E+06	(103)	5,12E+05	(17)	40	92 44	105.9	63.7	189.1
64	3,34E+06	(111)	6,63E+05	(22)	40	119 50	88.4	56.0	147.1
65	7,15E+06	(178)	1,24E+06	(31)	30	224 80	100.7	68.9	152.8
66	6,48E+06	(538)	1,05E+06	(87)	100	189 41	108.3	86.3	135.8
67	3,31E+06	(44)	1,51E+05	(2)	16	27 34	353.9	101.1	2682.8
68	5,37E+06	(312)	1,27E+06	(74)	70	230 53	73.9	57.4	95.3
69	7,15E+06	(178)	1,24E+06	(31)	30	224 80	100.7	68.9	152.8
70	6,48E+06	(538)	1,05E+06	(87)	100	189 41	108.3	86.3	135.8
71	3,31E+06	(44)	1,51E+05	(2)	16	27 34	353.9	101.1	2682.8
72	5,37E+06	(312)	1,27E+06	(74)	70	230 53	73.9	57.4	95.3
73	5,11E+06	(178)	1,15E+06	(40)	42	207 65	78.3	55.5	113.4
74	5,37E+06	(223)	5,30E+05	(22)	50	96 40	176.3	114.8	286.8
75	5,54E+06	(138)	9,64E+05	(24)	30	174 70	100.7	65.5	162.8
76	5,14E+06	(128)	8,03E+05	(20)	30	145 64	111.9	70.2	189.6
77	5,54E+06	(184)	1,11E+06	(37)	40	201 66	87.4	61.4	128.2
78	7,35E+06	(305)	1,01E+06	(42)	50	182 56	127.3	92.4	180.2
79	1,69E+06	(42)	3,61E+05	(9)	30	65 42	81.2	39.6	190.8
80	2,83E+06	(188)	6,02E+05	(40)	80	109 34	82.7	58.7	119.5
81	4,92E+06	(143)	8,26E+05	(24)	35	149 60	104.4	67.9	168.4
82	2,21E+06	(33)	1,34E+05	(2)	18	24 31	267.5	74.5	2127.5
83	7,08E+06	(94)	8,28E+05	(11)	16	149 88	148.1	80.5	307.1
84	5,35E+06	(71)	9,04E+05	(12)	16	163 92	103.1	56.3	209.5
85	9,49E+06	(126)	2,64E+06	(35)	16	475 160	63.4	43.5	95.2
86	8,18E+06	(285)	2,67E+06	(93)	42	481 100	53.9	42.6	68.2
87	7,68E+06	(255)	1,72E+06	(57)	40	309 82	78.3	58.8	104.3
88	2,94E+06	(171)	2,58E+05	(15)	70	47 24	197.4	118.2	359.6
89	7,93E+06	(79)	1,81E+06	(18)	12	326 152	76.9	46.1	136.8
90	6,35E+06	(369)	1,36E+06	(79)	70	245 55	81.9	64.2	104.5
91	2,82E+06	(164)	4,48E+05	(26)	70	81 31	110.5	73.3	174.3
92	3,90E+06	(97)	4,82E+05	(12)	30	87 49	140.3	78.0	281.3
93	5,79E+06	(202)	1,69E+06	(59)	42	305 79	60.0	44.9	80.2
94	8,19E+06	(204)	2,17E+06	(54)	30	391 106	66.6	49.2	91.8
95	6,53E+06	(271)	1,42E+06	(59)	50	256 67	80.4	60.7	106.5
96	5,23E+06	(152)	9,64E+05	(28)	35	174 65	95.2	63.7	148.3
97	5,49E+06	(114)	1,35E+06	(28)	25	243 91	71.6	47.3	112.7
POOLED	5,30E+06	(17574)	9,99E+05	(3314)	3996	180 8	93.5	88.7	98.5

CHI^2 PROBABILITY (%): 0.0

POOLED AGE W/ 68% CONF. INTERVAL(Ma): 93.5, 91.0 -- 96.0 (-2.5 +2.5)
 95% CONF. INTERVAL(Ma): 88.7 -- 98.5 (-4.8 +5.0)

CENTRAL AGE W/ 68% CONF. INTERVAL(Ma): 95.7, 92.2 -- 99.3 (-3.4 +3.6)
 95% CONF. INTERVAL(Ma): 89.1 -- 102.8 (-6.6 +7.1)
 AGE DISPERSION (%): 23.2

FIT OPTION: Best-fit peaks using the binomial model of Galbraith and Green

INITIAL GUESS FOR MODEL PARAMETERS (number of peaks to fit = 3)

Peak #.	Peak Age	Theta	Fraction(%)	Count
1.	93.50	0.841	41.1	39.89
2.	109.40	0.861	41.8	40.50
3.	139.60	0.888	18.9	18.33

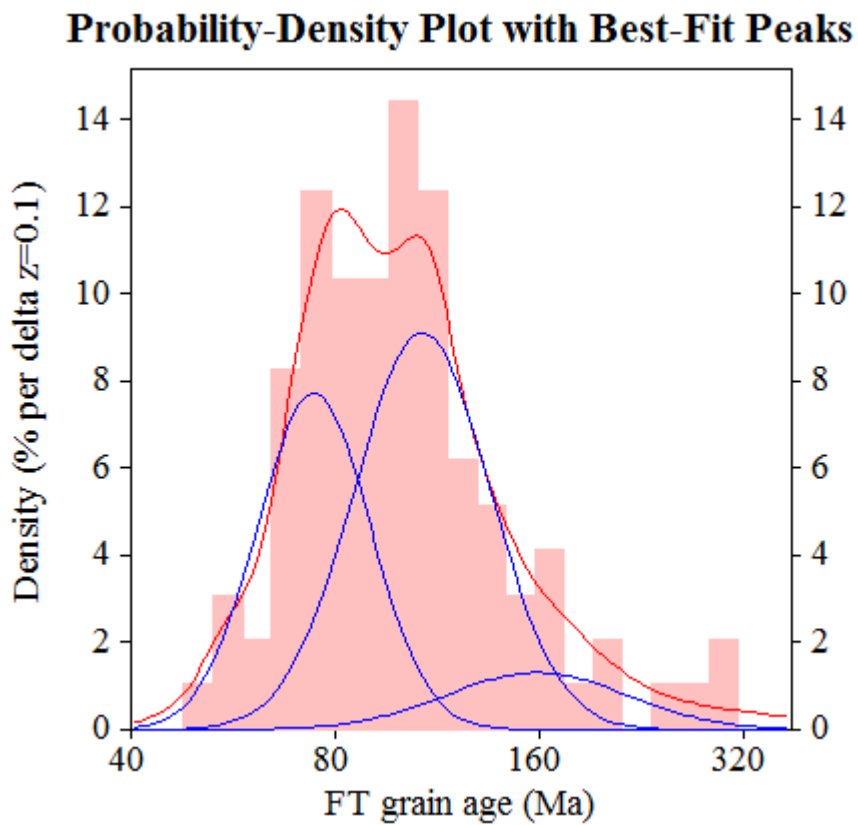
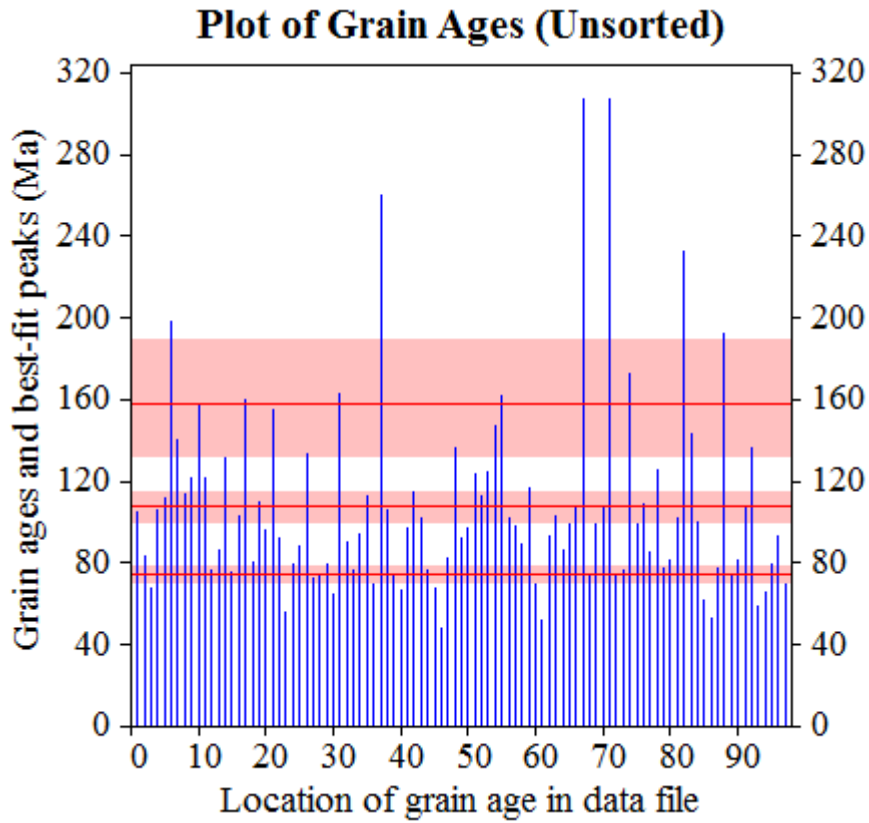
Total range for grain ages: 49,1 to 308,6 Ma
 Number of active grains (Num. used for fit): 97
 Number of removed grains: 0
 Degrees of freedom for fit: 92
 Average of the SE(Z)'s for the grains: 0,25
 Estimated width of peaks in PD plot in Z units: 0,3

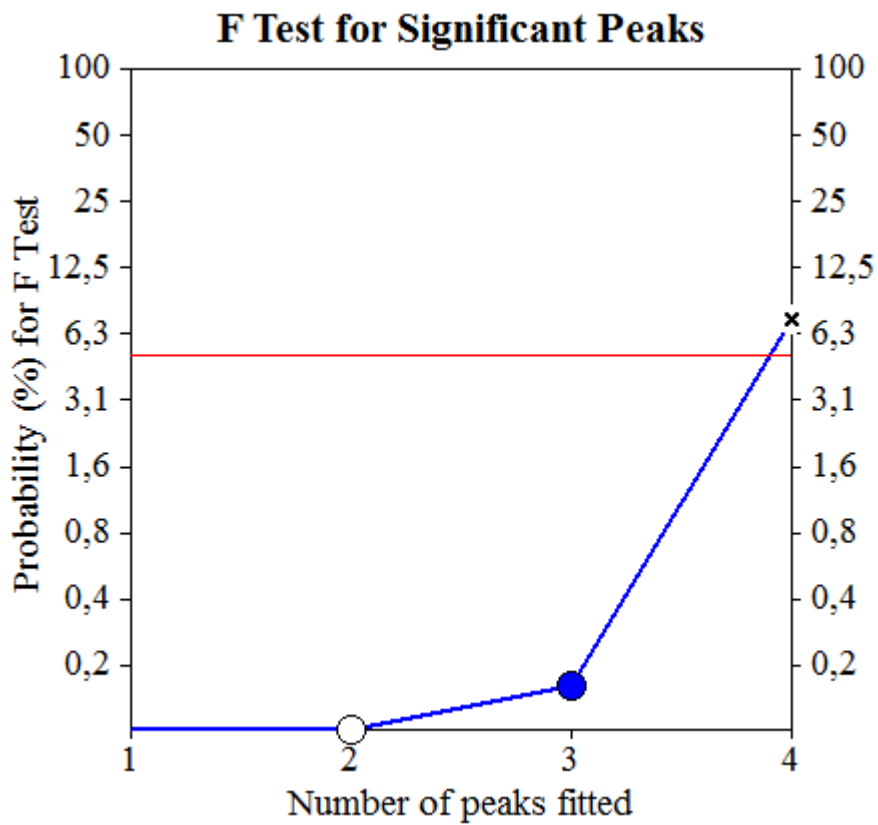
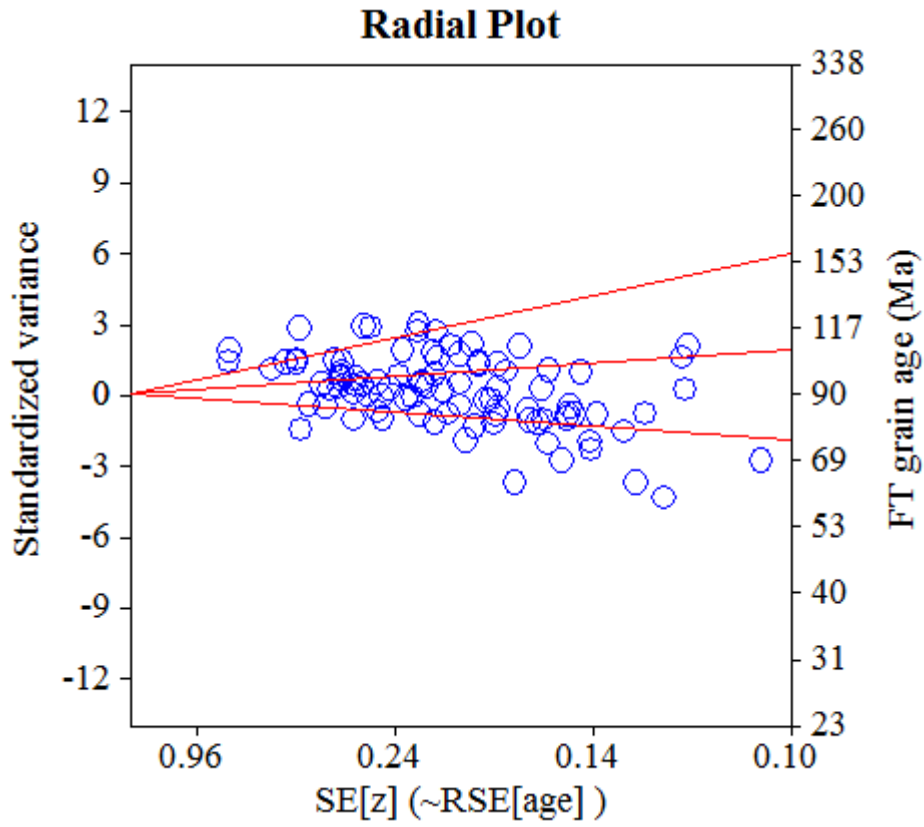
PARAMETERS FOR BEST-FIT PEAKS

- * Standard error for peak age includes group error
- * Peak width is for PD plot assuming a kernel factor = 0.60

#.	Peak Age(Ma)	68%CI	95%CI	W(Z)	Frac(%)	SE,%	Count
1.	74.4	-3,5 ...+3,7	-6,7 ...+7,4	0.19	37.1	10.7	35.9
2.	107.4	-7,2 ...+7,7	-13,6 ...+15,6	0.23	52.8	11.5	51.3
3.	158.4	-25,8 ...+30,7	-46,6 ...+65,6	0.31	10.1	9.0	9.8

Log-likelihood for best fit: -331,339
 Chi-squared value for best fit: 97,246
 Reduced chi-squared value: 1,057
 Probability for F test: 0%
 Condition number for COVAR matrix: 41,65
 Number of iterations: 22





Datafile: C:\Users\Administrador\Documents\PhD\Termocronologia resultados\Z\12-2015\EMP_35b\EMP_35b_12_2015_21.ftz
 Title: Sample No. EMP_35b Irr 12-2015-21

NEW PARAMETERS - ZETA METHOD

EFFECTIVE TRACK DENSITY FOR FLUENCE MONITOR (tracks/cm²): 2.76E+05
 RELATIVE ERROR (%): 1.22
 EFFECTIVE URANIUM CONTENT OF MONITOR (ppm): 50.00
 ZETA FACTOR AND STANDARD ERROR (yr cm²): 128.02 1.87
 SIZE OF COUNTER SQUARE (cm²): 8.30E-07

GRAIN AGES IN ORIGINAL ORDER

Grain no.	RhoS (cm ⁻²)	(Ns)	RhoI (cm ⁻²)	(Ni)	Squares	U+/-2s	Grain Age (Ma)		
							Age	--95% CI--	
1	2.61E+06	(26)	2.01E+05	(2)	12	36 46	211.0	57.3	1740.3
2	3.64E+06	(121)	1.51E+05	(5)	40	27 23	402.0	175.8	1218.6
3	5.70E+06	(142)	3.21E+05	(8)	30	58 40	300.7	153.0	701.5
4	5.22E+06	(39)	2.68E+05	(2)	9	48 61	313.3	88.6	2426.9
5	5.78E+06	(96)	5.42E+05	(9)	20	98 64	182.9	94.5	411.5
6	5.76E+06	(43)	1.20E+06	(9)	9	218 142	82.7	40.5	194.1
7	6.96E+06	(52)	1.20E+06	(9)	9	218 142	99.8	49.6	231.6
8	3.77E+06	(50)	4.52E+05	(6)	16	82 64	142.3	62.8	407.4
9	4.46E+06	(37)	3.61E+05	(3)	10	65 70	204.6	68.6	1020.2
10	7.71E+06	(192)	1.29E+06	(32)	30	233 82	104.7	72.2	157.6
11	6.96E+06	(52)	9.37E+05	(7)	9	170 124	127.4	59.1	333.9
12	5.34E+06	(93)	9.75E+05	(17)	21	177 85	95.2	57.0	170.8
13	3.09E+06	(77)	4.42E+05	(11)	30	80 47	121.0	65.1	253.2
14	6.27E+06	(156)	7.63E+05	(19)	30	138 63	142.4	89.2	243.0
15	4.96E+06	(247)	5.02E+05	(25)	60	91 36	171.3	114.4	269.6
16	6.57E+06	(229)	1.41E+06	(49)	42	255 73	81.9	60.1	114.0
17	5.22E+06	(52)	1.10E+06	(11)	12	200 118	82.1	42.9	175.3
18	3.09E+06	(77)	4.42E+05	(11)	30	80 47	121.0	65.1	253.2
19	5.14E+06	(427)	8.43E+05	(70)	100	153 37	106.2	82.5	136.6
20	3.42E+06	(142)	3.37E+05	(14)	50	61 32	175.0	102.5	328.0
21	8.70E+06	(65)	5.35E+05	(4)	9	97 91	271.0	106.0	1003.6
22	2.69E+06	(67)	6.83E+05	(17)	30	124 59	68.8	40.3	125.4
23	4.12E+06	(171)	4.34E+05	(18)	50	79 37	164.4	102.2	283.9
24	3.84E+06	(134)	4.88E+05	(17)	42	88 42	136.7	83.2	241.8
25	2.80E+06	(58)	6.27E+05	(13)	25	113 62	77.6	42.5	155.1
26	7.63E+06	(57)	1.07E+06	(8)	9	194 133	122.6	59.5	298.7
27	5.87E+06	(78)	1.28E+06	(17)	16	232 111	80.0	47.3	144.7
28	2.91E+06	(58)	2.01E+05	(4)	24	36 34	242.4	94.1	905.7
29	4.29E+06	(57)	8.28E+05	(11)	16	150 89	89.9	47.3	190.9
30	1.61E+06	(40)	1.61E+05	(4)	30	29 27	168.4	63.6	646.8
31	2.09E+06	(26)	1.61E+05	(2)	15	29 37	211.0	57.3	1740.3
32	2.13E+06	(53)	1.61E+05	(4)	30	29 27	222.0	85.7	834.8
33	5.85E+06	(102)	1.09E+06	(19)	21	197 90	93.5	57.5	162.1
34	8.99E+06	(179)	1.56E+06	(31)	24	282 101	100.8	69.0	153.0
35	5.92E+06	(59)	5.02E+05	(5)	12	91 78	199.5	83.9	633.9
36	2.77E+06	(46)	6.02E+05	(10)	20	109 67	79.8	40.3	178.3
37	5.66E+06	(169)	1.04E+06	(31)	36	188 67	95.2	65.0	144.7
38	4.30E+06	(107)	3.61E+05	(9)	30	65 43	203.4	105.7	455.7
39	7.53E+06	(75)	1.31E+06	(13)	12	236 129	100.1	55.9	197.4
40	4.15E+06	(62)	2.68E+05	(4)	18	48 46	258.8	100.9	961.8
41	3.82E+06	(57)	3.35E+05	(5)	18	61 52	192.8	80.9	614.2
42	4.19E+06	(327)	4.74E+05	(37)	94	86 28	153.7	109.9	222.1
43	1.93E+06	(32)	3.01E+05	(5)	20	55 47	109.2	43.6	361.6
44	1.00E+07	(75)	2.01E+06	(15)	9	364 185	87.0	50.1	163.7
45	2.71E+06	(36)	3.77E+05	(5)	16	68 58	122.6	49.6	402.7
46	2.29E+06	(95)	1.93E+05	(8)	50	35 24	202.8	101.4	481.9
47	5.16E+06	(214)	7.95E+05	(33)	50	144 50	113.1	78.6	168.7
48	3.77E+06	(50)	1.51E+05	(2)	16	27 35	398.7	115.0	2955.6
49	6.93E+06	(115)	4.22E+05	(7)	20	76 56	278.1	135.0	699.3
50	4.42E+06	(33)	1.34E+05	(1)	9	24 40	492.1	96.9	9950.2
51	3.68E+06	(110)	4.35E+05	(13)	36	79 43	146.2	83.3	283.6
52	2.33E+06	(116)	1.81E+05	(9)	60	33 21	220.2	114.8	491.7
53	8.51E+06	(113)	9.04E+05	(12)	16	164 93	162.4	90.9	323.6

Grain no.	RhoS (cm ⁻²)	(Ns)	RhoI (cm ⁻²)	(Ni)	Squares	U+/-2s	Grain Age (Ma)		
							Age	--95% CI--	
54	3.35E+06	(25)	1.47E+06	(11)	9	267 157	39.7	19.0	89.9
55	1.81E+06	(60)	2.41E+05	(8)	40	44 30	129.0	62.8	313.4
56	5.06E+06	(294)	4.82E+05	(28)	70	87 33	182.0	124.4	278.3
57	7.52E+06	(624)	1.28E+06	(106)	100	231 45	102.7	83.5	126.4
58	9.42E+06	(391)	1.40E+06	(58)	50	253 66	117.1	88.9	154.0
59	7.36E+06	(55)	6.69E+05	(5)	9	121 103	186.2	77.9	594.3
60	5.66E+06	(188)	7.23E+05	(24)	40	131 53	136.2	89.6	217.9
61	3.68E+06	(214)	6.37E+05	(37)	70	115 38	101.1	71.4	147.5
62	5.31E+06	(216)	7.62E+05	(31)	49	138 49	121.4	83.6	183.2
63	8.63E+06	(86)	1.10E+06	(11)	12	200 118	135.0	73.0	281.0
64	2.73E+06	(181)	5.42E+05	(36)	80	98 33	87.9	61.5	129.7
65	8.16E+06	(271)	7.83E+05	(26)	40	142 55	180.6	121.7	281.1
66	4.28E+06	(284)	3.61E+05	(24)	80	65 26	204.6	136.2	323.8
67	4.18E+06	(52)	5.62E+05	(7)	15	102 74	127.4	59.1	333.9
68	2.69E+06	(134)	3.41E+05	(17)	60	62 30	136.7	83.2	241.8
69	2.41E+06	(50)	1.45E+05	(3)	25	26 28	274.7	94.4	1329.8
70	7.53E+06	(250)	1.27E+06	(42)	40	229 71	104.0	75.1	148.0
71	3.73E+06	(93)	4.02E+05	(10)	30	73 45	160.0	84.9	344.8
72	4.02E+06	(30)	8.03E+05	(6)	9	145 114	85.9	36.1	254.7
73	4.82E+06	(280)	4.65E+05	(27)	70	84 32	179.8	122.0	277.3
74	4.63E+06	(192)	5.30E+05	(22)	50	96 41	151.4	98.2	247.3
75	3.03E+06	(151)	3.61E+05	(18)	60	65 30	145.4	90.0	252.1
76	4.93E+06	(368)	5.49E+05	(41)	90	99 31	156.2	113.6	221.1
77	3.18E+06	(264)	1.02E+06	(85)	100	185 40	54.4	42.5	69.5
78	4.88E+06	(405)	8.80E+05	(73)	100	159 37	96.7	75.3	124.0
79	7.92E+06	(657)	1.27E+06	(105)	100	229 45	109.1	88.7	134.2
80	6.46E+06	(193)	6.36E+05	(19)	36	115 52	175.7	110.8	297.9
81	6.70E+06	(139)	1.01E+06	(21)	25	183 79	115.2	73.1	192.2
82	3.37E+06	(28)	1.20E+06	(10)	10	218 135	48.8	23.3	113.2
83	6.31E+06	(262)	9.64E+05	(40)	50	175 55	114.3	82.2	163.8
84	2.37E+06	(63)	4.52E+05	(12)	32	82 46	91.1	49.4	186.4
85	6.59E+06	(547)	1.41E+06	(117)	100	255 48	81.8	66.8	100.0
86	3.13E+06	(260)	4.70E+05	(39)	100	85 27	116.4	83.3	167.5
87	3.04E+06	(252)	6.39E+05	(53)	100	116 32	83.3	61.9	114.4
88	5.30E+06	(220)	1.16E+06	(48)	50	209 60	80.3	58.7	112.3
89	7.08E+06	(147)	1.73E+06	(36)	25	314 104	71.5	49.6	106.2
90	5.35E+06	(160)	6.69E+05	(20)	36	121 54	138.9	87.9	233.6
91	6.18E+06	(154)	6.02E+05	(15)	30	109 55	177.2	105.7	324.0
92	6.67E+06	(554)	1.02E+06	(85)	100	185 40	113.5	90.3	142.7
93	3.36E+06	(209)	4.18E+05	(26)	75	76 29	139.8	93.5	219.0
94	6.73E+06	(335)	1.08E+06	(54)	60	196 53	107.7	80.9	143.4
95	5.88E+06	(244)	8.67E+05	(36)	50	157 52	118.2	83.6	172.9
96	3.15E+06	(157)	3.61E+05	(18)	60	65 30	151.1	93.7	261.7
97	7.43E+06	(259)	1.12E+06	(39)	42	203 65	115.9	83.0	166.8
98	7.45E+06	(99)	1.20E+06	(16)	16	218 108	107.5	63.8	195.9
99	6.96E+06	(52)	1.47E+06	(11)	9	267 157	82.1	42.9	175.3
POOLED	4.85E+06	(15535)	7.15E+05	(2288)	3856	129 6	118.9	112.3	125.8

CHI^2 PROBABILITY (%): 0.0

POOLED AGE W/ 68% CONF. INTERVAL(Ma): 118.9, 115.5 -- 122.4 (-3.4 +3.5)
 95% CONF. INTERVAL(Ma): 112.3 -- 125.8 (-6.6 +7.0)

CENTRAL AGE W/ 68% CONF. INTERVAL(Ma): 123.5, 118.4 -- 128.9 (-5.1 +5.3)
 95% CONF. INTERVAL(Ma): 113.7 -- 134.2 (-9.8 +10.7)
 AGE DISPERSION (%): 27.3

FIT OPTION: Best-fit peaks using the binomial model of Galbraith and Green

INITIAL GUESS FOR MODEL PARAMETERS (number of peaks to fit = 3)

Peak #.	Peak Age	Theta	Fraction(%)	Count
1.	39.10	0.690	1.2	1.18
2.	118.90	0.872	35.8	35.43
3.	311.20	0.947	5.2	5.12

Total range for grain ages: 39.1 to 383.1 Ma
 Number of active grains (Num. used for fit): 99
 Number of removed grains: 0
 Degrees of freedom for fit: 94
 Average of the SE(Z)'s for the grains: 0.33
 Estimated width of peaks in PD plot in Z units: 0.39

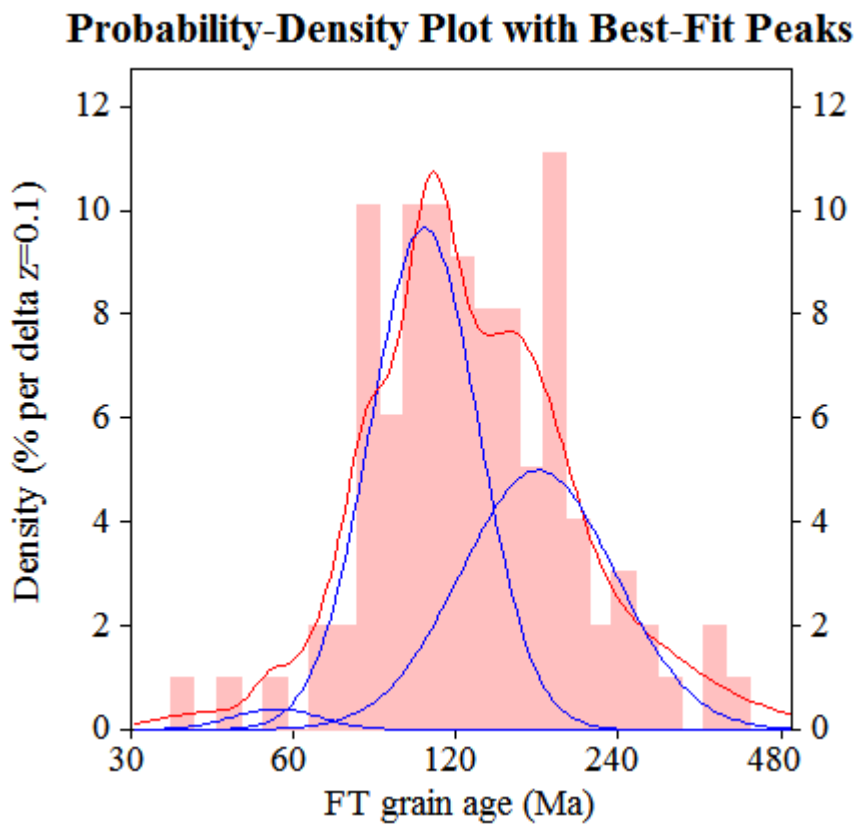
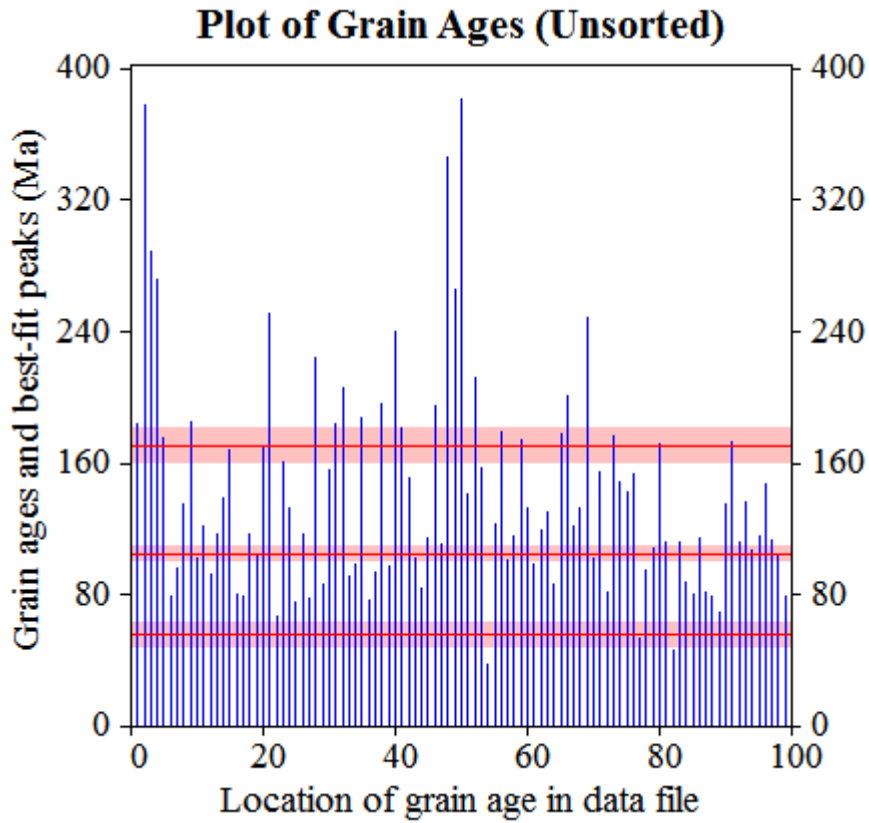
PARAMETERS FOR BEST-FIT PEAKS

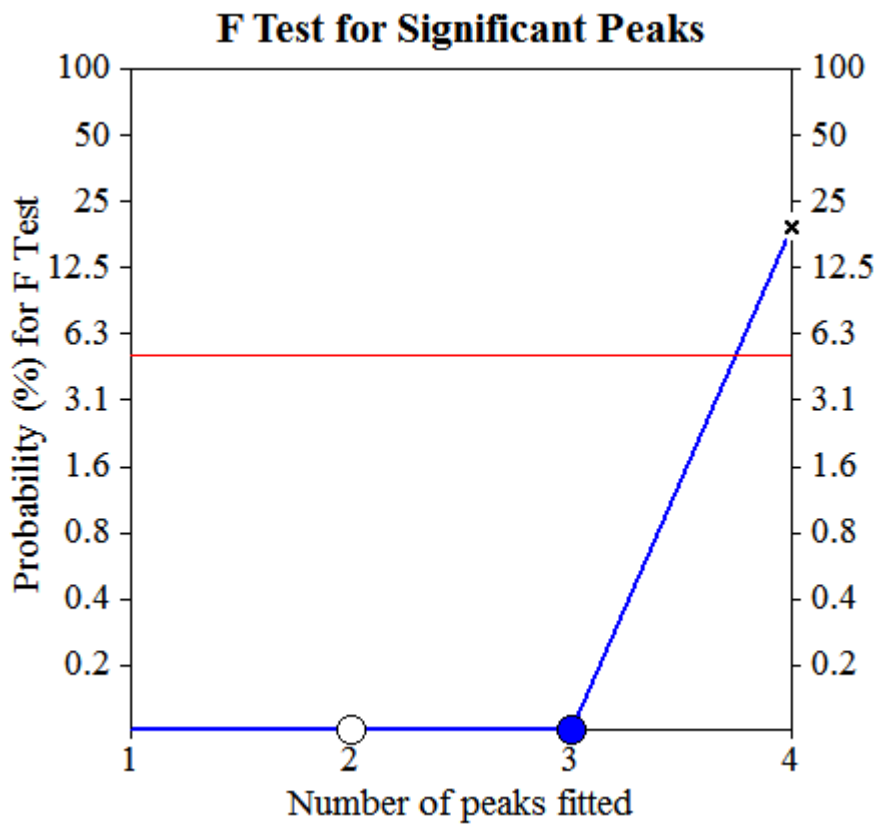
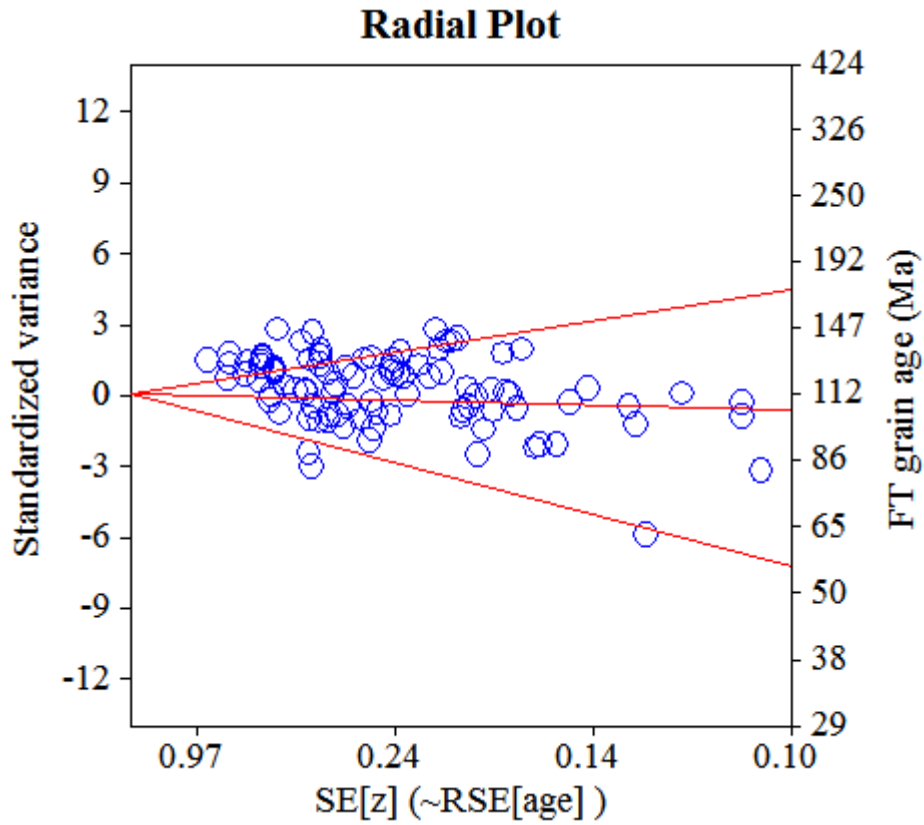
* Standard error for peak age includes group error

* Peak width is for PD plot assuming a kernel factor = 0.60

#.	Peak Age(Ma)	68%CI	95%CI	W(Z)	Frac(%)	SE,%	Count
1.	55.6	-7.1 ...+8.1	-13.0 ...+16.9	0.18	1.8	1.8	1.8
2.	104.9	-4.0 ...+4.2	-7.7 ...+8.3	0.23	56.4	8.4	55.8
3.	170.5	-10.3 ...+10.9	-19.6 ...+22.1	0.33	41.7	8.3	41.3

Log-likelihood for best fit: -306.293
 Chi-squared value for best fit: 103.646
 Reduced chi-squared value: 1.103
 Probability for F test: 0%
 Condition number for COVAR matrix: 64.72
 Number of iterations: 18





NEW PARAMETERS - ZETA METHOD

EFFECTIVE TRACK DENSITY FOR FLUENCE MONITOR (tracks/cm²): 2,77E+05
 RELATIVE ERROR (%): 1,23
 EFFECTIVE URANIUM CONTENT OF MONITOR (ppm): 50,00
 ZETA FACTOR AND STANDARD ERROR (yr cm²): 128,02 1,87
 SIZE OF COUNTER SQUARE (cm²): 8,30E-07

GRAIN AGES IN ORIGINAL ORDER

Grain no.	RhoS (cm ⁻²)	(Ns)	RhoI (cm ⁻²)	(Ni)	Squares	U+/-2s	Grain Age (Ma)		
							Age	--95% CI--	
1	7,28E+06	(302)	1,20E+06	(50)	50	217 61	106.1	78.7	146.2
2	6,60E+06	(274)	1,37E+06	(57)	50	248 66	84.1	63.2	111.8
3	5,76E+06	(478)	1,48E+06	(123)	100	267 49	68.4	56.0	83.5
4	3,80E+06	(63)	6,02E+05	(10)	20	109 67	109.4	56.7	240.1
5	6,07E+06	(247)	9,34E+05	(38)	49	168 54	114.0	81.2	165.0
6	4,73E+06	(165)	4,02E+05	(14)	42	72 38	203.8	120.0	380.0
7	4,03E+06	(164)	4,92E+05	(20)	49	89 39	143.0	90.6	240.3
8	9,04E+06	(225)	1,37E+06	(34)	30	246 84	116.0	81.1	171.7
9	6,55E+06	(87)	9,04E+05	(12)	16	163 92	126.0	69.7	253.8
10	5,54E+06	(230)	6,02E+05	(25)	50	109 43	160.4	106.9	252.9
11	5,24E+06	(87)	7,23E+05	(12)	20	130 74	126.0	69.7	253.8
12	8,07E+06	(134)	1,81E+06	(30)	20	326 118	78.5	52.8	121.1
13	7,67E+06	(191)	1,53E+06	(38)	30	275 89	88.3	62.4	128.8
14	5,75E+06	(191)	7,53E+05	(25)	40	136 54	133.5	88.5	211.6
15	7,23E+06	(210)	1,65E+06	(48)	35	298 86	77.0	56.2	107.8
16	8,29E+06	(344)	1,40E+06	(58)	50	252 66	103.6	78.5	136.7
17	2,37E+06	(59)	2,41E+05	(6)	30	43 34	168.3	75.1	477.0
18	6,60E+06	(219)	1,42E+06	(47)	40	255 74	82.0	59.8	115.0
19	2,74E+06	(91)	4,22E+05	(14)	40	76 40	113.3	65.0	215.9
20	8,06E+06	(107)	1,43E+06	(19)	16	258 117	98.5	60.7	170.4
21	7,95E+06	(66)	8,43E+05	(7)	10	152 111	162.0	76.4	418.9
22	5,73E+06	(476)	1,08E+06	(90)	100	195 41	92.7	73.9	116.2
23	4,56E+06	(265)	1,41E+06	(82)	70	254 56	56.8	44.3	72.8
24	3,93E+06	(261)	8,58E+05	(57)	80	155 41	80.1	60.2	106.6
25	7,11E+06	(118)	1,39E+06	(23)	20	250 103	89.9	57.6	147.7
26	3,32E+06	(248)	4,28E+05	(32)	90	77 27	135.6	94.2	202.6
27	4,82E+06	(168)	1,15E+06	(40)	42	207 65	73.9	52.3	107.3
28	8,76E+06	(109)	2,01E+06	(25)	15	362 144	76.6	49.5	123.7
29	7,20E+06	(299)	1,57E+06	(65)	50	282 70	80.6	61.6	105.3
30	7,44E+06	(247)	1,99E+06	(66)	40	358 88	65.6	50.0	86.1
31	4,22E+06	(210)	4,42E+05	(22)	60	80 34	166.2	108.0	270.7
32	6,40E+06	(85)	1,20E+06	(16)	16	217 107	92.9	54.6	170.3
33	6,69E+06	(50)	1,47E+06	(11)	9	265 157	79.3	41.3	169.8
34	8,55E+06	(213)	1,57E+06	(39)	30	282 90	95.9	68.3	138.8
35	6,90E+06	(573)	1,06E+06	(88)	100	191 41	114.0	91.0	142.7
36	7,04E+06	(146)	1,73E+06	(36)	25	313 104	71.4	49.5	106.0
37	4,50E+06	(112)	2,81E+05	(7)	30	51 37	272.2	132.0	685.4
38	5,93E+06	(246)	9,64E+05	(40)	50	174 55	107.9	77.4	154.9
39	6,51E+06	(216)	1,51E+06	(50)	40	271 77	76.1	55.9	105.8
40	5,92E+06	(59)	1,51E+06	(15)	12	271 138	68.9	39.0	131.3
41	3,06E+06	(142)	5,38E+05	(25)	56	97 38	99.5	65.2	159.2
42	9,12E+06	(227)	1,37E+06	(34)	30	246 84	117.0	81.8	173.2
43	4,05E+06	(84)	6,75E+05	(14)	25	122 64	104.7	59.8	200.2
44	5,53E+06	(225)	1,25E+06	(51)	49	226 63	77.7	57.3	107.6
45	6,14E+06	(255)	1,57E+06	(65)	50	282 70	68.8	52.4	90.3
46	4,13E+06	(137)	1,48E+06	(49)	40	266 76	49.4	35.5	70.0
47	2,30E+06	(153)	4,82E+05	(32)	80	87 31	84.0	57.4	127.3
48	3,34E+06	(222)	4,22E+05	(28)	80	76 29	138.6	94.0	213.3
49	6,10E+06	(76)	1,12E+06	(14)	15	203 107	94.8	53.8	182.1
50	7,13E+06	(142)	1,26E+06	(25)	24	226 90	99.5	65.2	159.2
51	2,82E+06	(187)	3,92E+05	(26)	80	71 27	125.8	83.8	197.7
52	8,03E+06	(80)	1,20E+06	(12)	12	217 123	116.0	63.8	234.5
53	8,76E+06	(218)	1,20E+06	(30)	30	217 79	127.2	87.2	193.1

Datafile: C:\BH2\Edna\12-2015\EMP_49\EMP_49_12_2015_22.ftz

Title: Sample No. EMP_49 Irr 12-2015-22

Grain no.	RhoS (cm ⁻²)	(Ns)	RhoI (cm ⁻²)	(Ni)	Squares	U+/-2s		Grain Age (Ma)		
								Age	--95% CI--	
54	1,54E+06	(46)	1,67E+05	(5)	36	30	26	156.9	64.8	506.6
55	3,33E+06	(69)	3,37E+05	(7)	25	61	44	169.2	80.0	436.6
56	2,71E+06	(90)	4,52E+05	(15)	40	81	41	104.7	61.0	195.3
57	4,82E+06	(64)	8,28E+05	(11)	16	149	88	101.3	53.8	213.8
58	2,31E+06	(94)	4,43E+05	(18)	49	80	37	91.4	55.3	161.3
59	3,33E+06	(83)	4,82E+05	(12)	30	87	49	120.3	66.3	242.7
60	3,13E+06	(78)	7,63E+05	(19)	30	138	62	72.0	43.5	126.3
61	2,81E+06	(28)	9,04E+05	(9)	12	163	106	54.3	25.3	131.8
62	7,98E+06	(265)	1,48E+06	(49)	40	266	76	95.1	70.1	131.8
63	3,10E+06	(103)	5,12E+05	(17)	40	92	44	105.9	63.7	189.1
64	3,34E+06	(111)	6,63E+05	(22)	40	119	50	88.4	56.0	147.1
65	7,15E+06	(178)	1,24E+06	(31)	30	224	80	100.7	68.9	152.8
66	6,48E+06	(538)	1,05E+06	(87)	100	189	41	108.3	86.3	135.8
67	3,31E+06	(44)	1,51E+05	(2)	16	27	34	353.9	101.1	2682.8
68	5,37E+06	(312)	1,27E+06	(74)	70	230	53	73.9	57.4	95.3
69	7,15E+06	(178)	1,24E+06	(31)	30	224	80	100.7	68.9	152.8
70	6,48E+06	(538)	1,05E+06	(87)	100	189	41	108.3	86.3	135.8
71	3,31E+06	(44)	1,51E+05	(2)	16	27	34	353.9	101.1	2682.8
72	5,37E+06	(312)	1,27E+06	(74)	70	230	53	73.9	57.4	95.3
73	5,11E+06	(178)	1,15E+06	(40)	42	207	65	78.3	55.5	113.4
74	5,37E+06	(223)	5,30E+05	(22)	50	96	40	176.3	114.8	286.8
75	5,54E+06	(138)	9,64E+05	(24)	30	174	70	100.7	65.5	162.8
76	5,14E+06	(128)	8,03E+05	(20)	30	145	64	111.9	70.2	189.6
77	5,54E+06	(184)	1,11E+06	(37)	40	201	66	87.4	61.4	128.2
78	7,35E+06	(305)	1,01E+06	(42)	50	182	56	127.3	92.4	180.2
79	1,69E+06	(42)	3,61E+05	(9)	30	65	42	81.2	39.6	190.8
80	2,83E+06	(188)	6,02E+05	(40)	80	109	34	82.7	58.7	119.5
81	4,92E+06	(143)	8,26E+05	(24)	35	149	60	104.4	67.9	168.4
82	2,21E+06	(33)	1,34E+05	(2)	18	24	31	267.5	74.5	2127.5
83	7,08E+06	(94)	8,28E+05	(11)	16	149	88	148.1	80.5	307.1
84	5,35E+06	(71)	9,04E+05	(12)	16	163	92	103.1	56.3	209.5
85	9,49E+06	(126)	2,64E+06	(35)	16	475	160	63.4	43.5	95.2
86	8,18E+06	(285)	2,67E+06	(93)	42	481	100	53.9	42.6	68.2
87	7,68E+06	(255)	1,72E+06	(57)	40	309	82	78.3	58.8	104.3
88	2,94E+06	(171)	2,58E+05	(15)	70	47	24	197.4	118.2	359.6
89	7,93E+06	(79)	1,81E+06	(18)	12	326	152	76.9	46.1	136.8
90	6,35E+06	(369)	1,36E+06	(79)	70	245	55	81.9	64.2	104.5
91	2,82E+06	(164)	4,48E+05	(26)	70	81	31	110.5	73.3	174.3
92	3,90E+06	(97)	4,82E+05	(12)	30	87	49	140.3	78.0	281.3
93	5,79E+06	(202)	1,69E+06	(59)	42	305	79	60.0	44.9	80.2
94	8,19E+06	(204)	2,17E+06	(54)	30	391	106	66.6	49.2	91.8
95	6,53E+06	(271)	1,42E+06	(59)	50	256	67	80.4	60.7	106.5
96	5,23E+06	(152)	9,64E+05	(28)	35	174	65	95.2	63.7	148.3
97	5,49E+06	(114)	1,35E+06	(28)	25	243	91	71.6	47.3	112.7
POOLED	5,30E+06	(17574)	9,99E+05	(3314)	3996	180	8	93.5	88.7	98.5

CHI^2 PROBABILITY (%): 0.0

POOLED AGE W/ 68% CONF. INTERVAL(Ma): 93.5, 91.0 -- 96.0 (-2.5 +2.5)
 95% CONF. INTERVAL(Ma): 88.7 -- 98.5 (-4.8 +5.0)

CENTRAL AGE W/ 68% CONF. INTERVAL(Ma): 95.7, 92.2 -- 99.3 (-3.4 +3.6)
 95% CONF. INTERVAL(Ma): 89.1 -- 102.8 (-6.6 +7.1)
 AGE DISPERSION (%): 23.2

FIT OPTION: Best-fit peaks using the binomial model of Galbraith and Green

INITIAL GUESS FOR MODEL PARAMETERS (number of peaks to fit = 3)

Peak #.	Peak Age	Theta	Fraction(%)	Count
1.	93.50	0.841	41.1	39.89
2.	109.40	0.861	41.8	40.50
3.	139.60	0.888	18.9	18.33

Total range for grain ages: 49,1 to 308,6 Ma
 Number of active grains (Num. used for fit): 97
 Number of removed grains: 0
 Degrees of freedom for fit: 92
 Average of the SE(Z)'s for the grains: 0,25
 Estimated width of peaks in PD plot in Z units: 0,3

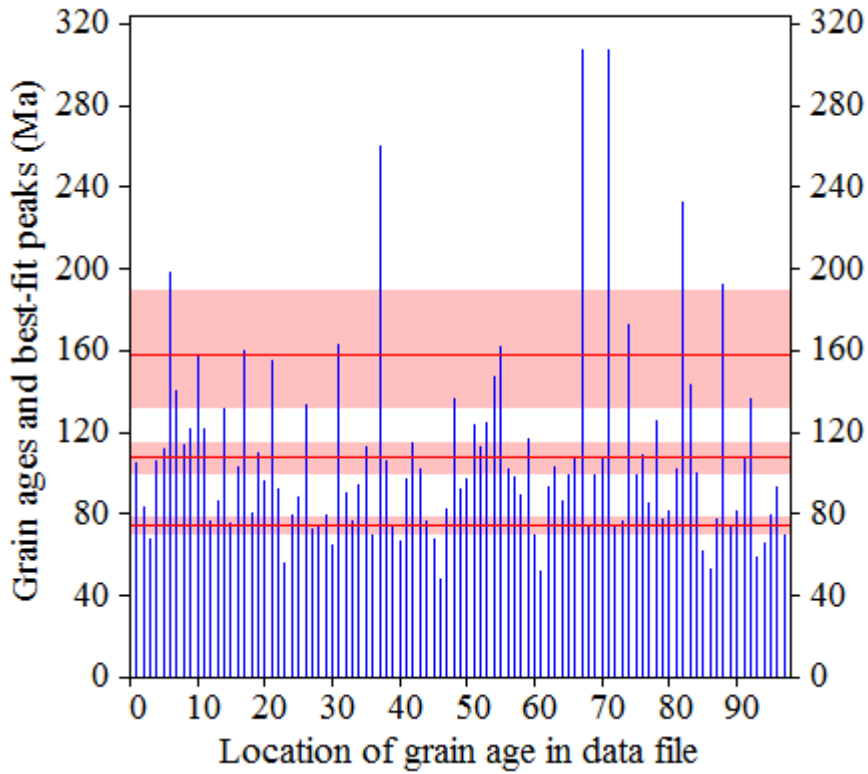
PARAMETERS FOR BEST-FIT PEAKS

- * Standard error for peak age includes group error
- * Peak width is for PD plot assuming a kernel factor = 0.60

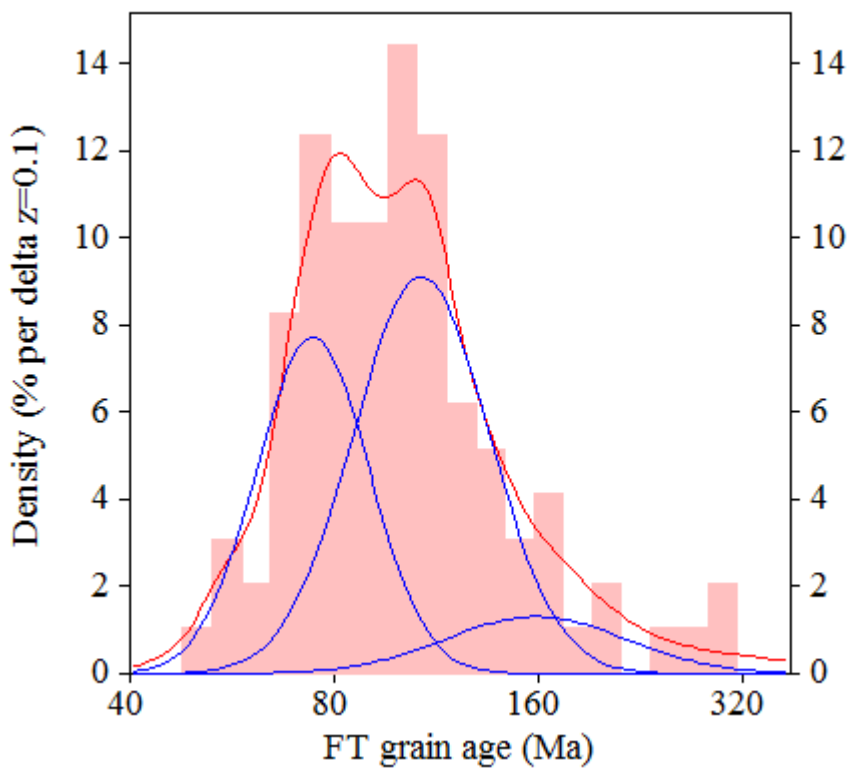
#.	Peak Age(Ma)	68%CI	95%CI	W(Z)	Frac(%)	SE,%	Count
1.	74.4	-3,5 ...+3,7	-6,7 ...+7,4	0.19	37.1	10.7	35.9
2.	107.4	-7,2 ...+7,7	-13,6 ...+15,6	0.23	52.8	11.5	51.3
3.	158.4	-25,8 ...+30,7	-46,6 ...+65,6	0.31	10.1	9.0	9.8

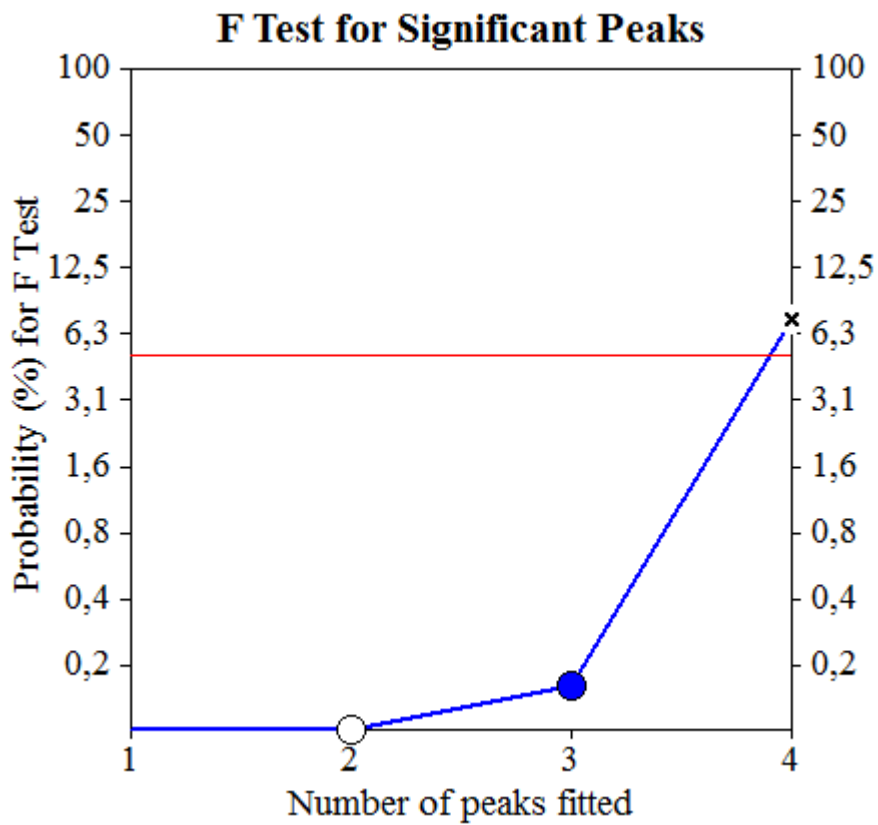
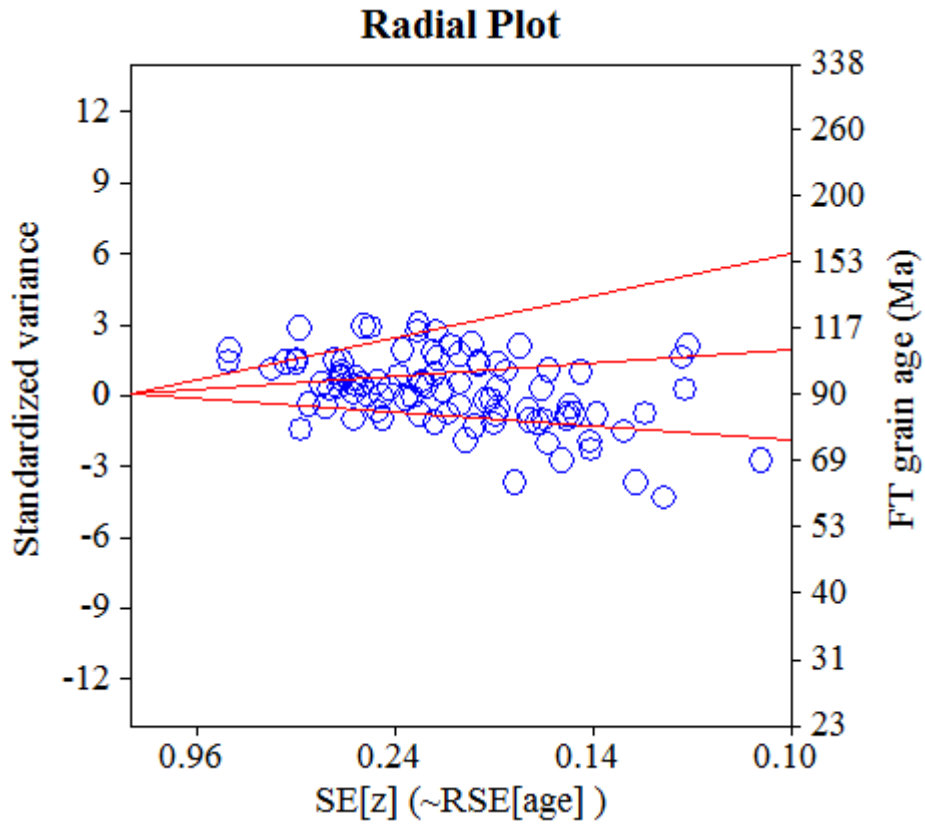
Log-likelihood for best fit: -331,339
 Chi-squared value for best fit: 97,246
 Reduced chi-squared value: 1,057
 Probability for F test: 0%
 Condition number for COVAR matrix: 41,65
 Number of iterations: 22

Plot of Grain Ages (Unsorted)



Probability-Density Plot with Best-Fit Peaks





Datafile: C:\BH2\Alejandro\AFT\11-2015\AP_045\AP_045.ftz

Title: AP_045 Irr 11-2015-9

NEW PARAMETERS - ZETA METHOD

EFFECTIVE TRACK DENSITY FOR FLUENCE MONITOR (tracks/cm²): 1,01E+06
 RELATIVE ERROR (%): 1,16
 EFFECTIVE URANIUM CONTENT OF MONITOR (ppm): 15,00
 ZETA FACTOR AND STANDARD ERROR (yr cm²): 284,52 5,65
 SIZE OF COUNTER SQUARE (cm²): 6,39E-07

GRAIN AGES IN ORIGINAL ORDER

Grain no.	RhoS (cm ⁻²)	(Ns)	RhoI (cm ⁻²)	(Ni)	Squares	U+/-2s	Grain Age (Ma)		
							Age	--95% CI--	
1	1,56E+05	(6)	3,13E+05	(12)	60	5 3	72.3	22.0	203.3
2	9,39E+04	(3)	2,19E+05	(7)	50	3 2	63.0	10.2	263.8
3	1,88E+05	(12)	5,01E+05	(32)	100	7 3	54.0	25.2	106.4
4	2,35E+05	(15)	7,36E+05	(47)	100	11 3	45.9	23.7	82.8
5	1,25E+05	(4)	1,56E+05	(5)	50	2 2	114.6	22.7	512.3
6	1,17E+05	(6)	2,15E+05	(11)	80	3 2	78.7	23.7	226.8
7	7,82E+04	(3)	2,09E+05	(8)	60	3 2	55.3	9.2	220.3
8	3,35E+05	(15)	1,12E+06	(50)	70	17 5	43.2	22.4	77.4
9	1,96E+05	(10)	3,52E+05	(18)	80	5 2	79.8	32.8	179.6
10	2,19E+05	(7)	5,63E+05	(18)	50	8 4	56.3	19.7	138.5
11	6,96E+05	(16)	2,22E+06	(51)	36	33 9	45.2	23.9	79.7
12	1,10E+05	(7)	2,50E+05	(16)	100	4 2	63.2	21.8	159.3
13	2,90E+05	(10)	8,40E+05	(29)	54	13 5	49.8	21.5	103.6
14	1,88E+05	(12)	4,54E+05	(29)	100	7 2	59.5	27.5	118.9
15	1,88E+05	(12)	4,69E+05	(30)	100	7 3	57.6	26.7	114.4
16	4,69E+04	(3)	9,39E+04	(6)	100	1 1	73.1	11.6	327.4
17	9,39E+04	(6)	2,82E+05	(18)	100	4 2	48.5	15.5	124.6
18	1,56E+04	(1)	4,69E+04	(3)	100	1 1	51.9	0.9	569.6
19	1,56E+05	(4)	4,69E+05	(12)	40	7 4	48.9	11.2	155.9
20	8,94E+04	(4)	2,46E+05	(11)	70	4 2	53.2	12.1	173.7
21	9,78E+04	(5)	6,85E+05	(35)	80	10 3	21.0	6.3	52.3
22	1,37E+05	(7)	4,30E+05	(22)	80	6 3	46.2	16.4	109.7
23	1,72E+05	(11)	5,01E+05	(32)	100	7 3	49.6	22.4	99.6
24	5,63E+05	(18)	1,82E+06	(58)	50	27 7	44.6	24.6	76.2
25	1,17E+05	(6)	5,48E+05	(28)	80	8 3	31.3	10.4	75.2
26	3,76E+05	(24)	1,82E+06	(116)	100	27 5	29.8	18.2	46.2
27	6,26E+04	(2)	1,56E+05	(5)	50	2 2	59.6	5.5	341.3
28	3,76E+05	(12)	7,82E+05	(25)	50	12 5	69.0	31.4	140.7
29	9,39E+04	(3)	1,88E+05	(6)	50	3 2	73.1	11.6	327.4
30	1,17E+05	(6)	3,91E+05	(20)	80	6 3	43.7	14.1	110.2
31	2,19E+05	(14)	7,82E+05	(50)	100	12 3	40.4	20.5	73.4
32	1,30E+05	(5)	3,39E+05	(13)	60	5 3	56.0	15.4	162.9
33	6,26E+04	(4)	3,44E+05	(22)	100	5 2	26.9	6.5	76.4
34	1,04E+05	(4)	4,69E+05	(18)	60	7 3	32.8	7.8	96.1
35	2,61E+05	(6)	9,13E+05	(21)	36	14 6	41.6	13.5	104.2
36	1,41E+05	(9)	5,16E+05	(33)	100	8 3	39.5	16.4	83.1
37	2,03E+05	(13)	1,22E+06	(78)	100	18 4	24.1	12.2	43.2
38	9,39E+05	(54)	2,07E+06	(119)	90	31 6	64.9	46.1	89.9
39	4,69E+04	(3)	2,82E+05	(18)	100	4 2	24.9	4.5	81.4
40	4,07E+05	(13)	1,44E+06	(46)	50	21 6	40.8	20.0	75.9
41	6,52E+05	(15)	3,65E+06	(84)	36	54 12	25.8	13.7	44.6
42	9,39E+04	(6)	5,63E+05	(36)	100	8 3	24.4	8.2	57.0
43	4,69E+04	(3)	1,41E+05	(9)	100	2 1	49.3	8.3	188.8
44	3,13E+04	(2)	3,44E+05	(22)	100	5 2	13.9	1.5	52.8
45	4,47E+04	(2)	3,13E+05	(14)	70	5 2	21.8	2.3	88.6
46	4,56E+05	(14)	2,18E+06	(67)	48	33 8	30.2	15.5	53.7
47	1,41E+05	(9)	2,97E+05	(19)	100	4 2	68.2	27.0	155.7
48	1,25E+05	(8)	4,07E+05	(26)	100	6 2	44.6	17.2	99.6
49	8,94E+04	(4)	8,72E+05	(39)	70	13 4	15.2	3.8	40.7
50	1,10E+05	(7)	2,97E+05	(19)	100	4 2	53.4	18.7	130.0
51	4,54E+05	(29)	2,43E+06	(155)	100	36 6	26.9	17.4	40.0
52	8,94E+04	(4)	1,56E+05	(7)	70	2 2	82.8	17.6	314.6
53	1,56E+05	(9)	4,35E+05	(25)	90	6 3	52.0	21.2	113.4
54	4,69E+04	(3)	5,63E+05	(36)	100	8 3	12.5	2.4	37.7

Grain no.	RhoS (cm ⁻²)	(Ns)	RhoI (cm ⁻²)	(Ni)	Squares	U+/-2s	Grain Age (Ma)		
							Age	--95% CI--	
55	3,13E+04	(1)	5,01E+05	(16)	50	7 4	10.2	0.2	57.5
56	2,09E+05	(8)	2,87E+05	(11)	60	4 3	104.0	36.3	278.7
57	7,82E+04	(5)	1,22E+06	(78)	100	18 4	9.5	2.9	22.4
58	1,96E+05	(6)	3,59E+05	(11)	48	5 3	78.7	23.7	226.8
59	9,78E+04	(5)	2,35E+05	(12)	80	3 2	60.6	16.5	179.8
60	2,66E+05	(17)	1,36E+06	(87)	100	20 4	28.2	15.6	47.3
61	1,30E+05	(5)	3,13E+05	(12)	60	5 3	60.6	16.5	179.8
62	1,25E+05	(8)	2,50E+05	(16)	100	4 2	72.0	26.5	175.2
63	5,01E+05	(16)	8,14E+05	(26)	50	12 5	88.0	44.1	168.7
64	1,03E+06	(46)	2,73E+06	(122)	70	41 7	54.0	37.5	76.1
65	3,13E+05	(6)	5,22E+05	(10)	30	8 5	86.4	25.6	256.2
66	1,88E+05	(6)	4,38E+05	(14)	50	7 3	62.1	19.3	168.1
67	5,16E+05	(33)	1,35E+06	(86)	100	20 4	55.0	35.6	82.6
68	1,10E+05	(7)	6,10E+05	(39)	100	9 3	26.2	9.7	58.0
69	6,26E+04	(2)	1,25E+05	(4)	50	2 2	73.9	6.5	482.0
70	2,50E+05	(16)	8,14E+05	(52)	100	12 3	44.3	23.5	78.1
71	5,87E+04	(3)	5,48E+05	(28)	80	8 3	16.1	3.0	49.6
72	1,56E+05	(5)	2,82E+05	(9)	50	4 3	80.3	20.9	259.5
POOLED	1,96E+05	(685)	6,77E+05	(2369)	5478	10 0	41.3	37.5	45.5

CHI² PROBABILITY (%): 0.3

>>> Beware: possible upward bias in Chi² probability due to low counts <<<

POOLED AGE W/ 68% CONF. INTERVAL (Ma): 41.3, 39.4 -- 43.4 (-2.0 +2.1)
 95% CONF. INTERVAL (Ma): 37.5 -- 45.5 (-3.8 +4.2)

CENTRAL AGE W/ 68% CONF. INTERVAL (Ma): 42.6, 40.0 -- 45.3 (-2.6 +2.7)
 95% CONF. INTERVAL (Ma): 37.7 -- 48.1 (-4.9 +5.5)
 AGE DISPERSION (%): 27.1

FIT OPTION: Best-fit peaks using the binomial model of Galbraith and Green

INITIAL GUESS FOR MODEL PARAMETERS (number of peaks to fit = 2)

Peak #.	Peak Age	Theta	Fraction(%)	Count
1.	41.30	0.224	23.4	16.81
2.	63.80	0.309	27.3	19.63

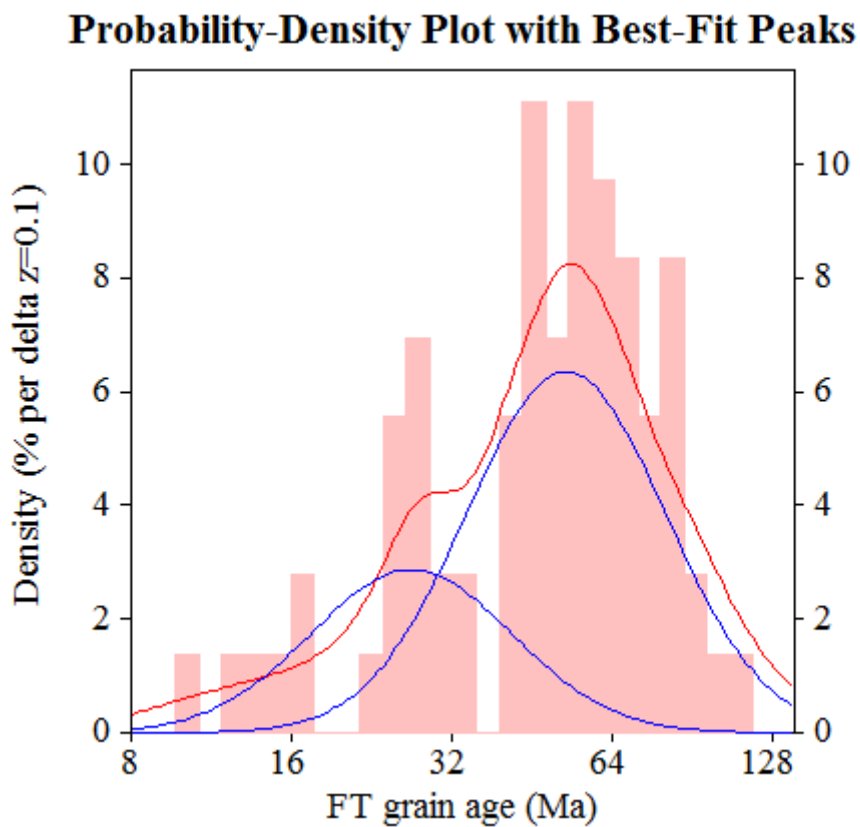
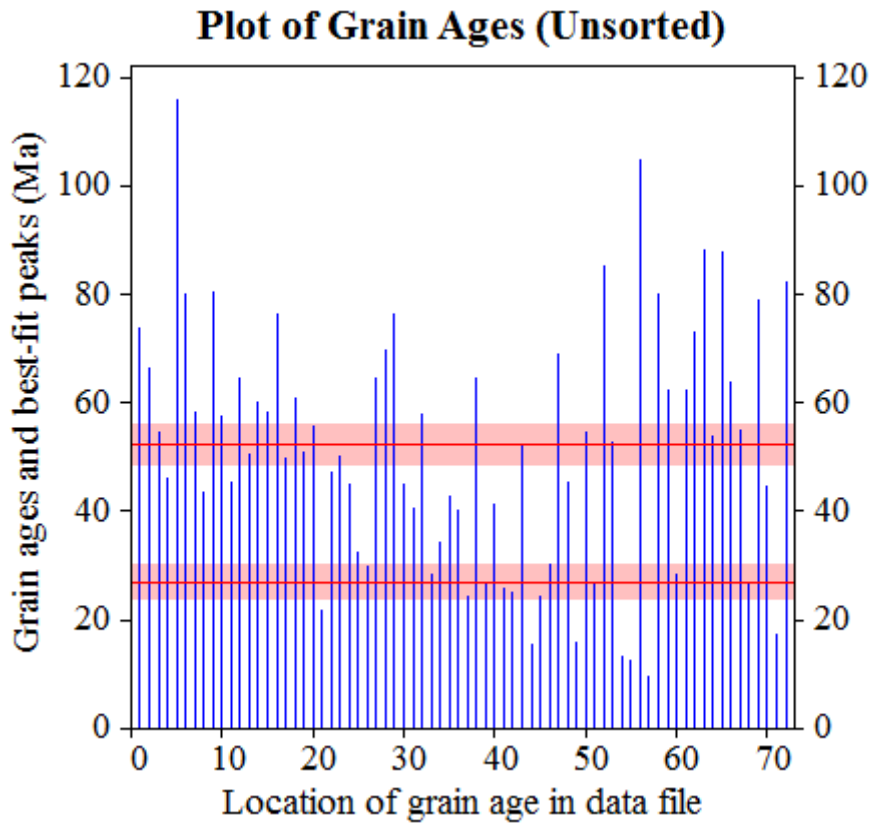
Total range for grain ages: 10,0 to 116,3 Ma
 Number of active grains (Num. used for fit): 72
 Number of removed grains: 0
 Degrees of freedom for fit: 69
 Average of the SE(Z)'s for the grains: 0,48
 Estimated width of peaks in PD plot in Z units: 0,56

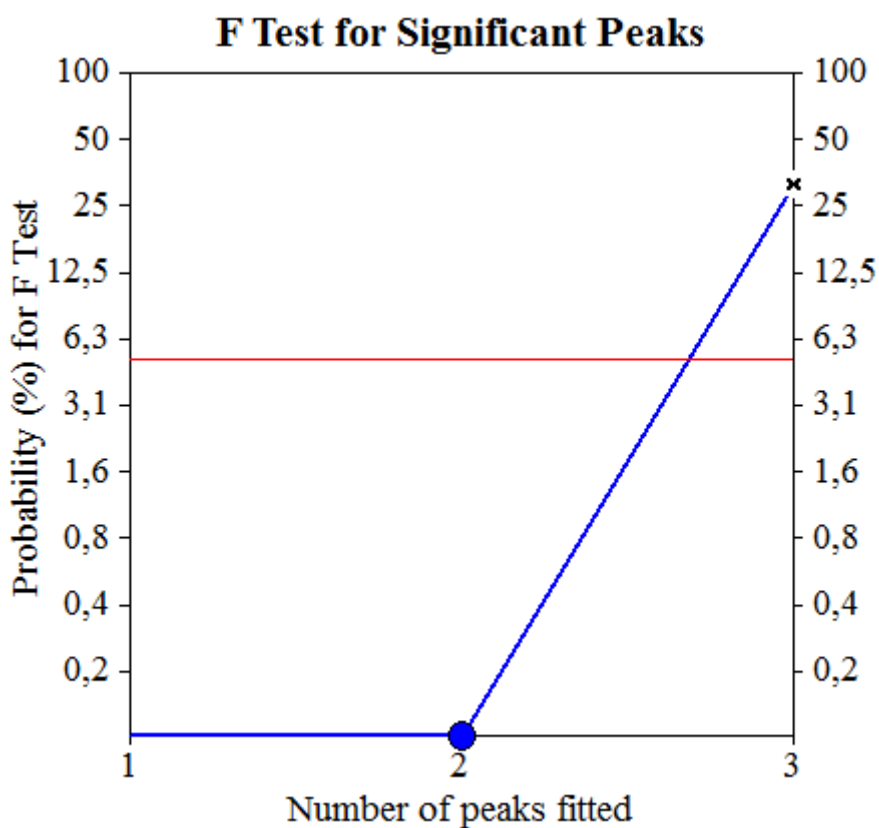
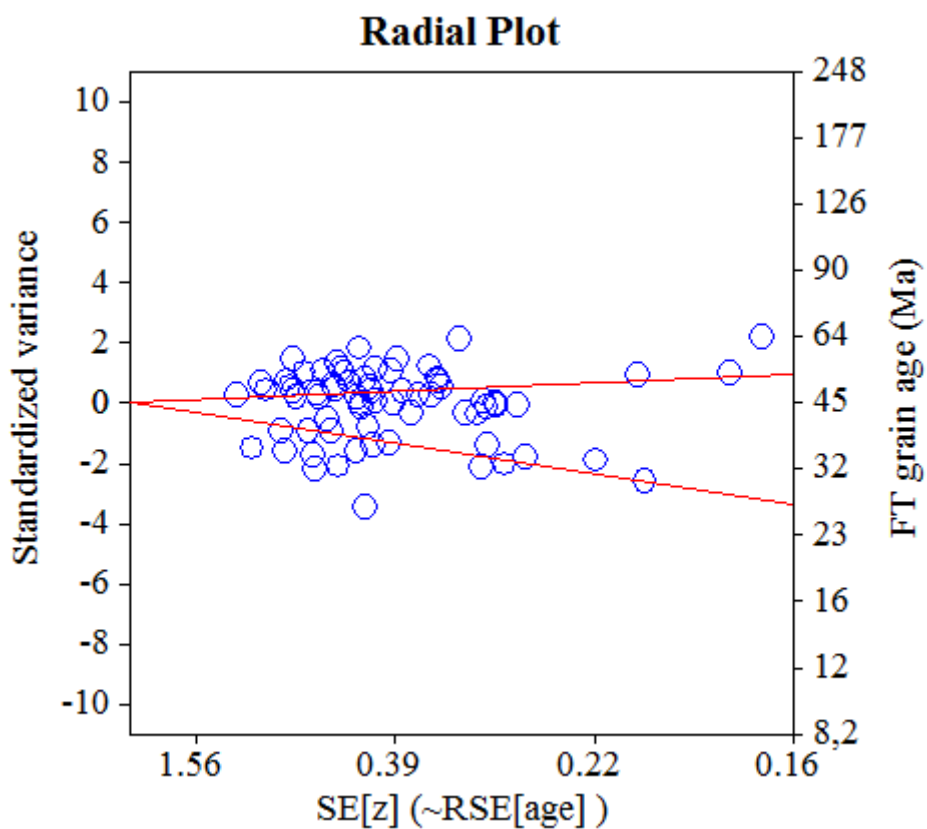
PARAMETERS FOR BEST-FIT PEAKS

- * Standard error for peak age includes group error
- * Peak width is for PD plot assuming a kernel factor = 0.60

#.	Peak Age (Ma)	68%CI	95%CI	W(Z)	Frac (%)	SE, %	Count
1.	26.8	-3,0 ...+3,4	-5,6 ...+7,1	0.44	31.3	10.5	22.5
2.	52.3	-3,6 ...+3,9	-6,8 ...+7,8	0.43	68.7	10.5	49.5

Log-likelihood for best fit: -173,701
 Chi-squared value for best fit: 61,407
 Reduced chi-squared value: 0,890
 Probability for F test: 0%
 Condition number for COVAR matrix: 6,12
 Number of iterations: 21





NEW PARAMETERS - ZETA METHOD

EFFECTIVE TRACK DENSITY FOR FLUENCE MONITOR (tracks/cm²): 1,31E+06
 RELATIVE ERROR (%): 1,11
 EFFECTIVE URANIUM CONTENT OF MONITOR (ppm): 15,00
 ZETA FACTOR AND STANDARD ERROR (yr cm²): 270,07 6,38
 SIZE OF COUNTER SQUARE (cm²): 6,39E-07

GRAIN AGES IN ORIGINAL ORDER

Grain no.	RhoS (cm ⁻²)	(Ns)	RhoI (cm ⁻²)	(Ni)	Squares	U+/-2s	Grain Age (Ma)		
							Age	--95% CI--	
1	6,26E+05	(40)	3,80E+06	(243)	100	44 6	29.0	20.2	40.6
2	4,23E+05	(27)	2,27E+06	(145)	100	26 4	32.9	20.9	49.6
3	5,48E+05	(35)	3,68E+06	(235)	100	42 6	26.3	17.8	37.5
4	3,60E+05	(23)	1,92E+06	(123)	100	22 4	33.1	20.1	51.6
5	3,76E+05	(24)	3,19E+06	(204)	100	37 5	20.8	13.0	31.7
6	6,57E+05	(42)	3,99E+06	(255)	100	46 6	29.1	20.4	40.3
7	1,33E+06	(85)	5,90E+06	(377)	100	68 7	39.6	31.1	50.3
8	1,75E+06	(112)	7,59E+06	(485)	100	87 8	40.5	32.8	50.0
9	5,48E+05	(35)	2,57E+06	(164)	100	30 5	37.6	25.3	54.3
10	4,07E+05	(26)	3,76E+06	(240)	100	43 6	19.2	12.2	28.6
11	1,03E+06	(66)	5,93E+06	(379)	100	68 7	30.6	23.5	39.9
12	5,95E+05	(38)	2,97E+06	(190)	100	34 5	35.3	24.1	50.0
13	5,63E+05	(36)	3,87E+06	(247)	100	44 6	25.7	17.5	36.5
14	2,35E+05	(15)	1,38E+06	(88)	100	16 3	30.2	16.1	52.1
15	1,72E+05	(11)	8,92E+05	(57)	100	10 3	34.3	16.1	65.2
16	4,38E+05	(28)	2,55E+06	(163)	100	29 5	30.3	19.5	45.3
17	7,51E+05	(48)	2,82E+06	(180)	100	32 5	46.9	33.3	64.7
18	4,23E+05	(27)	3,62E+06	(231)	100	42 6	20.7	13.3	30.7
19	9,55E+05	(61)	3,90E+06	(249)	100	45 6	43.0	32.4	57.0
20	3,91E+05	(25)	3,51E+06	(224)	100	40 5	19.8	12.4	29.8
21	8,14E+05	(52)	4,10E+06	(262)	100	47 6	35.0	25.4	47.1
22	6,57E+05	(42)	3,47E+06	(222)	100	40 5	33.4	23.3	46.4
23	5,16E+05	(33)	3,43E+06	(219)	100	39 5	26.6	17.8	38.3
24	7,20E+05	(46)	7,75E+06	(495)	100	89 8	16.4	11.8	22.2
25	5,01E+05	(16)	2,88E+06	(92)	50	33 7	30.8	16.8	52.3
26	7,82E+05	(50)	2,54E+06	(162)	100	29 5	54.3	38.6	74.7
27	2,97E+05	(19)	2,49E+06	(159)	100	29 5	21.2	12.3	33.9
28	3,60E+05	(23)	3,63E+06	(232)	100	42 6	17.6	10.8	26.8
29	1,10E+06	(42)	3,50E+06	(134)	60	40 7	55.1	38.0	78.2
30	5,39E+05	(31)	3,55E+06	(204)	90	41 6	26.8	17.7	39.1
31	1,03E+06	(66)	5,77E+06	(369)	100	66 7	31.4	24.1	41.0
32	7,04E+05	(45)	3,13E+06	(200)	100	36 5	39.6	28.0	54.8
33	1,27E+06	(81)	4,74E+06	(303)	100	55 6	46.9	36.6	60.1
34	5,32E+05	(17)	2,50E+06	(80)	50	29 6	37.6	20.8	63.5
35	2,35E+05	(15)	7,82E+06	(50)	100	9 3	53.1	27.5	95.0
POOLED	6,46E+05	(1382)	3,58E+06	(7662)	3350	41 1	31.7	29.4	34.2

CHI² PROBABILITY (%): 0.0

POOLED AGE W/ 68% CONF. INTERVAL (Ma): 31.7, 30.5 -- 33.0 (-1.2 +1.3)
 95% CONF. INTERVAL (Ma): 29.4 -- 34.2 (-2.3 +2.5)

CENTRAL AGE W/ 68% CONF. INTERVAL (Ma): 31.9, 30.1 -- 33.9 (-1.8 +1.9)
 95% CONF. INTERVAL (Ma): 28.4 -- 35.8 (-3.5 +3.9)
 AGE DISPERSION (%): 25.3

FIT OPTION: Best-fit peaks using the binomial model of Galbraith and Green

INITIAL GUESS FOR MODEL PARAMETERS (number of peaks to fit = 3)

Peak #.	Peak Age	Theta	Fraction(%)	Count
1.	31.70	0.153	48.8	17.09
2.	48.60	0.217	20.3	7.11
3.	55.40	0.240	12.9	4.52

Total range for grain ages: 16,5 to 55,4 Ma
 Number of active grains (Num. used for fit): 35
 Number of removed grains: 0
 Degrees of freedom for fit: 30
 Average of the SE(Z)'s for the grains: 0,2
 Estimated width of peaks in PD plot in Z units: 0,23

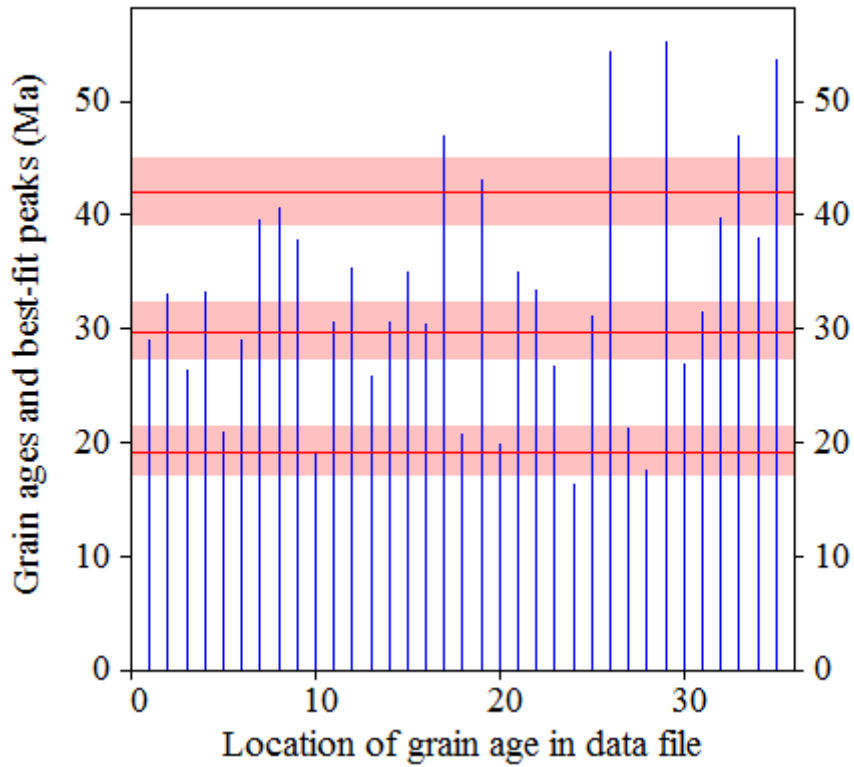
PARAMETERS FOR BEST-FIT PEAKS

- * Standard error for peak age includes group error
- * Peak width is for PD plot assuming a kernel factor = 0.60

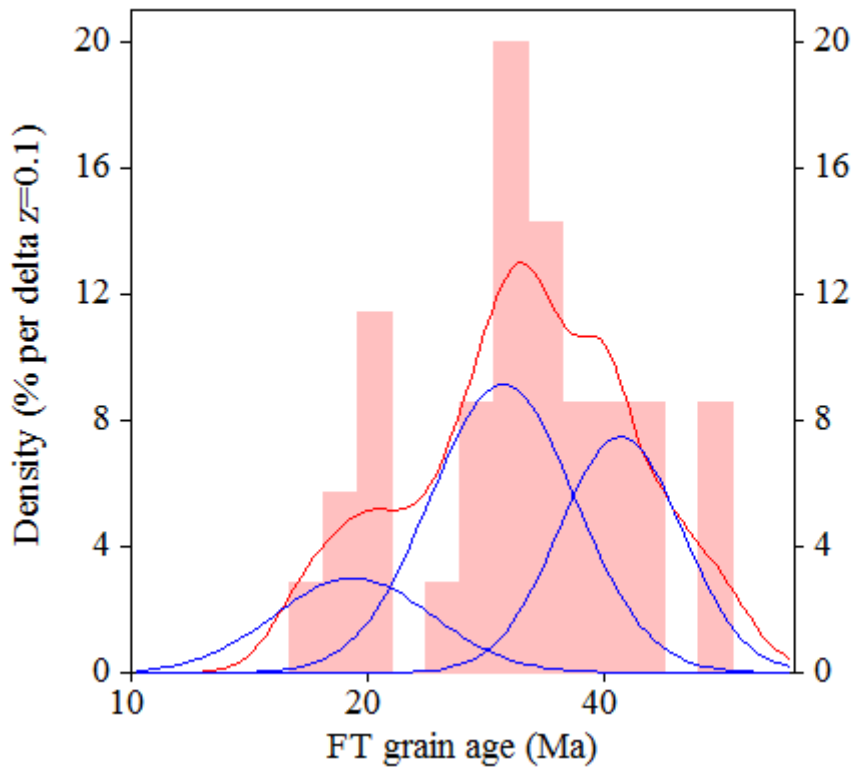
#.	Peak Age (Ma)	68%CI	95%CI	W(Z)	Frac(%)	SE, %	Count
1.	19.2	-2,0 ...+2,3	-3,8 ...+4,7	0.23	17.3	10.1	6.1
2.	29.8	-2,3 ...+2,5	-4,4 ...+5,1	0.21	48.8	14.5	17.1
3.	42.0	-2,8 ...+3,0	-5,4 ...+6,1	0.18	33.9	13.0	11.9

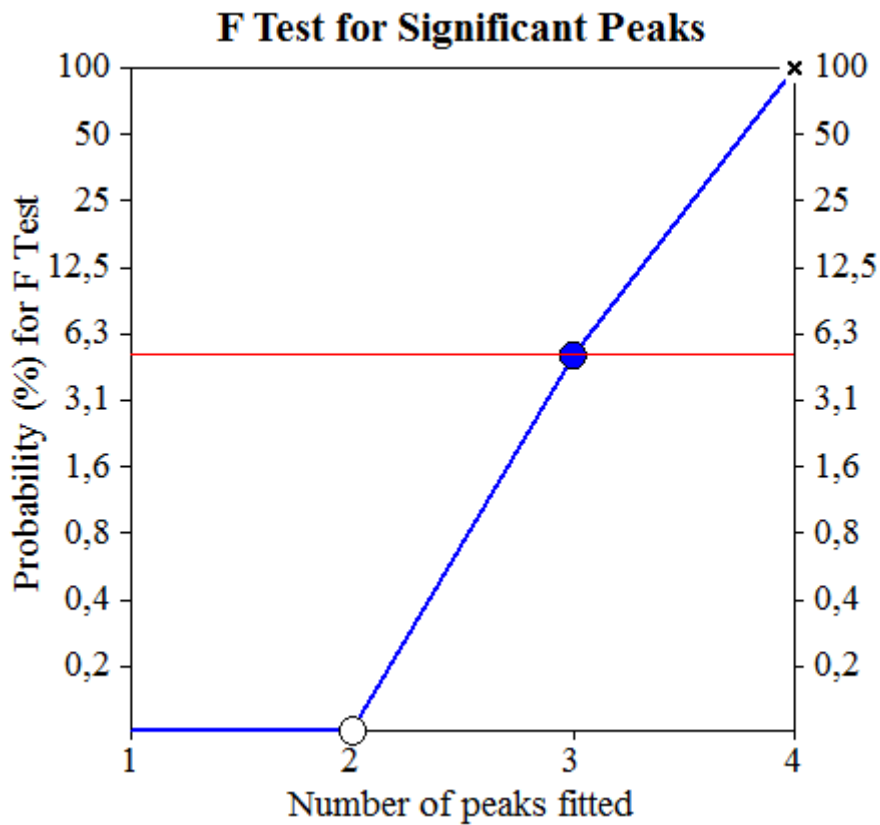
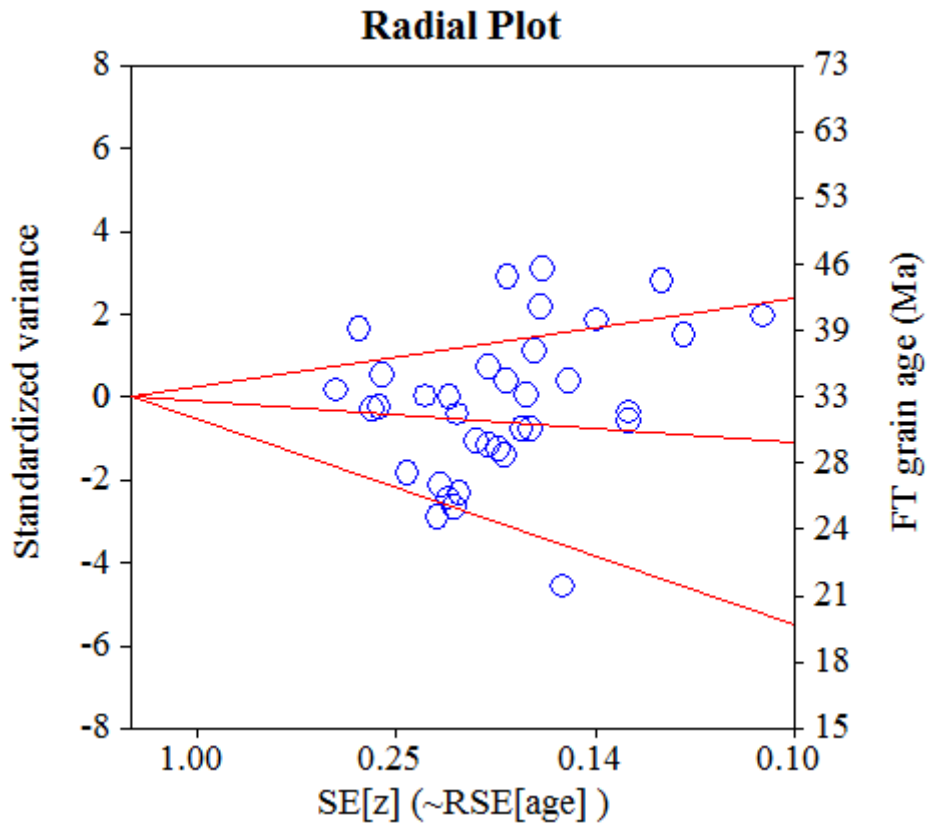
Log-likelihood for best fit: -126,208
 Chi-squared value for best fit: 33,775
 Reduced chi-squared value: 1,126
 Probability for F test: 5%
 Condition number for COVAR matrix: 11,08
 Number of iterations: 48

Plot of Grain Ages (Unsorted)



Probability-Density Plot with Best-Fit Peaks





NEW PARAMETERS - ZETA METHOD

EFFECTIVE TRACK DENSITY FOR FLUENCE MONITOR (tracks/cm²): 1,02E+06
 RELATIVE ERROR (%): 1,24
 EFFECTIVE URANIUM CONTENT OF MONITOR (ppm): 15,00
 ZETA FACTOR AND STANDARD ERROR (yr cm²): 284,52 5,65
 SIZE OF COUNTER SQUARE (cm²): 6,39E-07

GRAIN AGES IN ORIGINAL ORDER

Grain no.	RhoS (cm ⁻²)	(Ns)	RhoI (cm ⁻²)	(Ni)	Squares	U+/-2s	Grain Age (Ma)		
							Age	--95% CI--	
1	3,29E+05	(21)	5,32E+05	(34)	100	8 3	89.2	49.2	157.0
2	1,96E+05	(10)	1,33E+06	(68)	80	20 5	21.6	9.8	41.5
3	2,03E+05	(13)	4,69E+05	(30)	100	7 3	63.0	30.0	123.0
4	6,26E+04	(4)	1,88E+05	(12)	100	3 2	49.4	11.3	157.6
5	1,56E+05	(10)	7,36E+05	(47)	100	11 3	31.2	13.9	61.6
6	2,82E+05	(18)	4,23E+05	(27)	100	6 2	96.2	49.9	179.6
7	2,97E+05	(19)	6,73E+05	(43)	100	10 3	64.1	35.1	111.4
8	4,69E+04	(3)	2,97E+05	(19)	100	4 2	23.9	4.3	77.3
9	2,35E+05	(9)	4,43E+05	(17)	60	7 3	76.9	30.1	179.6
10	4,69E+04	(3)	2,35E+05	(15)	100	3 2	30.1	5.4	101.7
11	1,41E+05	(9)	1,00E+06	(64)	100	15 4	20.7	8.9	41.1
12	2,03E+05	(13)	5,01E+05	(32)	100	7 3	59.1	28.3	114.4
13	2,03E+05	(13)	3,60E+05	(23)	100	5 2	81.9	38.0	166.6
14	3,80E+05	(17)	7,82E+05	(35)	70	12 4	70.4	36.9	127.9
15	1,25E+05	(8)	7,51E+05	(48)	100	11 3	24.5	9.9	51.4
16	2,24E+05	(10)	4,92E+05	(22)	70	7 3	66.1	27.8	143.5
17	6,71E+04	(3)	6,04E+05	(27)	70	9 3	16.8	3.1	52.2
18	7,82E+04	(5)	5,32E+05	(34)	100	8 3	21.9	6.5	54.6
19	6,26E+04	(4)	3,13E+05	(20)	100	5 2	29.9	7.2	86.0
20	4,38E+05	(28)	3,16E+06	(202)	100	47 7	20.2	13.0	29.9
21	2,66E+05	(17)	7,98E+05	(51)	100	12 3	48.5	26.1	84.6
22	5,48E+05	(35)	9,23E+05	(59)	100	14 4	85.6	54.7	131.6
23	3,13E+05	(12)	1,98E+06	(76)	60	29 7	23.1	11.3	42.2
24	8,76E+05	(56)	7,10E+06	(454)	100	105 10	17.9	13.5	23.6
25	2,19E+05	(14)	4,23E+05	(27)	100	6 2	75.2	36.3	146.9
26	2,82E+05	(18)	9,08E+05	(58)	100	13 4	45.1	24.9	77.1
27	5,09E+05	(26)	7,63E+05	(39)	80	11 4	96.1	56.2	160.9
28	1,56E+05	(10)	5,16E+05	(33)	100	8 3	44.3	19.3	90.7
29	1,25E+05	(8)	7,67E+05	(49)	100	11 3	24.0	9.7	50.2
30	4,54E+05	(29)	1,35E+06	(86)	100	20 4	48.9	30.8	74.9
31	2,66E+05	(17)	1,13E+06	(72)	100	17 4	34.4	18.9	58.5
32	1,56E+05	(8)	1,21E+06	(62)	80	18 5	19.0	7.7	39.1
33	3,13E+05	(20)	1,03E+06	(66)	100	15 4	44.0	25.2	73.0
34	1,72E+05	(11)	8,92E+05	(57)	100	13 3	28.3	13.2	53.7
35	1,41E+05	(9)	5,95E+05	(38)	100	9 3	34.7	14.6	71.9
36	1,56E+05	(10)	5,48E+05	(35)	100	8 3	41.8	18.3	85.0
37	2,97E+05	(19)	8,14E+05	(52)	100	12 3	53.0	29.5	90.5
38	2,82E+05	(18)	5,63E+05	(36)	100	8 3	72.4	38.6	129.8
POOLED	2,44E+05	(557)	9,51E+05	(2169)	3570	14 1	37.1	33.5	41.2

CHI² PROBABILITY (%): 0.0

>>> Beware: possible upward bias in Chi² probability due to low counts <<<

POOLED AGE W/ 68% CONF. INTERVAL (Ma): 37.1, 35.2 -- 39.1 (-1.9 +2.0)
 95% CONF. INTERVAL (Ma): 33.5 -- 41.2 (-3.6 +4.0)

CENTRAL AGE W/ 68% CONF. INTERVAL (Ma): 43.3, 39.5 -- 47.4 (-3.7 +4.1)
 95% CONF. INTERVAL (Ma): 36.2 -- 51.6 (-7.0 +8.4)
 AGE DISPERSION (%): 43.3

FIT OPTION: Best-fit peaks using the binomial model of Galbraith and Green

INITIAL GUESS FOR MODEL PARAMETERS (number of peaks to fit = 2)

Peak #.	Peak Age	Theta	Fraction(%)	Count
1.	22.50	0.135	19.4	7.37
2.	37.10	0.204	17.7	6.71

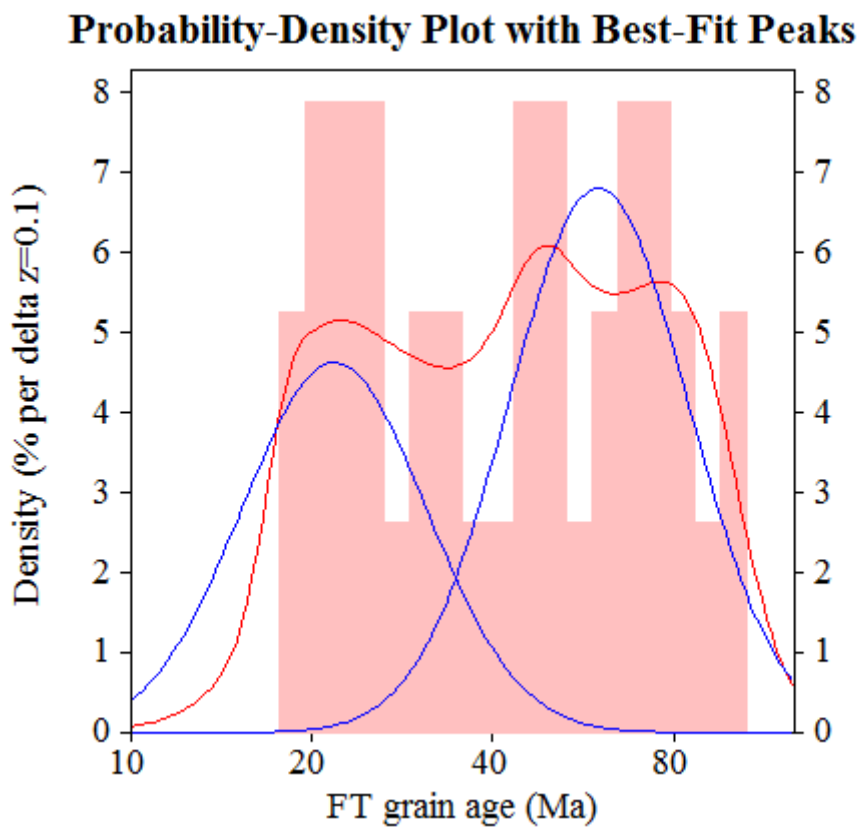
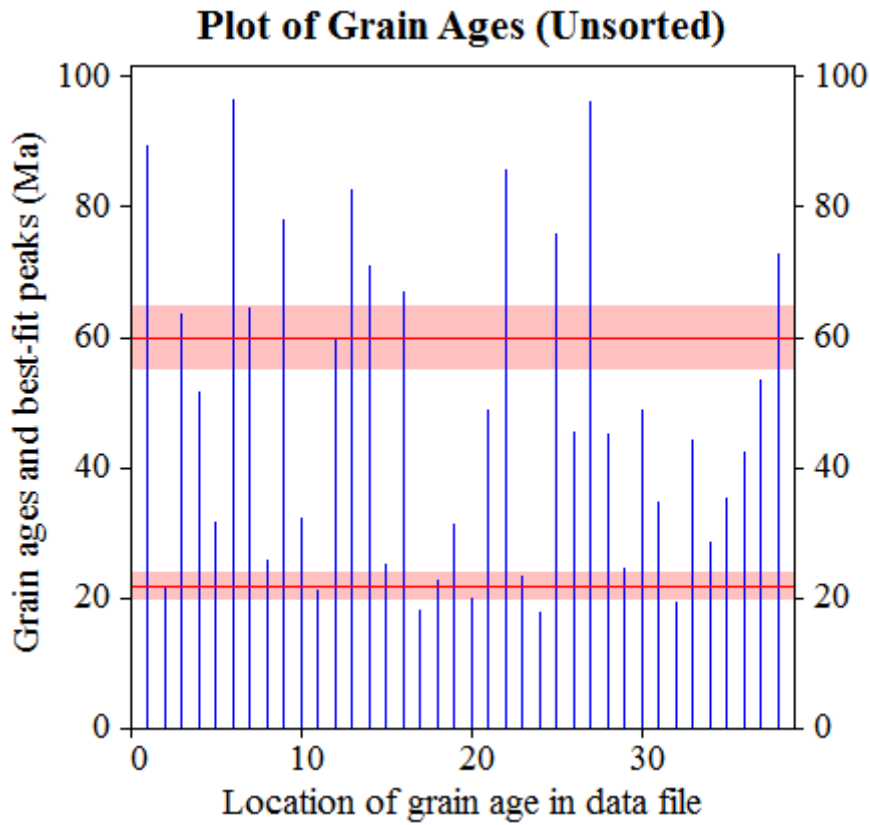
Total range for grain ages: 18,0 to 96,8 Ma
 Number of active grains (Num. used for fit): 38
 Number of removed grains: 0
 Degrees of freedom for fit: 35
 Average of the SE(Z)'s for the grains: 0,36
 Estimated width of peaks in PD plot in Z units: 0,42

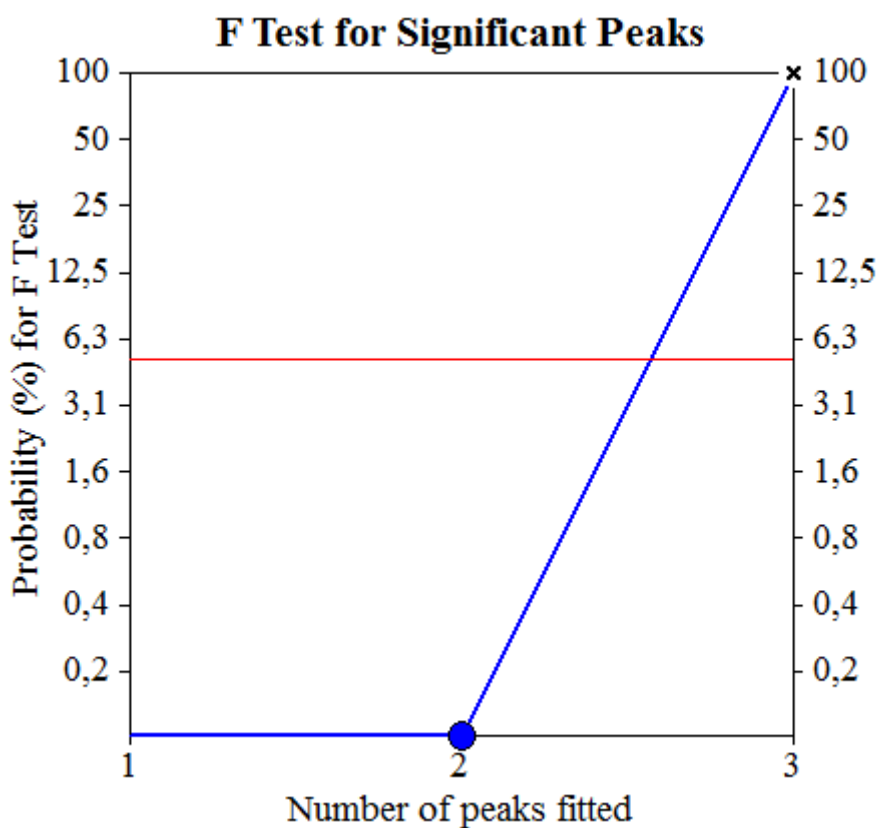
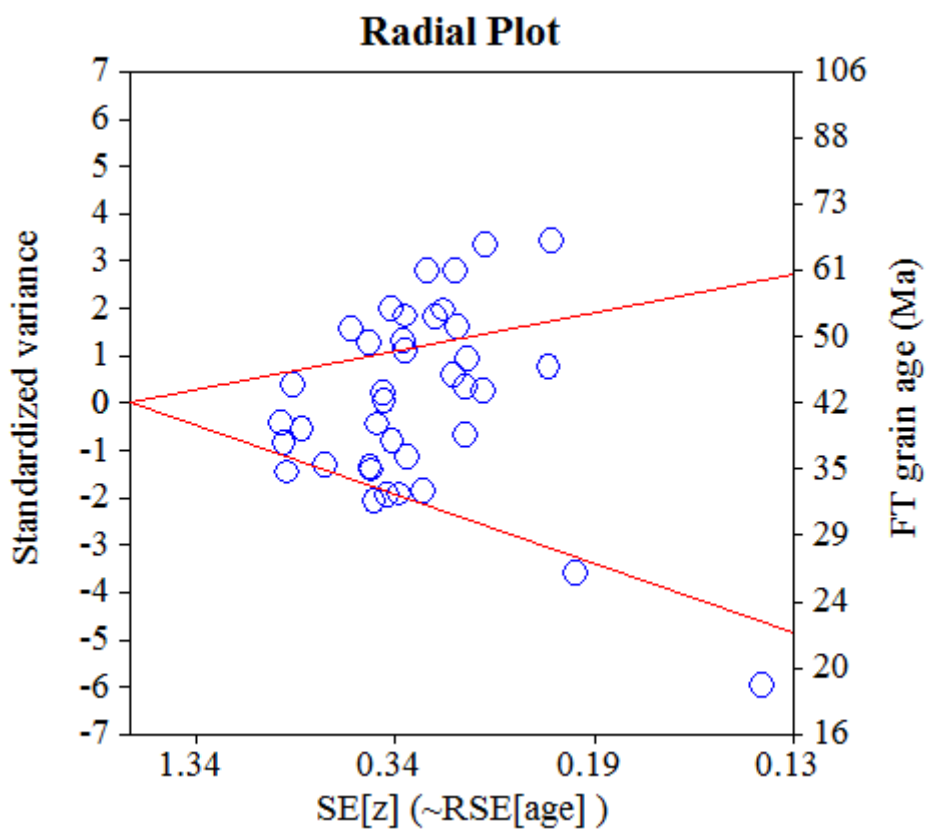
PARAMETERS FOR BEST-FIT PEAKS

- * Standard error for peak age includes group error
- * Peak width is for PD plot assuming a kernel factor = 0.60

#.	Peak Age (Ma)	68%CI	95%CI	W(Z)	Frac (%)	SE, %	Count
1.	21.8	-1,9 ...+2,1	-3,6 ...+4,3	0.35	41.0	9.9	15.6
2.	59.9	-4,6 ...+5,0	-8,6 ...+10,1	0.35	59.0	9.9	22.4

Log-likelihood for best fit: -113,730
 Chi-squared value for best fit: 34,537
 Reduced chi-squared value: 0,987
 Probability for F test: 0%
 Condition number for COVAR matrix: 2,68
 Number of iterations: 9





NEW PARAMETERS - ZETA METHOD

EFFECTIVE TRACK DENSITY FOR FLUENCE MONITOR (tracks/cm²): 1,26E+06
 RELATIVE ERROR (%): 1,15
 EFFECTIVE URANIUM CONTENT OF MONITOR (ppm): 15,00
 ZETA FACTOR AND STANDARD ERROR (yr cm²): 288,88 16,55
 SIZE OF COUNTER SQUARE (cm²): 6,39E-07

GRAIN AGES IN ORIGINAL ORDER

Grain no.	RhoS (cm ⁻²)	(Ns)	RhoI (cm ⁻²)	(Ni)	Squares	U+/-2s	Grain Age (Ma)		
							Age	--95% CI--	
1	2,35E+05	(15)	6,26E+05	(40)	100	7 2	68.2	34.8	124.8
2	1,56E+05	(8)	8,22E+05	(42)	80	10 3	35.1	14.0	74.2
3	3,13E+05	(20)	1,50E+06	(96)	100	18 4	38.0	22.1	61.4
4	5,95E+05	(38)	1,22E+06	(78)	100	15 3	88.0	58.1	130.7
5	2,50E+05	(16)	9,23E+05	(59)	100	11 3	49.4	26.4	86.2
6	4,23E+05	(27)	1,36E+06	(87)	100	16 3	56.3	35.0	87.1
7	1,72E+05	(11)	1,11E+06	(71)	100	13 3	28.4	13.4	53.2
8	1,25E+05	(8)	3,29E+05	(21)	100	4 2	69.7	26.4	160.6
POOLED	2,87E+05	(143)	9,91E+05	(494)	780	12 1	52.3	42.1	65.0

CHI² PROBABILITY (%): 3.3

POOLED AGE W/ 68% CONF. INTERVAL (Ma): 52.3, 46.8 -- 58.4 (-5.5 +6.1)
 95% CONF. INTERVAL (Ma): 42.1 -- 65.0 (-10.2 +12.7)

CENTRAL AGE W/ 68% CONF. INTERVAL (Ma): 51.9, 44.9 -- 60.1 (-7.1 +8.2)
 95% CONF. INTERVAL (Ma): 39.0 -- 69.2 (-13.0 +17.3)
 AGE DISPERSION (%): 26.2

FIT OPTION: Best-fit peaks using the binomial model of Galbraith and Green

INITIAL GUESS FOR MODEL PARAMETERS (number of peaks to fit = 2)

Peak #.	Peak Age	Theta	Fraction(%)	Count
1.	29.10	0.139	17.2	1.38
2.	52.30	0.224	32.7	2.61

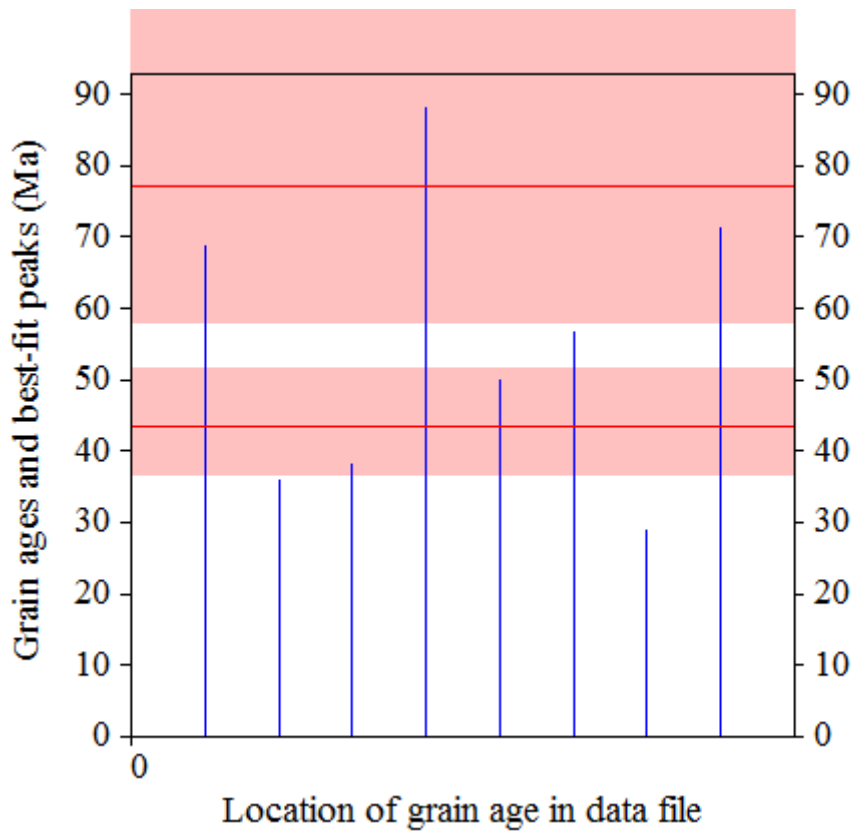
Total range for grain ages: 29,1 to 88,4 Ma
 Number of active grains (Num. used for fit): 8
 Number of removed grains: 0
 Degrees of freedom for fit: 5
 Average of the SE(Z)'s for the grains: 0,3
 Estimated width of peaks in PD plot in Z units: 0,35

PARAMETERS FOR BEST-FIT PEAKS

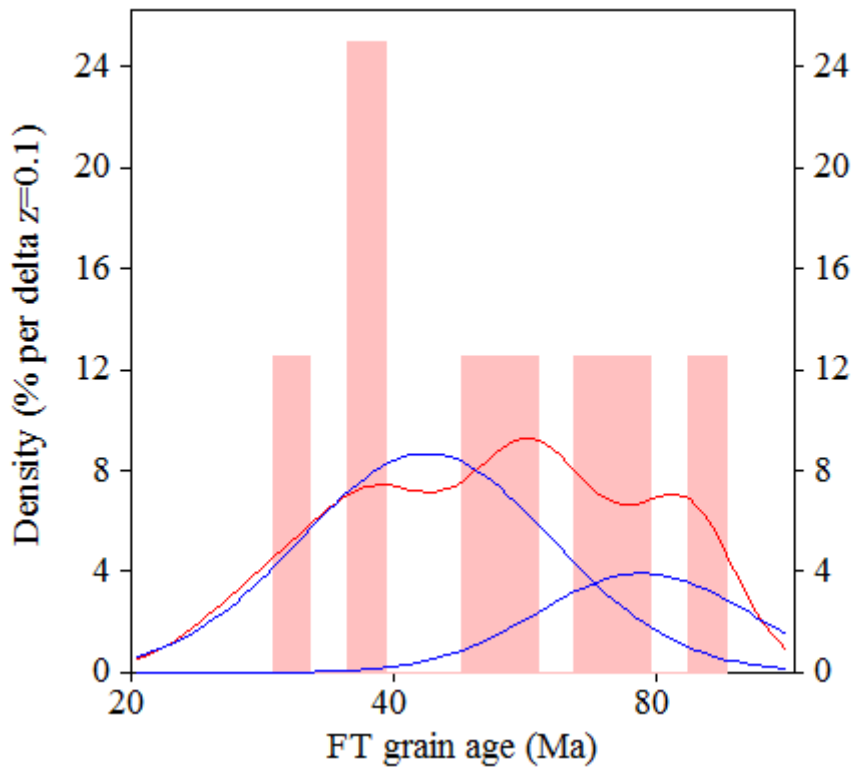
- * Standard error for peak age includes group error
- * Peak width is for PD plot assuming a kernel factor = 0.60

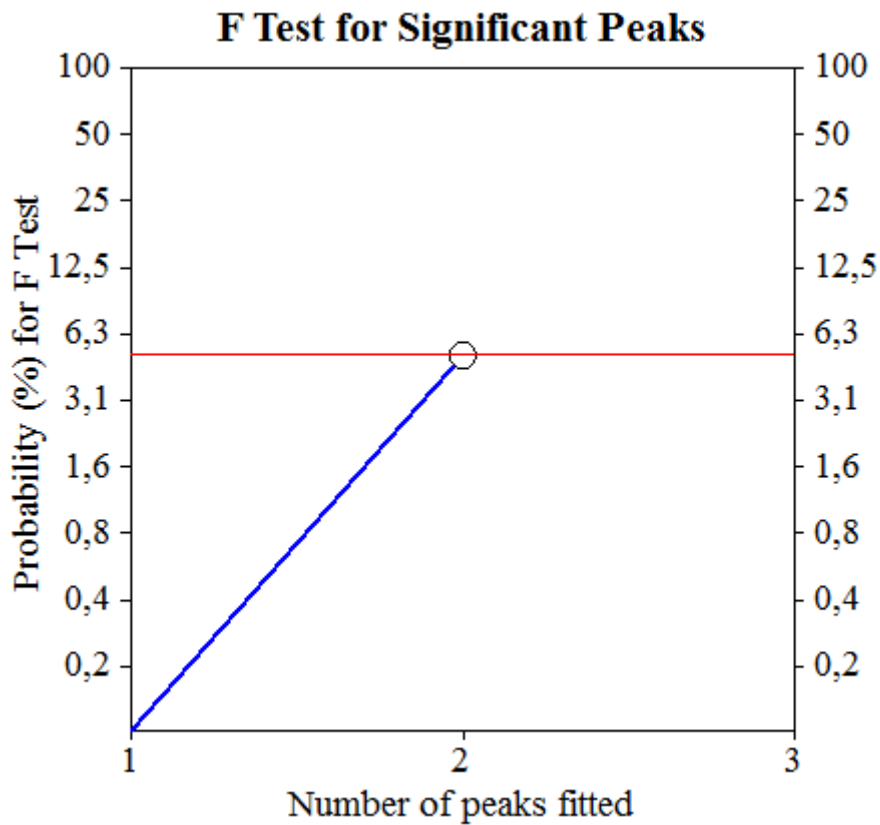
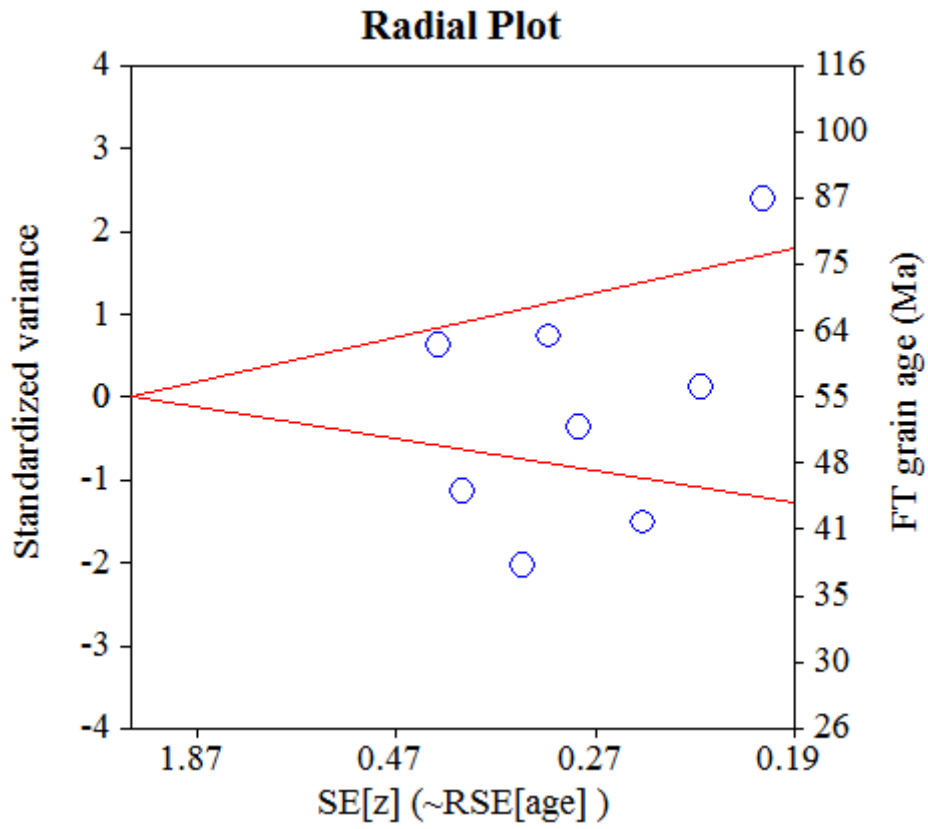
#.	Peak Age (Ma)	68%CI	95%CI	W(Z)	Frac (%)	SE, %	Count
1.	43.6	-6,7 ...+7,9	-12,2 ...+16,9	0.33	72.6	31.3	5.8
2.	77.2	-19,0 ...+25,2	-32,9 ...+57,1	0.28	27.4	31.3	2.2

Log-likelihood for best fit: -23,278
 Chi-squared value for best fit: 7,077
 Reduced chi-squared value: 1,415
 Probability for F test: 5%
 Condition number for COVAR matrix: 9,94
 Number of iterations: 32

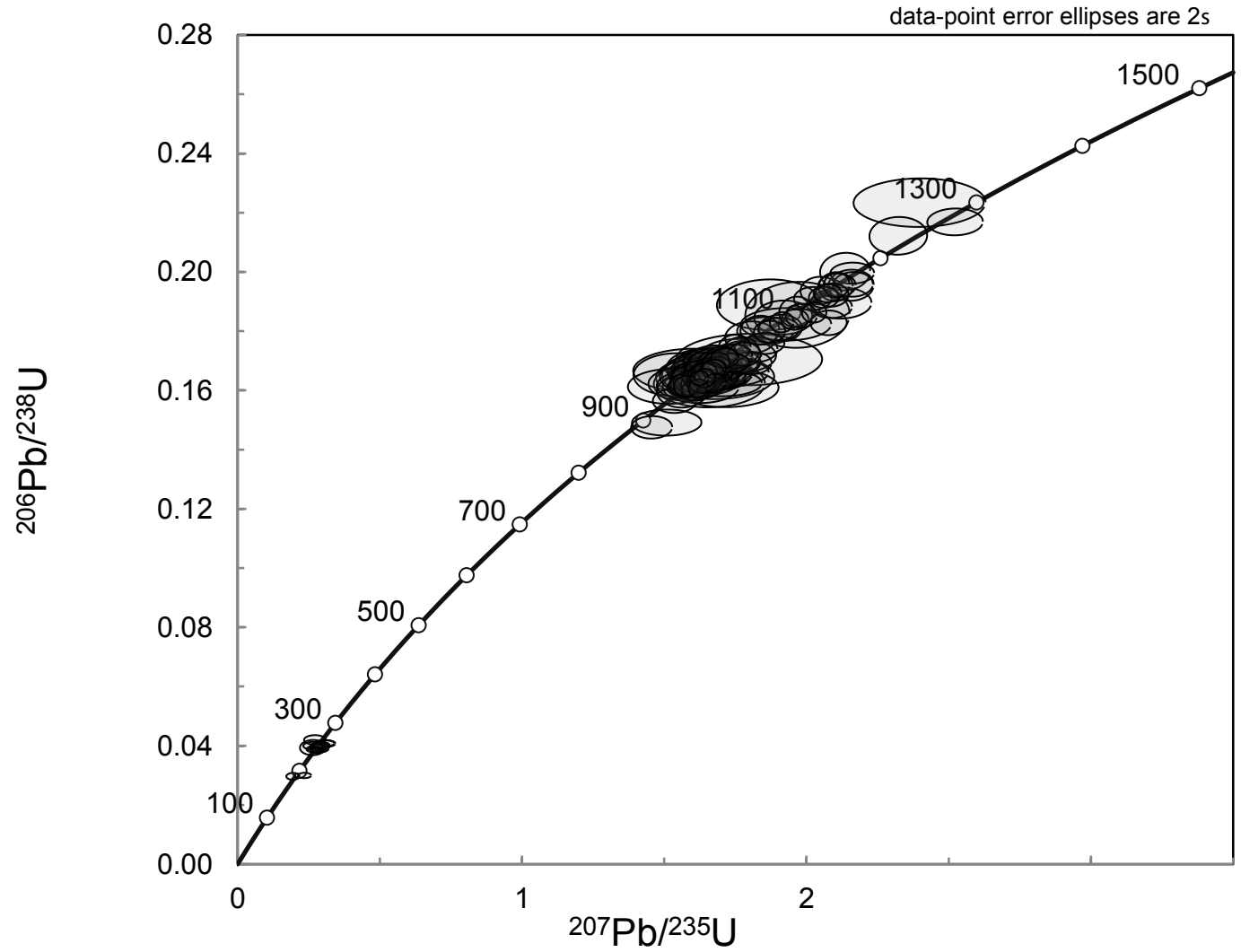


Probability-Density Plot with Best-Fit Peaks

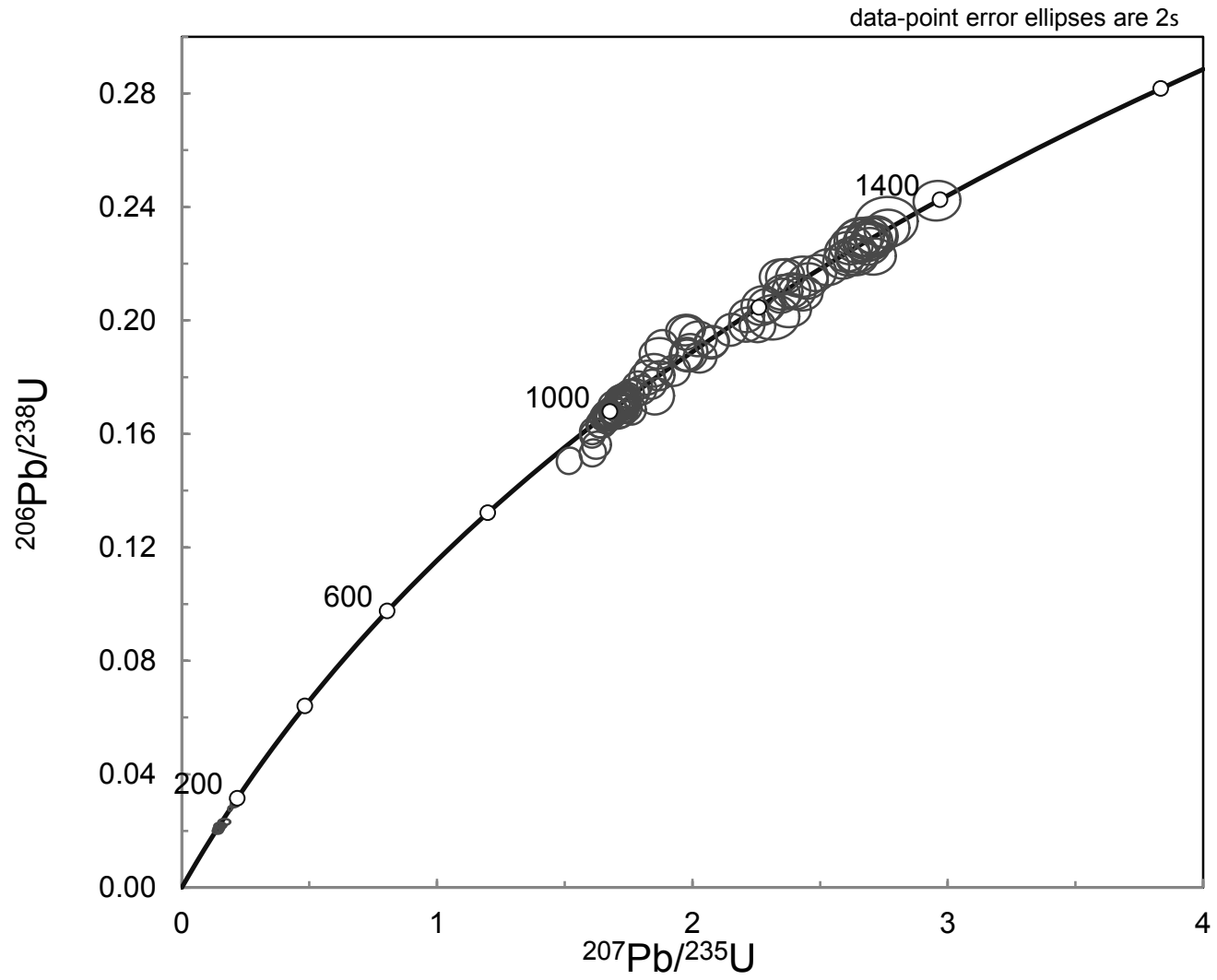




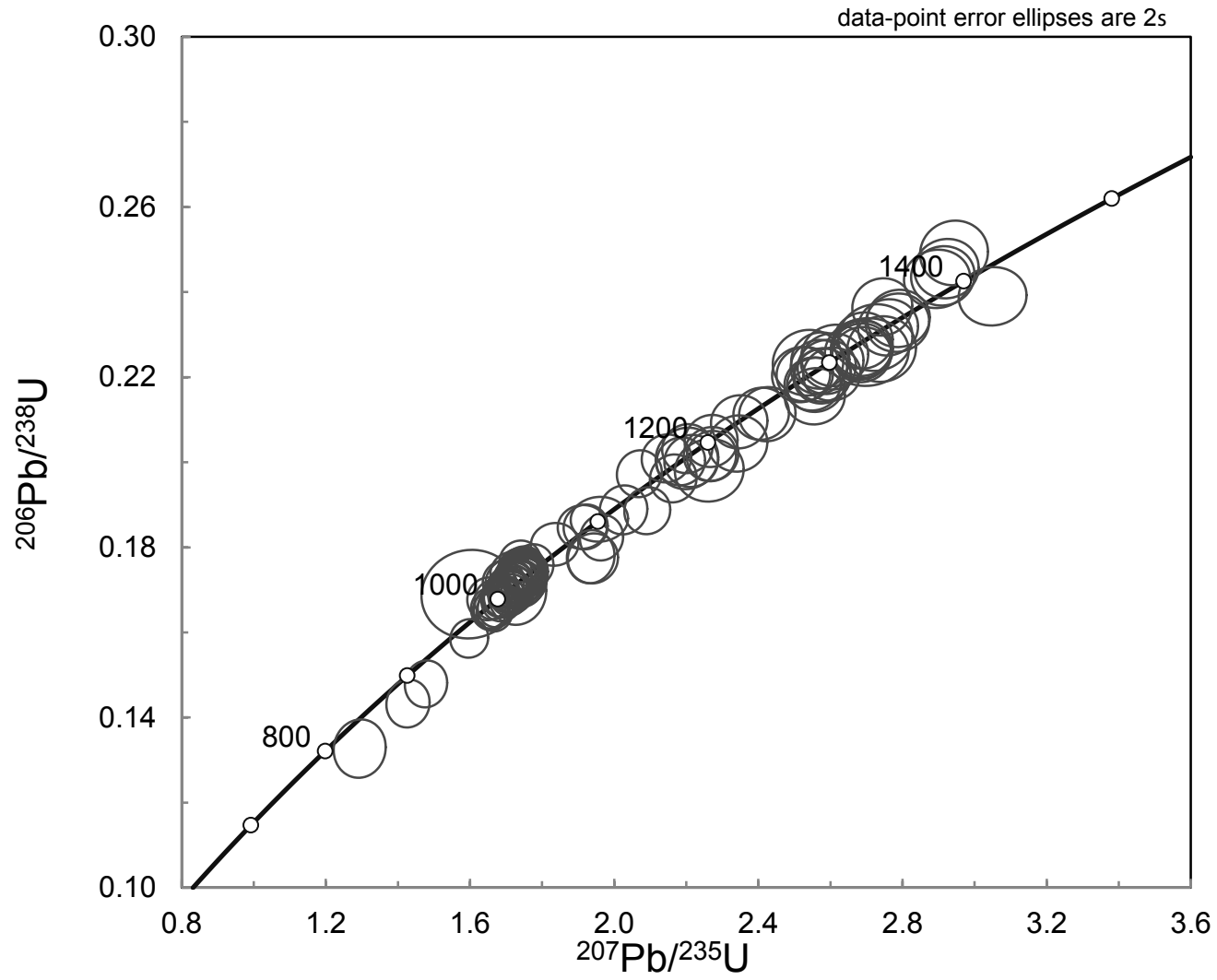
AP-045



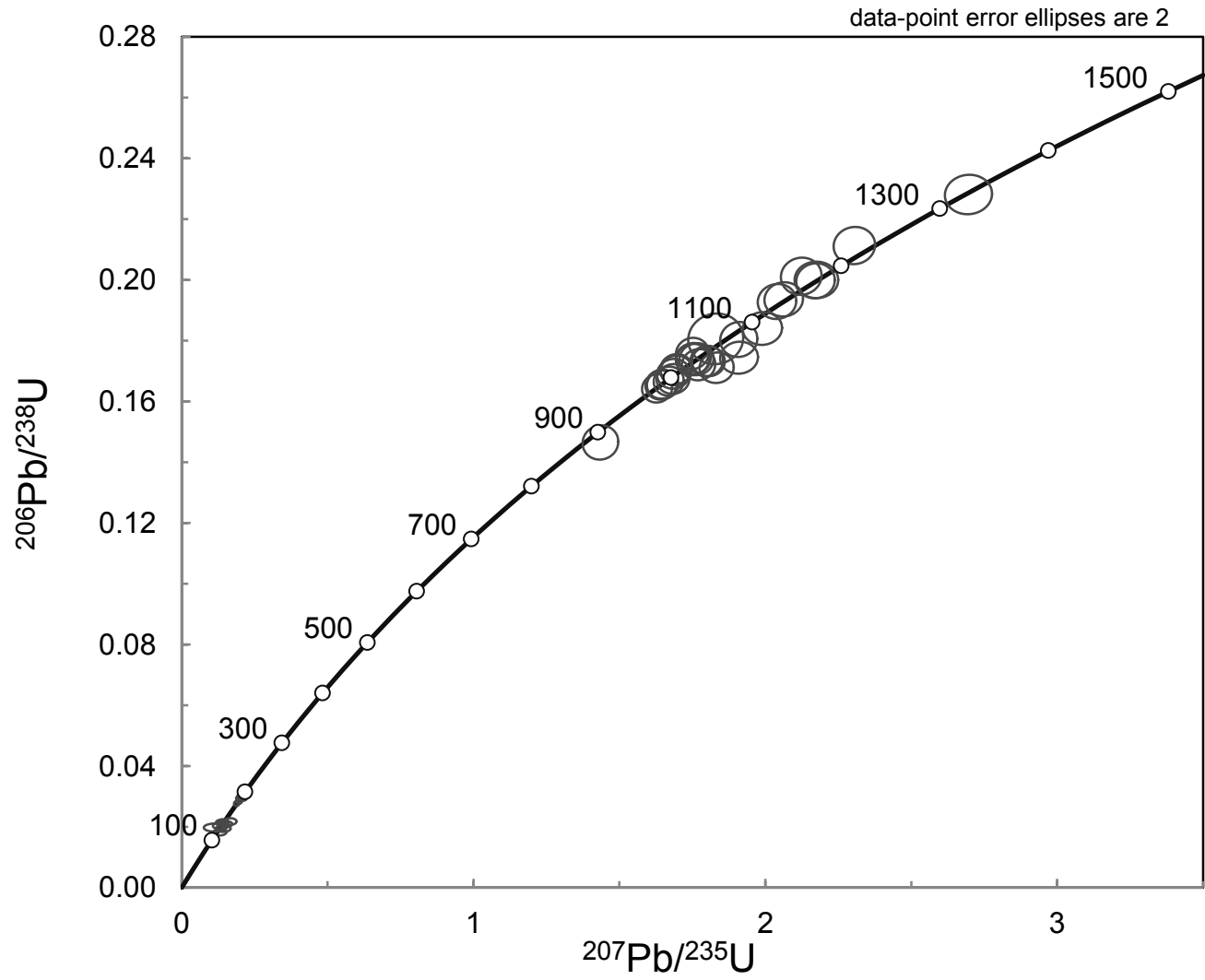
CP-088



CVI13-02

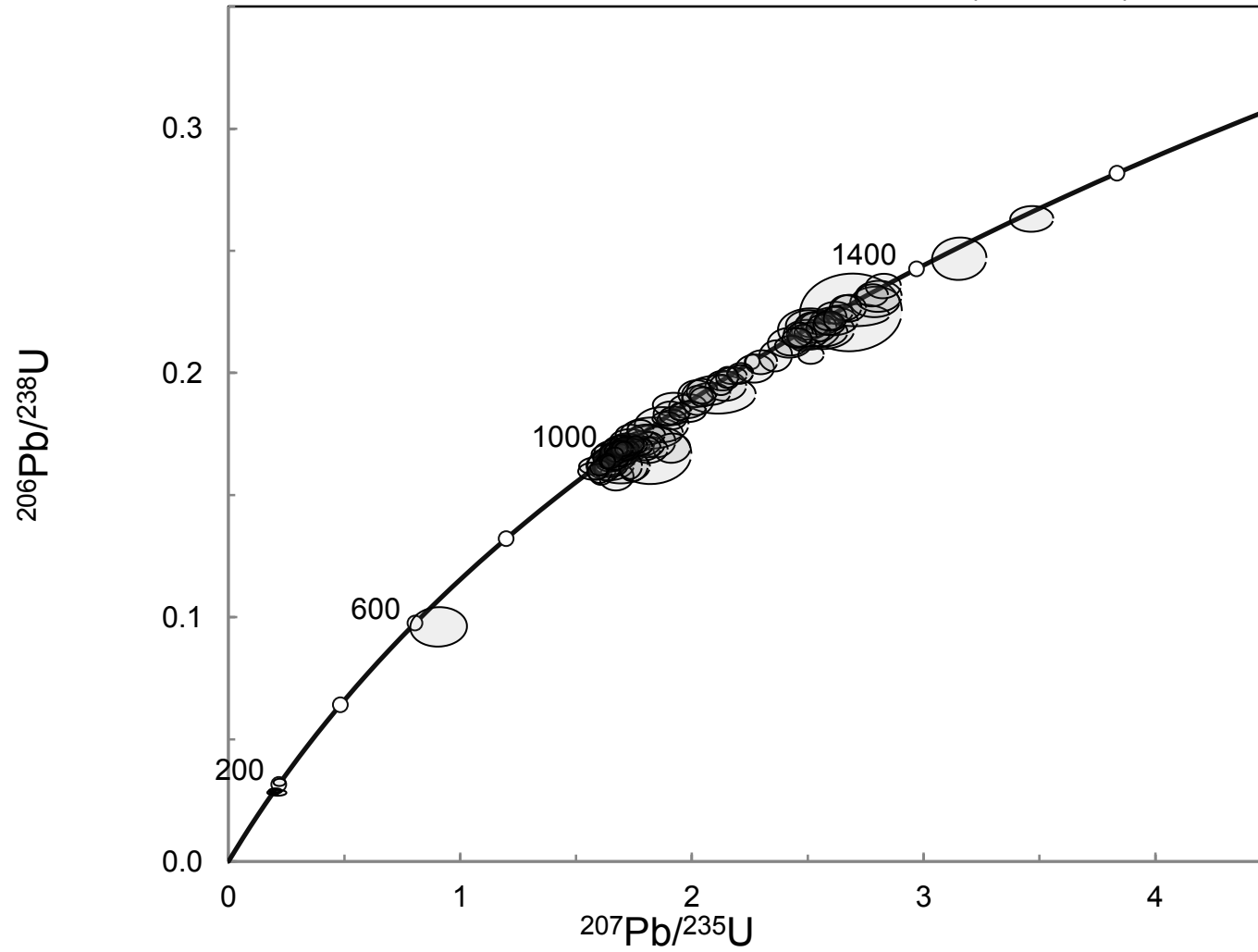


CVI13-115



EMP-16

data-point error ellipses are 2s



EMP-49A

data-point error ellipses are 2s

